

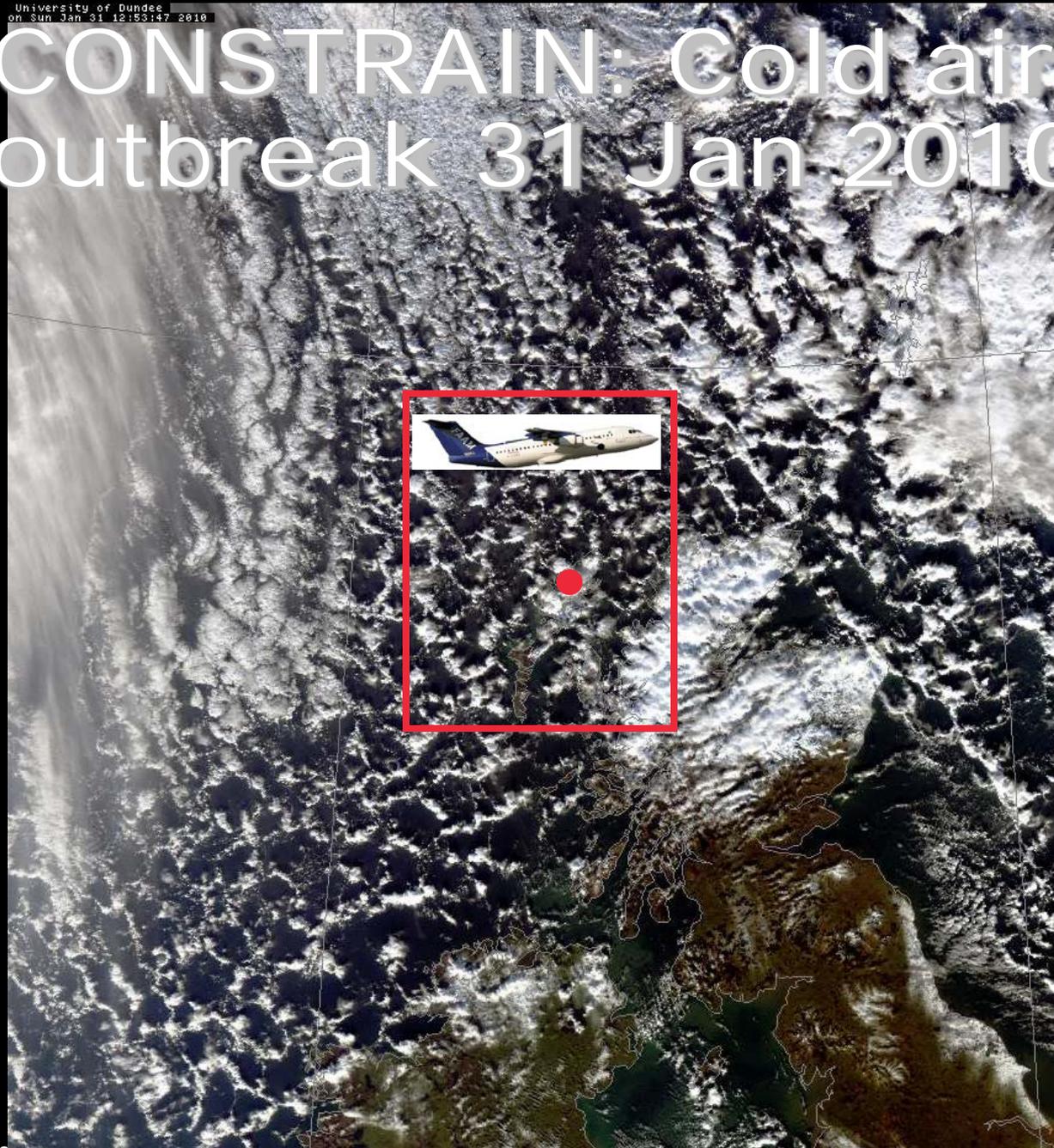


Mixed-phase

Paul Field

Richard Cotton, Adrian Lock, Kirsty McBeath, Richard Allen, Stuart Webster, Adrian Hill,
Kalli Furtado, Alexei Korolev, Andy Heymsfield, Arron Bansemer.

