

# **Assimilation of Infrared Sounder Radiances**

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# Overview

Introduction

The Infrared Spectrum

Infrared Hyperspectral Sounders and their Performance

Channel Selection

O-B Comparisons

Clouds

Humidity Assimilation

Overview of IR Sounder Performance in NWP

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Conclusions

# Introduction

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# Scope of Talk

- Infrared satellite sounders have long been part of the satellite observing system: The first HIRS was launched on NIMBUS-6 in 1975.
- However, the launch of EOS-Aqua AIRS in 2002, marked a paradigm shift in the use of these data resulting in the infrared sounders (AIRS, IASI, CrIS) being among the most important sensors in global data assimilation systems
- For this reason, this talk will focus exclusively on the use of hyperspectral sounders.

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# Scope of Talk

- Important issues in the assimilation of infrared radiances are discussed in other talks in this seminar and so will not be discussed here in detail:
- Observation errors (including spectrally correlated errors):  
Niels Bormann
- Instrument characterisation: Dieter Klaes and Bill Bell
- Cloudy radiance assimilation:  
Jean-Francois Mahfouf, Alan Geer, Jérôme Vidot
- Convective scale DA: Thomas Auligné
- Principal components and reconstructed radiances:  
Marco Matricardi
- Surface Emissivity: Fatima Karbou

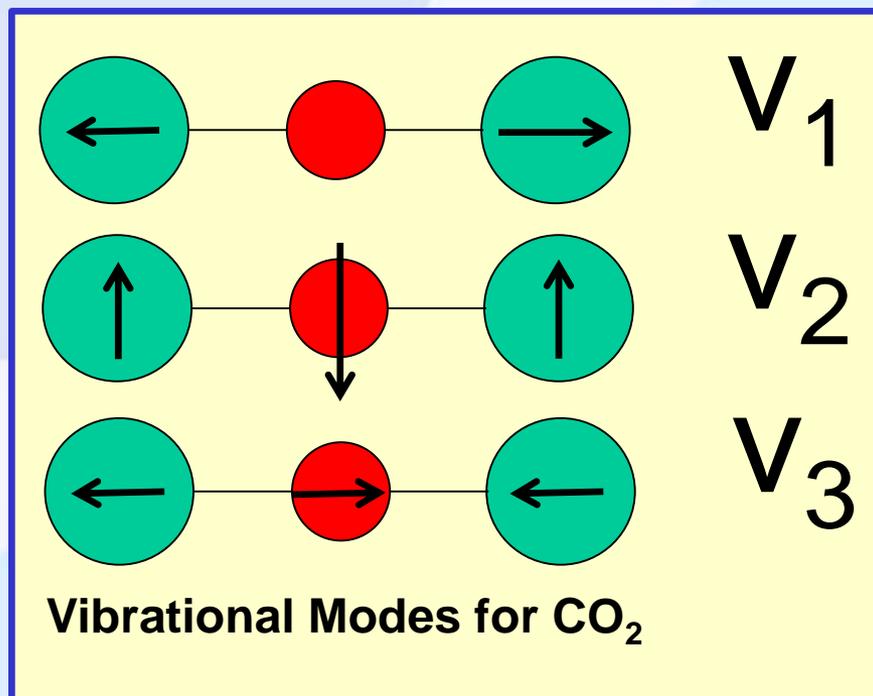
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# The Infrared Spectrum

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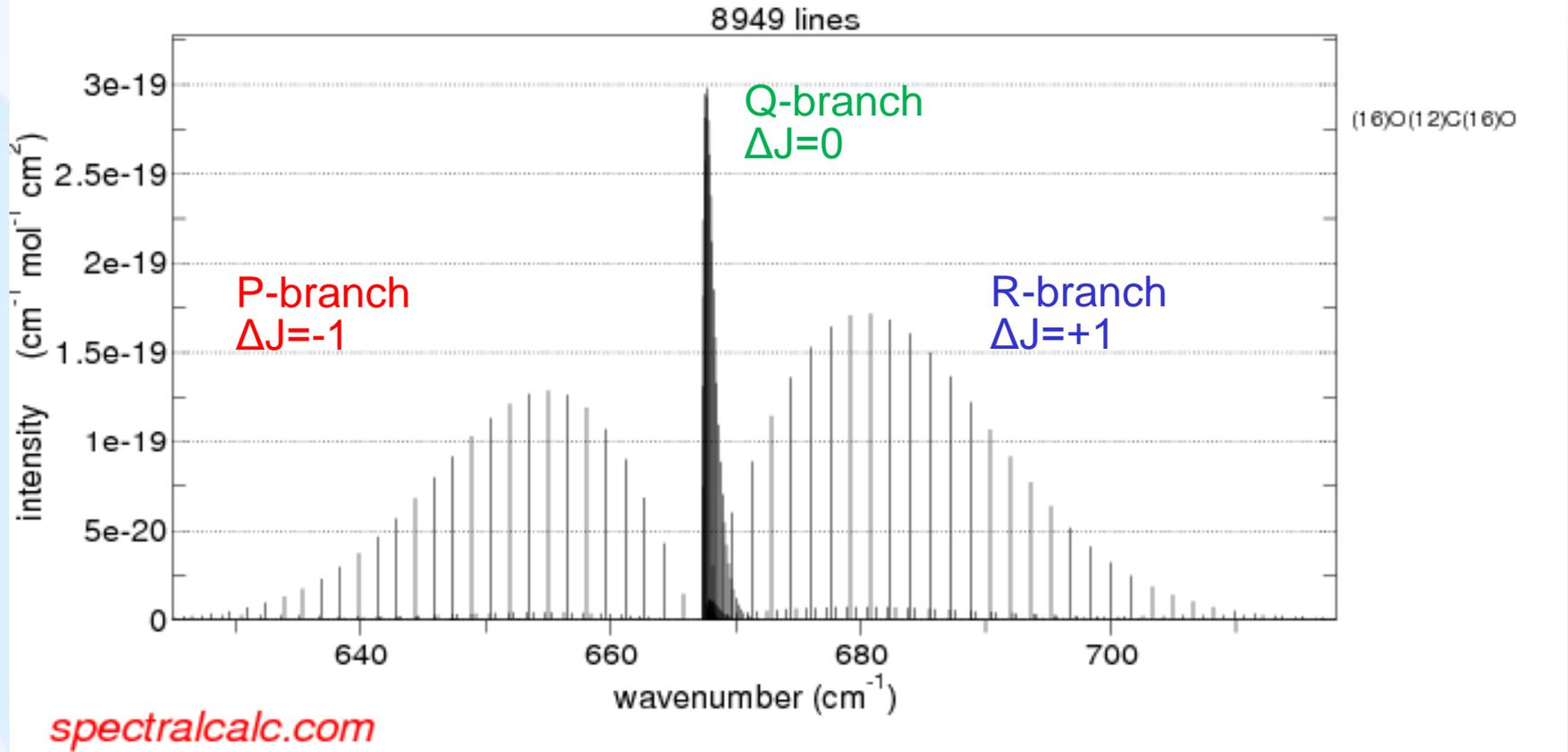
# The clear sky infrared spectrum is dominated by emission and absorption by atmospheric molecules

- Molecules in the atmosphere have energy stored as rotational, vibrational and electronic components
- The energy states are quantised and may be transformed through emission or absorption of electromagnetic radiation. This results in discrete spectral emission/absorption features in the spectrum.



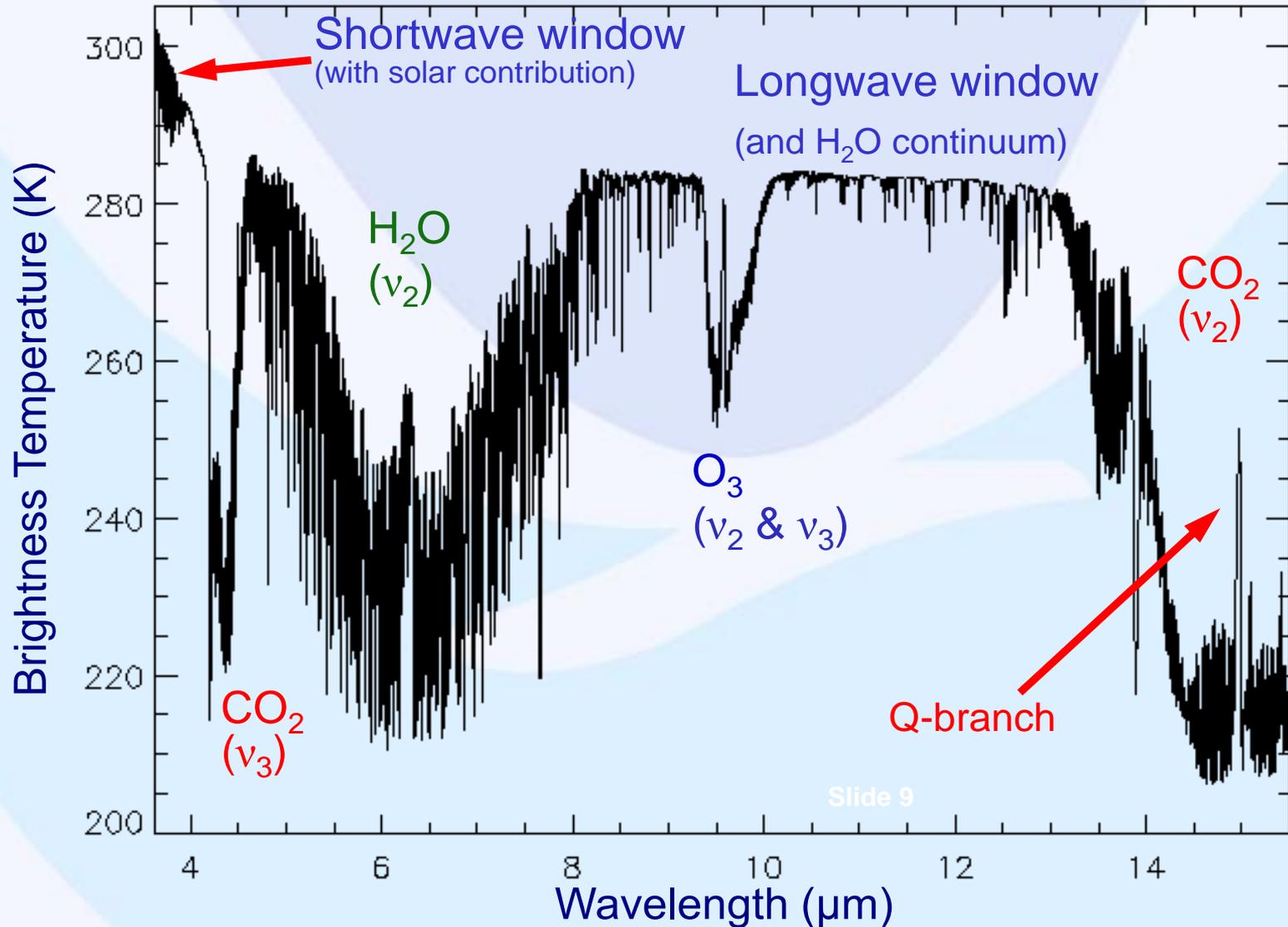
- In the microwave these are due to rotational transitions
- In the infrared these are rotational and vibrational transitions
- Electronic transitions manifest themselves in the visible and ultraviolet

# Vibration-Rotation Spectrum Ground→v2 transition for CO<sub>2</sub>



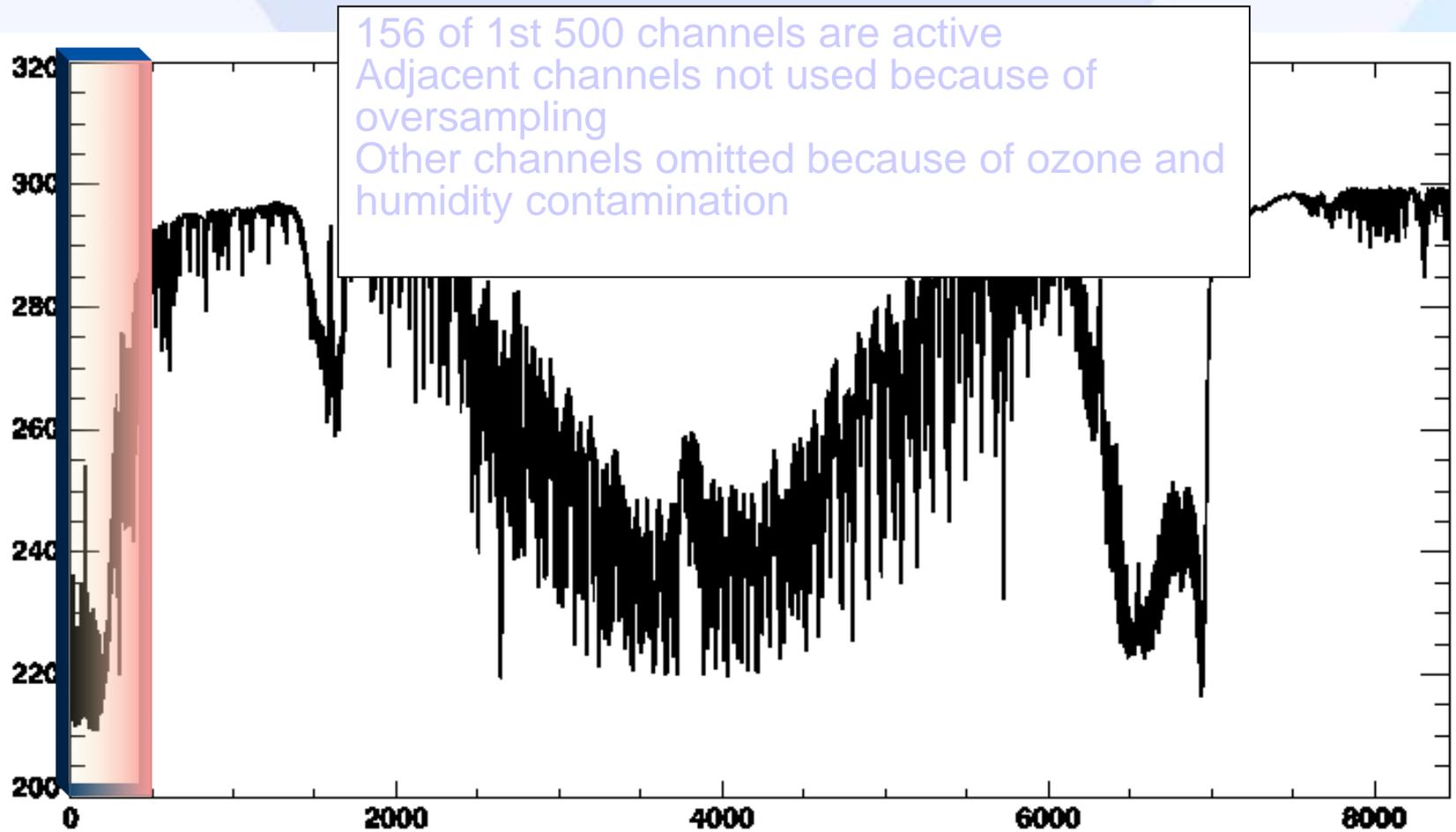
An example of a vibration-rotation band in the infrared CO<sub>2</sub> spectrum. Due to considerations of angular momentum, only changes in the rotational quantum number,  $J$ , of -1, 0 or 1 are optically active, producing the characteristic three branch structure to the band (some linear molecules have the Q-branch missing).

