

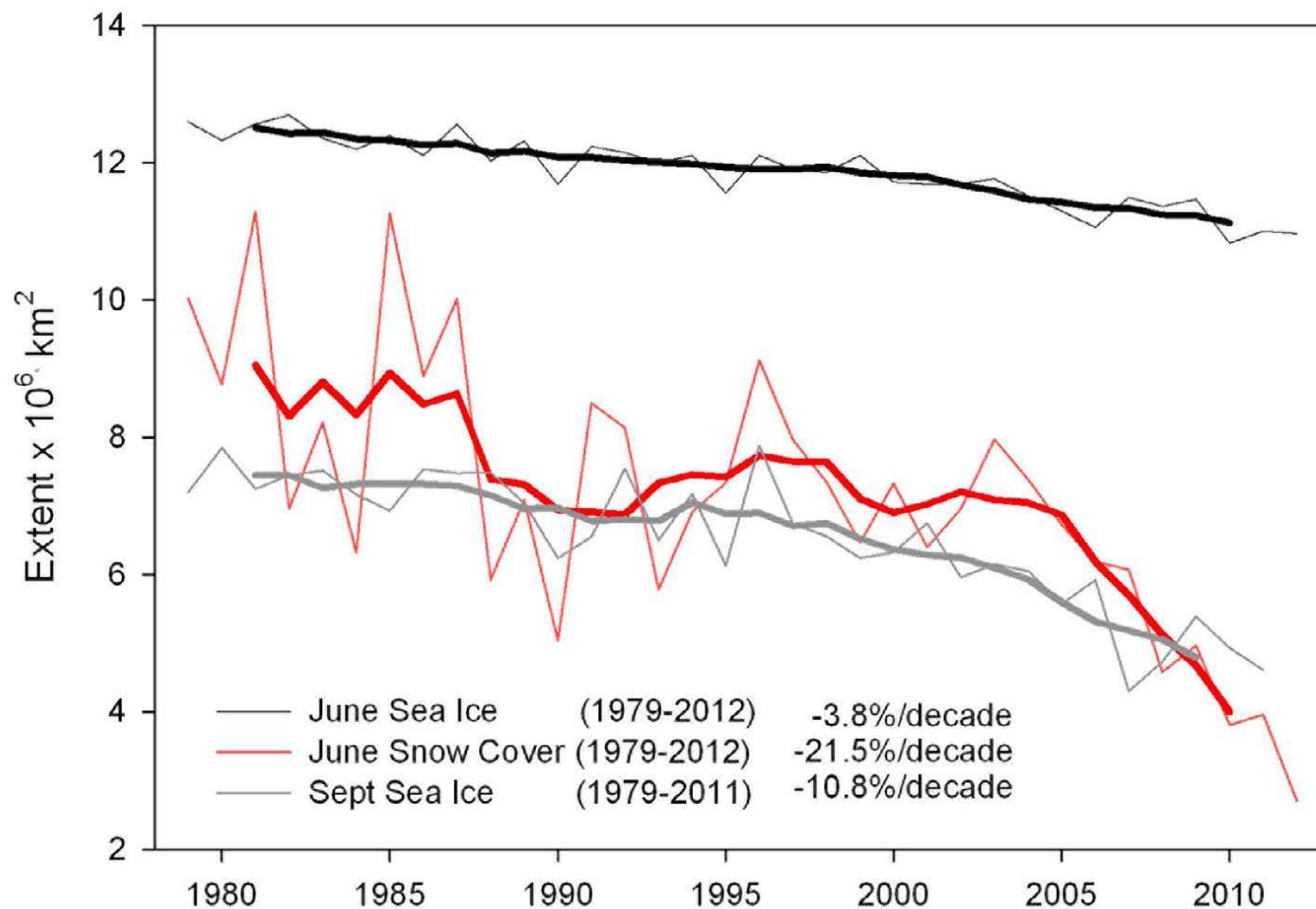


Snowpack modelling and data assimilation

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ECMWF-WWRP/THORPEX workshop on polar prediction

Northern Hemisphere ice and snow extents



Derkzen and Brown, 2012. *GRL*, **39**, doi:10.1029/2012GL053387

Snowpack model state variables

N_{snow} model layers, each with state variables:

m_j mass (kg m^{-2})

H_j heat content (J m^{-2}) → temperature, liquid water content

ρ_j density (kg m^{-3})

...

Bulk variables:

snow mass $M = \sum_{j=1}^{N_{snow}} m_j$, snow depth $d = \sum_{j=1}^{N_{snow}} m_j / \rho_j$

	model	layers	variables / layer
	UKMO	1	1
	ISBA-FR	1	2
Surfex	ISBA-ES	3	3
	Crocus	≤ 50	6

Prognostic equations for state variables

Mass balance:

$$\frac{dM}{dt} = S_f - E - R$$

Liquid water drainage:

$$\frac{dw_j}{dt} = Q_{j-1} - Q_j + f_j$$

Energy balance:

$$C \frac{dT_j}{dt} = G_{j-1} - G_j - Lf_j$$

Parametrizations required for all flux and capacity terms

Parametrizations of snow compaction

Constant density (e.g. MOSES, SiB):

$$\rho = \text{constant}$$

Bulk compaction (e.g. CLASS, ISBA-FR, TESSEL):

$$\frac{d\rho}{dt} = \tau^{-1}(\rho_{\max} - \rho)$$

Overburden profile (e.g. CLM, ISBA-ES, HTESSEL, JULES):

$$\frac{1}{\rho} \frac{\partial \rho}{\partial t} = \frac{M(z)g}{\eta(\rho, T)}$$

Experimental snow data assimilation

Ensemble Kalman Filter

Andreadis and Lettenmaier (2006)

Dechant and Moradkhani (2011)

Durand and Margulis (2006)

Durand et al. (2009)

Kumar et al. (2008)

Slater and Clark (2006)

Su et al. (2008)

Toure et al. (2011)

Extended Kalman Filter

Dong et al. (2007)

Sun et al. (2004)

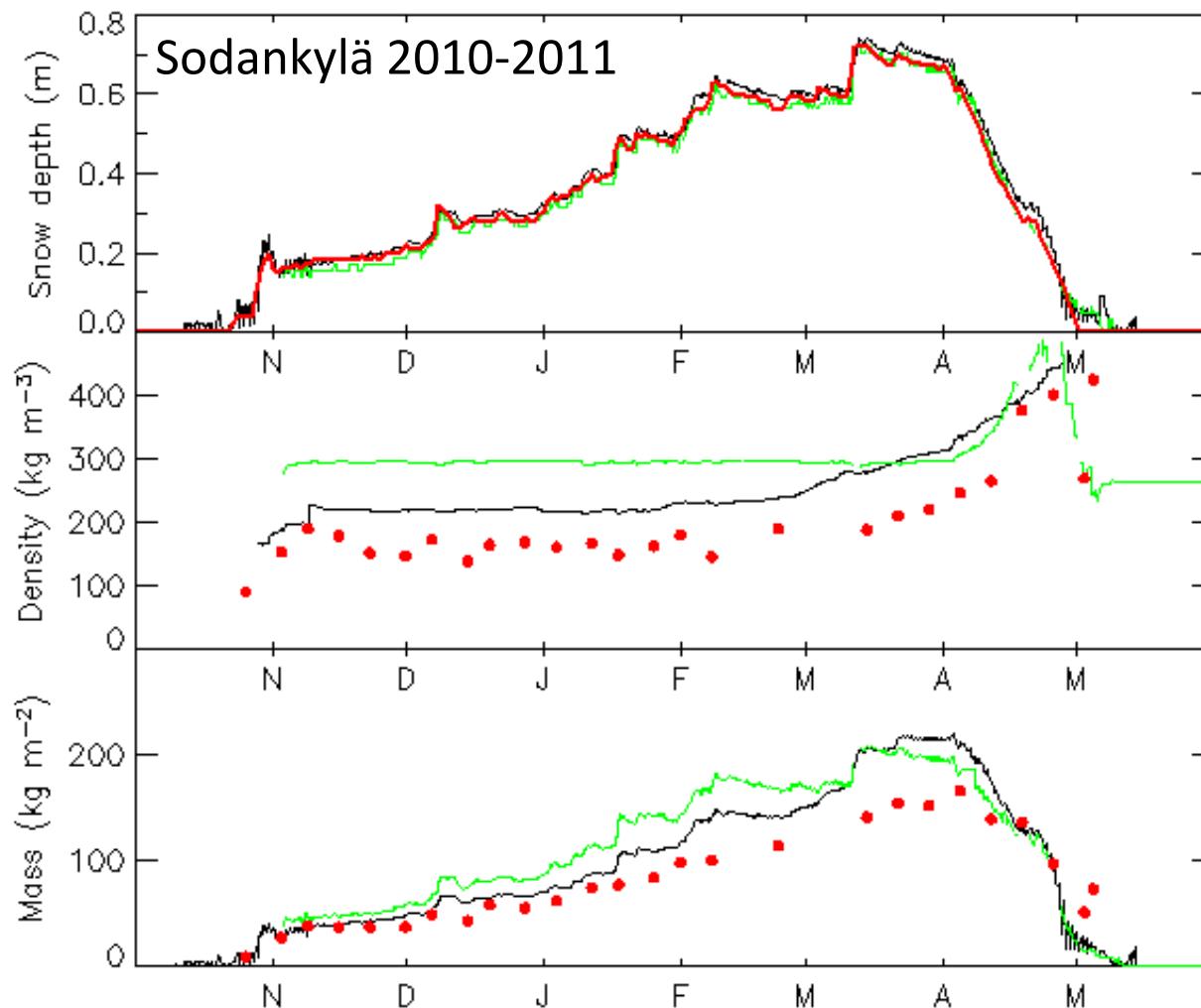
Particle Filter

Dechant and Moradkhani (2011)

Operational snow analyses

Model	Observations	Assimilation	Operational
CMC	SYNOP	OI	1999
ECMWF	SYNOP	Cressman	1987
	IMS	Cressman	2004
		OI	2010
HARMONIE	SYNOP	OI	2010
HIRLAM	SYNOP	Cressman	1995
	SYNOP	OI	2004
	Globsnow	OI	Experimental
Met Office	IMS	Update	2009