CONSTRAIN: Cold air outbreak 31 Jan 2010



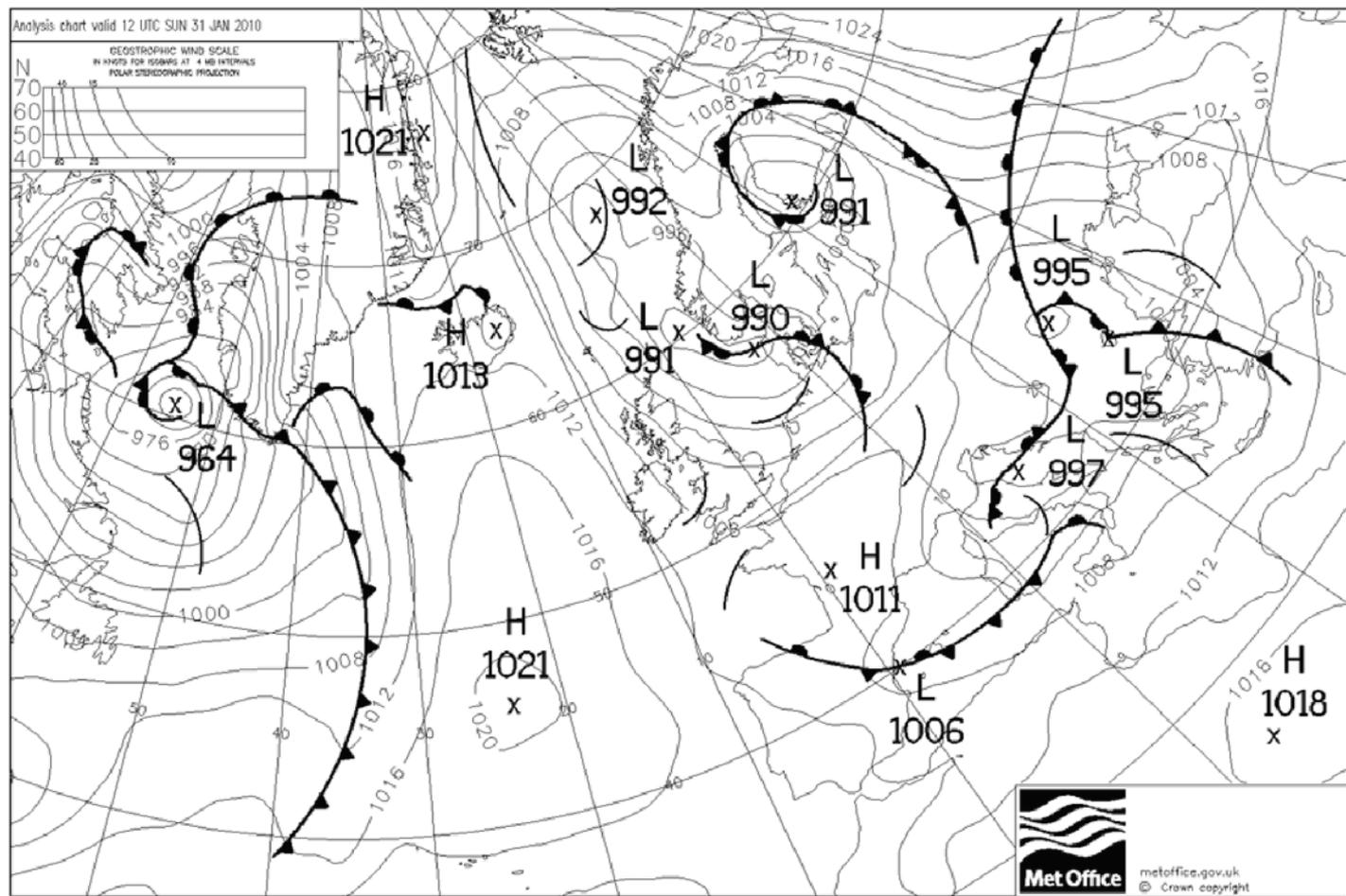