# An Infrared Spectrum



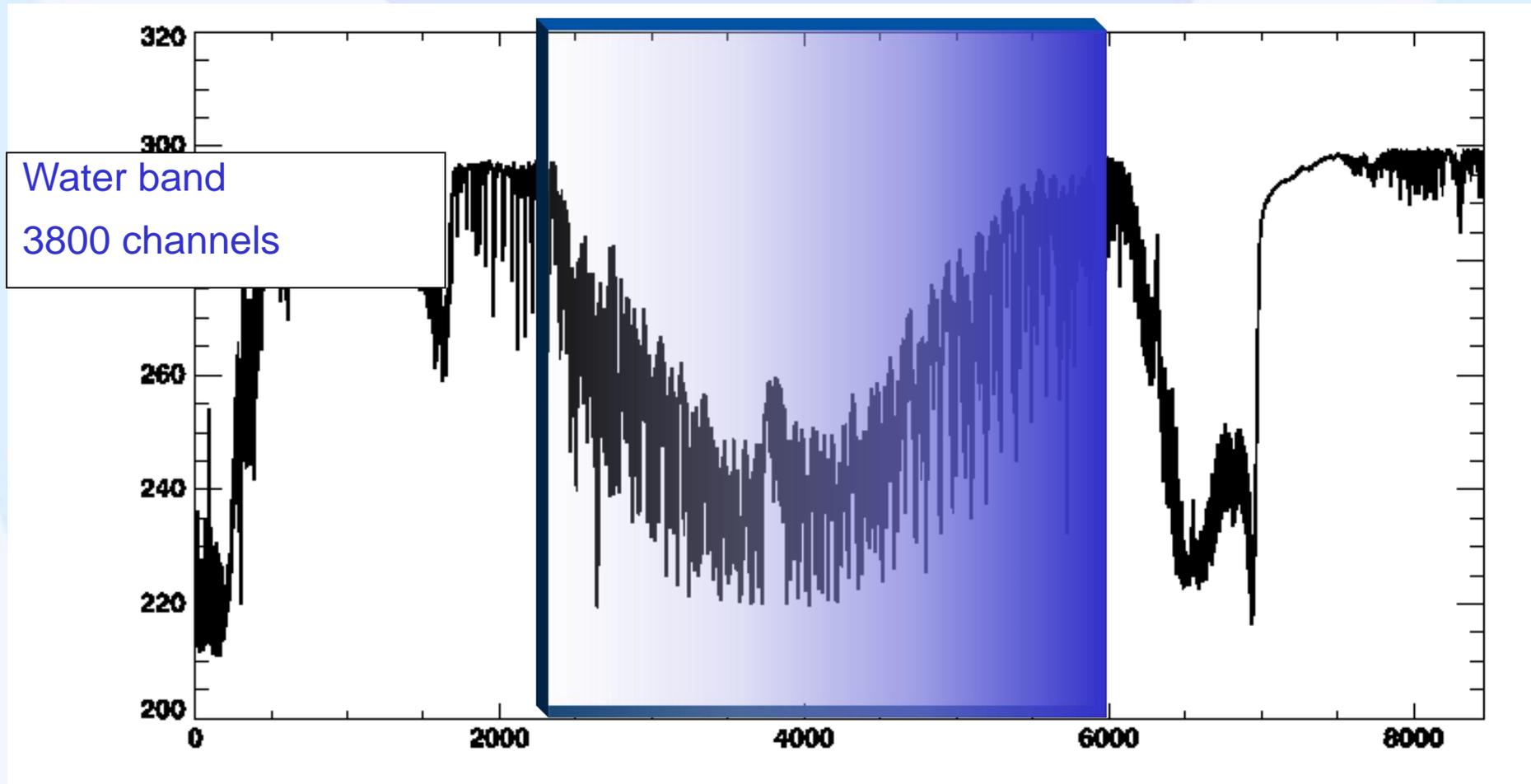
# Regions of the Infrared Spectrum

## Longwave CO<sub>2</sub> Band



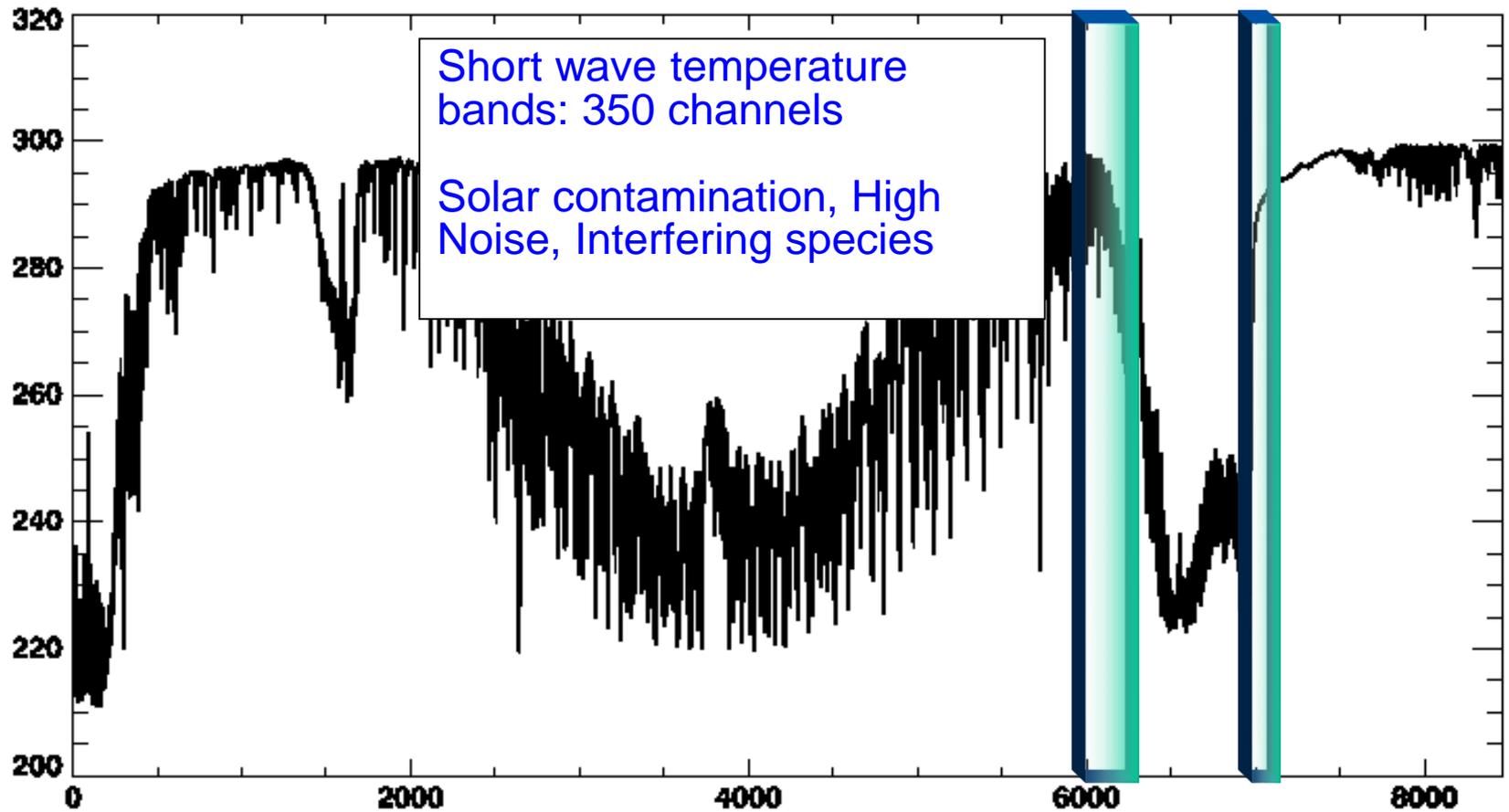
# Regions of the Infrared Spectrum

## The 6.3 $\mu\text{m}$ Water Band



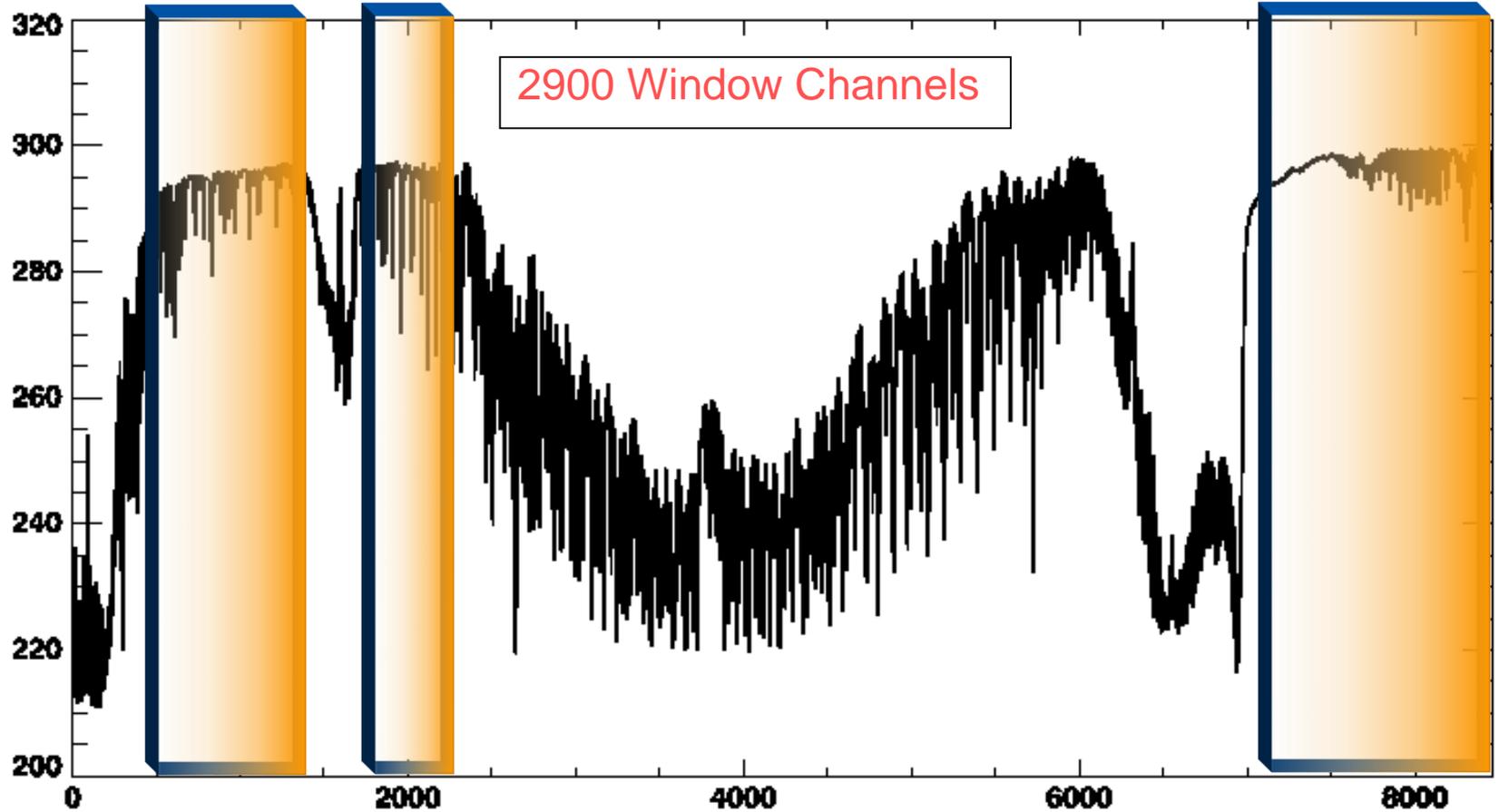
# Regions of the Infrared Spectrum

## Shortwave CO<sub>2</sub> Band



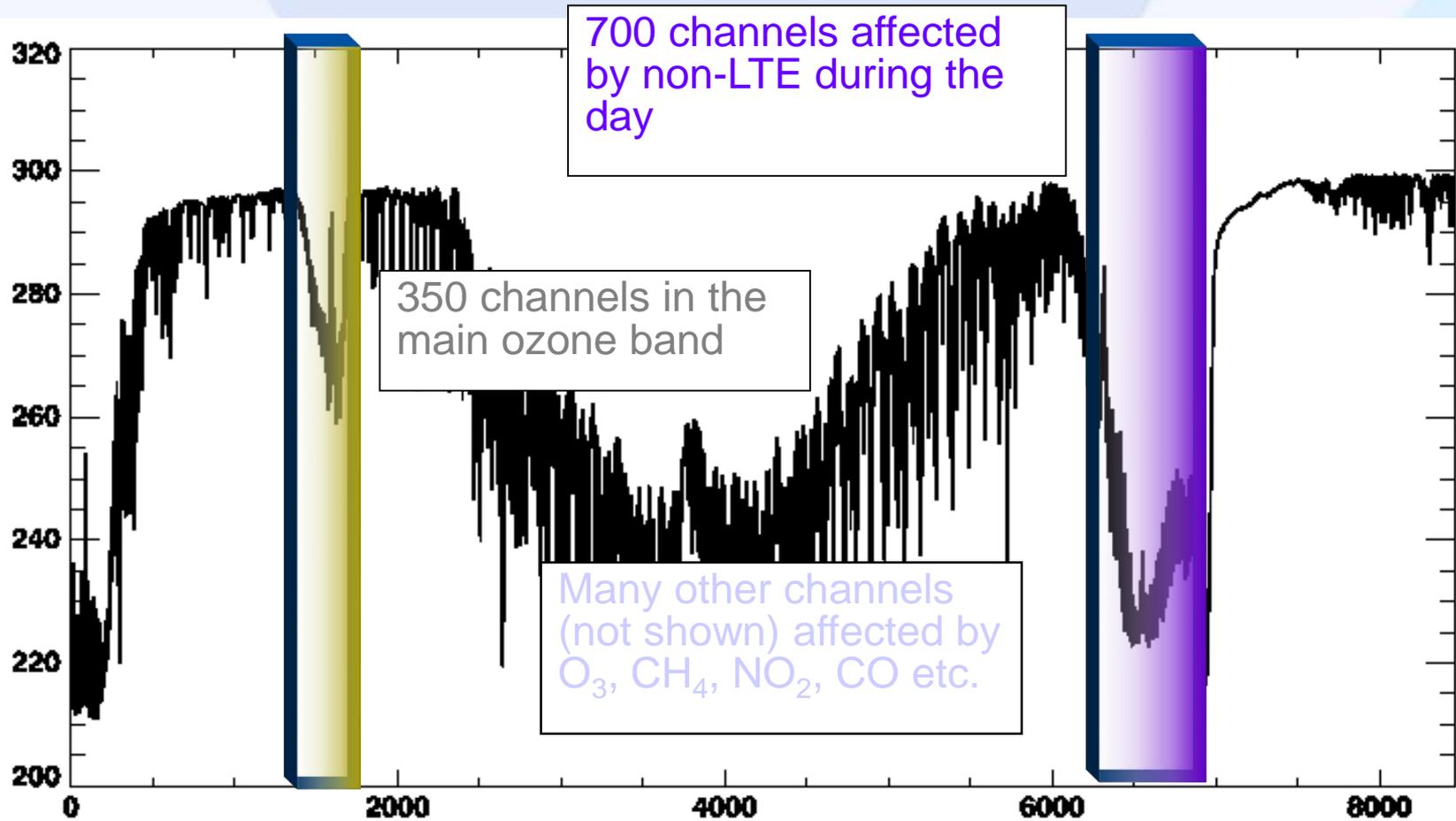
# Regions of the Infrared Spectrum

## Channels Primarily Sensitive to the Surface



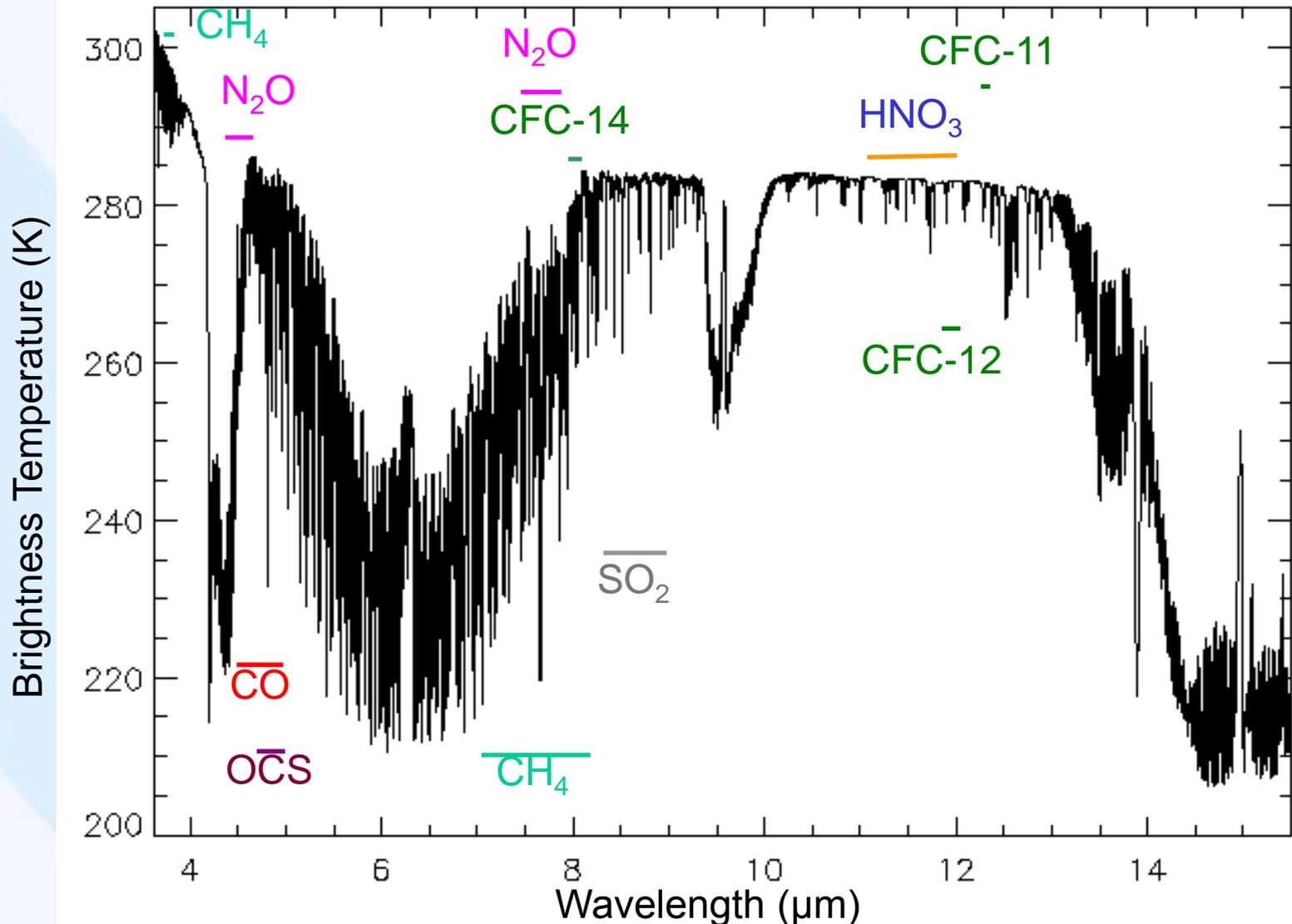
# Regions of the Infrared Spectrum

## Trace Gases and RT Challenges



# Regions of the Infrared Spectrum

## Trace Gases



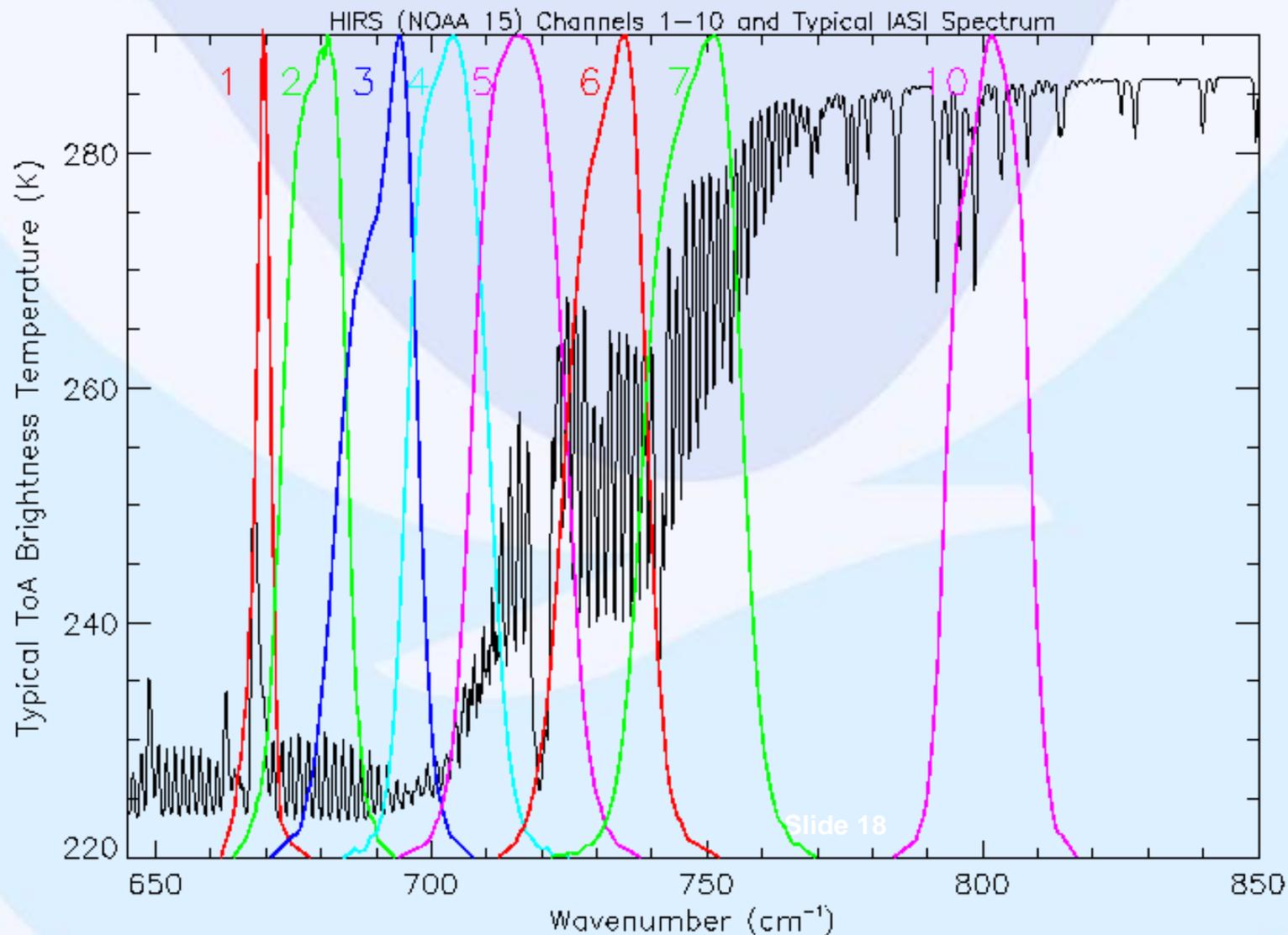
# **Infrared Hyperspectral Sounders and their Performance**

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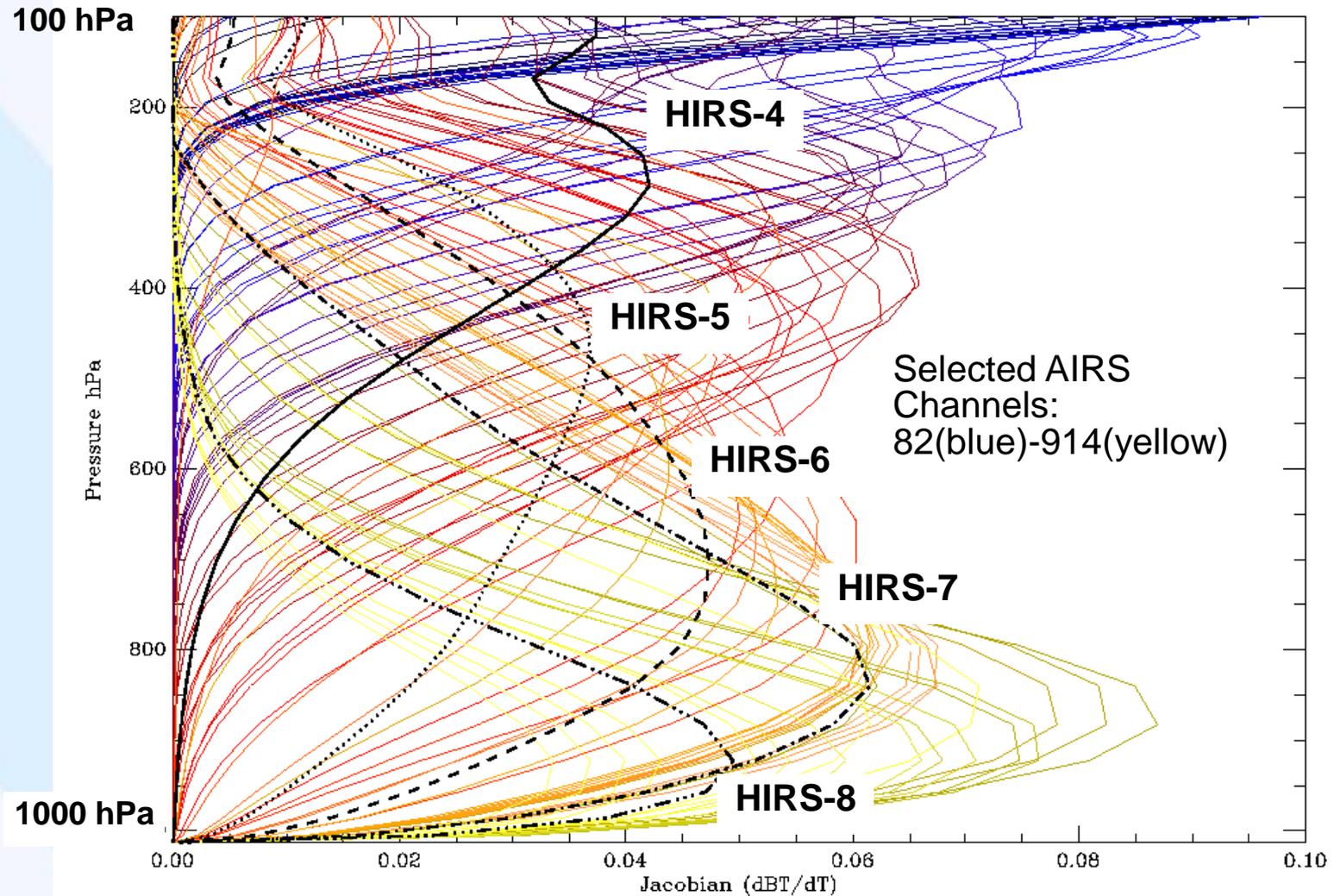
# Current and Future High-Spectral Resolution InfraRed Sounders

<b>Instrument/ Satellite/ Launch</b>	<b>No. of Channels</b>	<b>Spectral Range</b>	<b>Spectral Resolution</b>	<b>IFOV</b>	<b>Type/ Orbit</b>
<b>AIRS/ Aqua(EOS-PM)/ May 2002</b>	<b>2378</b>	<b>650-2760cm<sup>-1</sup></b>	<b>~1cm<sup>-1</sup></b>	<b>13.5km</b>	<b>Grating Spectrometer/ Polar</b>
<b>IASI/ MetOp/ October 2006</b>	<b>8461</b>	<b>645-2760cm<sup>-1</sup></b>	<b>0.5cm<sup>-1</sup></b>	<b>12km</b>	<b>Interferometer/ Polar</b>
<b>CrIS/ NPP &amp; JPSS/ October 2011</b>	<b>1400</b>	<b>635-2450cm<sup>-1</sup></b>	<b>1.125- 4.5cm<sup>-1</sup></b>	<b>12km</b>	<b>Interferometer/ Polar</b>
<b>MTG-IRS 2018</b>	<b>1720</b>	<b>700-2175cm<sup>-1</sup></b>	<b>0.5- 0.625cm<sup>-1</sup></b>	<b>4km</b>	<b>Interferometer/ Geostationary</b>
<b>IASI-NG 2021</b>	<b>16920</b>	<b>645-2760cm<sup>-1</sup></b>	<b>0.25 cm<sup>-1</sup></b>	<b>12km</b>	<b>Interferometer/ Polar</b>