Operational snow analyses



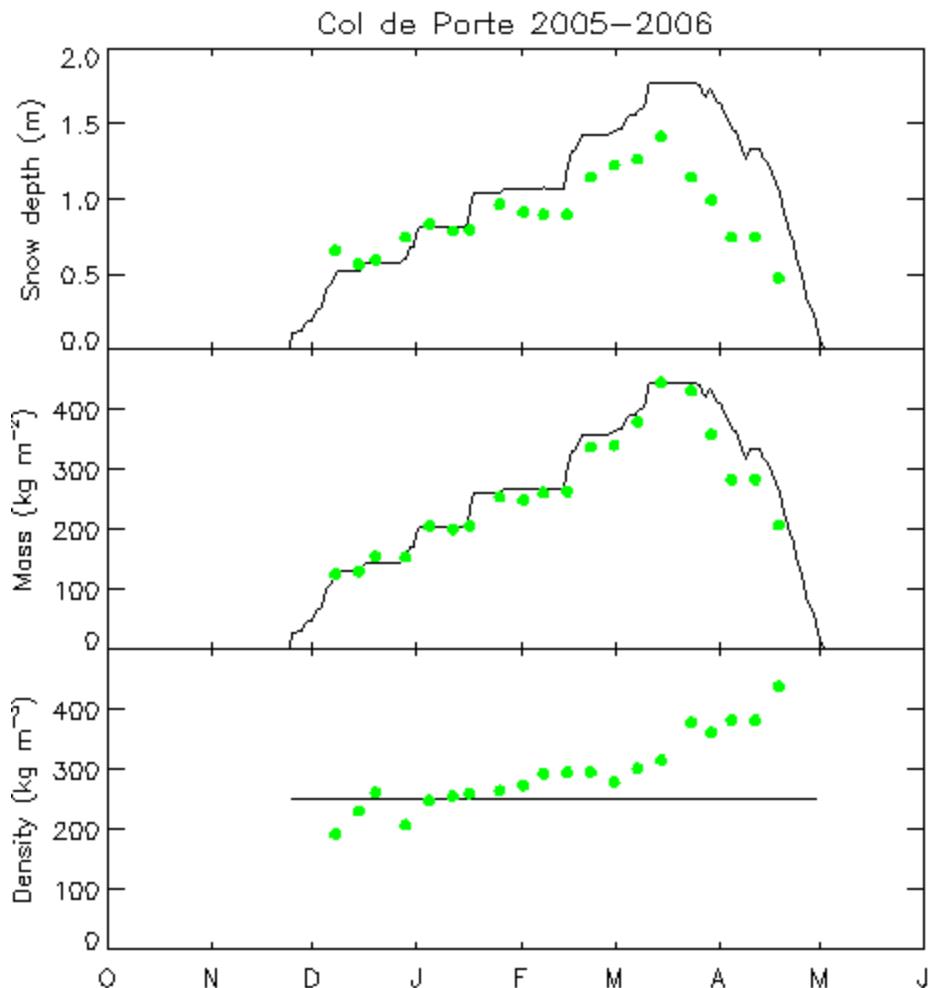
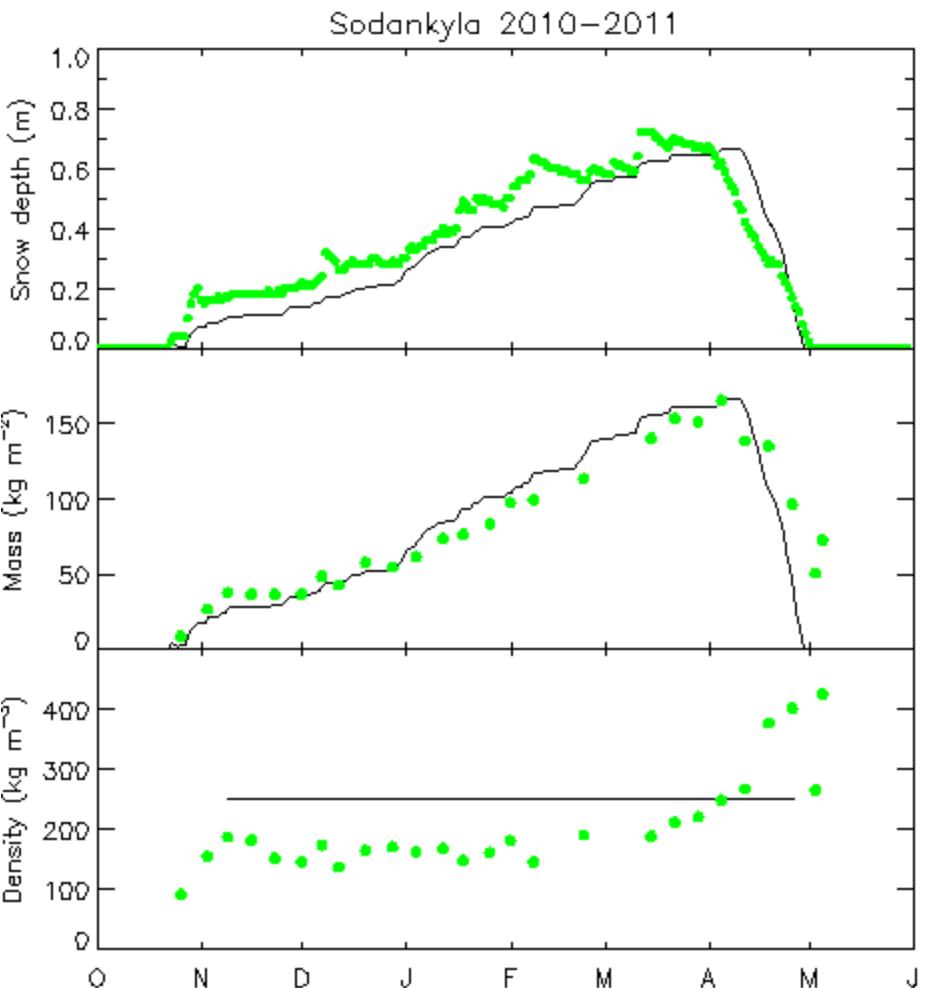
SYNOP snow depths and FMI snow pits (from Timo Ryyppö)

Hirlam snow analyses (from Laura Rontu)

ECMWF snow analyses (from Patricia de Rosnay)

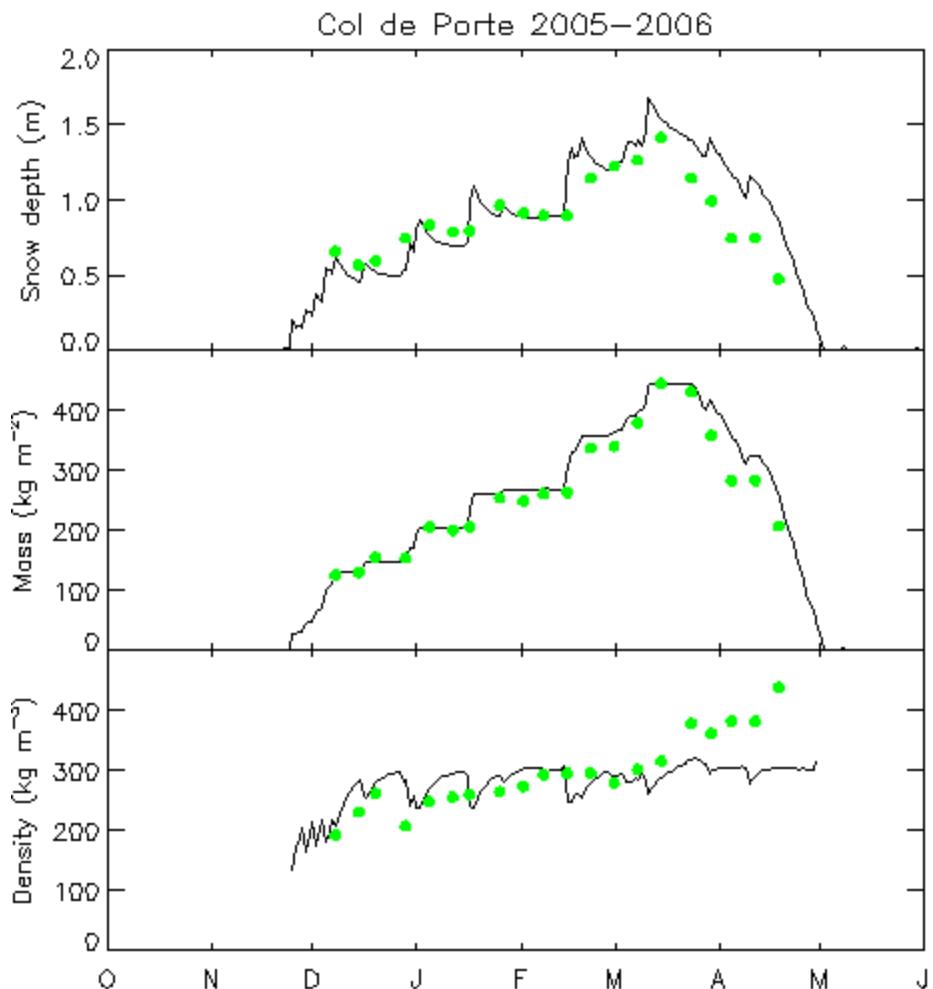
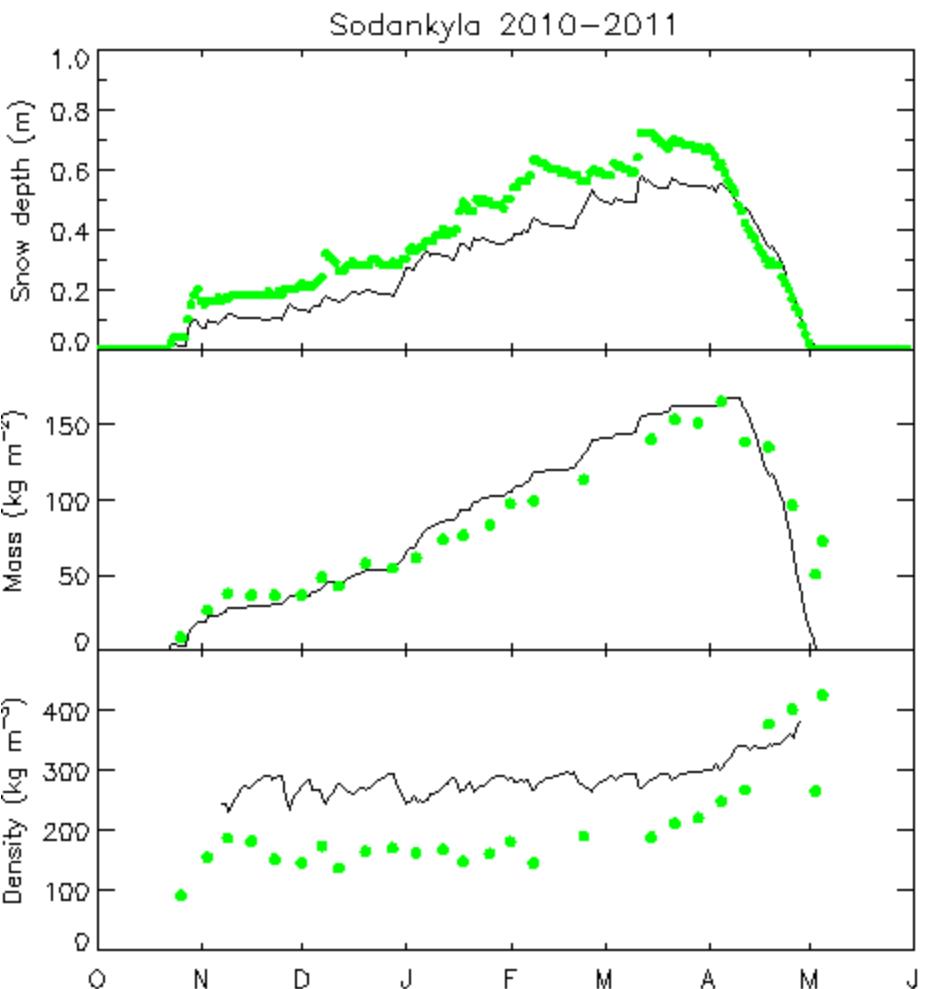
Open loop simulations with in situ driving data