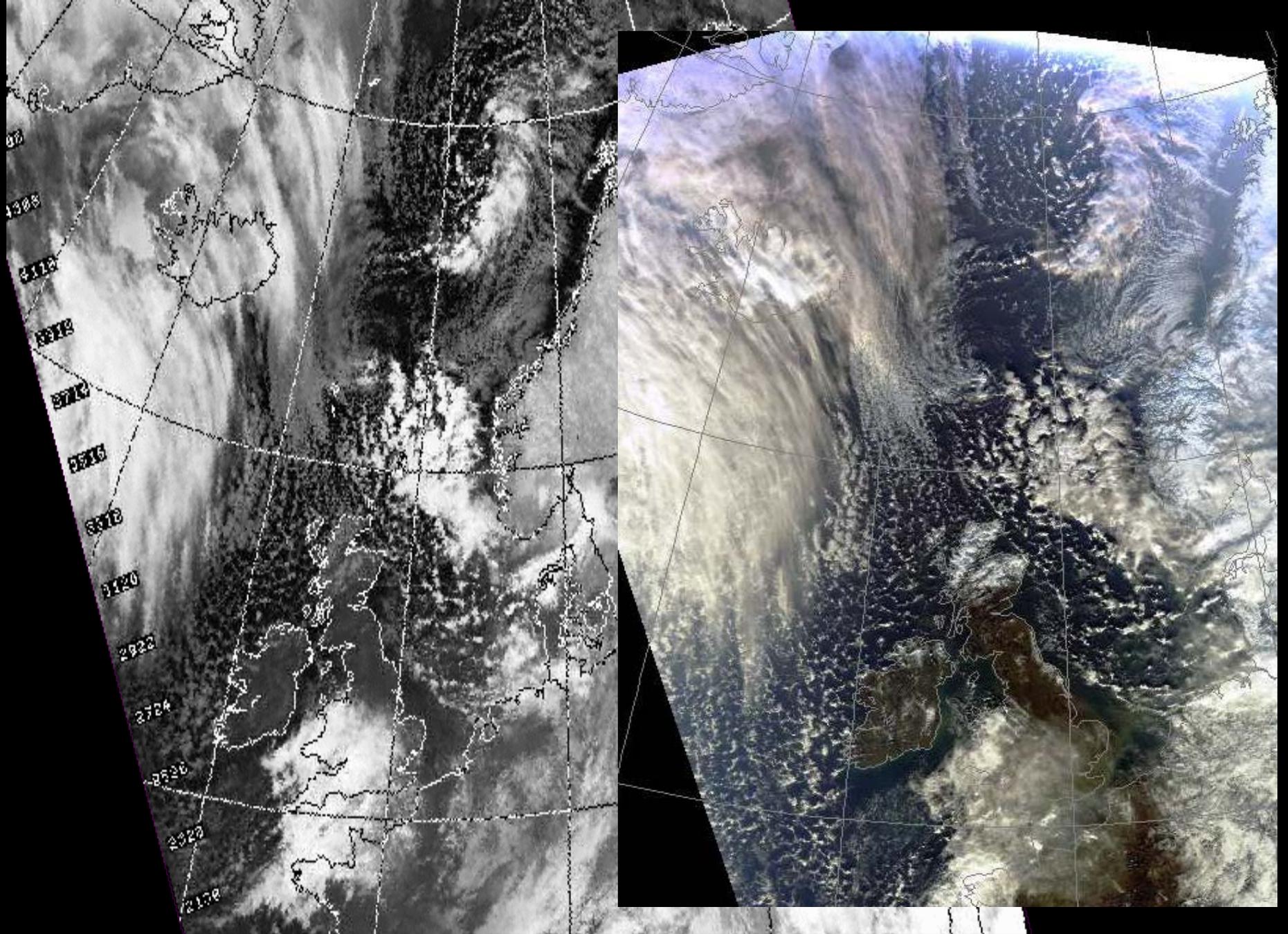
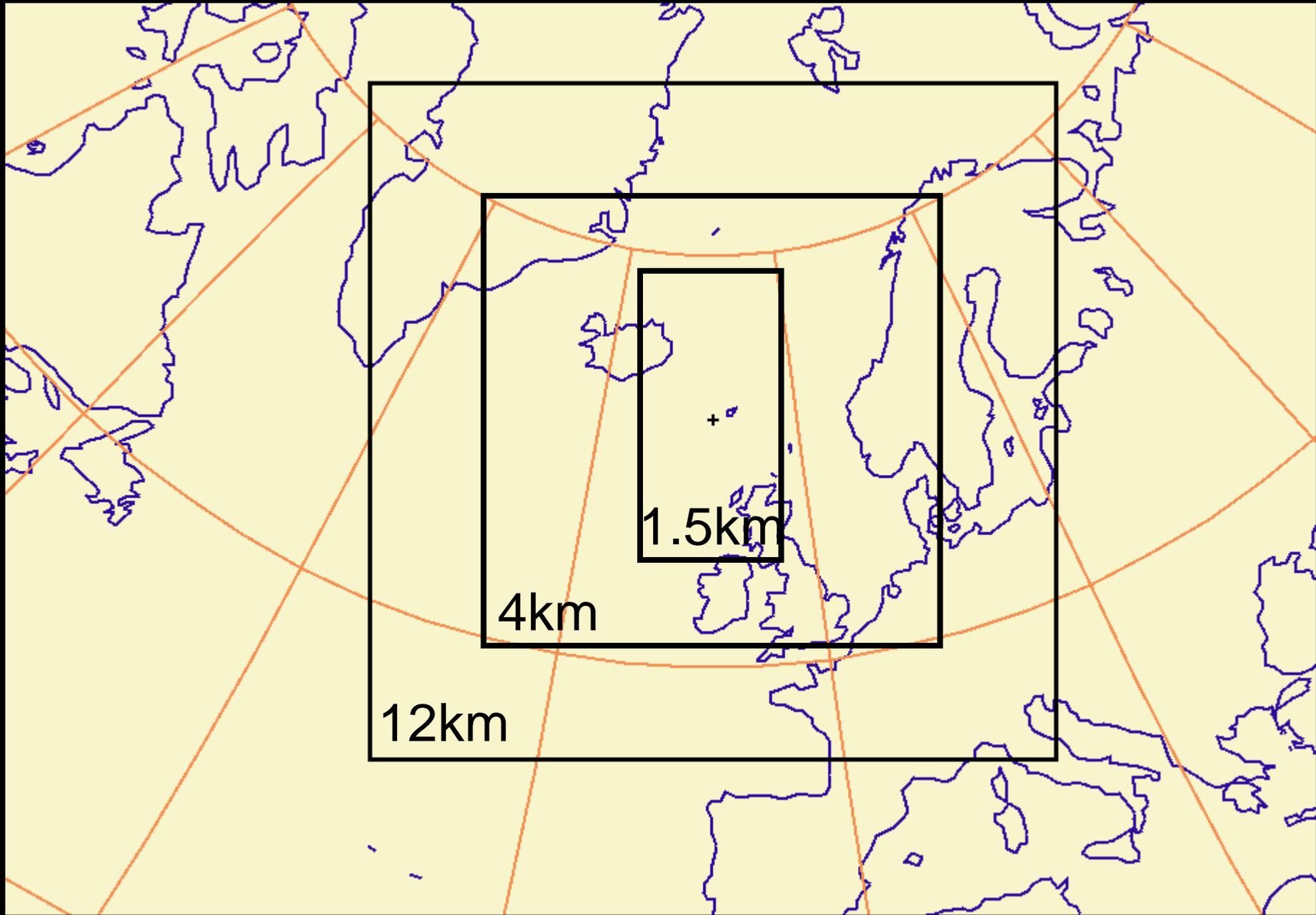
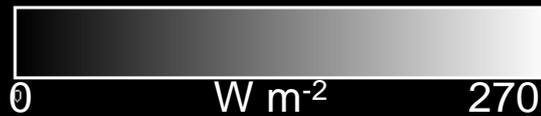