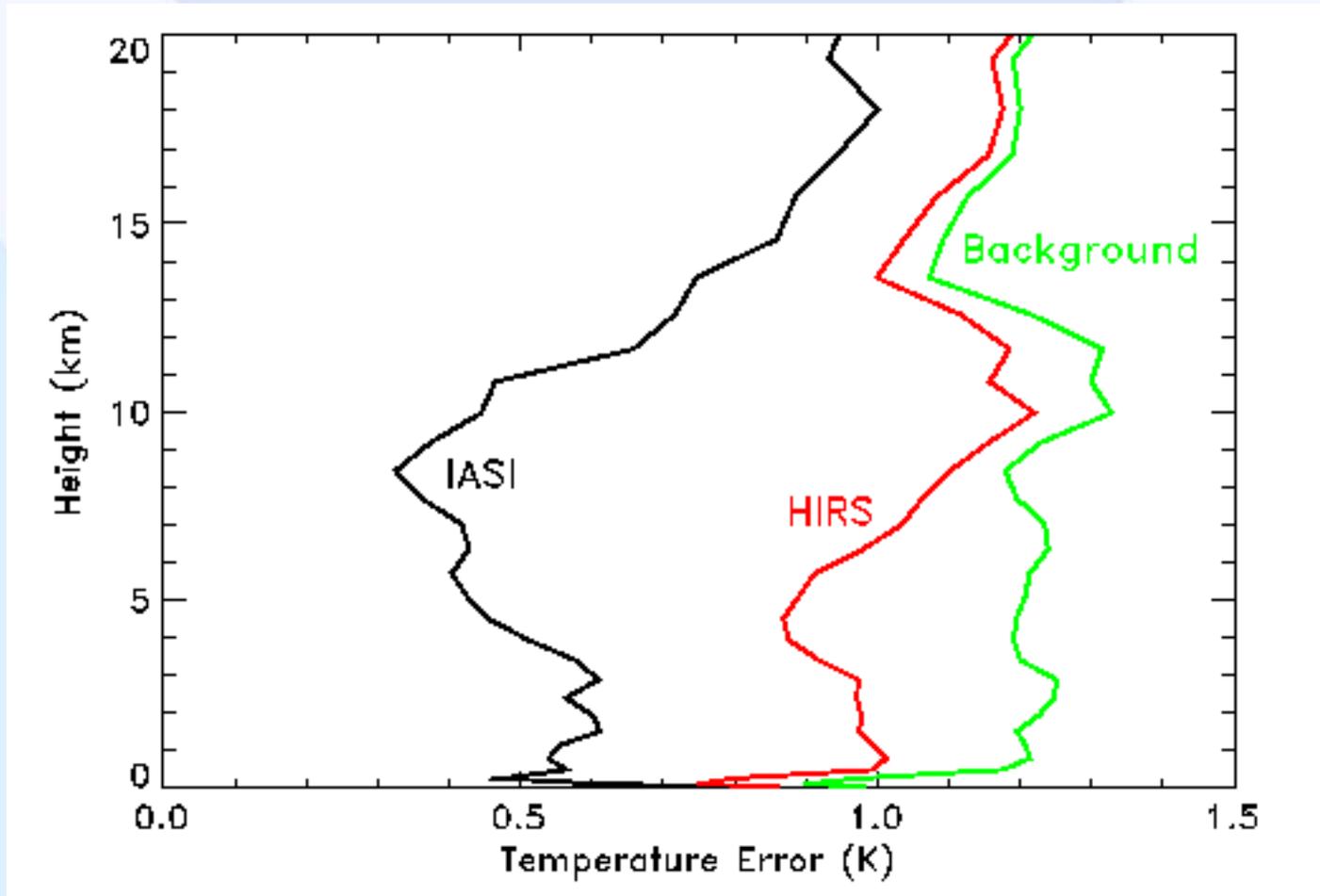
# IASI vs HIRS: The Thermal InfraRed



# AIRS vs HIRS Jacobians in the $15\mu\text{m}$ $\text{CO}_2$ band

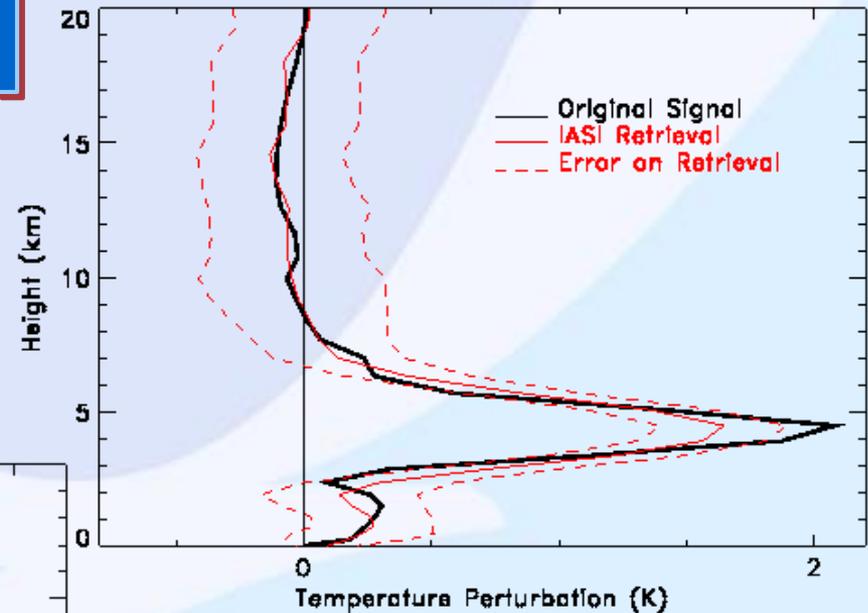


# HIRS vs IASI: Temperature Retrieval Accuracy

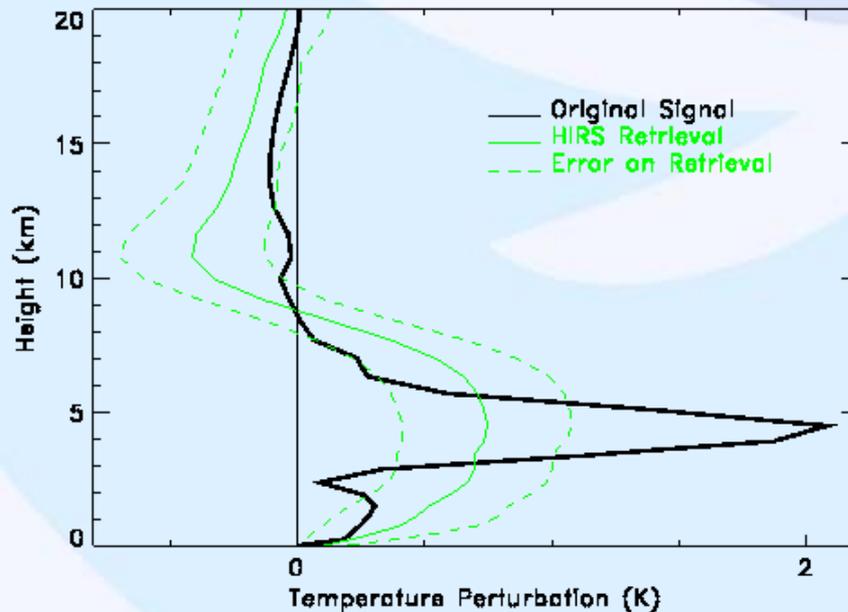


# HIRS vs IASI: Response to Important Atmospheric Structure

IASI



HIRS



**Response to a structure the observation of which would have improved the forecast of the reintensification of Hurricane Floyd over SW France and SW England on 12<sup>th</sup> September 1993. (Rabier et al., 1996)**

# Channel Selection

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# Why Select Channels?

The volume of hyperspectral IR data available is such that we do not have the computational resources to simulate and assimilate all these data in an operational timeframe

Not all channels are of equal use when assimilated into an NWP system

We choose channels that we wish to monitor (often with a view to future use)

We choose a subset of these channels which we actively assimilate

The following describes the channel selection performed for IASI (Collard, 2007, based on Rodgers, 2000)

# IASI Channel Selection

## Pre-screen channels

Ignore channels with large contribution from un-assimilated trace gases.

## Use the channel selection method of Rodgers (1996)

Iterative method which adds each channel to the selection based on its ability to improve a chosen figure of merit (in this case degrees of freedom for signal).

Determine the channels which contribute most information to a number of atmospheric states and view angles.

Use multiple runs to reduce the effect of non-linearity and to focus on particular species.

Impose additional selection criterion: No channel adjacent to one already chosen may be selected to reduce the impact of inter-channel correlated errors due to apodisation

## Add extra channels that the Rogers method cannot choose

E.g. Cloud detection channels.

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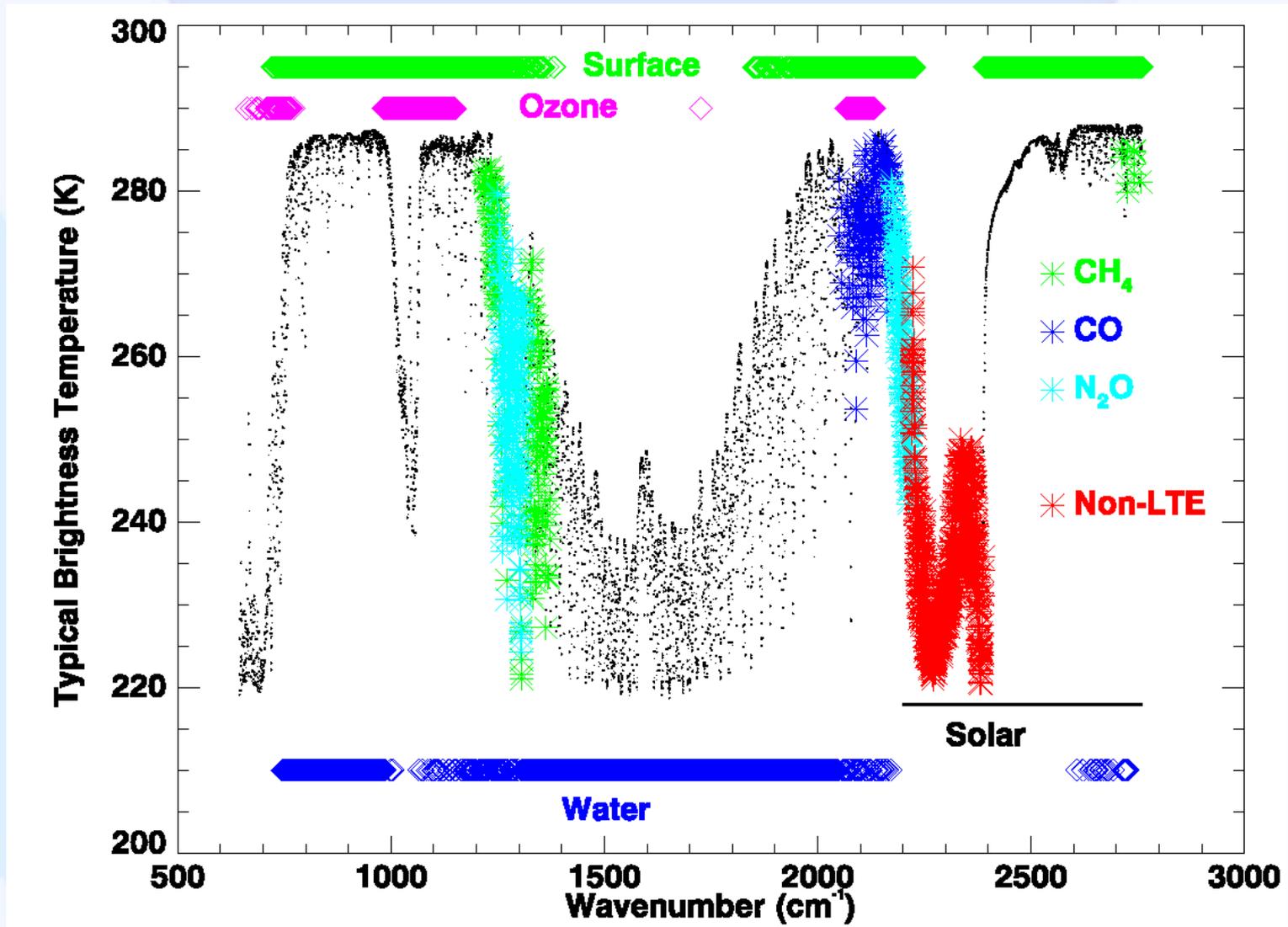
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# Pre-screened channels



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# Figure of merit for channel selection

- Rogers (2000) suggests two possible figures of merit for channel selection:
- The degrees of freedom for signal (DFS) for the retrieval is given by:

$$\text{DFS} = \text{Tr}(\mathbf{I} - \mathbf{A}\mathbf{B}^{-1})$$

- and, the entropy reduction

$$\text{ER} = -\frac{1}{2} \text{Ln}|\mathbf{A}\mathbf{B}^{-1}|$$

- Past experience has shown very similar results from these and in this study the former is used.

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# Efficient calculation of A-matrix

- The Rogers (2000) channel selection technique requires repeated calculation of the A matrix every time a new channel is being tested. Rogers notes that for a diagonal observation error covariance, the change in A matrix on adding a new channel  $i$  is calculated efficiently thus:

$$\mathbf{A}_{i-1} = \mathbf{A}_{i-1} \left( 1 + \frac{\mathbf{h}_i (\mathbf{A}_{i-1} \mathbf{h}_i)^T}{1 + (\mathbf{A}_{i-1} \mathbf{h}_i)^T \mathbf{h}_i} \right)$$

- where  $\mathbf{h}_i$  is the Jacobian of the  $i$ th channel in observation noise-normalised units.

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# IASI Channel Selection

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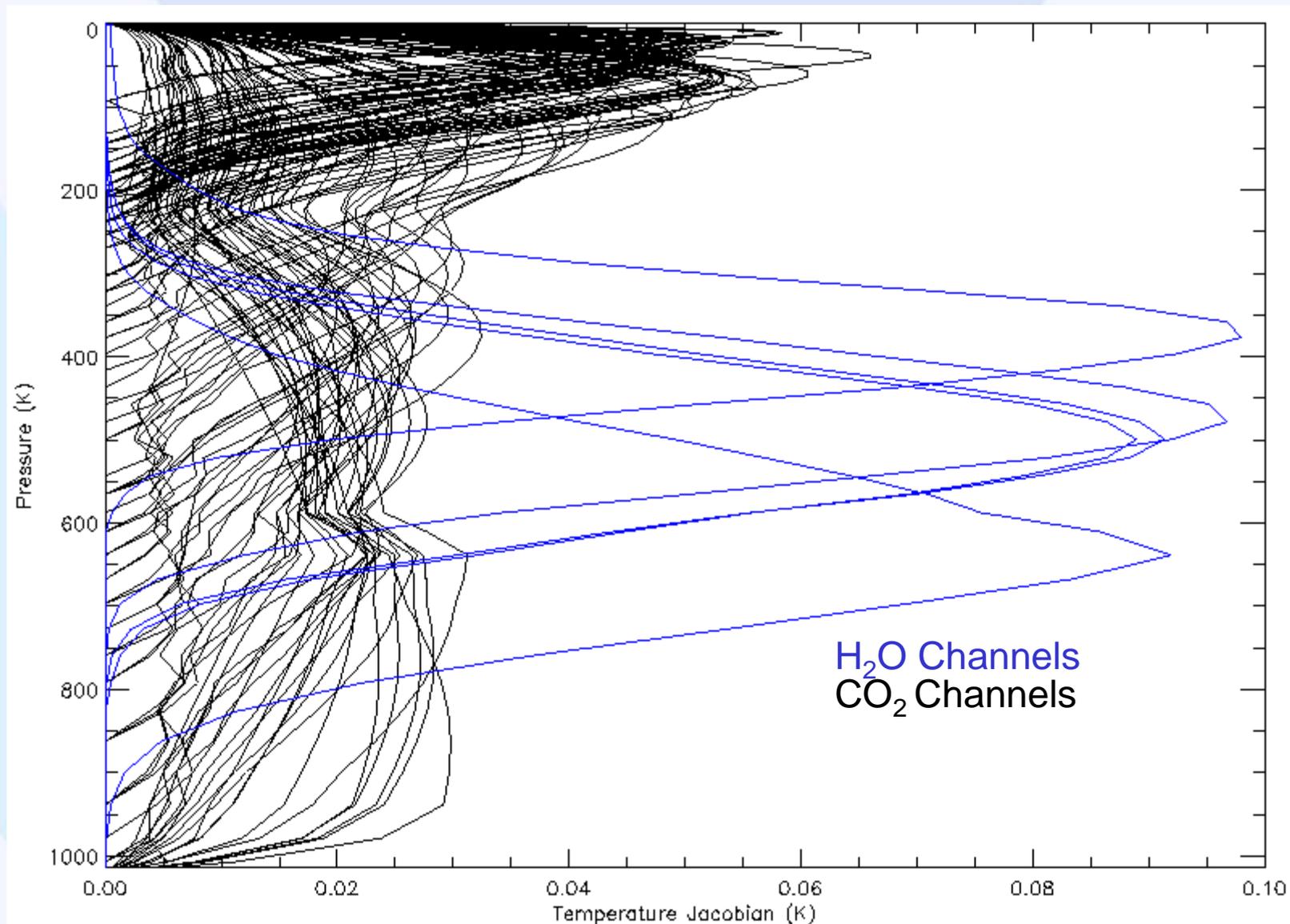
**Use multiple runs to reduce the effect of non-linearity and to focus on particular species.**

Impose additional selection criterion: No channel adjacent to one already chosen may be selected to reduce the impact of inter-channel correlated errors due to apodisation

## Add extra channels that the Rogers method cannot choose

E.g. Cloud detection channels.

# Temperature Jacobians of Used Channels



# IASI Channel Selection

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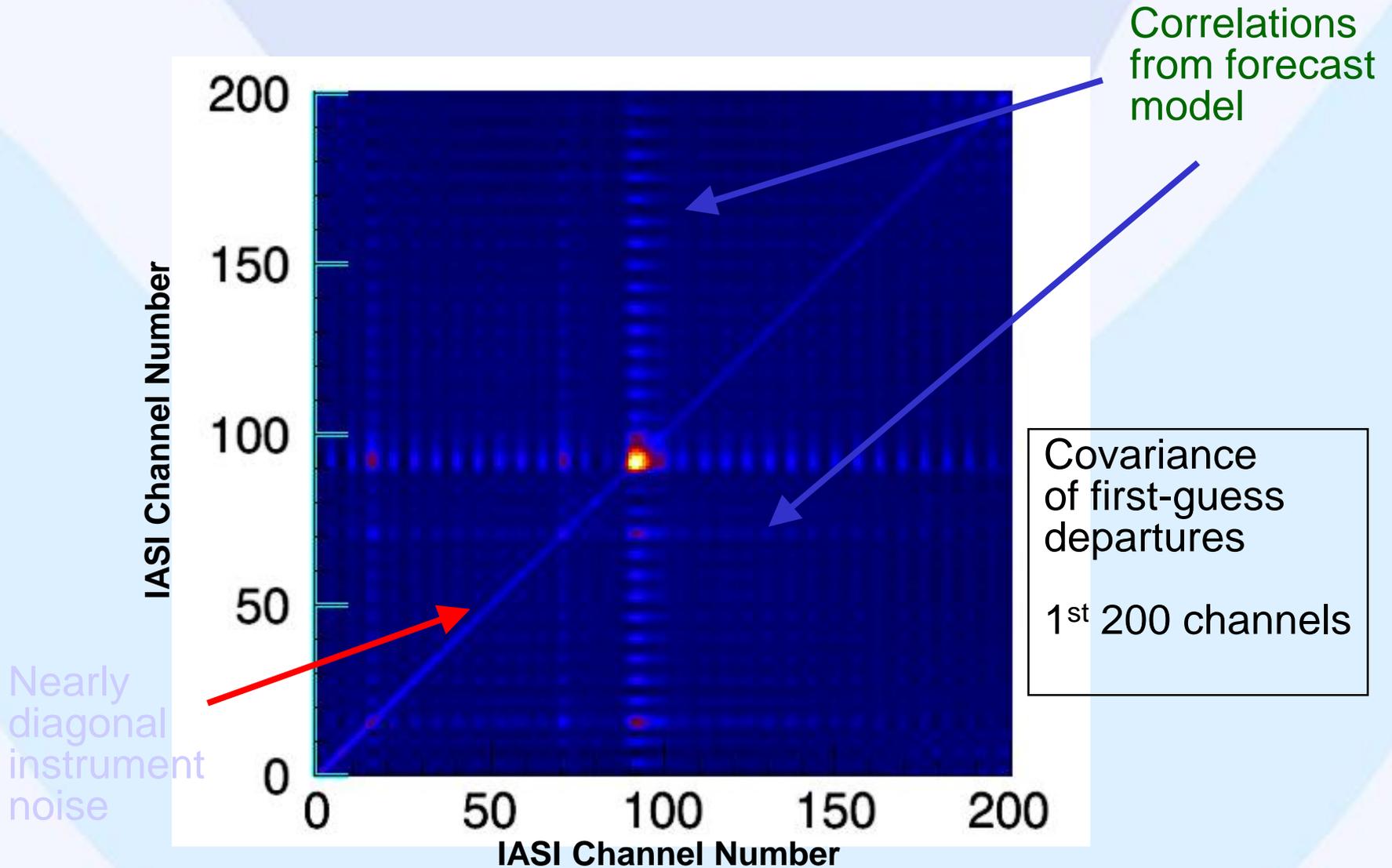
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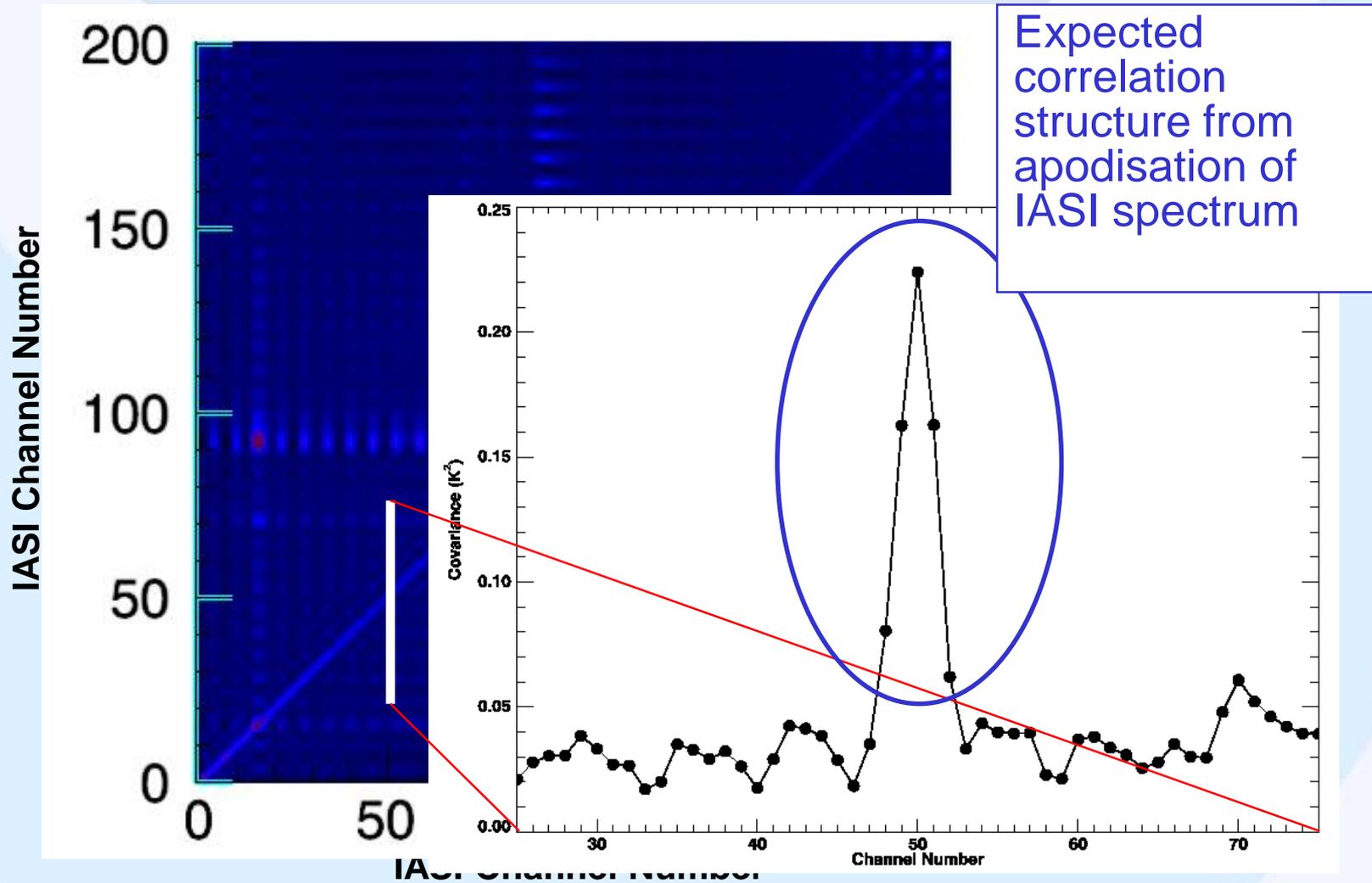
## Add extra channels that the Rogers method cannot choose

E.g. Cloud detection channels.