Constant density



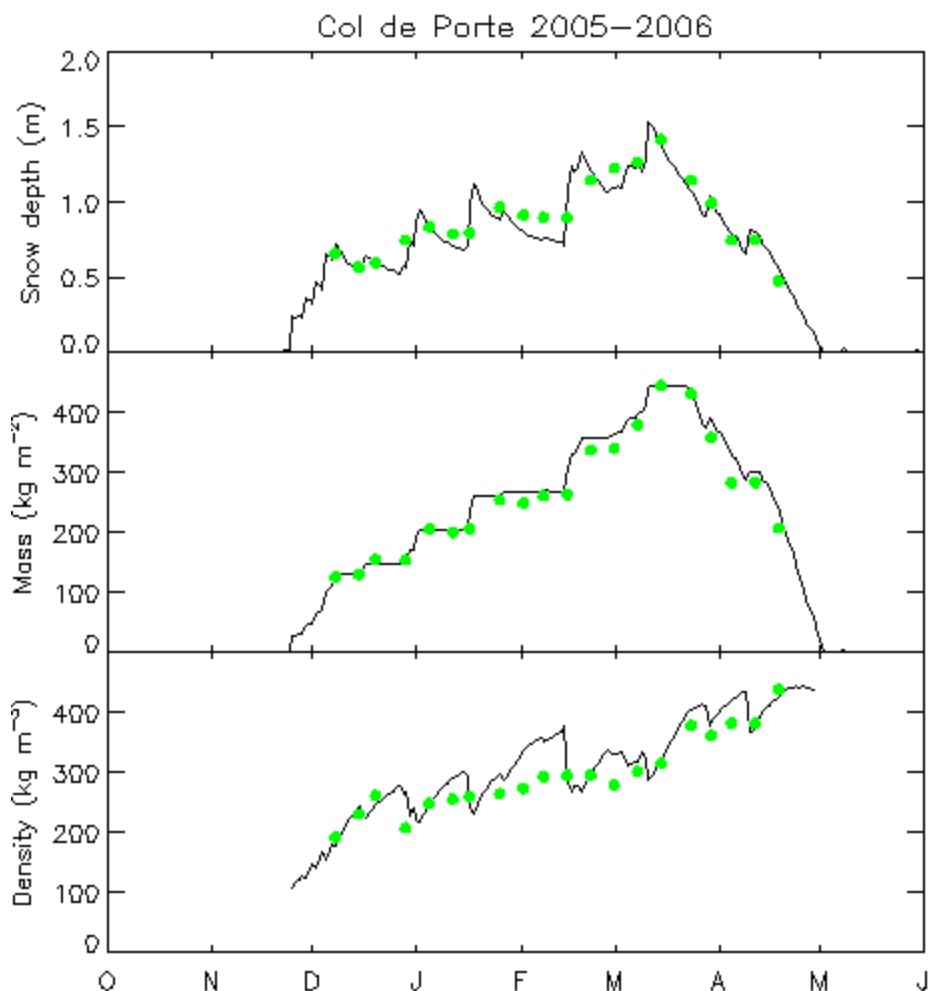
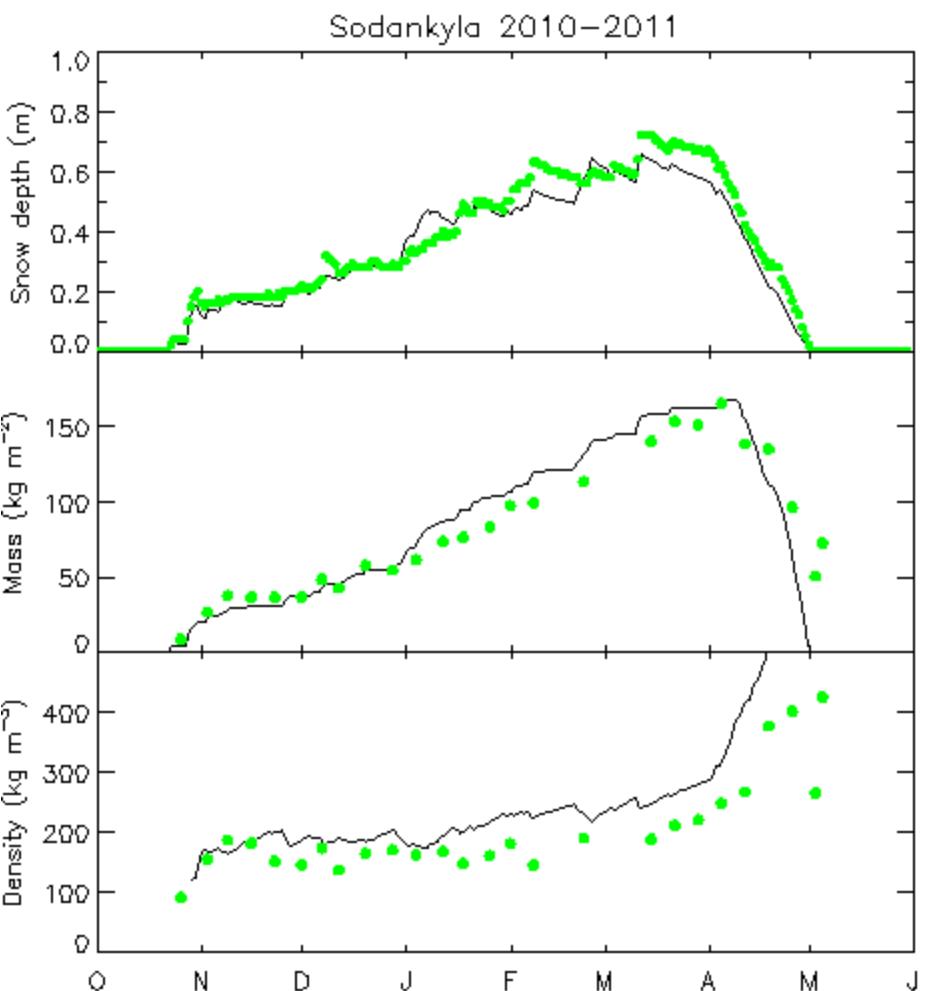
Open loop simulations with in situ driving data

Bulk density



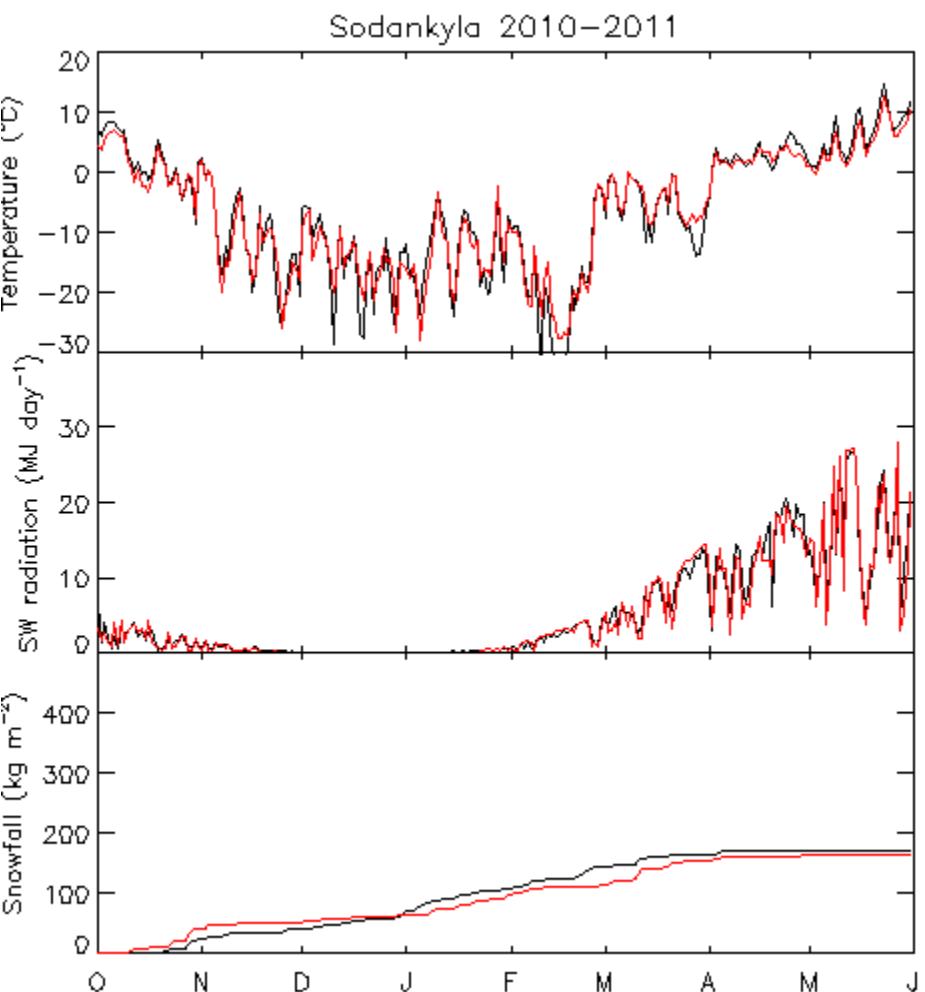
Open loop simulations with in situ driving data