# IASI Spectral Correlation



# IASI Spectral Correlation



# Selected Channels (1)

30 channels chosen from  $15\mu\text{m CO}_2$  band considering temperature assimilation only

36 channels from  $707\text{-}760\text{cm}^{-1}$  region – found to be particularly important when assimilating AIRS.

252 channels considering temperature and water vapour together

15 ozone channels

13 Channels in the solar-affected shortwave region

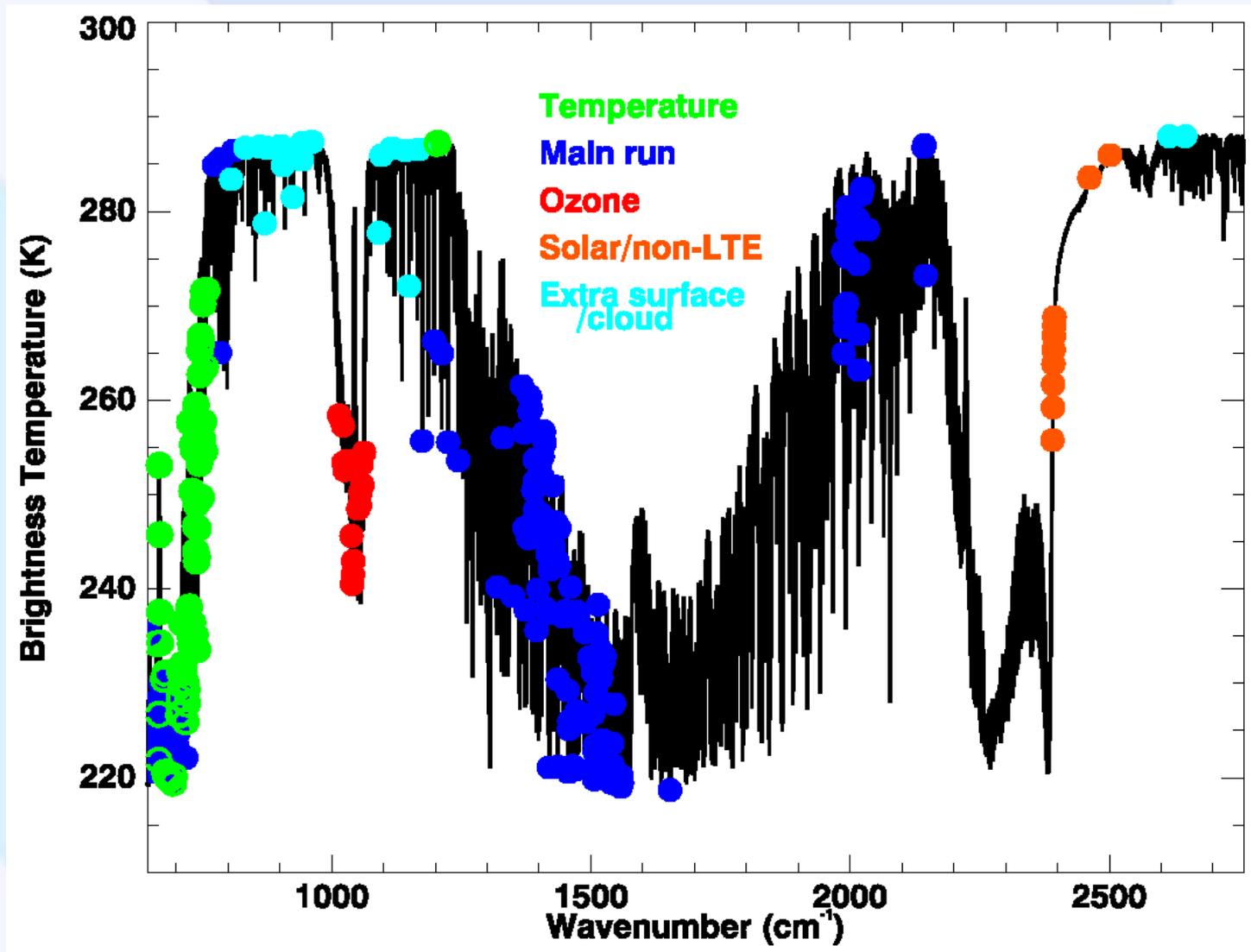
In ECMWF selection only:

22 channels used for monitoring (HIRS analogues and requested by CNES)

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Another 44 channels in the  $707\text{-}760\text{cm}^{-1}$  region

## Selected Channels (2)



# Improved treatment of correlated error

- In their 2013 paper, Ventress and Dudhia, are able to extend the calculation of the A matrix – and thus the degrees of freedom for signal – to the more realistic situation where the assumed observation error covariance is diagonal but the real error covariance is correlated.
- This allows one, for example, to explicitly allow for the correlated observation error term arising from uncertain molecular abundances and for channels to be chosen that are more robust against these sources of error.

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# Extension to cloudy scenes

- Martinet et al. (2014) has extended the selection of 366 IASI channels that are currently used to 500, with the other channels being chosen based on their use in obtaining cloud properties with particular reference to the AROME regional model.
  - A range of situations covering low liquid clouds, opaque ice clouds and semi-transparent ice clouds were considered with RTTOV-CLD used to generate the spectra.
- Migliorini is also considering channel selection in cloudy situations, with particular focus on correlated observation errors and a flow-dependent estimate of forecast uncertainty.

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# Extension to reconstructed radiances.

- As part of her thesis work, Fiona Hilton has considered channel selection in the context of reconstructed radiances.
- The transformation of instrument noise into reconstructed radiances will result in a singular R-matrix. We therefore need to consider whether the R-matrix associated with a particular channel selection is suitably well-conditioned.

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# More efficient ways of representing the spectrum

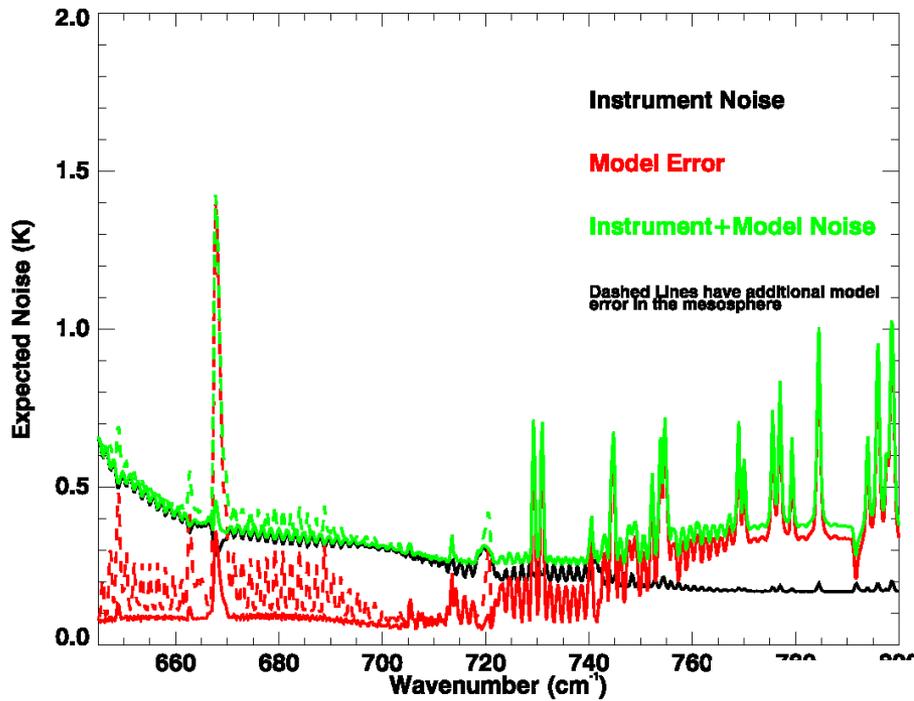
- It is often pointed out that channel selection where we select ~5-10% of the channels for monitoring and even fewer for assimilation is an inefficient way of representing the hyperspectral infrared spectrum.
- More sophisticated methods such as Principal Component Analysis and Reconstructed Radiances will be the subject of the following talk by Marco Matricardi.

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# **O-B Comparisons**

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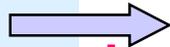
# First –Guess Departure Standard Deviations in 15 $\mu\text{m}$ CO<sub>2</sub> Band



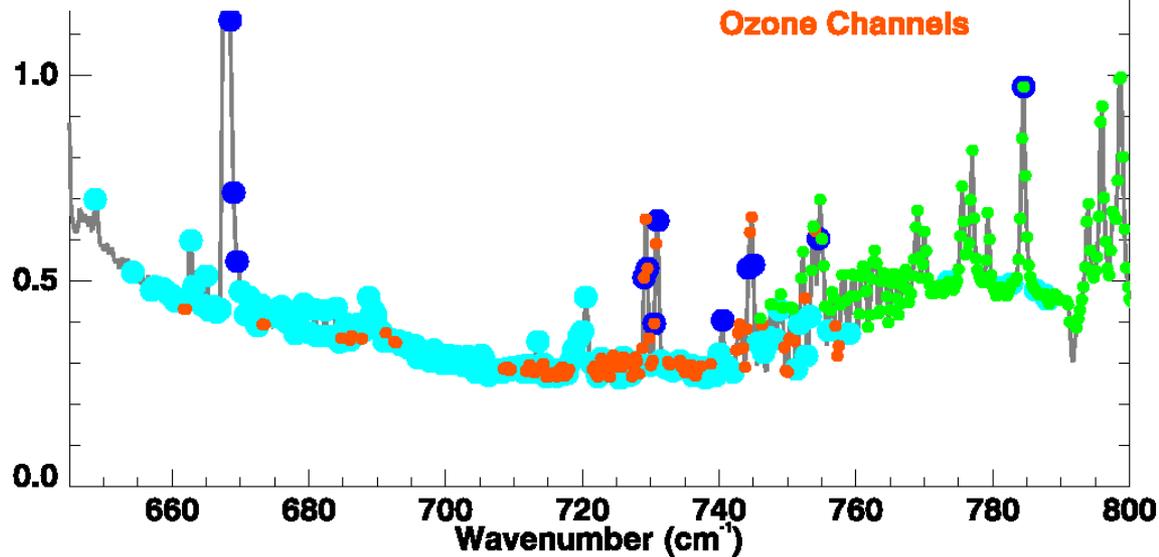
Calculated  
Std. Dev.



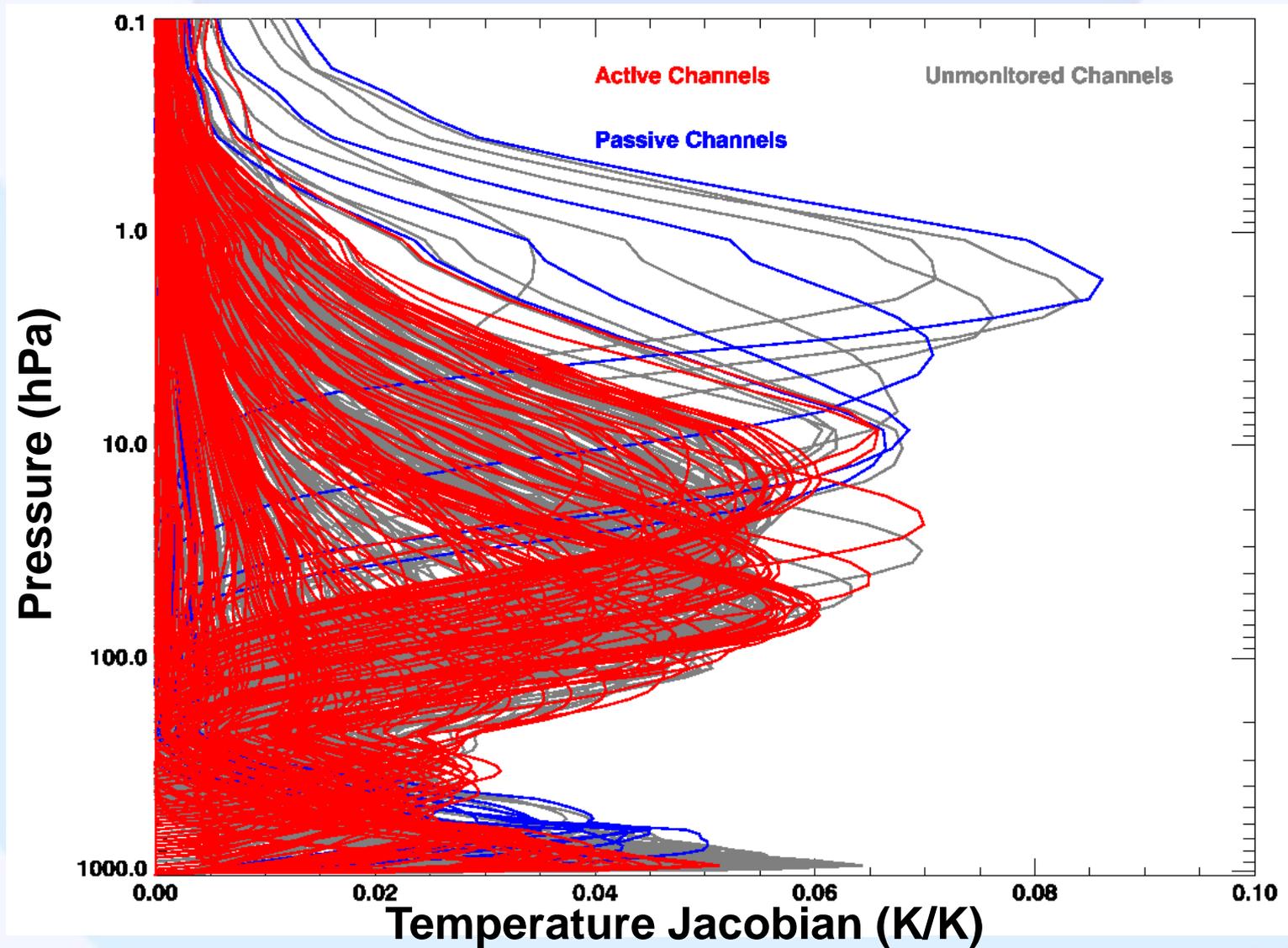
Observed  
Std. Dev.



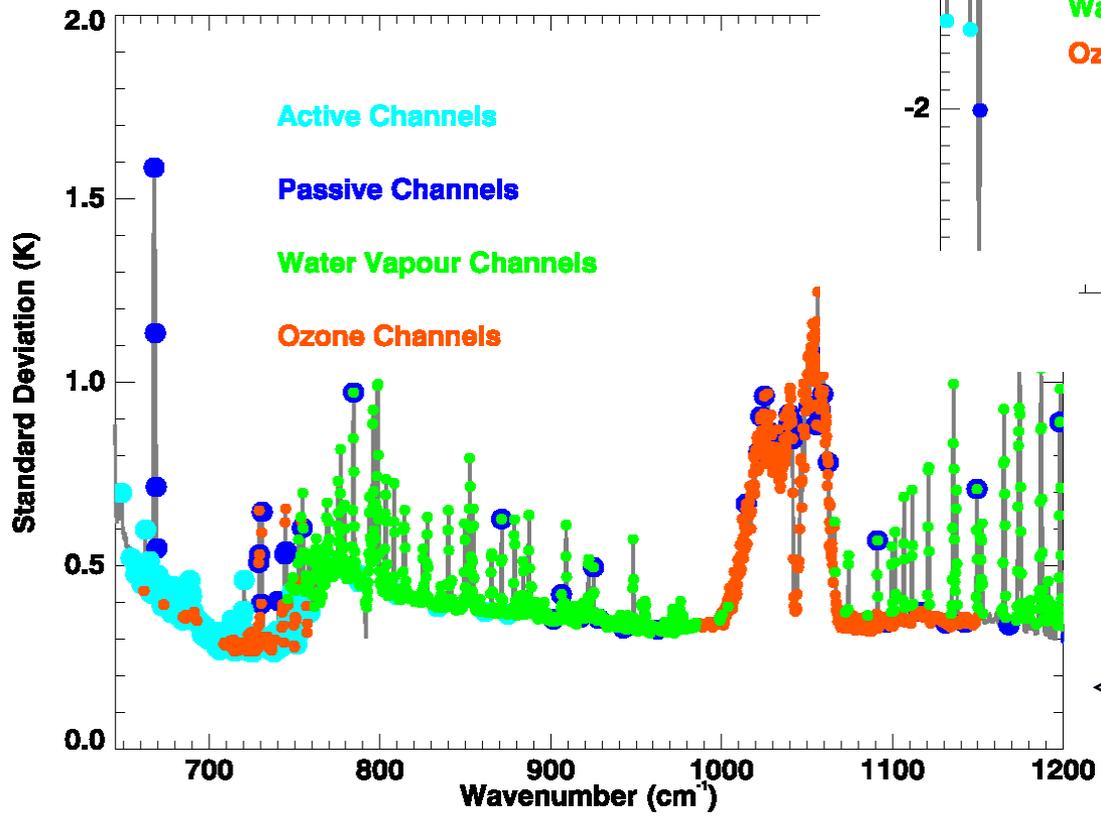
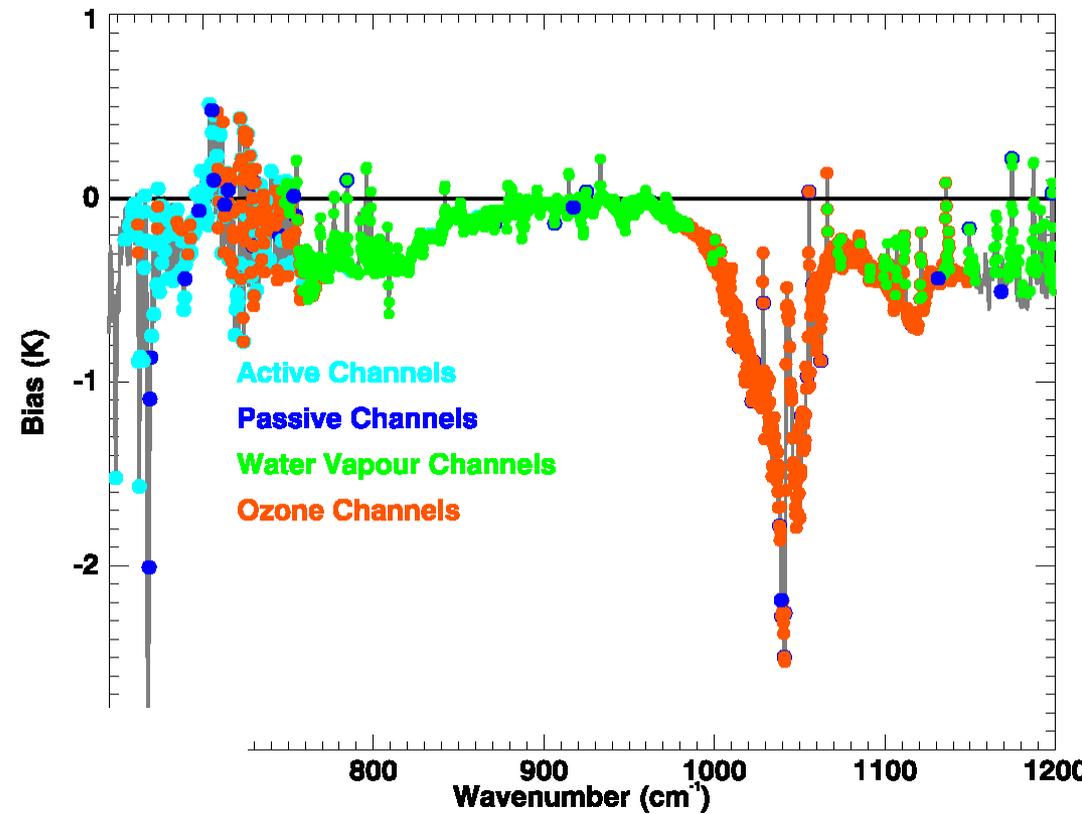
Standard Deviatric



# Jacobians of 15 $\mu$ m CO<sub>2</sub> Band

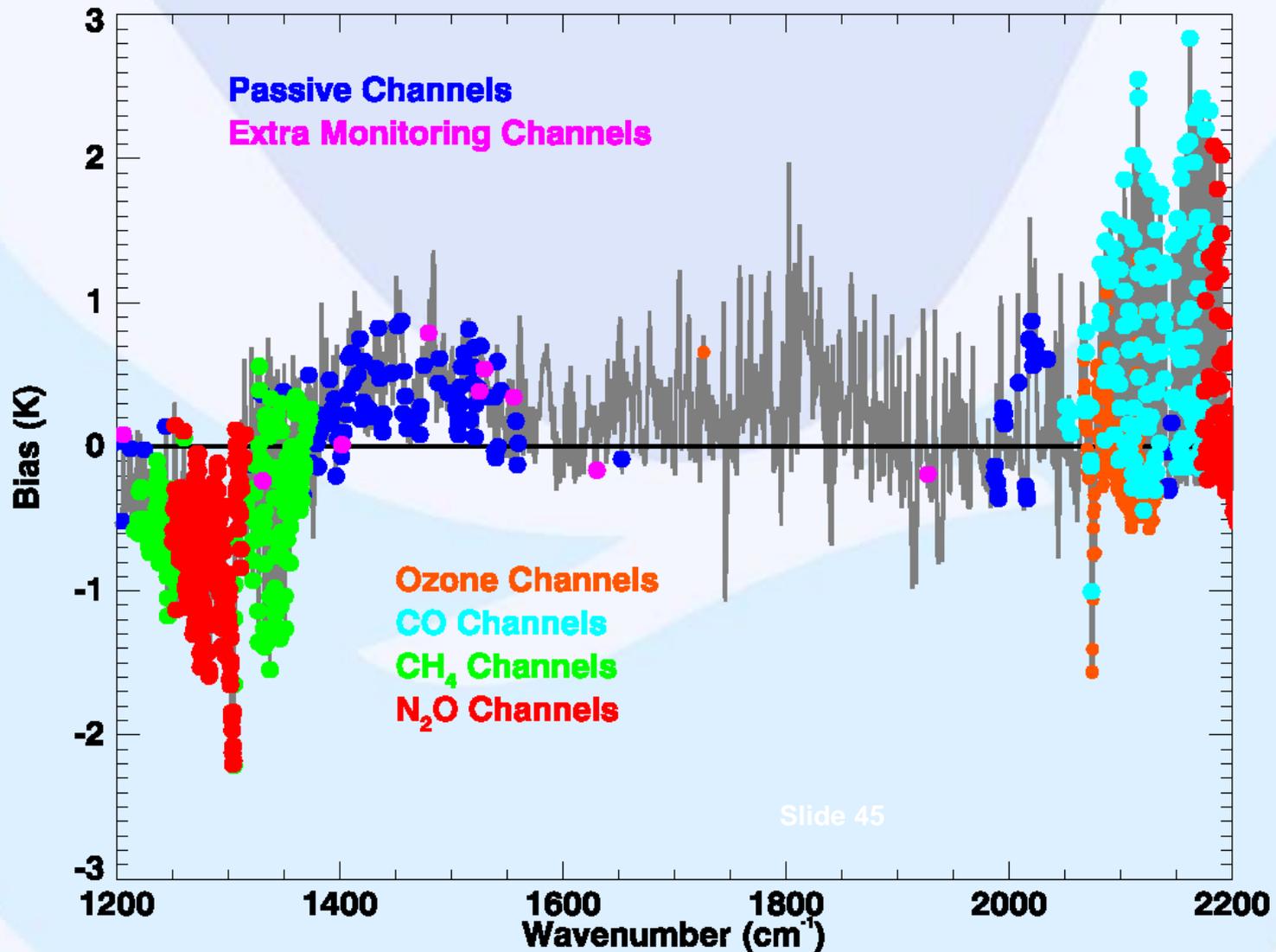


# First Guess Departure Standard Deviations and Biases in the Longwave Window

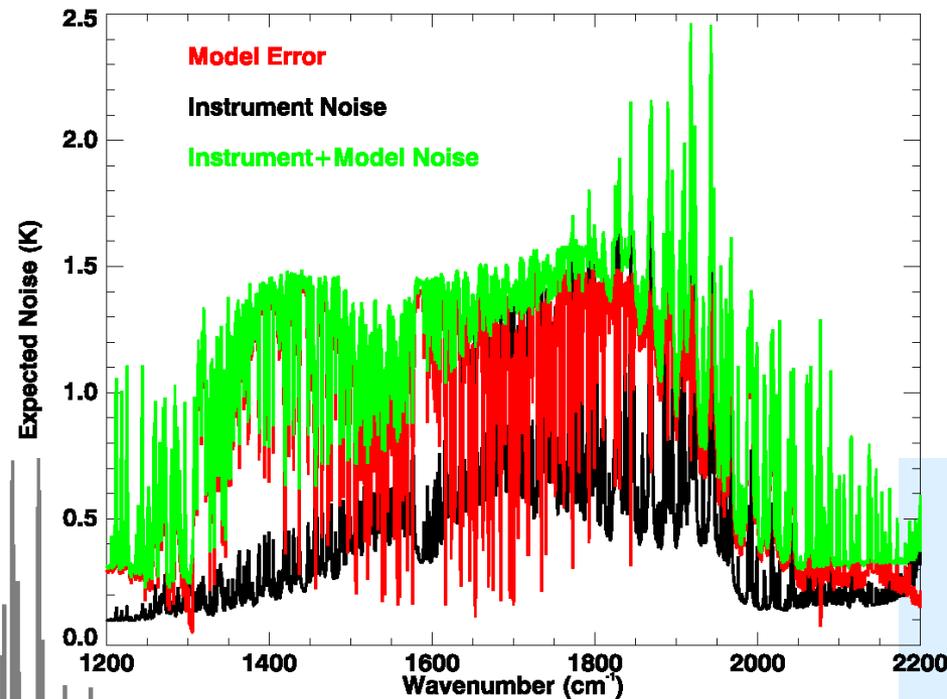
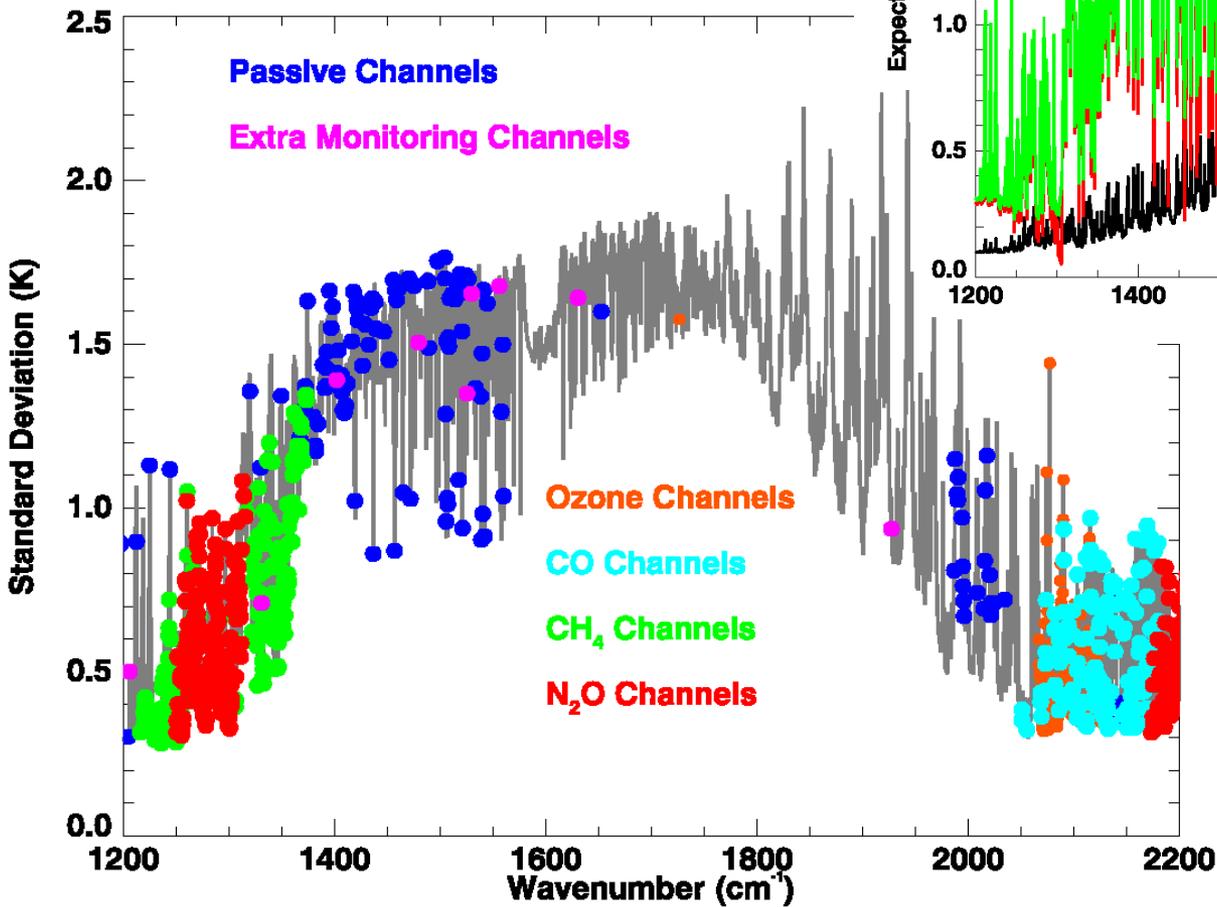


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# First-Guess Departure Biases in Water Band



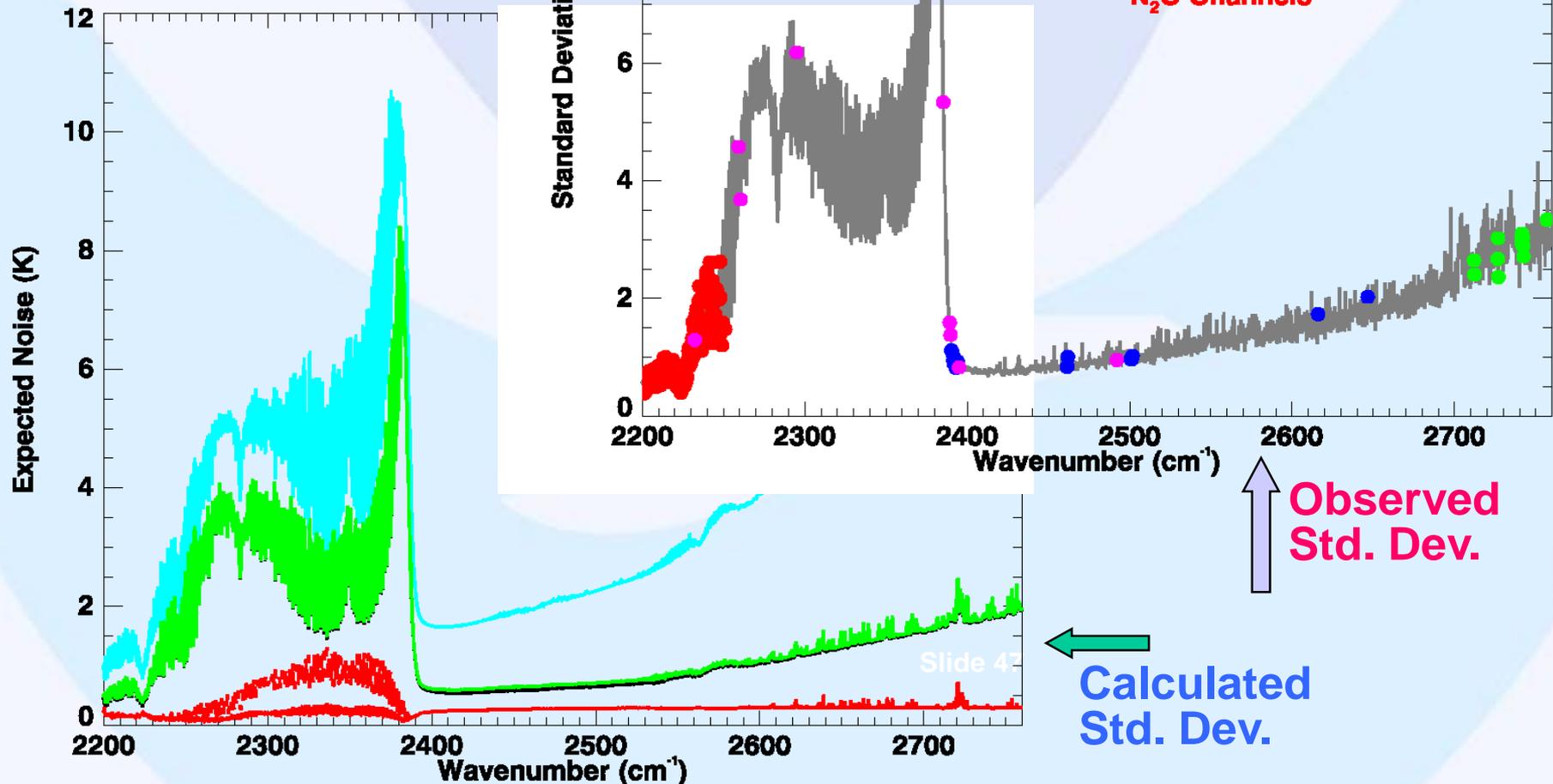
# First-Guess Departure Standard Deviations in Water Band



Calculated Std. Dev. ↑

← Observed Std. Dev.

# First –Guess Departure Standard Deviations in Shortwave Band



# Clouds

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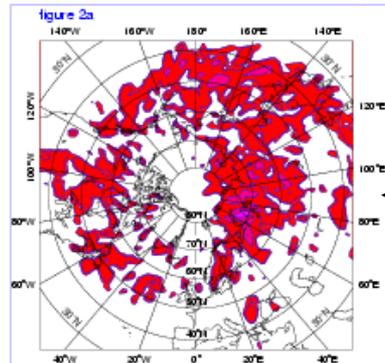
# Dealing with Cloud

Cloud can be treated in five ways:

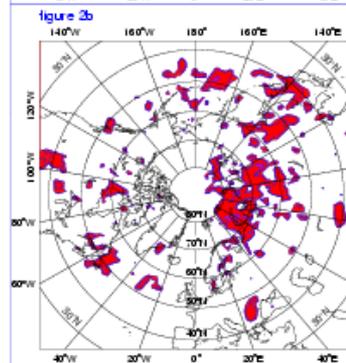
- 1) Avoid all FOVs with cloud (“hole hunting”)
- 2) Only assimilate channels that are insensitive to cloud
- 3) Correct the observations to remove the effect of clouds (“cloud-clearing”)
- 4) Explicitly model the effect of cloud on the radiances either during pre-processing or as a sink variable. But DO NOT assimilate the cloud properties.
- 5) Initialise model cloud variables from the cloudy radiances.

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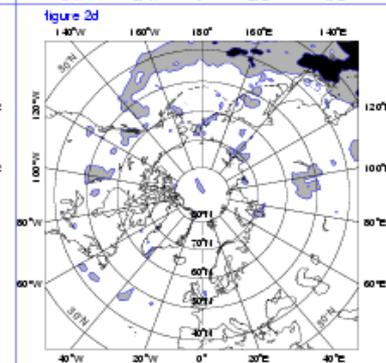
# Sensitive areas and cloud cover



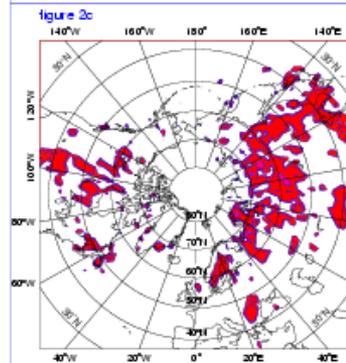
Location of sensitive regions  
← Summer-2001 (no clouds)



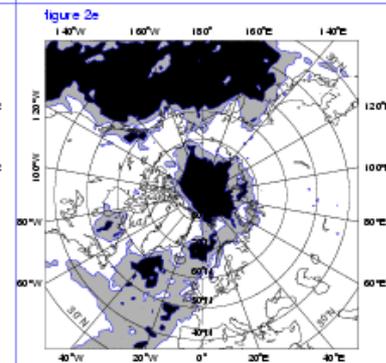
sensitivity surviving **high** cloud cover



monthly mean **high** cloud cover



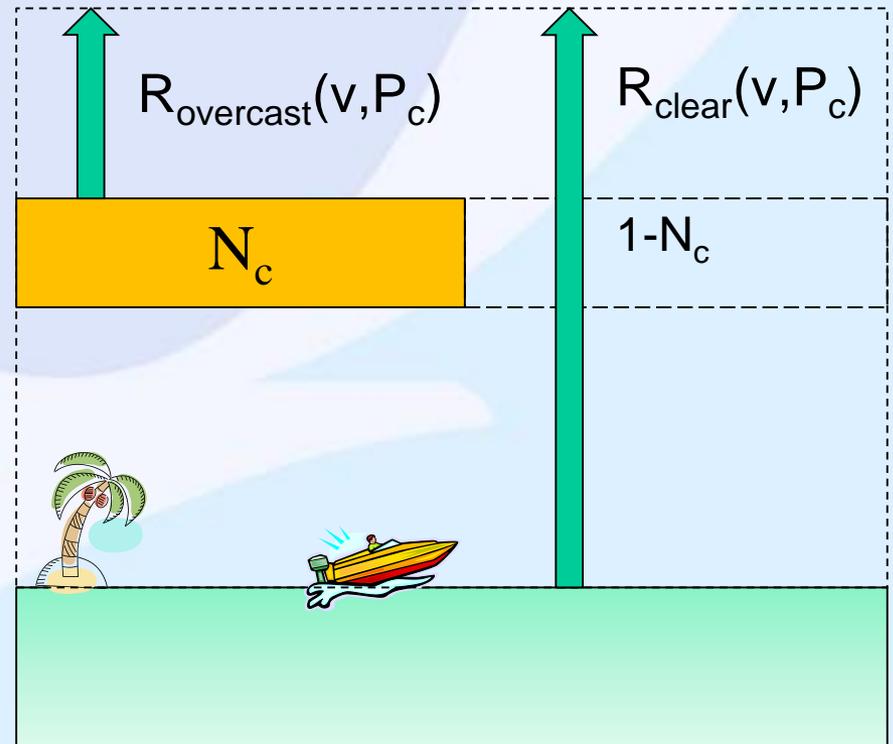
sensitivity surviving **low** cloud cover



monthly mean **low** cloud cover

# Cloud Detection in the GSI

- Assume the cloud is a single layer at pressure  $P_c$  and with unit emissivity and coverage within the FOV,  $N_c$ .
- $0 \leq N_c \leq 1$
- $P_c$  is below the tropopause and above the ground
- Find  $P_c$  and  $N_c$  so that the RMS deviation,  $J(N_c, P_c)$ , of the calculated cloud from the model (over a number of channels) is minimized.
- Remove all channels that would be radiatively affected by this cloud.



$$R_{\text{clid}}(v, P_c) = N_c R_{\text{overcast}}(v, P_c) + (1 - N_c) R_{\text{clear}}(v, P_c)$$

$$J(N_c, P_c) = \sum_v (R_{\text{clid}}(v, P_c) - R_{\text{obs}}(v) / \sigma(v))^2$$

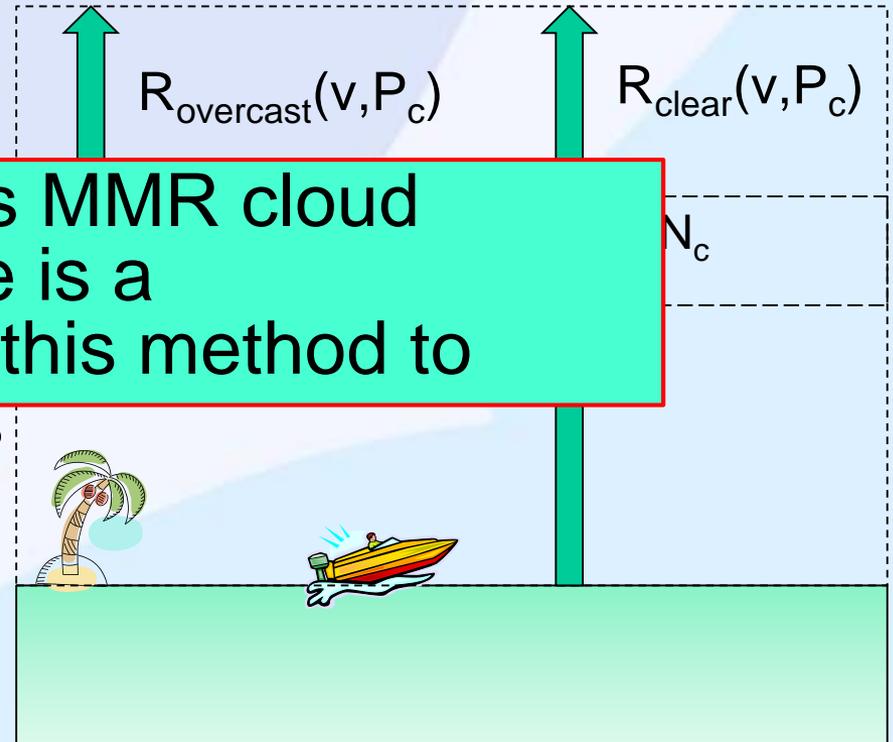
$\sigma(v)$  is the assumed observation error for channel  $v$

This calculation should be done in radiance, not brightness temperature space.