Density profile

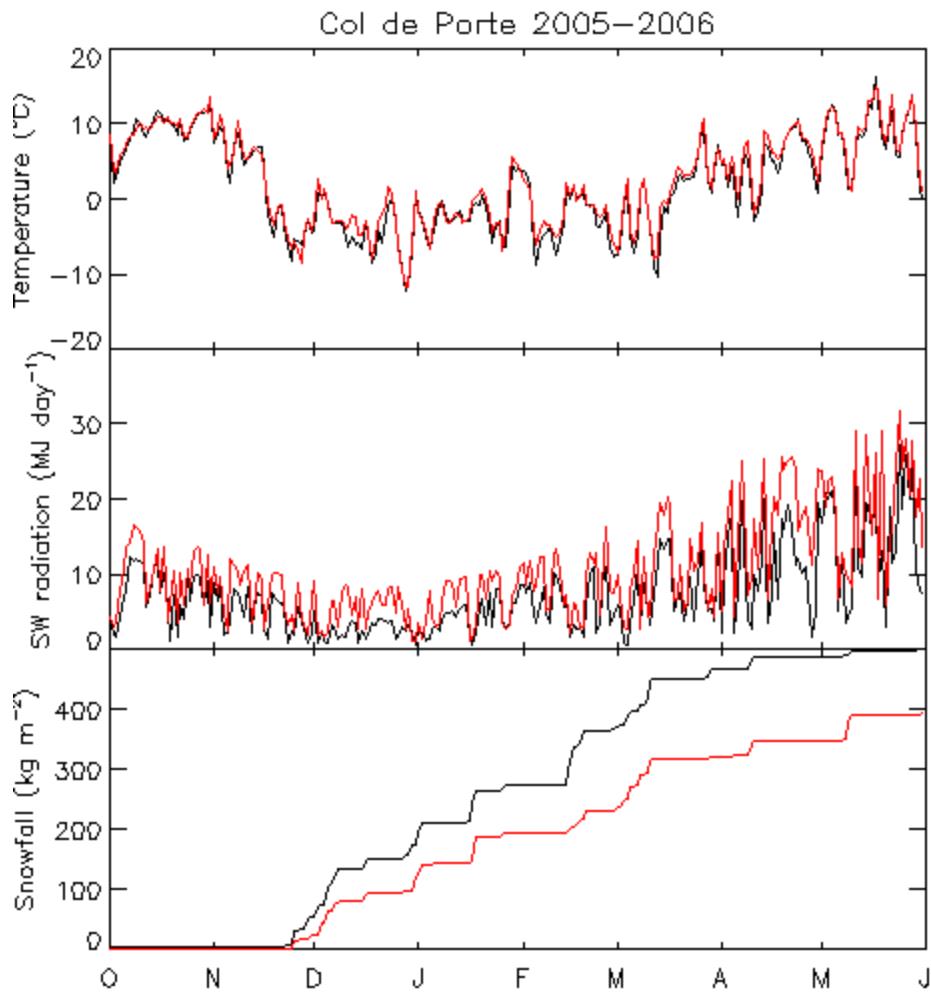


In situ and NWP driving data

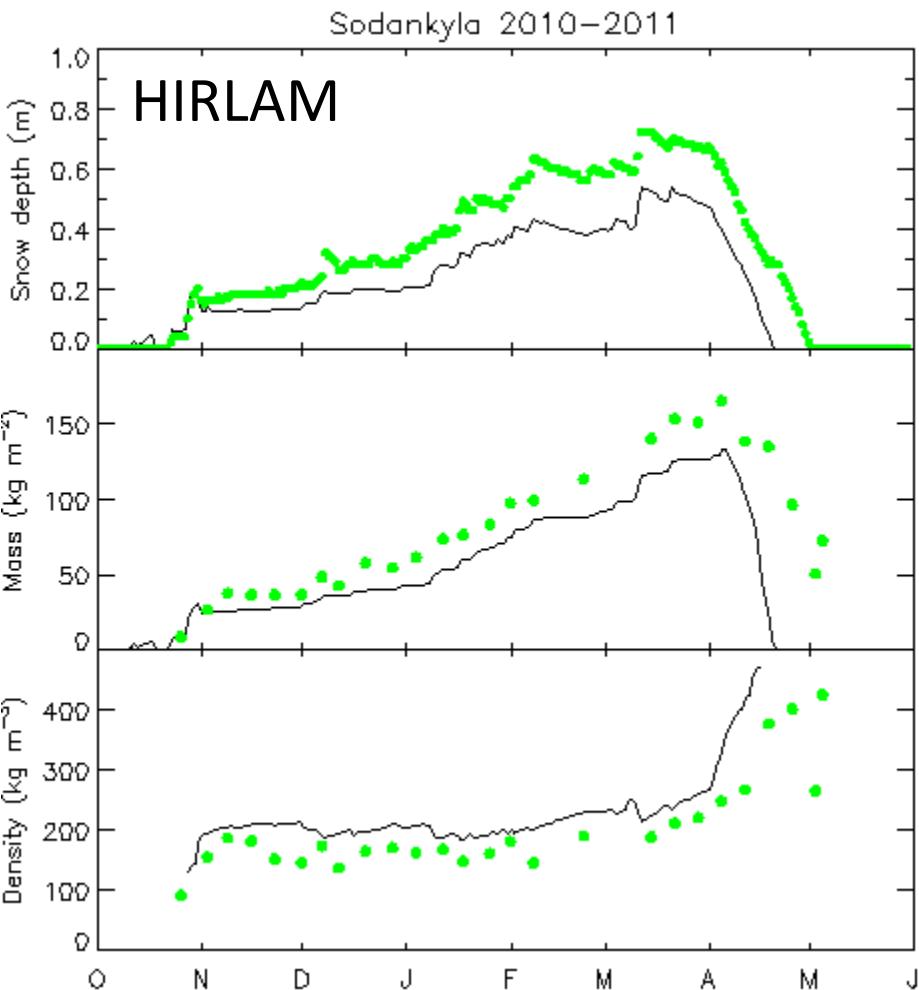
HIRLAM



SAFRAN



Open loop simulations with NWP driving data



rms errors

snow depth

snow mass

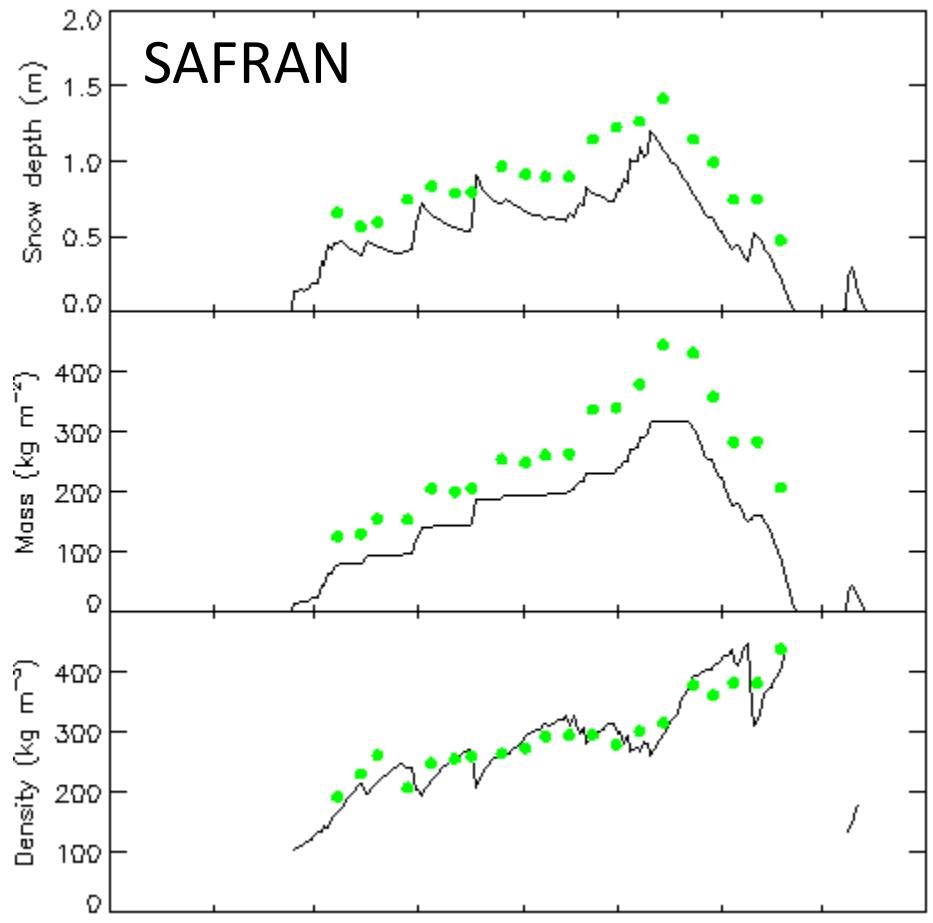
Sodankylä

13 cm

13 kg m^{-2}

Col de Porte 2005–2006

SAFRAN



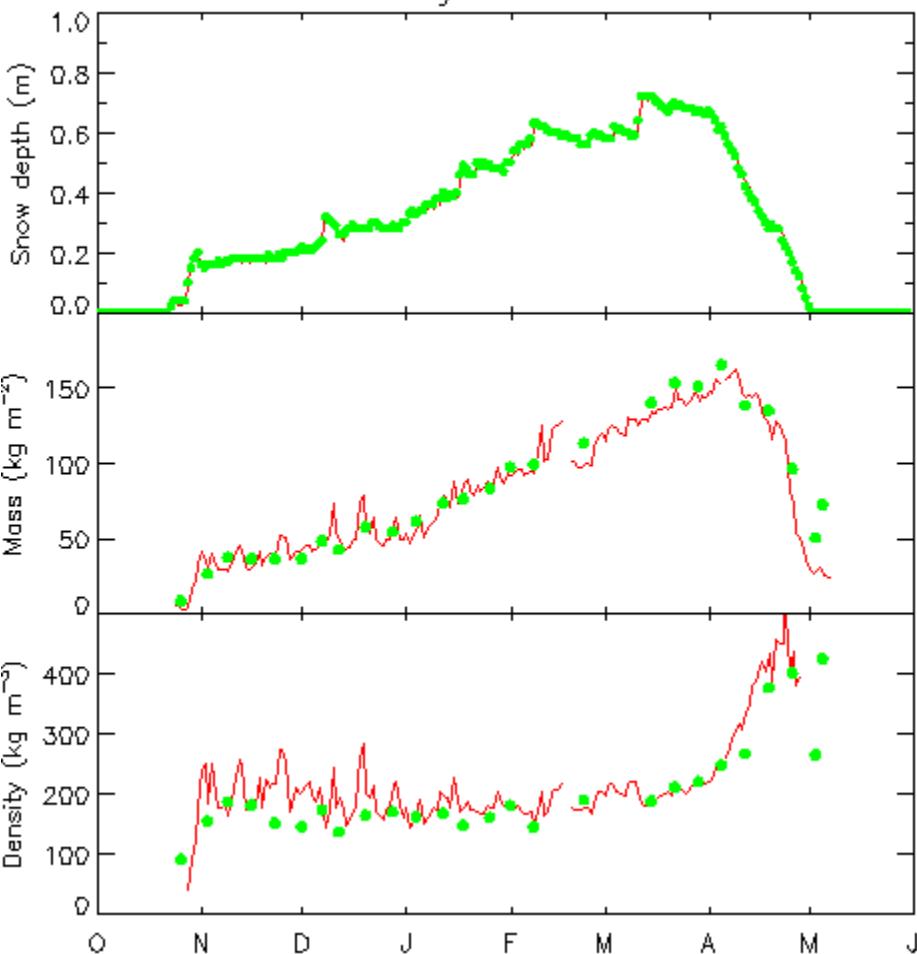
Col de Porte

28 cm

86 kg m^{-2}

Independent snow measurements

Sodankylä 2010–2011



rms differences

snow depth

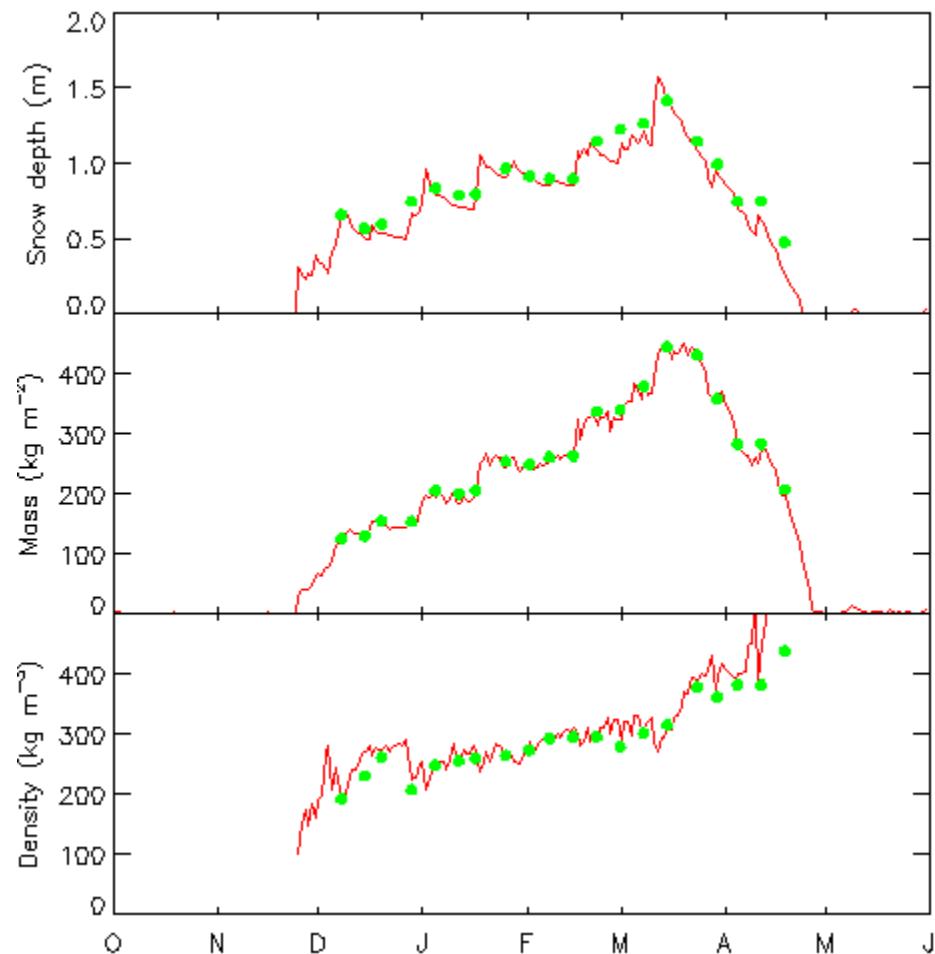
snow mass

Sodankylä

0.7 cm

4.2 kg m^{-2}

Col de Porte 2005–2006



Col de Porte

7.6 cm

8.8 kg m^{-2}

Data assimilation

Model state vector \mathbf{x}_k and inputs \mathbf{u}_k at time k

Model operator $\mathbf{x}_{k+1} = \mathbf{f}(\mathbf{x}_k, \mathbf{u}_k)$

Background state \mathbf{x}_b estimates true state \mathbf{x}_t , error covariance matrix

$$\mathbf{B} = \overline{(\mathbf{x}_b - \mathbf{x}_t)^T (\mathbf{x}_b - \mathbf{x}_t)}$$

Model error covariance for forecast from true initial state

$$\mathbf{Q} = \overline{[\mathbf{f}(\mathbf{x}_{t,k}) - \mathbf{x}_{t,k+1}]^T [\mathbf{f}(\mathbf{x}_{t,k}) - \mathbf{x}_{t,k+1}]}$$

Observation operator $\mathbf{y} = \mathbf{h}(\mathbf{x})$

Observed value \mathbf{y}_o estimates true value \mathbf{y}_t , error covariance matrix