# Cloud Detection in the GSI

- Assume the cloud is a single layer at pressure  $P_c$  and with unit emissivity and coverage within the FOV,  $N_c$
- $0 \leq N_c \leq 1$
- $P_c$  is below the surface and above the surface
- Find  $P_c$  and  $N_c$  such that the deviation, calculated from the model (over a number of channels) is minimized.
- Remove all channels that would be radiatively affected by this cloud.

Thomas Auligne's MMR cloud detection scheme is a generalisation of this method to multi-level clouds



$$R_{\text{clid}}(v, P_c) = N_c R_{\text{overcast}}(v, P_c) + (1 - N_c) R_{\text{clear}}(v, P_c)$$

$$J(N_c, P_c) = \sum_v (R_{\text{clid}}(v, P_c) - R_{\text{obs}}(v) / \sigma(v))^2$$

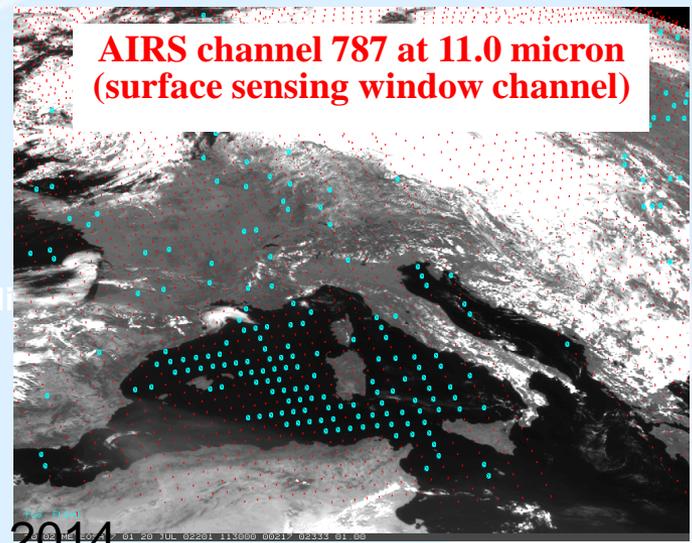
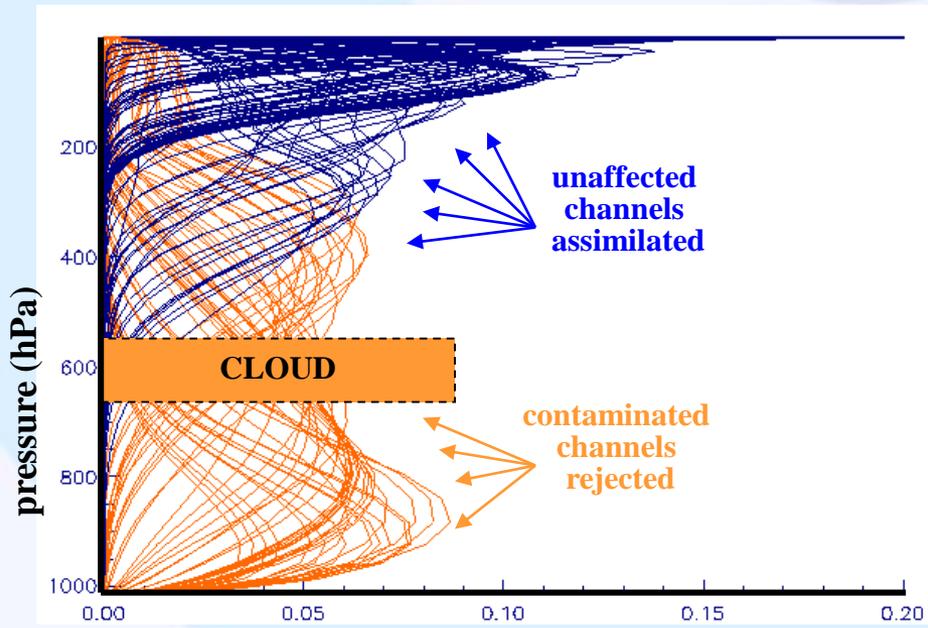
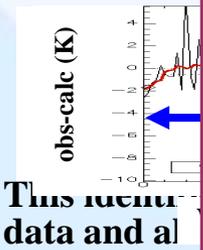
$\sigma(v)$  is the assumed observation error for channel  $v$   
This calculation should be done in radiance, not brightness temperature space.

# Cloud detection scheme for Advanced Sounders

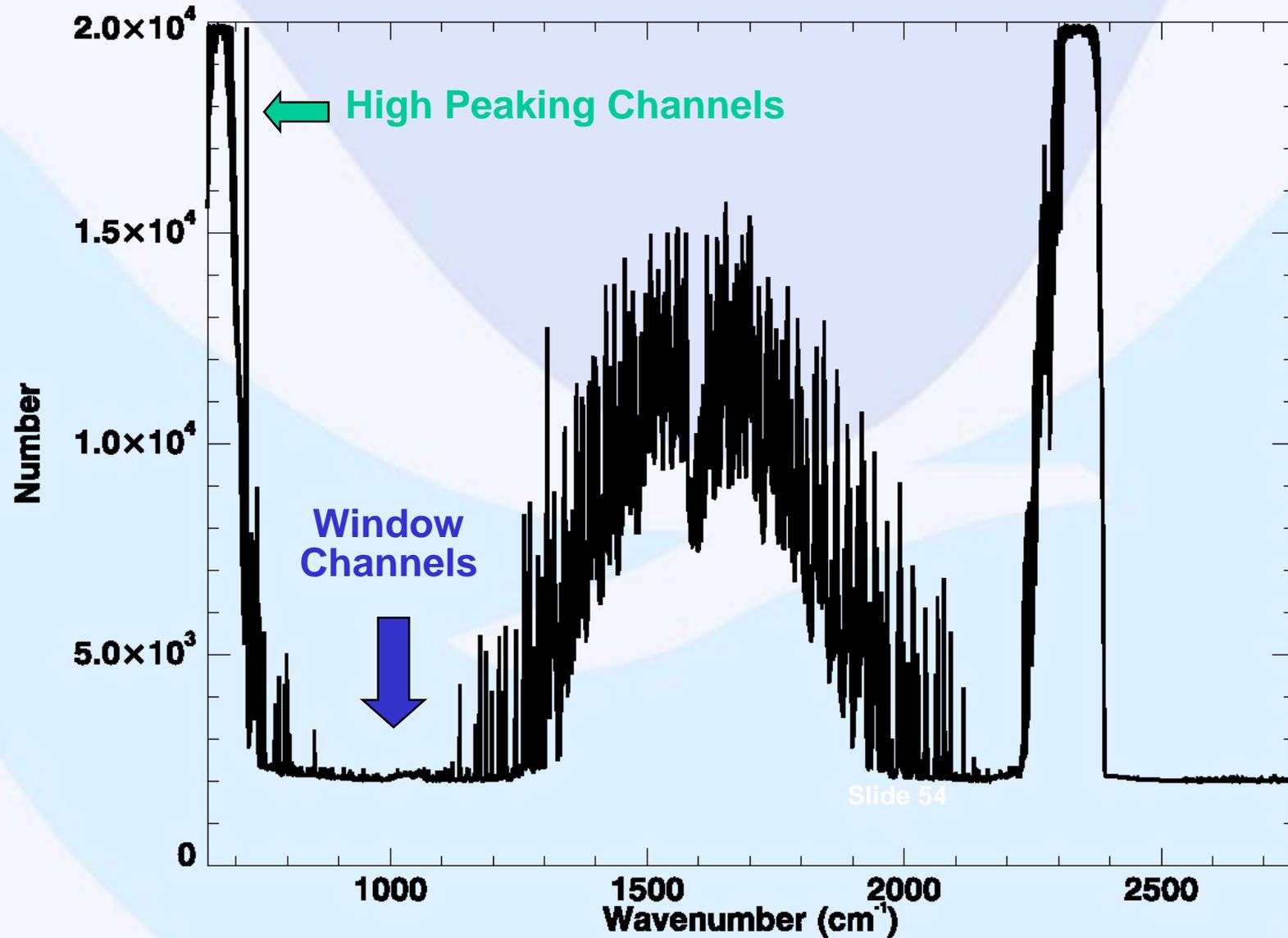
A non-linear pattern recognition algorithm is applied to departures of the observed radiance spectra from a computed clear-sky background spectra.



The large number of AIRS or IASI channels allows improved measurement of the cloud-top height compared to HIRS

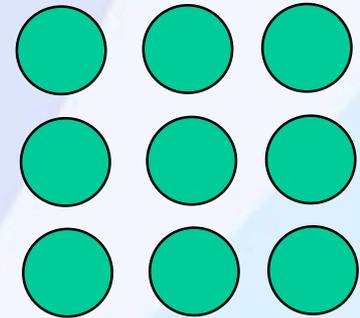


# Number of Clear Channels



# Cloud Cleared Radiances

Cloud Cleared Radiances derive a single “clear” spectrum from an array of partially cloudy fields-of-view (9 in the case of AIRS and CrIS)



Assumes the cloud height in each FOV is identical and only cloud fraction varies between the FOVs.

Needs to be initialised with a high quality first guess, usually either a regression from AMSU-A radiances or an NWP model field.

Can calculate a noise amplification factor which is the basis of the QC flag

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## Cloud-clearing methodology

Following Chahine (1977) and Joiner and Rokke (2000), the observed radiance at channel  $i$  from field of view (FOV)  $j$  can be written as

$$R_j^i = \left( 1 - \sum_{k=1}^K \alpha_{j,k} \right) \times R_{clr}^i + \sum_{k=1}^K \alpha_{j,k} \times R_{cld}^{i,k}$$

$\alpha_{j,k}$  is the cloud type  $k$  fraction in FOV  $j$ ,  $R_{clr}^i$  is the clear-sky radiance at channel  $i$  and  $R_{cld}^{i,k}$  is the radiance at channel  $i$  from the cloud type  $k$  with fraction  $\alpha_{j,k}$ .

The cloud-cleared radiance can be derived if assuming  $R_{clr}^i$  and  $R_{cld}^{i,k}$  are the same in the 3x3 FOVs in a single CrIS field of regard (FOR) and eliminating  $R_{cld}^{i,k}$ ,

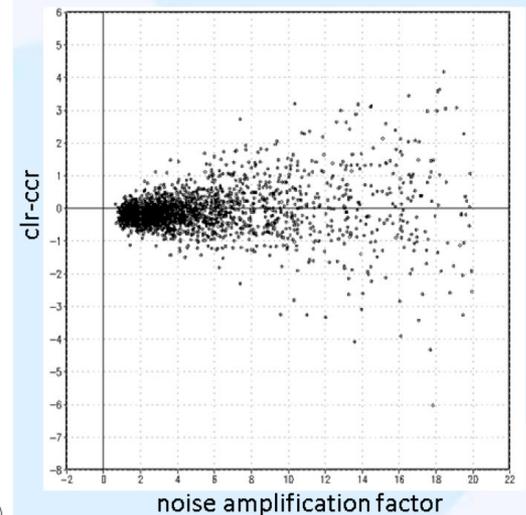
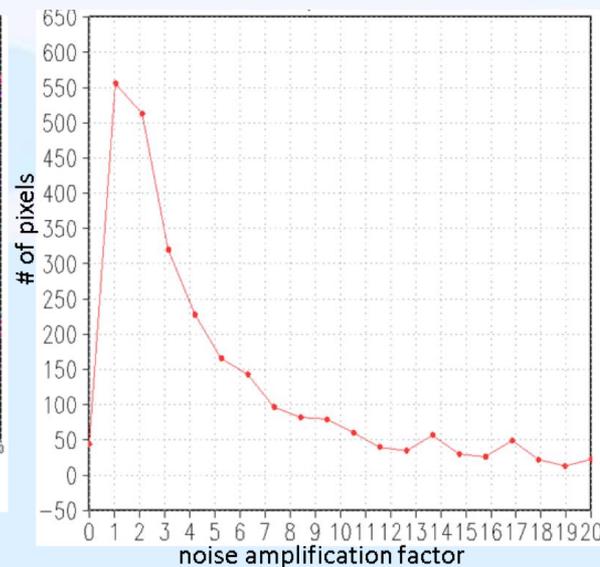
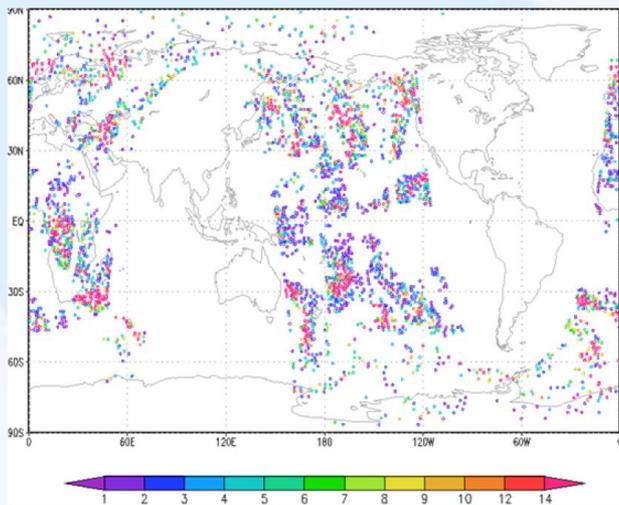
$$R_{ccr}^i = R_1^i + \eta_1 \times (R_1^i - R_2^i) + \eta_2 \times (R_1^i - R_3^i) + \dots + \eta_k \times (R_1^i - R_{k+1}^i)$$

$\eta_k$  are the cloud-clearing parameters. They depend on the cloud fraction  $\alpha_{j,k}$  only, so they are channel-independent. These parameters can be solved by using model guess simulated clear-sky radiances and observed radiances in adjacent pixels at cloud sounding channels in a least-square sense.

# Noise amplification factor

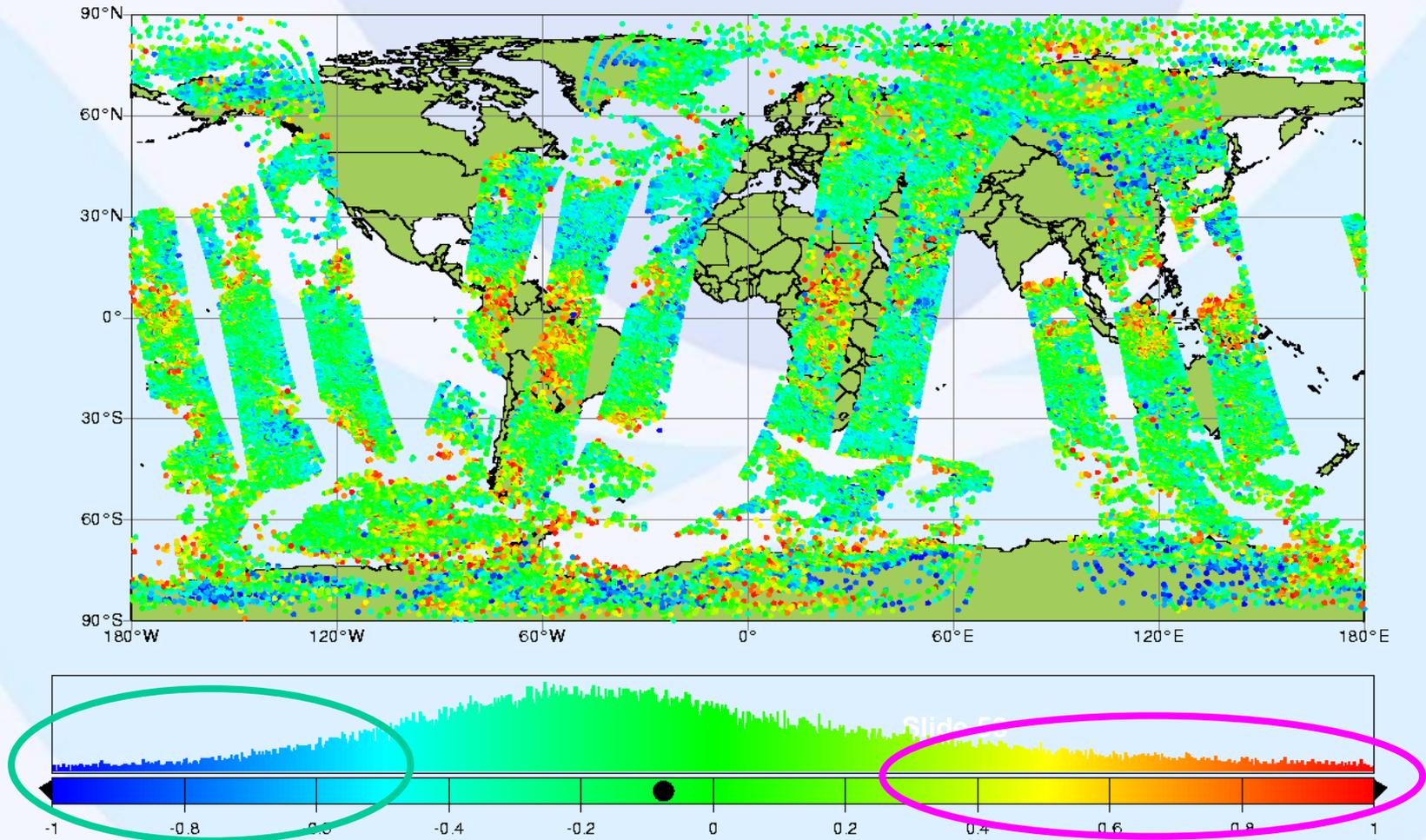
The reconstructed clear-column radiances will be treated as other IR clear-sky radiances when being assimilated except for amplifying the observation error

by a factor of  $A = \sqrt{\left(1 + \sum_{k=1}^K \eta_k\right)^2 + \sum_{k=1}^K \eta_k^2}$ .