$$\mathbf{R} = \overline{(\mathbf{y}_o - \mathbf{y}_t)^T (\mathbf{y}_o - \mathbf{y}_t)}$$

Find analysed state \mathbf{x}_a to minimize quadratic cost function

$$J(\mathbf{x}_a) = (\mathbf{x}_a - \mathbf{x}_b)^T \mathbf{B}^{-1} (\mathbf{x}_a - \mathbf{x}_b) + [\mathbf{y}_o - \mathbf{h}(\mathbf{x}_a)]^T \mathbf{R}^{-1} [\mathbf{y}_o - \mathbf{h}(\mathbf{x}_a)]$$

Analysis equation

Find \mathbf{x}_a such that

$$\frac{\partial J}{\partial \mathbf{x}} = 0$$

$$\Rightarrow \mathbf{x}_a = \mathbf{x}_b + \mathbf{B}\mathbf{H}^T(\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1}[\mathbf{y}_o - \mathbf{h}(\mathbf{x}_b)]$$

with Jacobian $\mathbf{H} = \frac{\partial \mathbf{h}}{\partial \mathbf{x}}$

Extended Kalman Filter (EKF):

Background error forecast $\mathbf{B}_{k+1} = \mathbf{F}\mathbf{B}_k\mathbf{F}^T + \mathbf{Q}$

and analysis $\mathbf{B}_a = \mathbf{B}_k + \mathbf{B}_k\mathbf{H}^T(\mathbf{H}\mathbf{B}_k\mathbf{H}^T + \mathbf{R})^{-1}\mathbf{H}\mathbf{B}_k$

using tangent linear model $\mathbf{F} = \frac{\partial \mathbf{f}}{\partial \mathbf{x}}$

Simplification (SEKF):

$$F_{ij} = \frac{\partial f_i}{\partial x_j} \approx \frac{f_i(\mathbf{x} + \delta x_j) - f_i(\mathbf{x})}{\delta x_j} \quad H_{ij} = \frac{\partial h_i}{\partial x_j} \approx \frac{h_i(\mathbf{x} + \delta x_j) - h_i(\mathbf{x})}{\delta x_j}$$

– control and one perturbed forecast for each state variable

Single observable state variable

$$\mathbf{x} = [x], \mathbf{y} = [x_o], \mathbf{H} = [1]$$

Analysis

$$x_a = x_b + \frac{\sigma_b^2}{\sigma_b^2 + \sigma_r^2} [x_o - x_b]$$

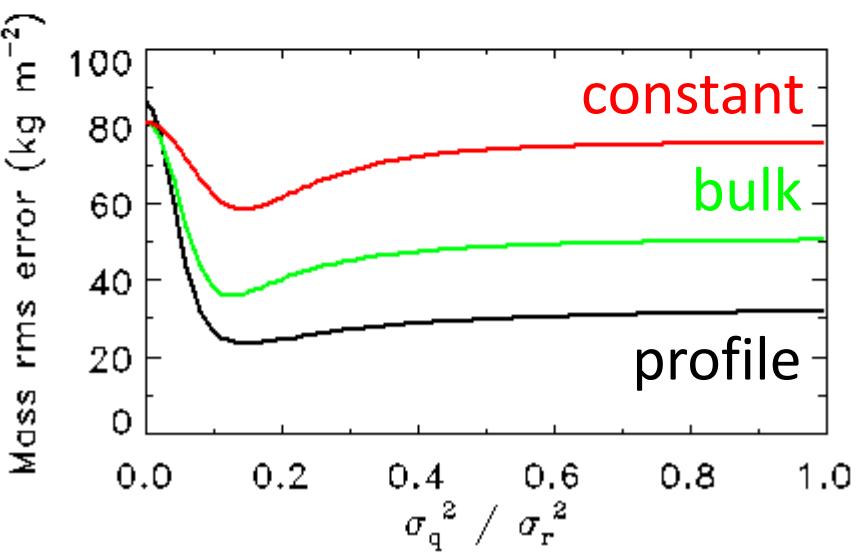
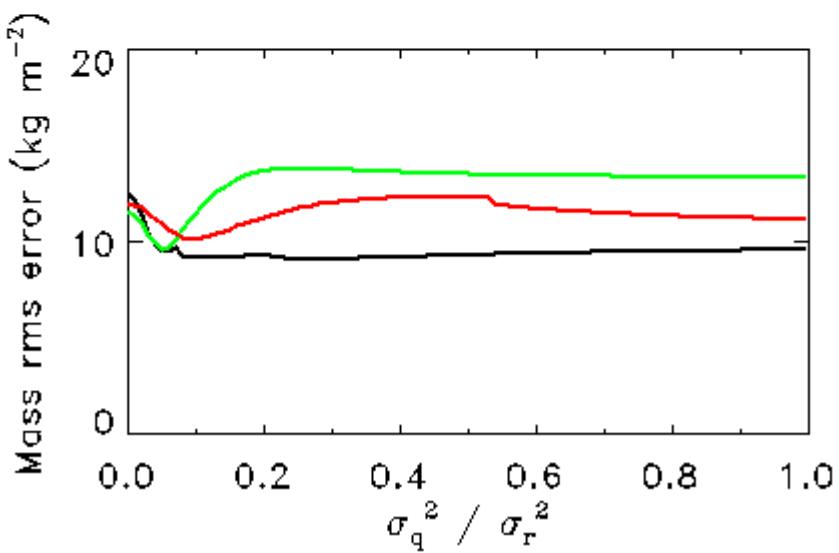
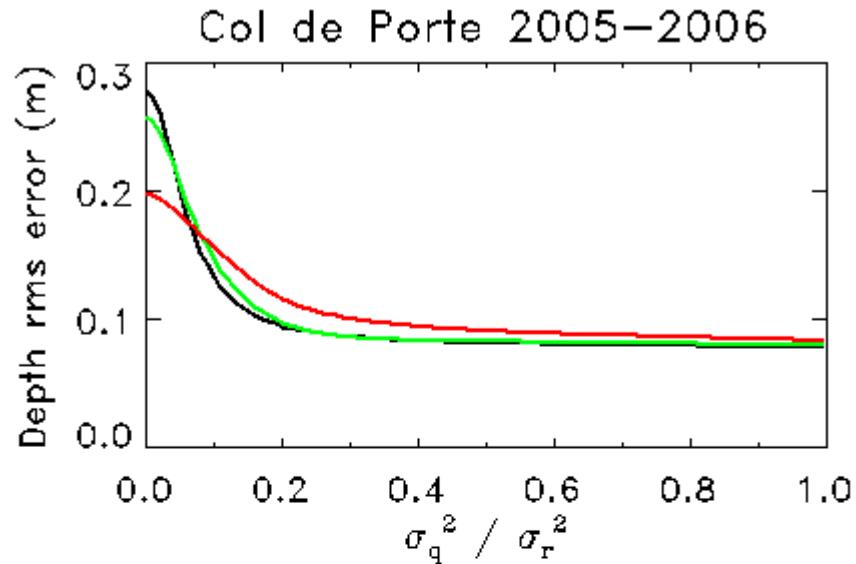
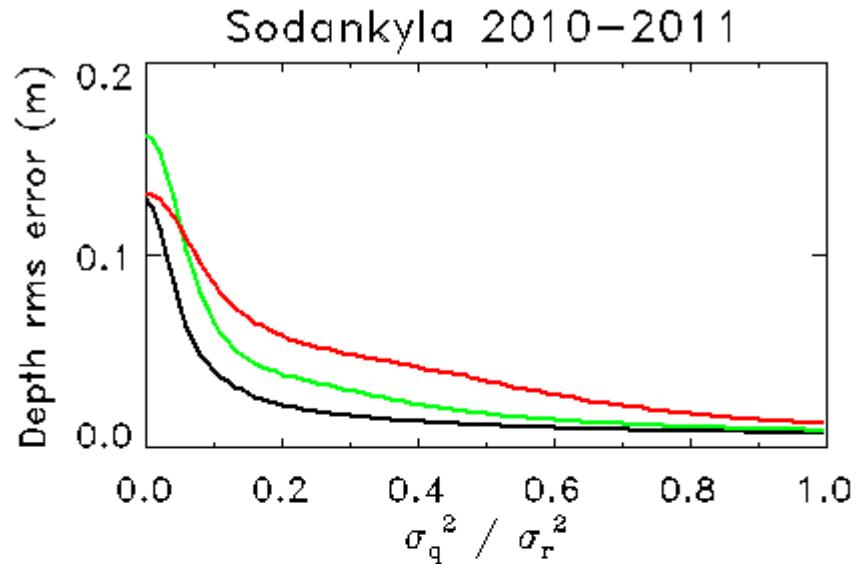
$\sigma_b^2 \gg \sigma_r^2 \Rightarrow x_a = x_o$ (direct insertion of observation)
 $\sigma_b^2 \ll \sigma_r^2 \Rightarrow x_a = x_b$ (open loop)

Background error forecast $\sigma_b^2 \rightarrow F^2 \sigma_b^2 + \sigma_q^2$

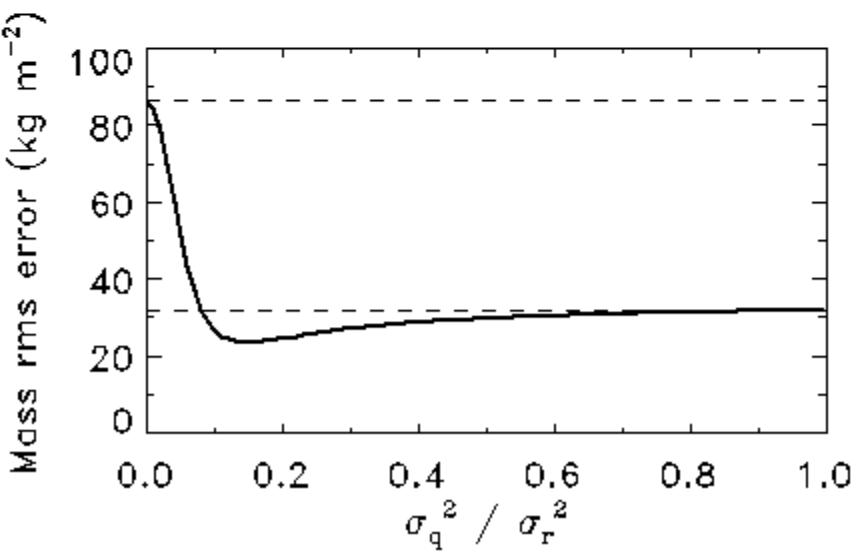
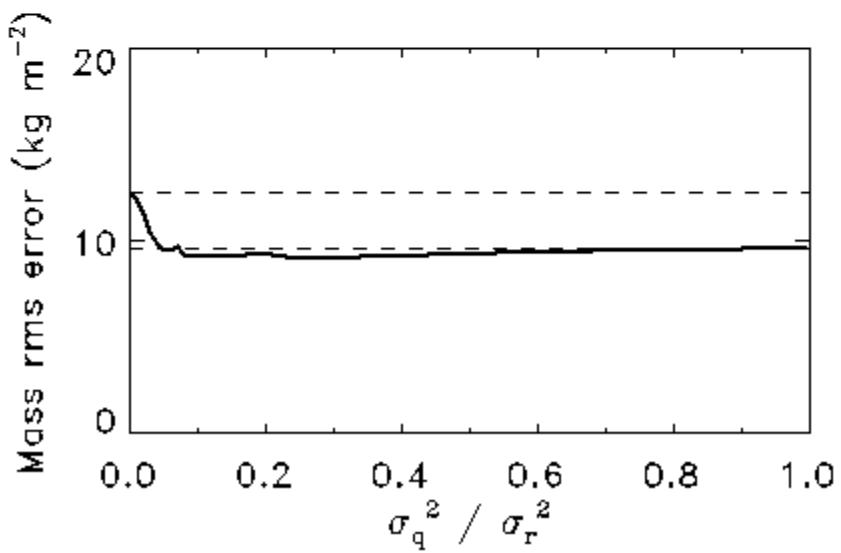
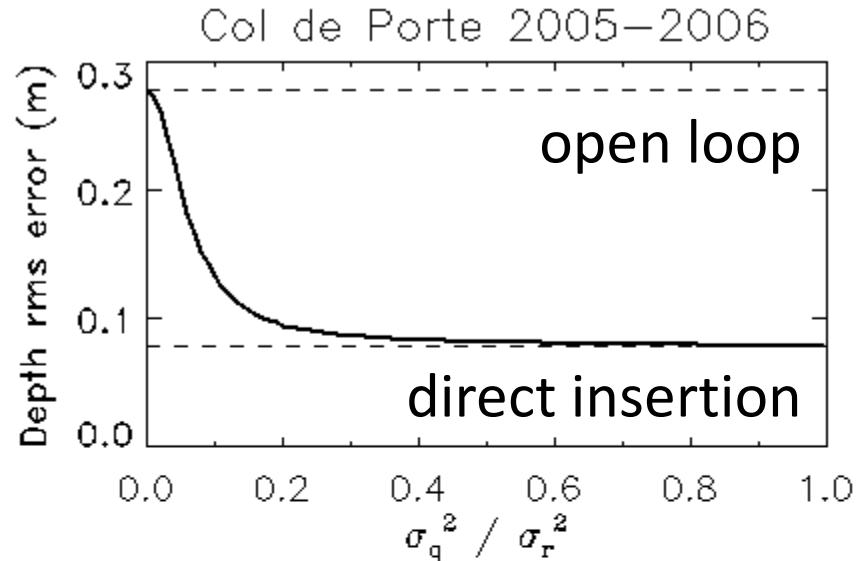
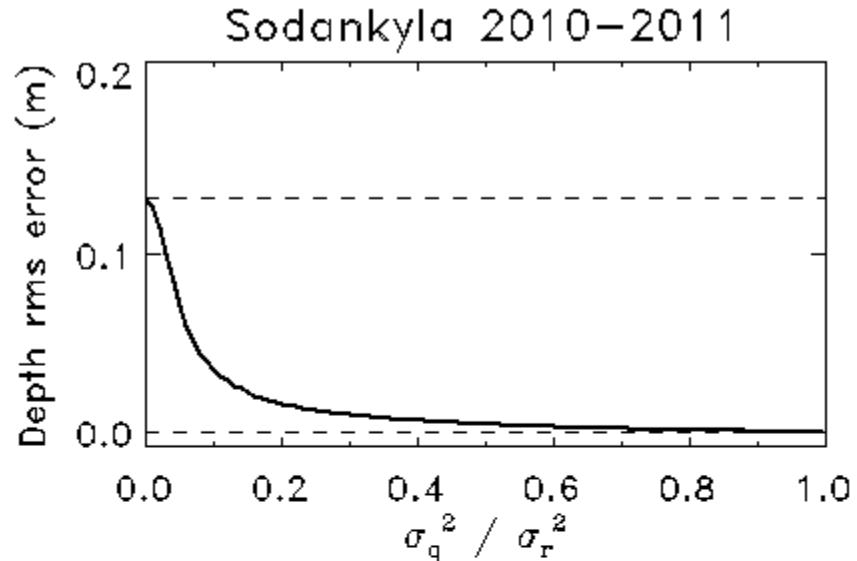
and error analysis