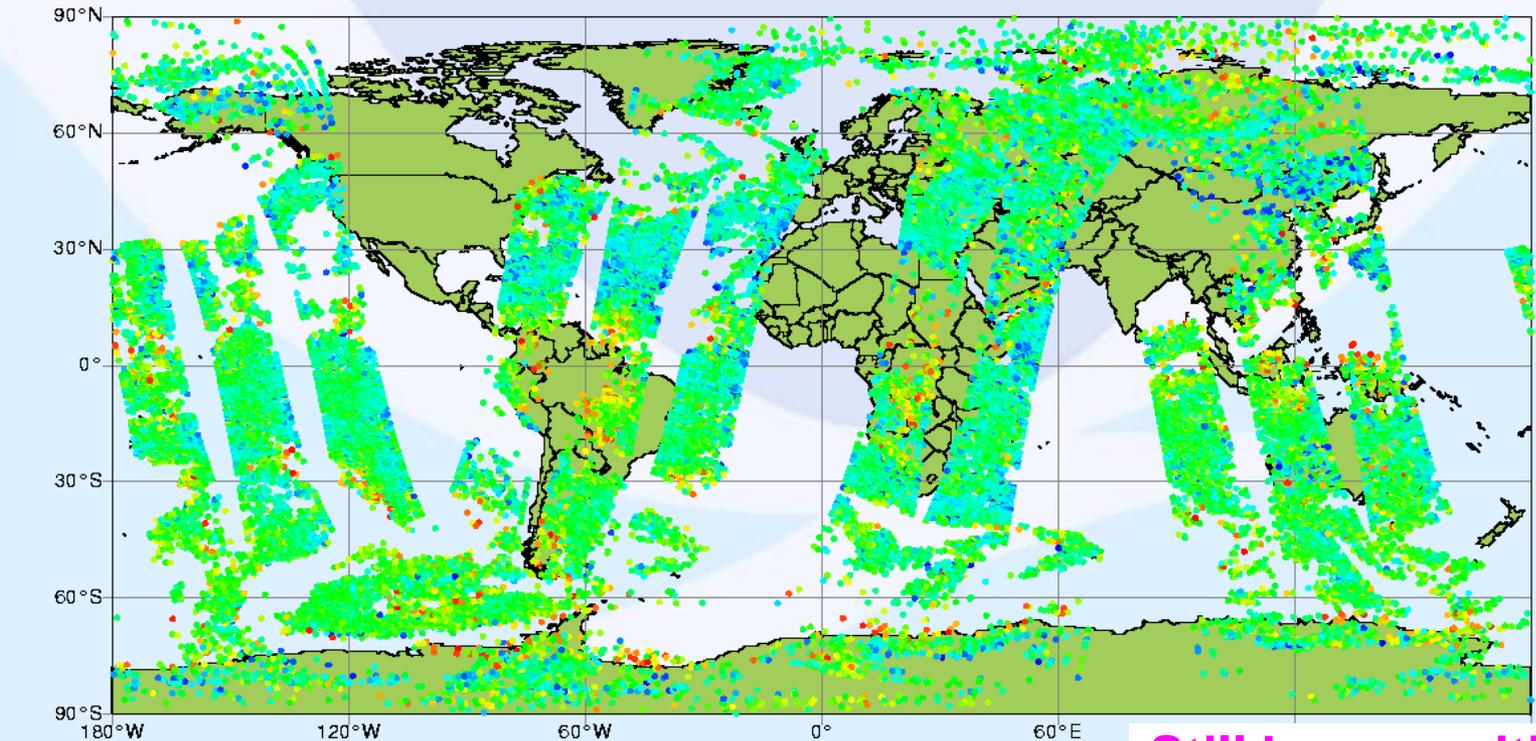
# Cloud-Cleared Radiances

**AIRS Ch. 221. QC Flag Used. No Cloud Detection**

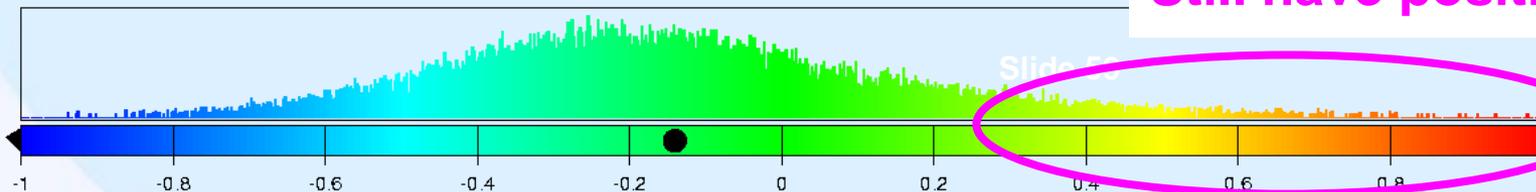


# Cloud Cleared Radiances

AIRS Ch. 221. QC Flag Used. With Cloud Detection

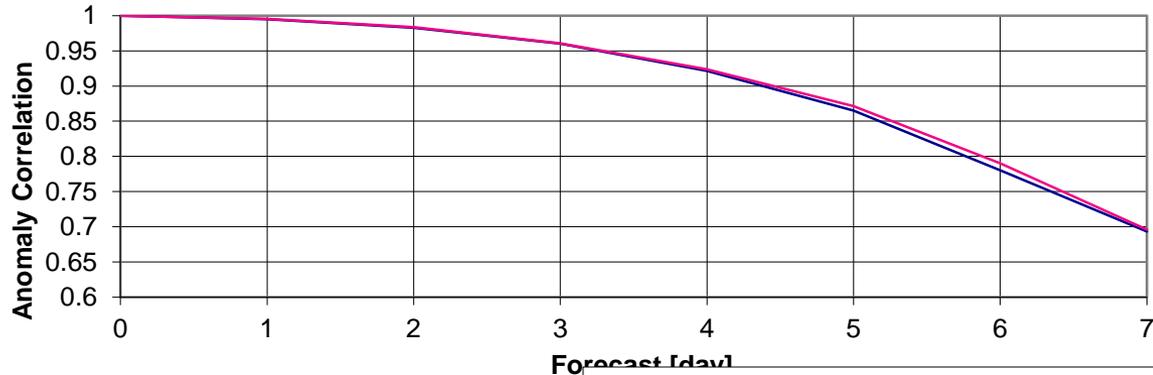


Still have positive tail

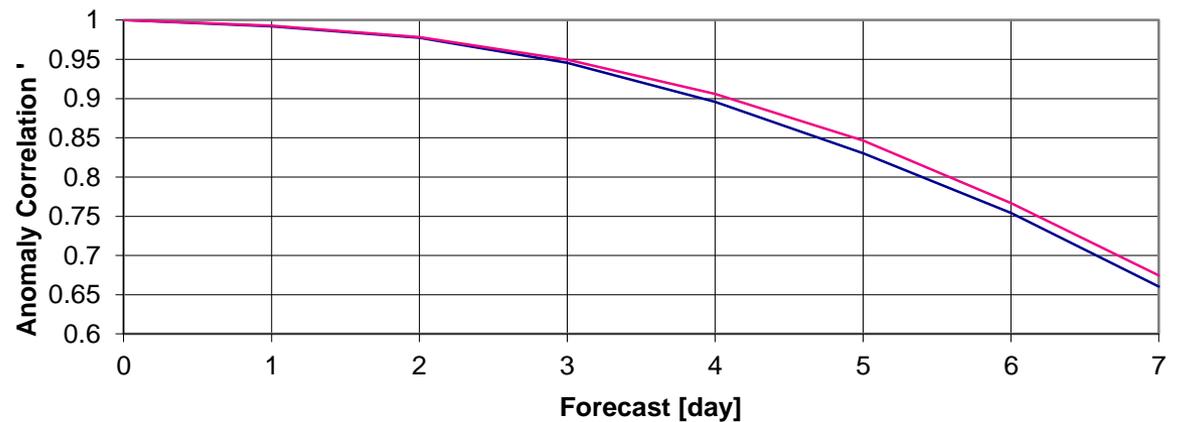


# Impact of CCRs on Forecast?

**N. Hemisphere 500 hPa AC Z**  
**20N - 80N Waves 1-20**  
**1 Jan - 24 Feb '07**



**S. Hemisphere 500 hPa AC Z**  
**20S - 80S Waves 1-20**  
**1 Jan - 24 Feb '07**



**Results from Jim Jung using NCEP GFS model.**

# Simultaneous Analysis of Cloud Properties (Tony McNally)

Derive a cloud-top pressure (CTP) and cloud fraction from observed radiances with a 2-D least-squares fit in the screening run

For overcast FOVs use all channels (that are currently used for clear sky case)

This has the advantage of reducing the degrees of freedom

For other FOVs revert to operational cloud-detection scheme to identify clear channels

Assimilate these radiances with CTP as a sink variable

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# Using data in cloudy areas

Tony McNally

Clear data coverage of mid/lower tropospheric soundings radiances:

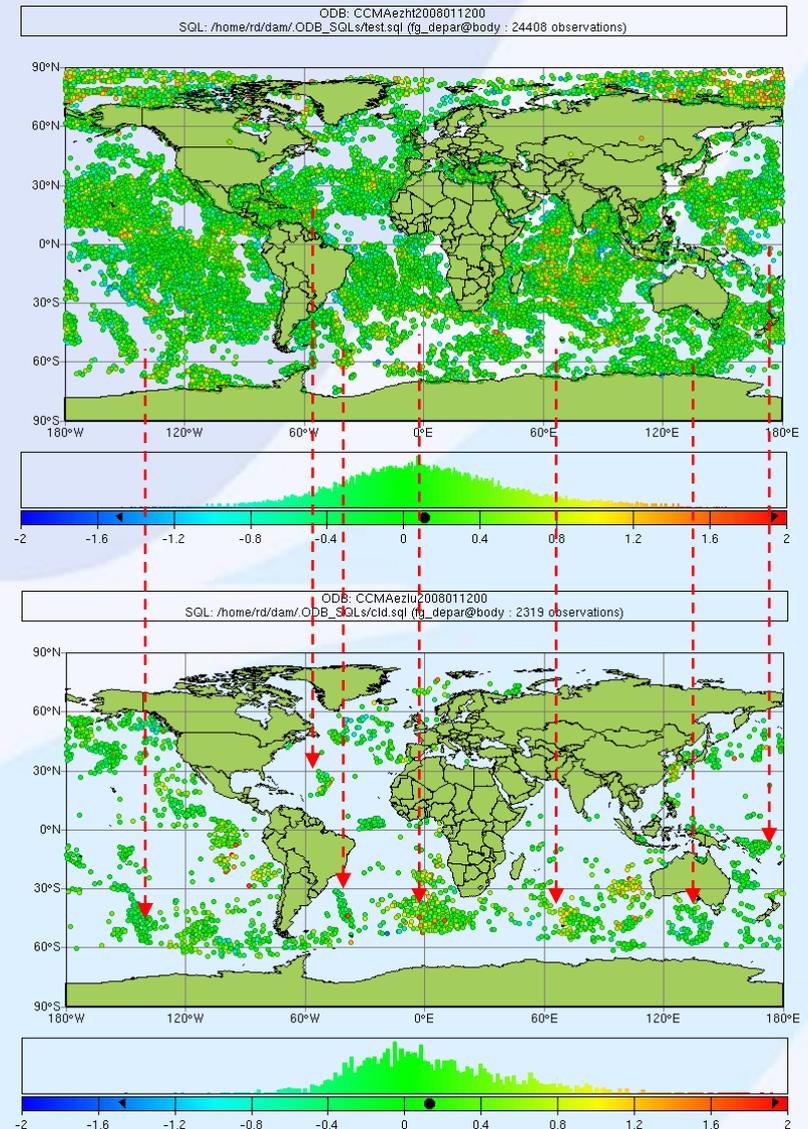
IASI 434 (METOP-A)  
AIRS 355 (AQUA)  
HIRS 7 (NOAA-17 / METOP-A)

Colour indicates first guess departure

Additional overcast locations where cloudy radiance analysis fills gaps due to cloud detection rejections:

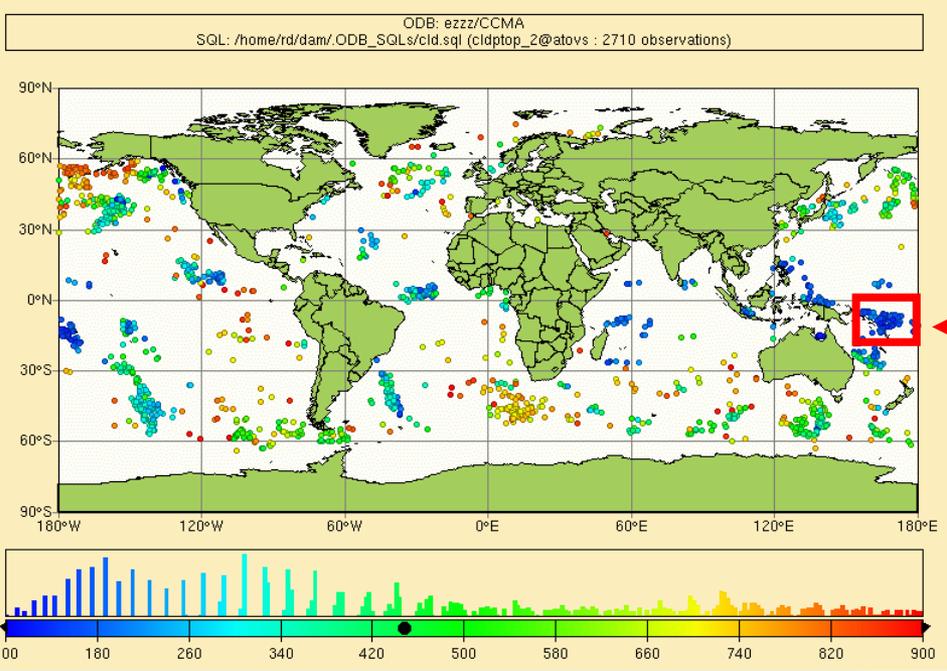
IASI 434 (METOP-A)  
AIRS 355 (AQUA)  
HIRS 7 (NOAA-17 / METOP-A)

First guess departures similar to clear data after QC of *complex* clouds



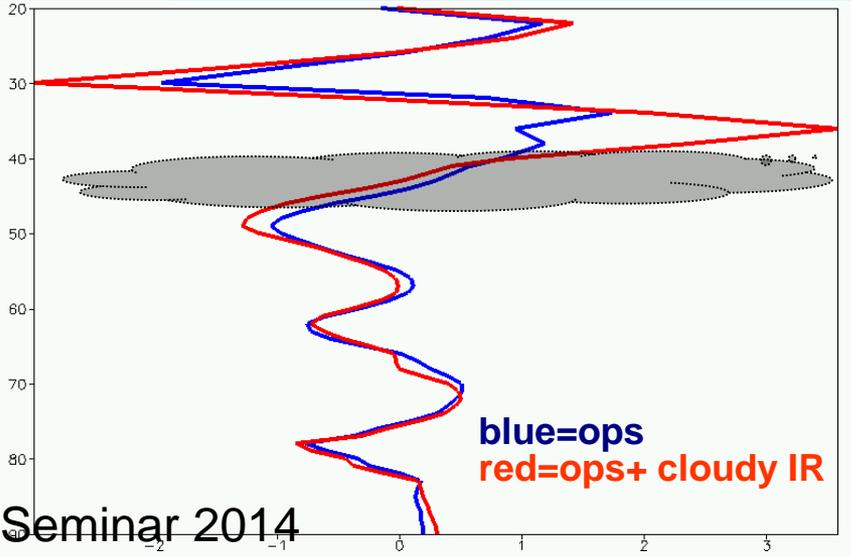
# Temperature increments at the cloud top

Tony McNally



Cell of very high overcast clouds off the coast of PNG

Temperature increments (IASI)



All channels collapse to near delta-functions at the cloud top giving very high vertical resolution temperature increments just above the diagnosed cloud

# Direct Assimilation of IR Cloudy Radiances

- The direct use of the cloudy information in IR cloudy radiances is the ultimate aim. A number of centres are starting to get encouraging results in this area.
- The main challenges are:
  - Non-linearity of the observation operator (particularly for opaque clouds where the Jacobian is essentially a delta-function at the cloud top).
  - Radiative transfer accuracy
  - Partially cloudy scenes
  - Propagation of cloudy errors into other fields (such as temperature).

# Direct Assimilation of IR Cloudy Radiances

- Pauline Martinet is exploring the direct assimilation of IR radiances in the AROME model. Analyses are limited to schemes where the model and observations agree that the cloud field is homogeneous. Some initial encouraging results.
- Stefano Migliorini has shown positive impact through the assimilation of cloudy radiances in the ECMWF IFS. To minimize the risk of cloudy signal adversely affecting the temperature field, this work focuses on the use of humidity-sensitive channels.

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# Humidity Assimilation

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# Humidity assimilation (1)

Some centres have demonstrated positive impact from assimilating H<sub>2</sub>O channels (with reduced weight) to the analysis and 1-2 day forecast

NWP models have a hard time keeping impact of assimilation after 1-2 days.

Ambiguity with humidity Jacobians - the water vapor (WV) channels have strong sensitivity to humidity and temperature

Humidity Jacobians are non-linear; i.e., the Jacobians themselves are a function of the humidity field

Representivity error is probably the most important contributor to the observation error covariance in the water band (at least for global DA). This takes the form of a highly correlated error covariance with amplitude of 1-2K, which therefore dominates the instrument noise.

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## Humidity assimilation (2)

Large bias relative to NWP model (model bias). Bias correction algorithms remove this bias.

RT model errors/biases may contribute as well.

Variational bias correction algorithms need to have suitable anchoring observations.

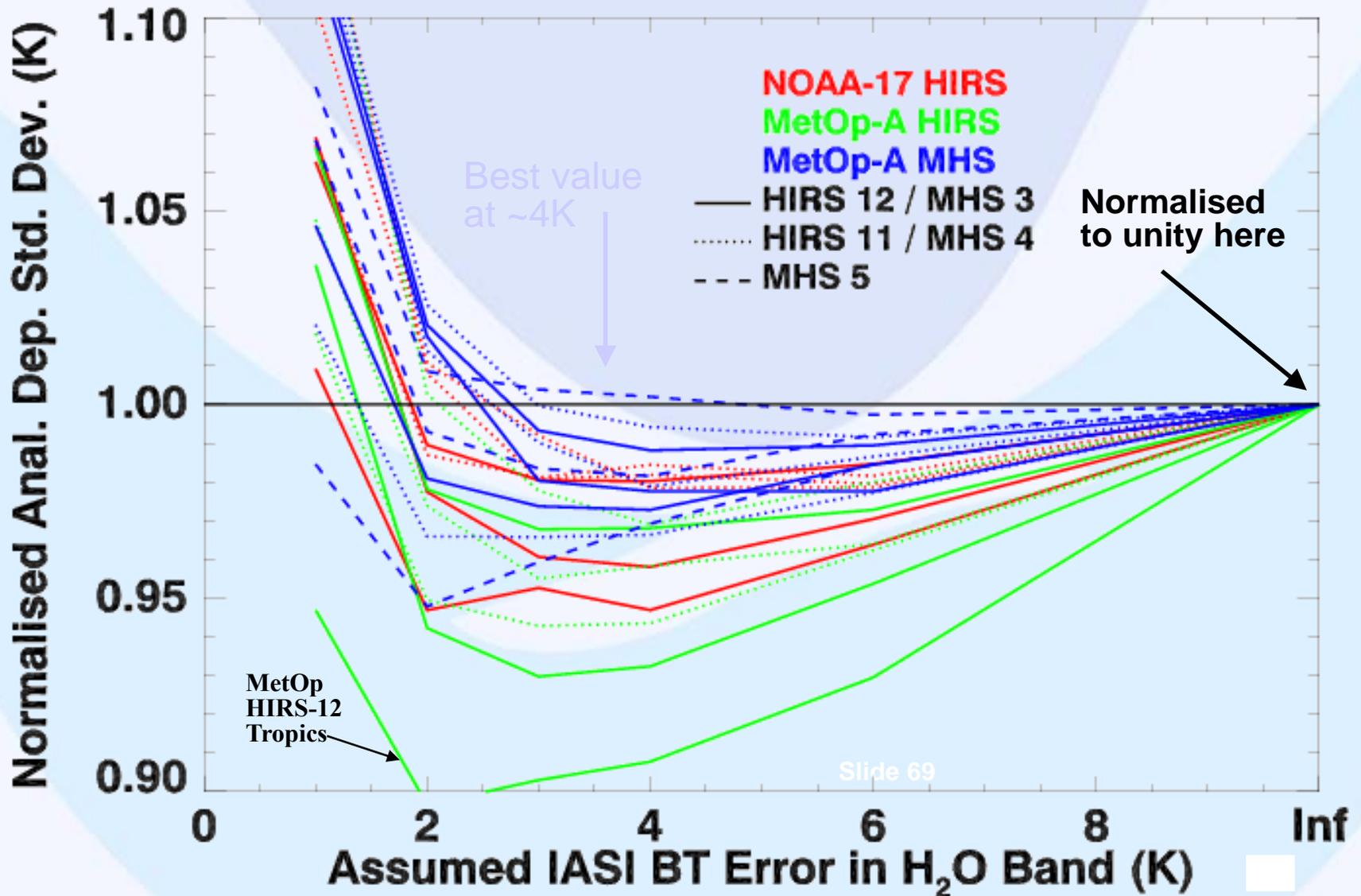
Radiance departures should be used to inform forecast model changes (Hilton et al., 2012)

Above issues are mitigated through inflated observation errors; reduced number of channels, tight QC and explicit representation of correlated observation errors.

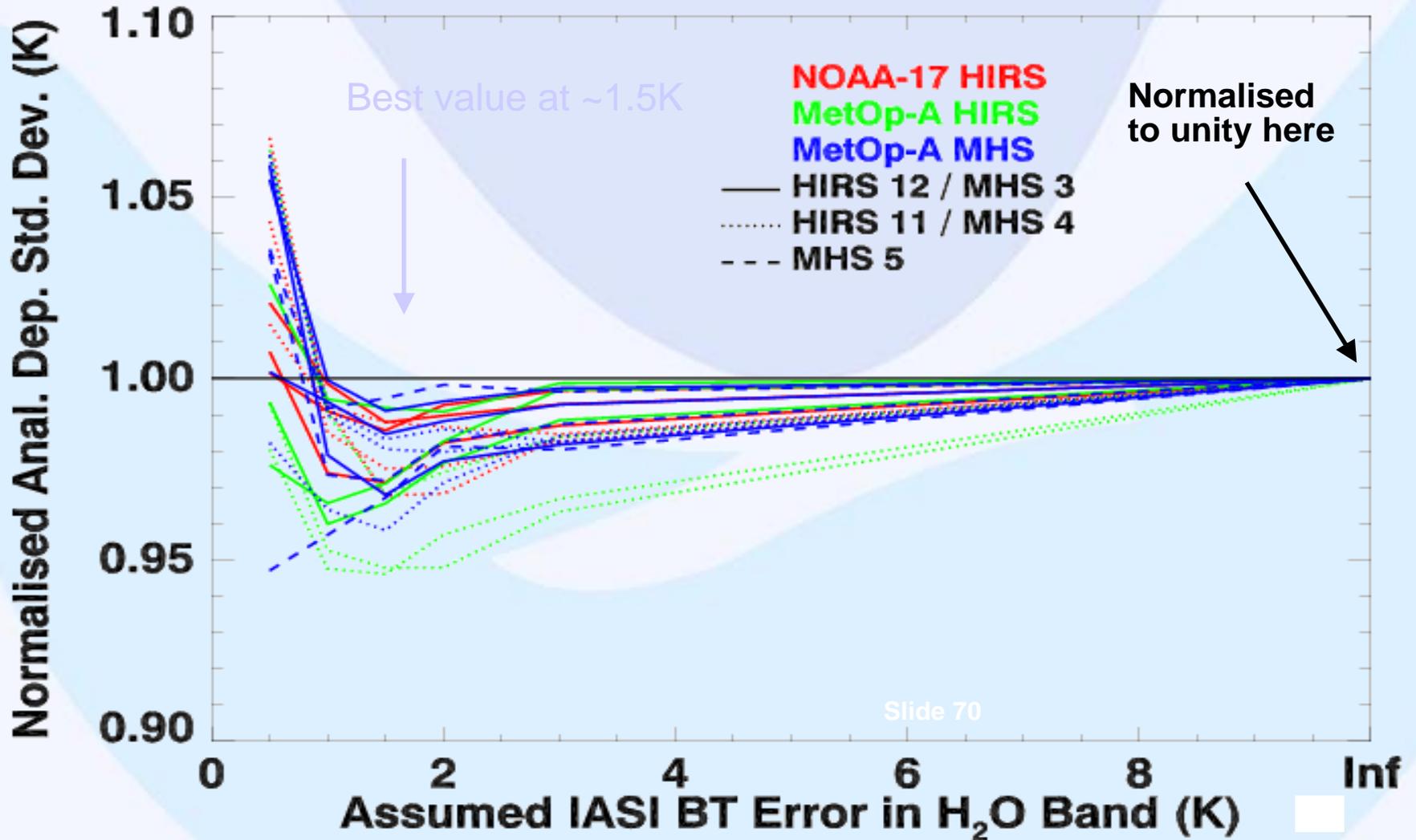
NCEP use tight QC (~1K) but increase data useage through re-evaluation of QC every outer loop.

The Met Office explicitly specifies the<sup>6</sup> correlated observation error for IASI (Bormann talk).