$$\sigma_b^2 \rightarrow \frac{\sigma_r^2}{\sigma_b^2 + \sigma_r^2} \sigma_b^2$$

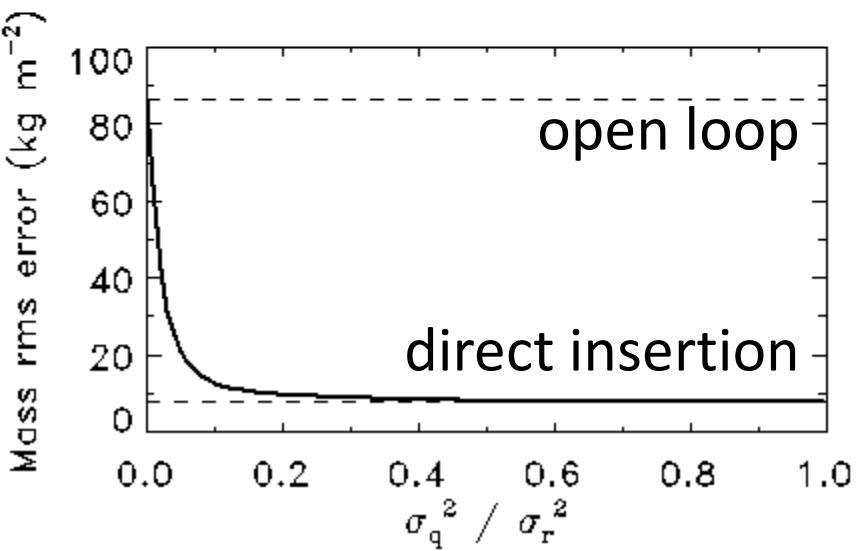
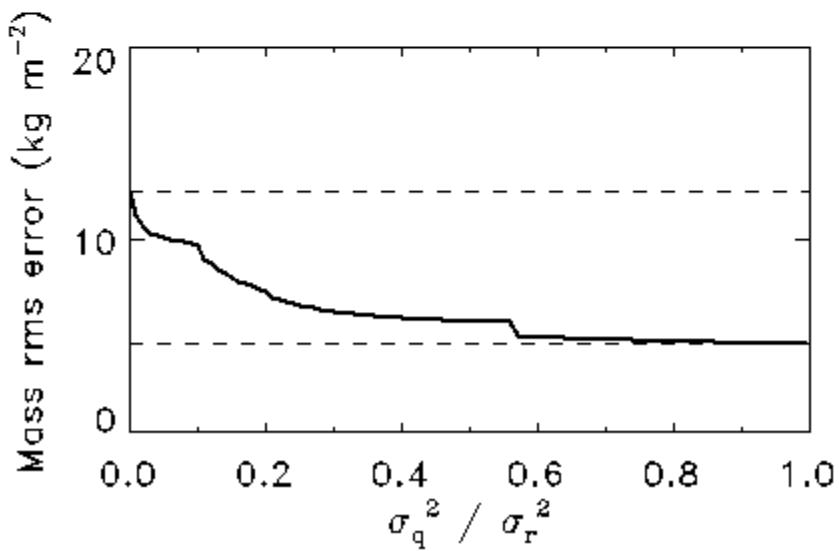
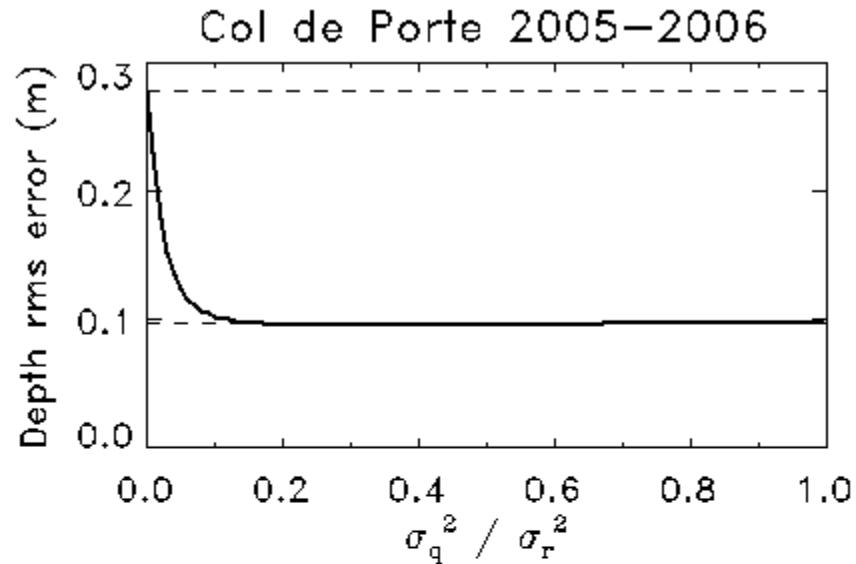
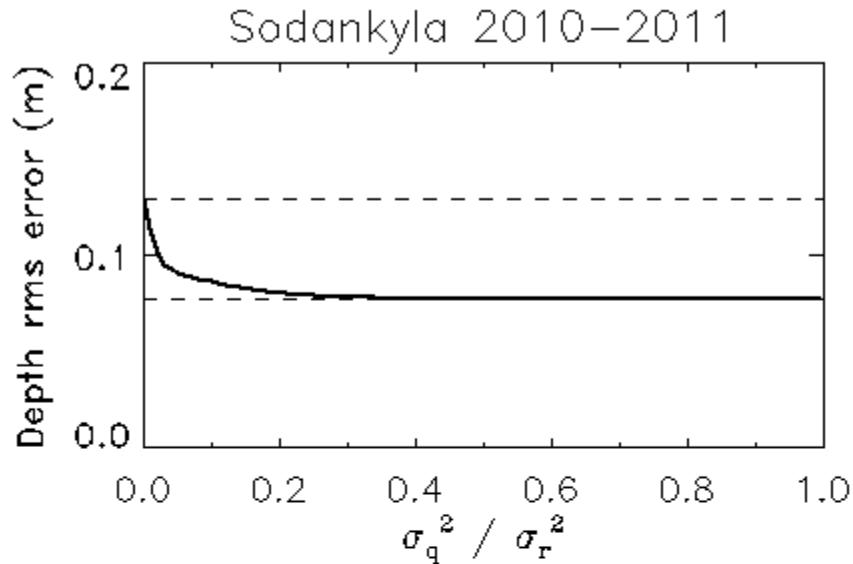
Daily assimilation of snow depth



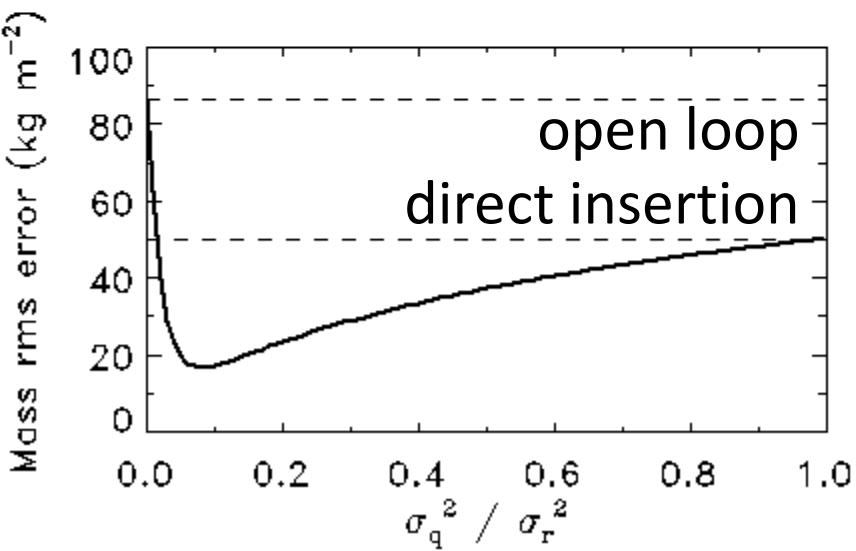
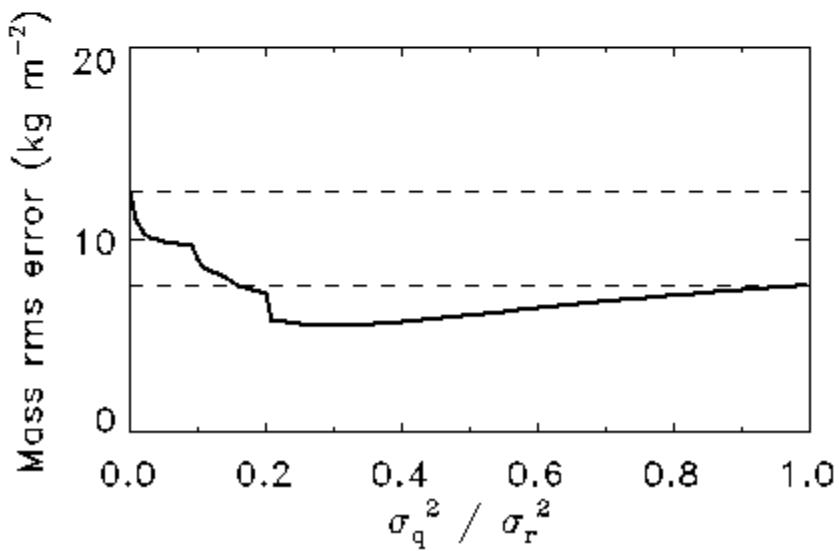
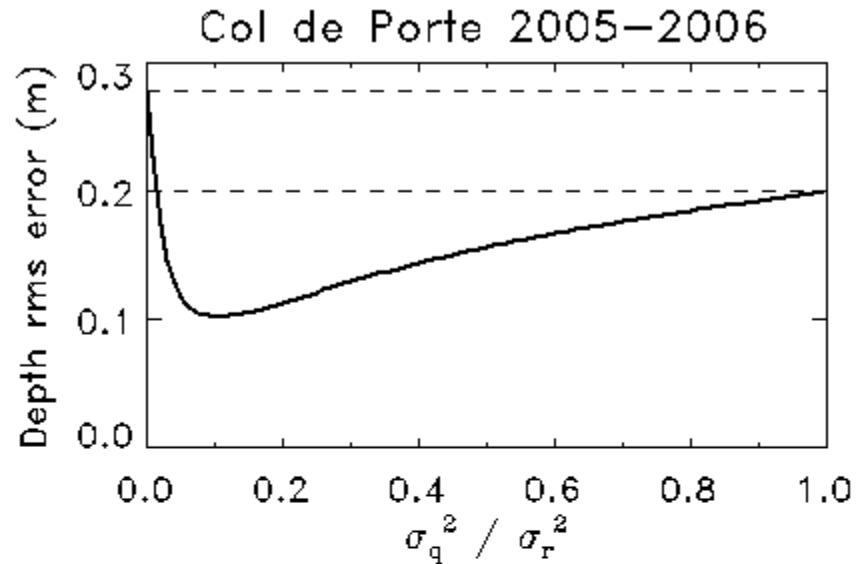
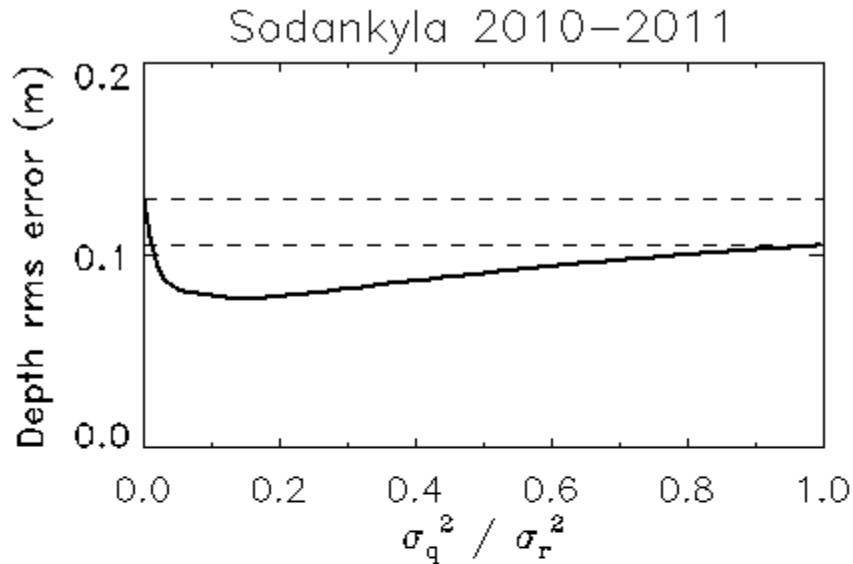
Daily assimilation of snow depth



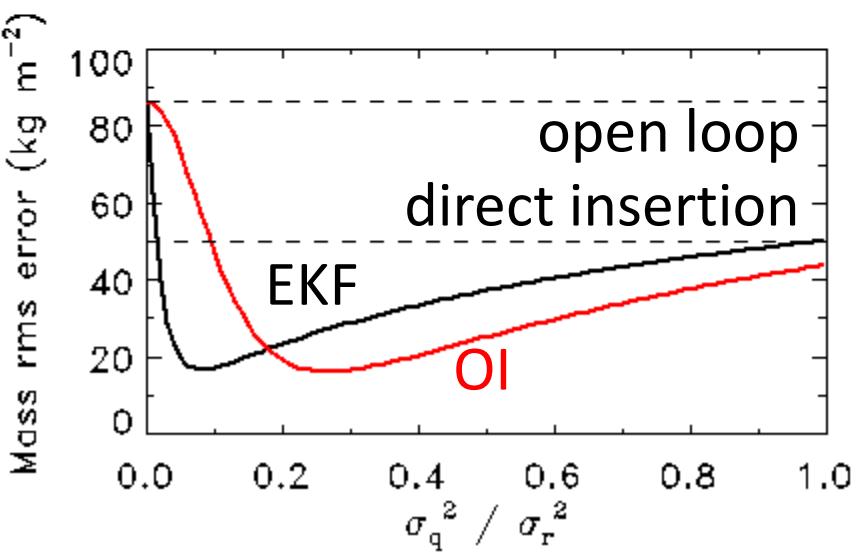
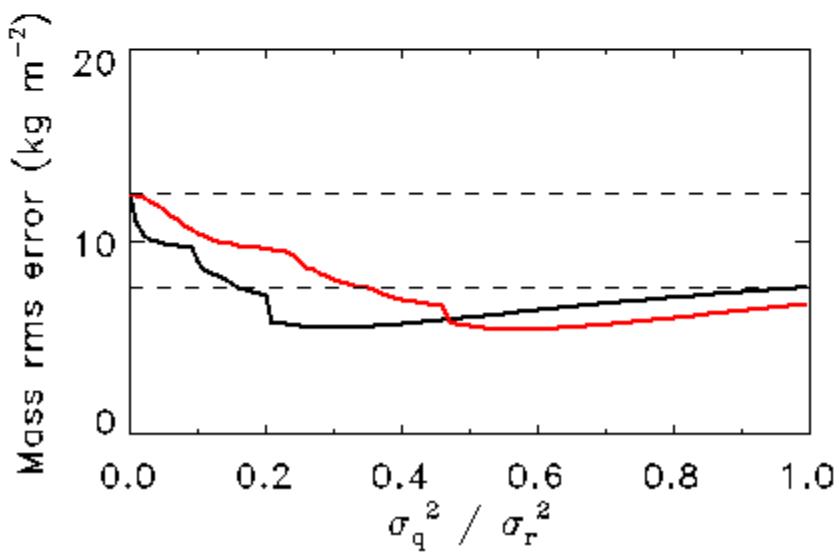
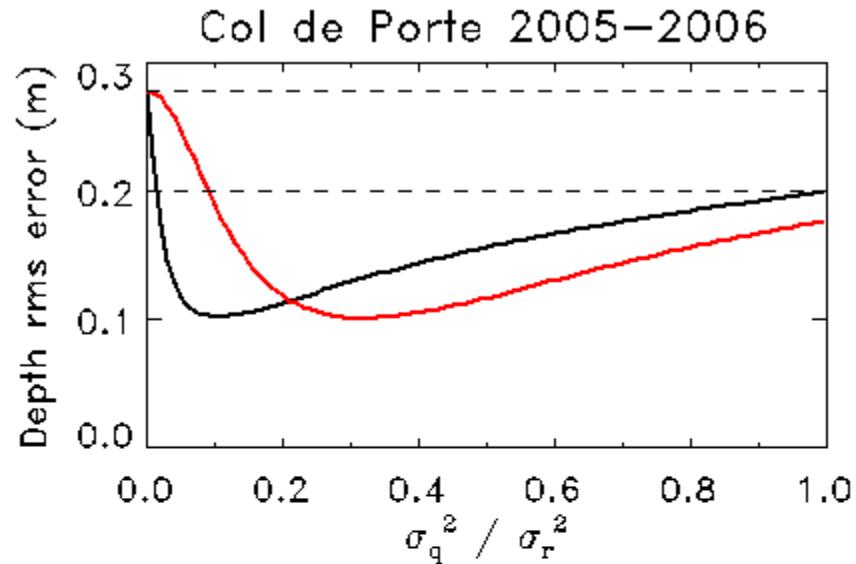
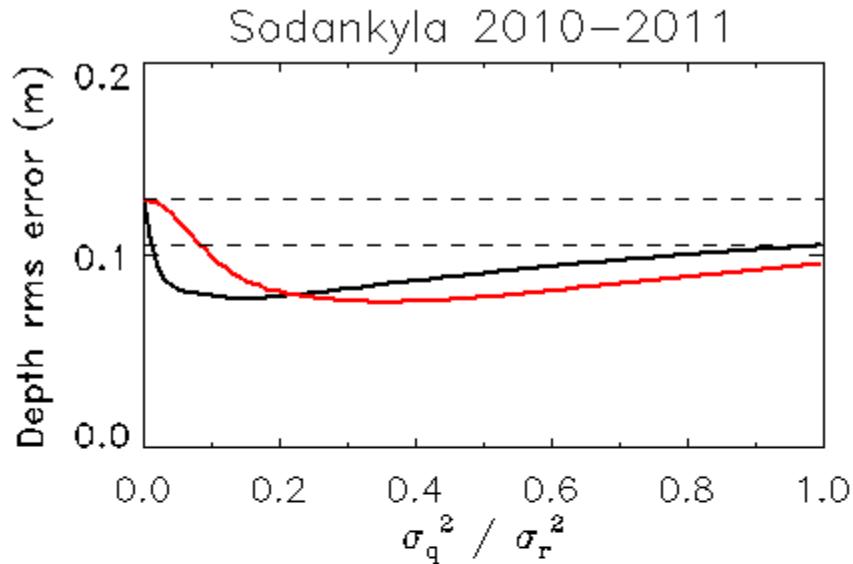
Daily assimilation of snow mass



Daily assimilation of degraded snow mass



Daily assimilation of degraded snow mass



Why is EKF no better than OI?

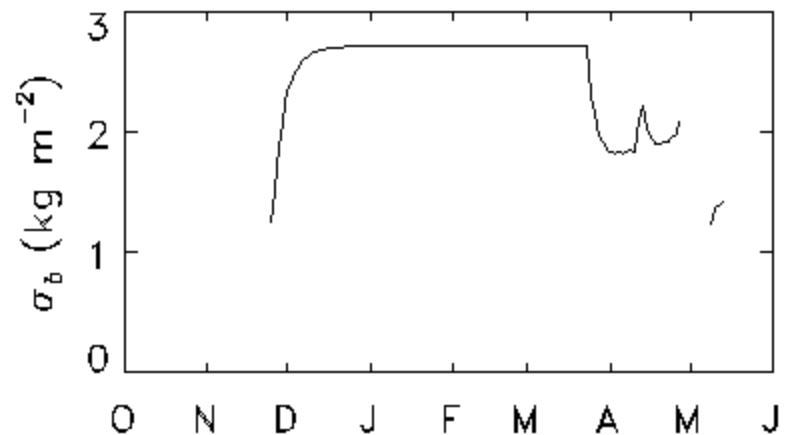
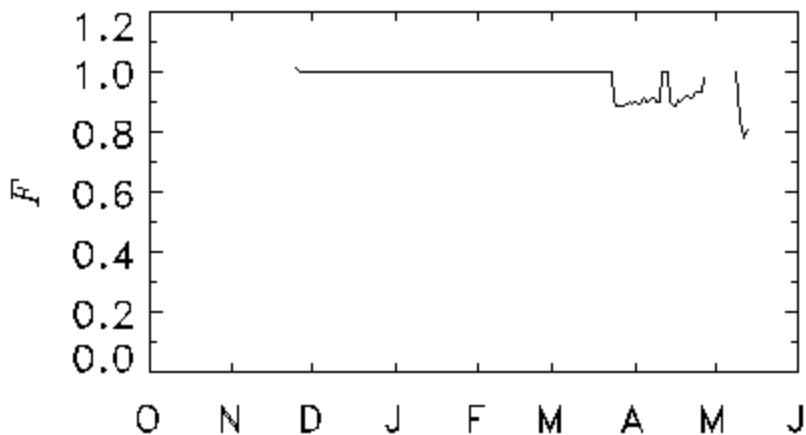
Mass balance $M_{k+1} = M_k + (S_f - E - R)\delta t$

– almost linear for deep snow

$$F = \frac{dm_{k+1}}{dm_k} \approx 1$$

$\Rightarrow \sigma_b^2 \rightarrow \text{constant}$

$\Rightarrow \text{SEKF} \approx \text{OI}$



Conclusions

Assimilation of ground-based snow data requires:

- good background estimate of snow density
- good estimates of observation and model errors
(underestimation of model / observation error ratio is worse than overestimation)
- may not require advanced data assimilation techniques