# Fit to other observations 84 IASI Water Channels



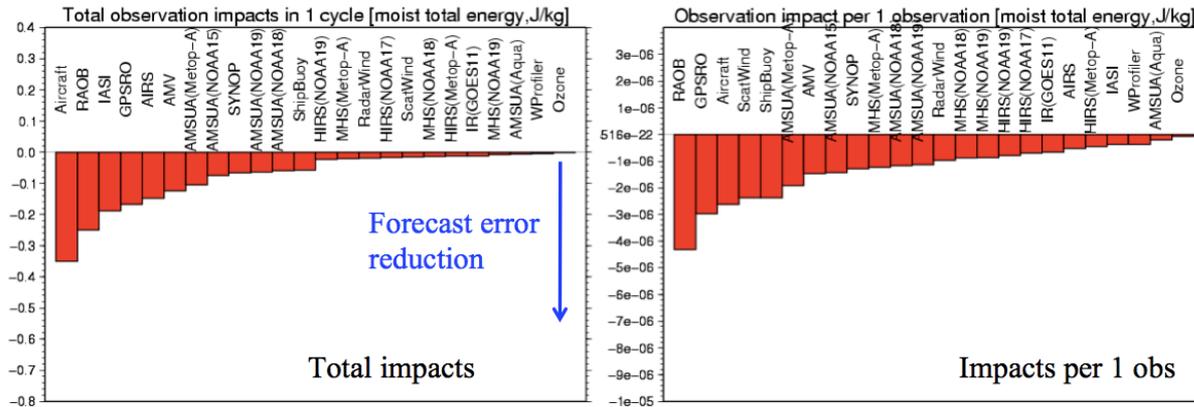
# Fit to other observations: 10 IASI Water Channels



# **Overview of IR Sounder Performance in NWP**

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# Global NWP Data Impacts

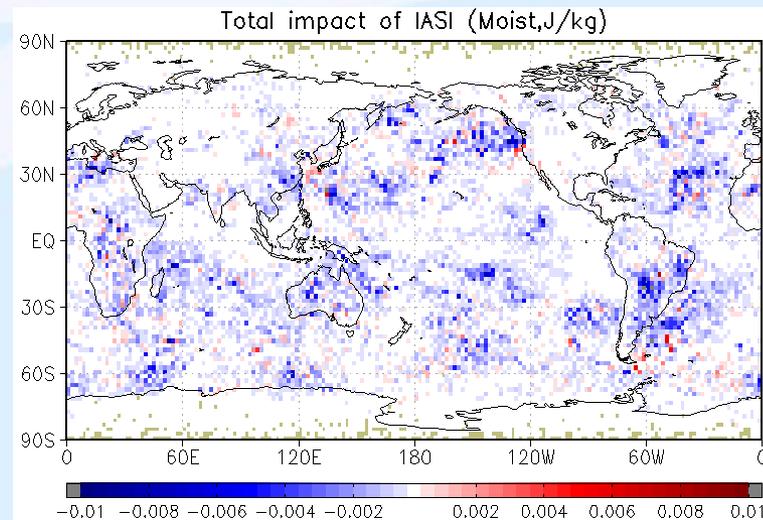
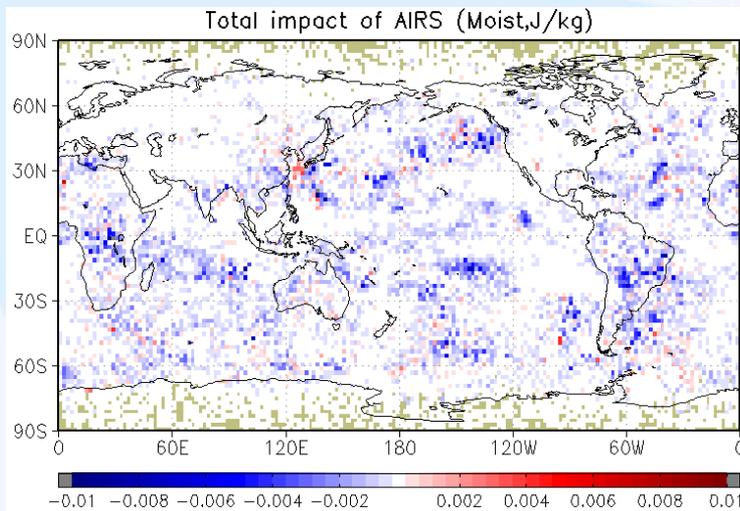


Ensemble  
FSO at  
NCEP

All observation types have positive impacts on average.

For the total impact, 1: aircraft, 2: AMSU-A, 3: radiosonde, 4: IASI, 5: GPSRO

For impact per 1 obs., 1: radiosonde, 2: GPSRO, 3: aircraft, 4: Scattrometer wind, 5: marine surface observation



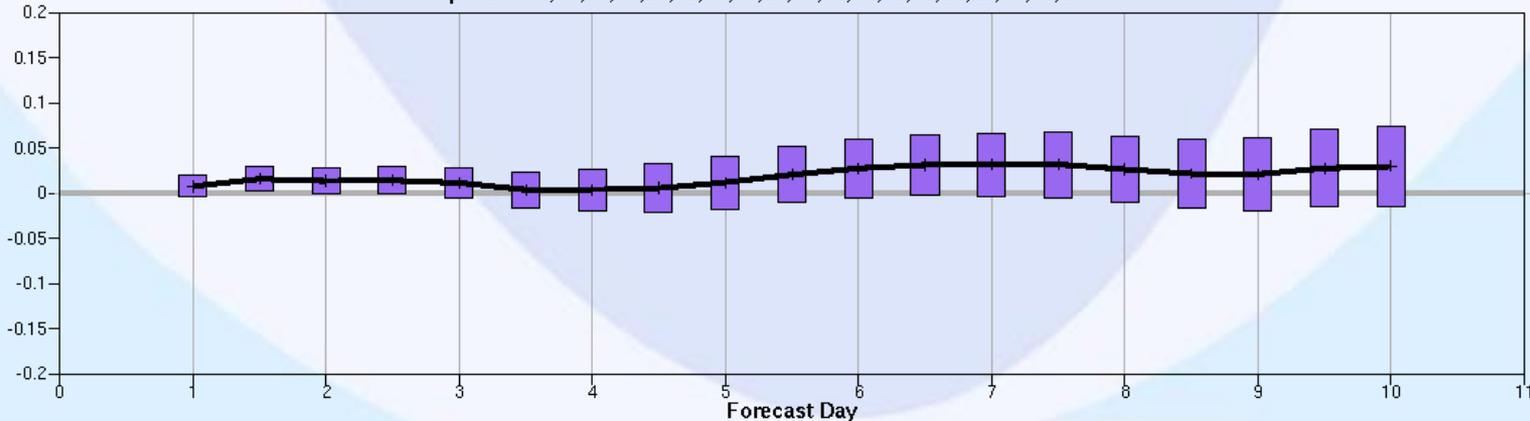
# IASI Forecast Scores: 500hPa Geopot. AC

8<sup>th</sup> March-5<sup>th</sup> May 2007

control normalised evto minus evng  
 Anomaly correlation forecast  
 N.hem Lat 20.0 to 90.0 Lon -180.0 to 180.0  
 Date: 20070308 00UTC to 20070505 00UTC  
 500hPa Geopotential 00UTC  
 Confidence: 90%

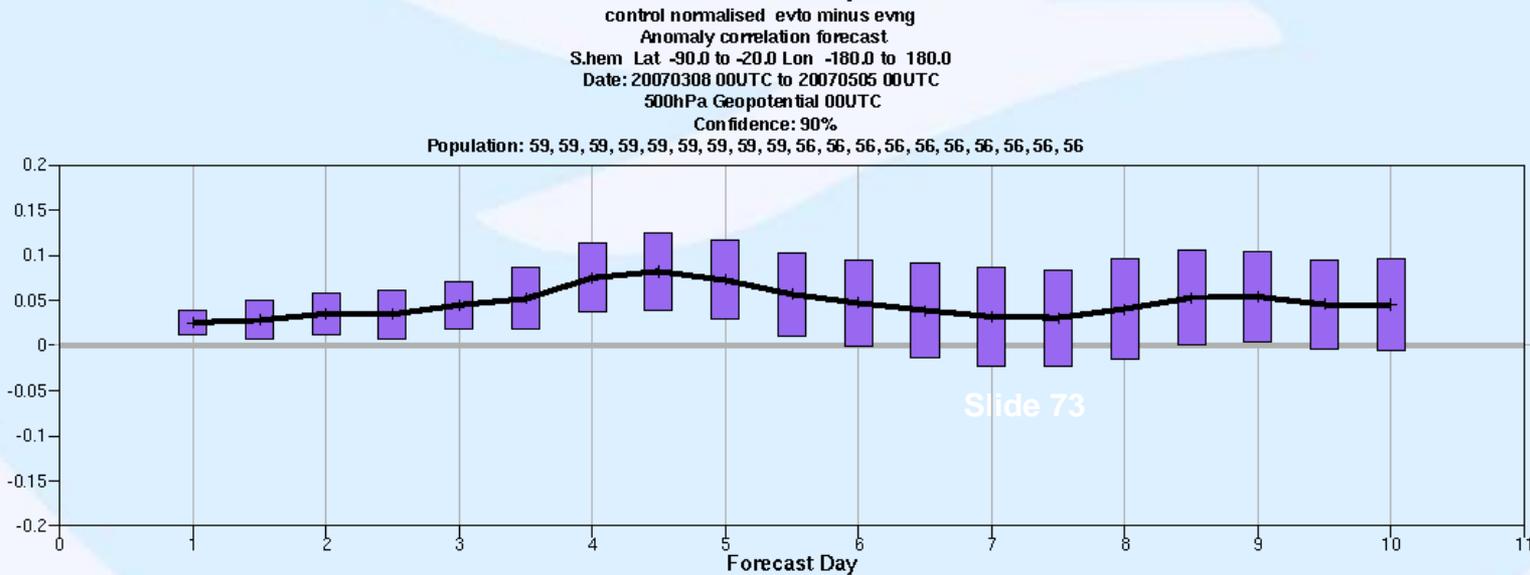
Population: 59, 59, 59, 59, 59, 59, 59, 59, 59, 56, 56, 56, 56, 56, 56, 56, 56, 56, 56

NH



IASI  
Better

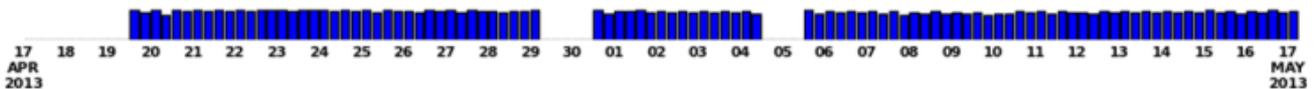
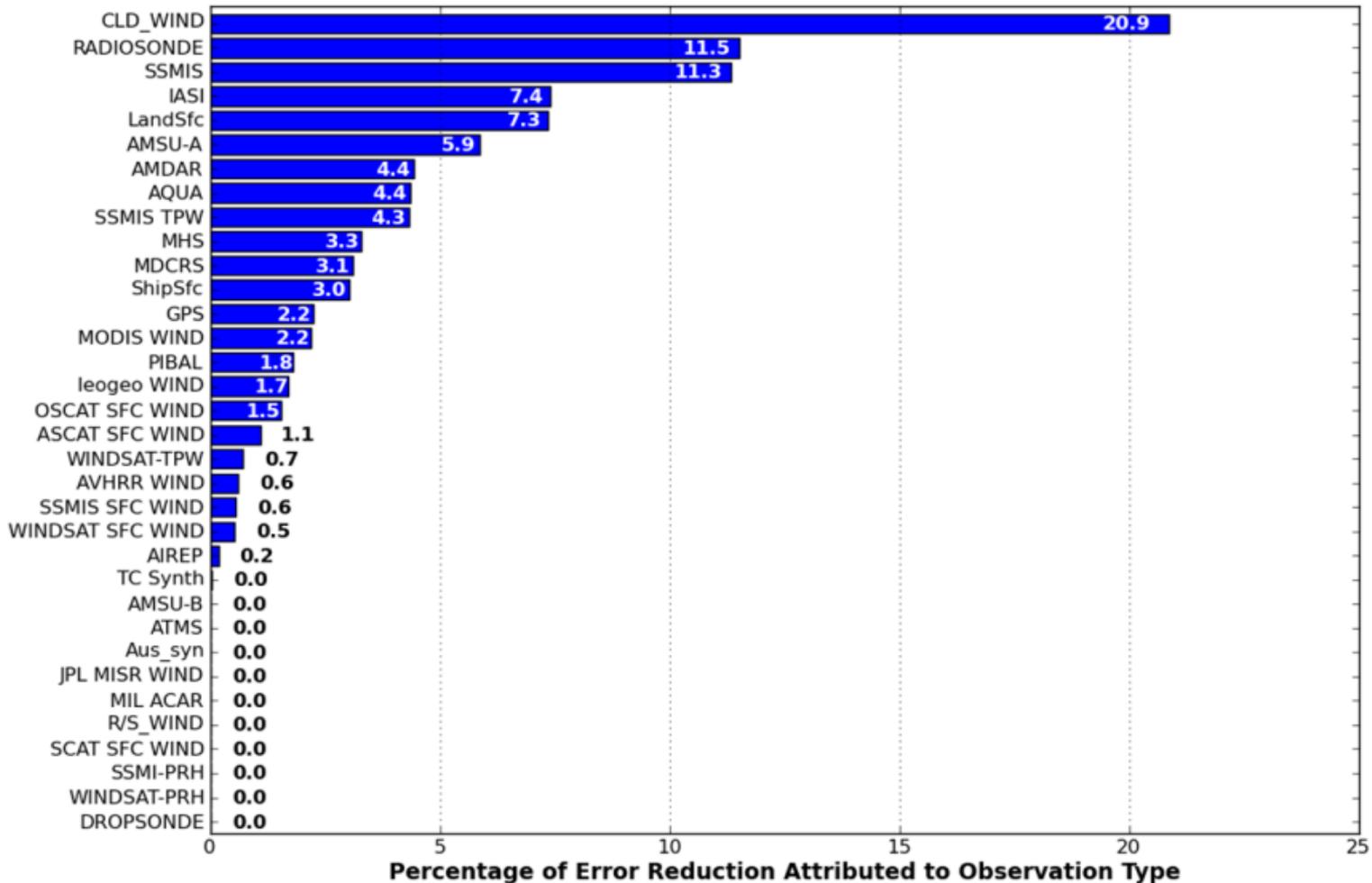
SH



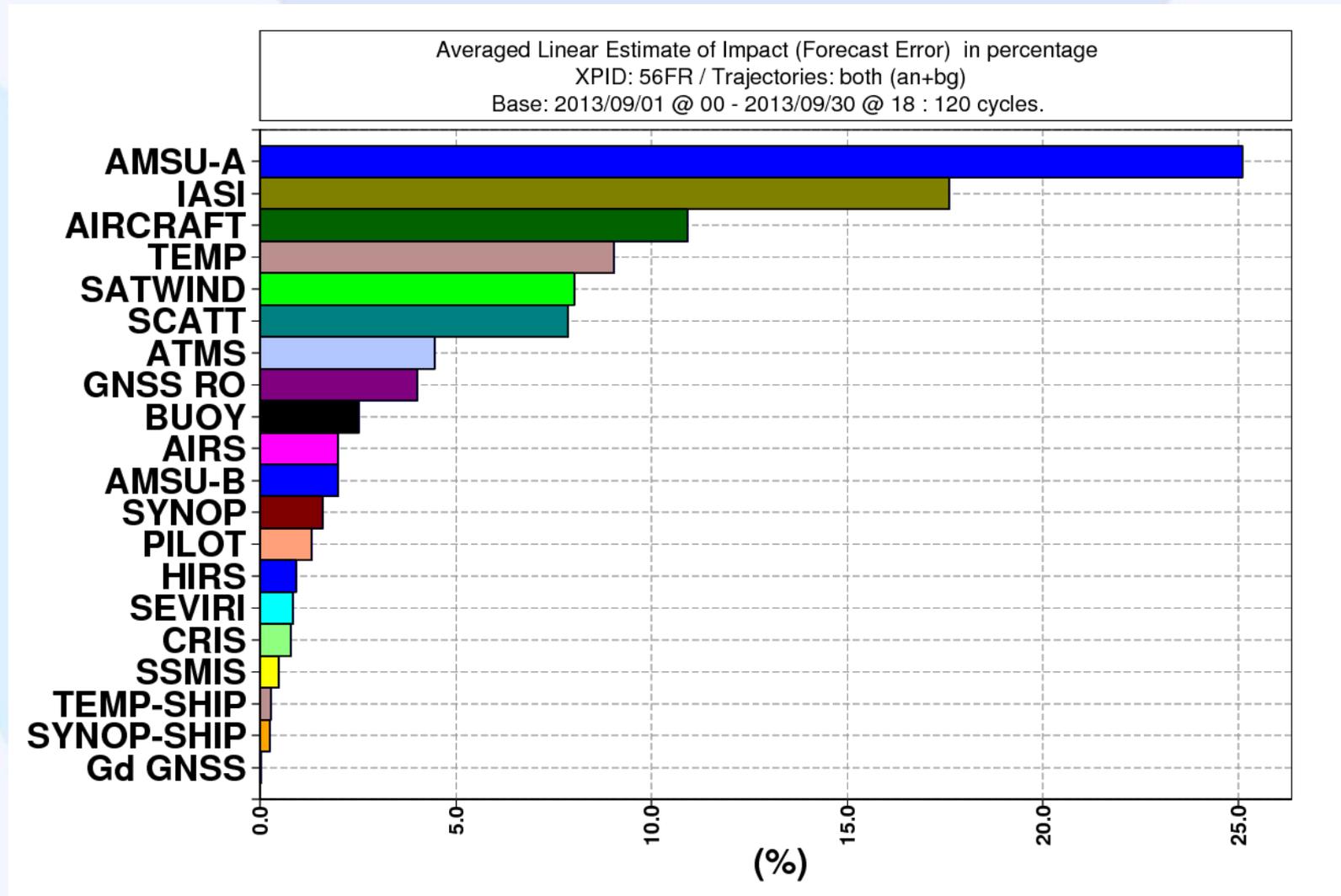
IASI  
Worse

# Global NWP Data Impacts: NRL

NAVGEM Observation Sensitivity



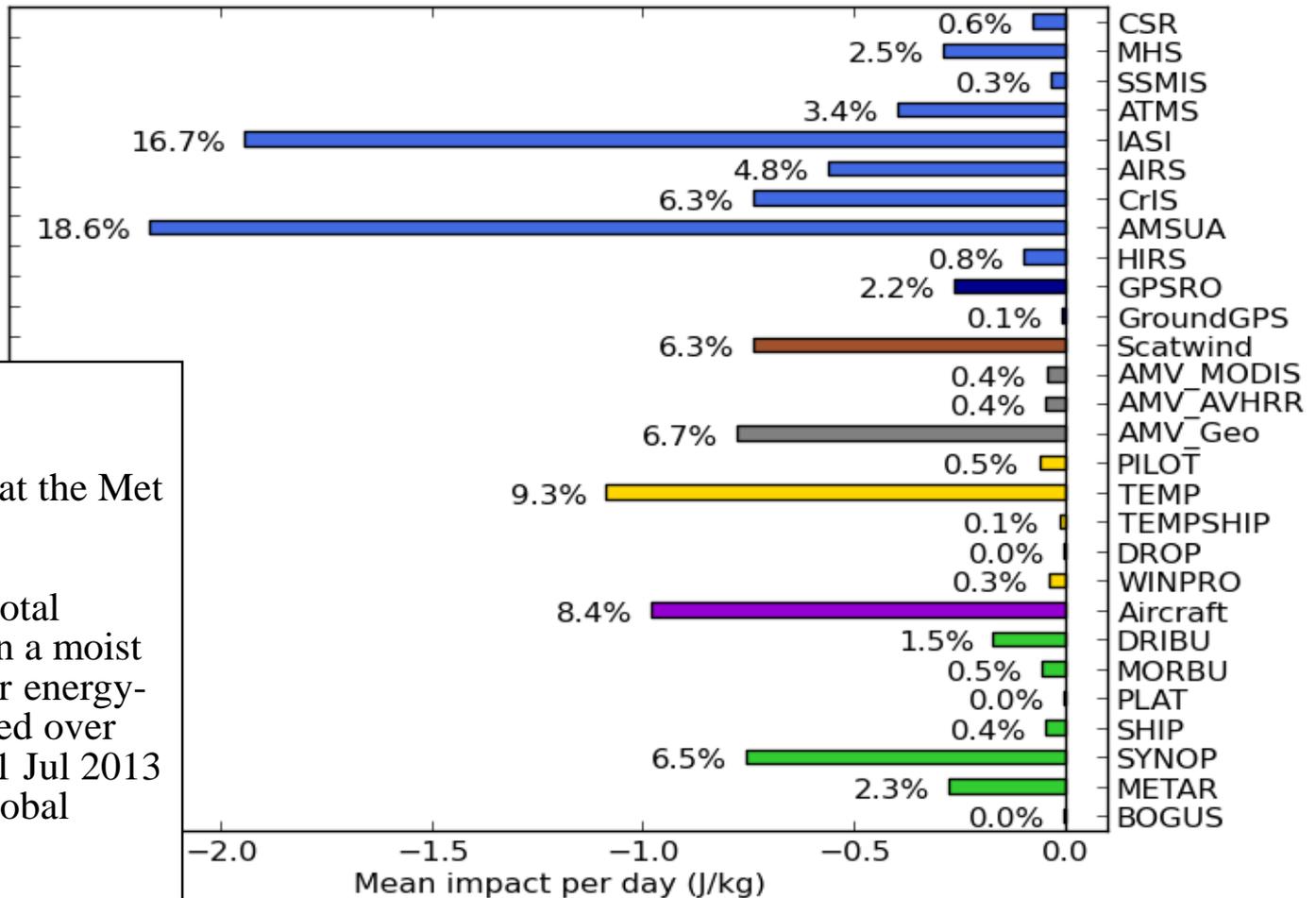
# Global NWP Data Impacts: Météo-France



# Global NWP Data Impacts: Met Office

Observation impacts at the Met Office – FSO results

Contributions to the total observation impact on a moist 24-hour forecast-error energy-norm per day, averaged over the period 01 Apr - 31 Jul 2013 for the Met Office Global Model



# Conclusions

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# Conclusions

- Since the launch of AIRS in 2002, hyperspectral infrared sounders have become some of the leading contributors to forecast accuracy.
- Most NWP centres assimilate a large number of  $15\mu\text{m}$   $\text{CO}_2$  channels with a smaller number of humidity and ozone channels.
- Areas of study include:
  - Use of cloudy radiances
  - Efficient use of information (channel selection, PCs, Reconstructed Radiances)
  - Specification of observation errors (including correlations).

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