Figure 1. Analysis chart for 12 UTC 31st January 2010.

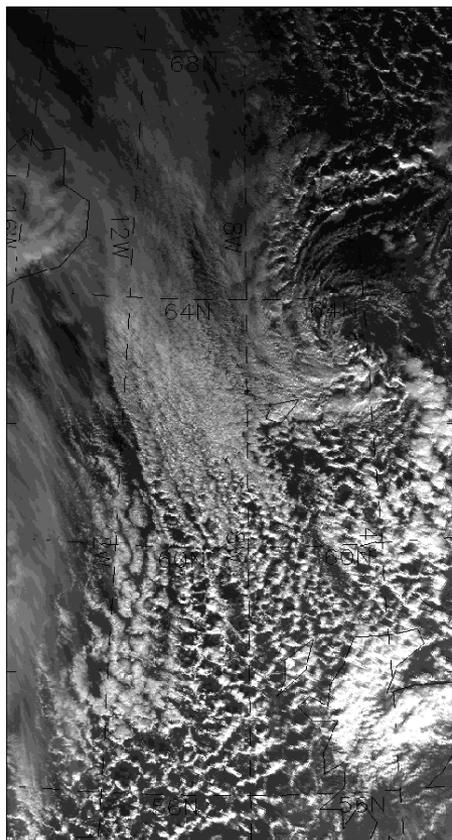




DIAFJ Top of atmosphere Atmos outgoing sw rad flux (toa)
At 09Z on 31/ 1/2010, from 05Z on 31/ 1/2010



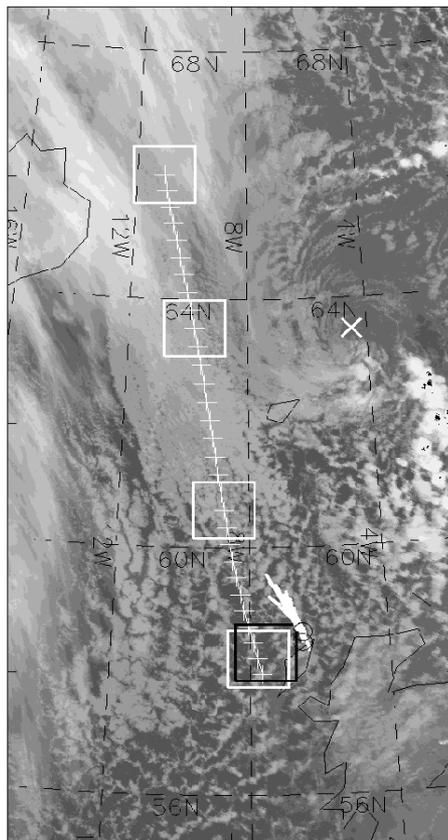
MODIS ch4



0 20 40 60 80

$W m^{-2} sr^{-1}$

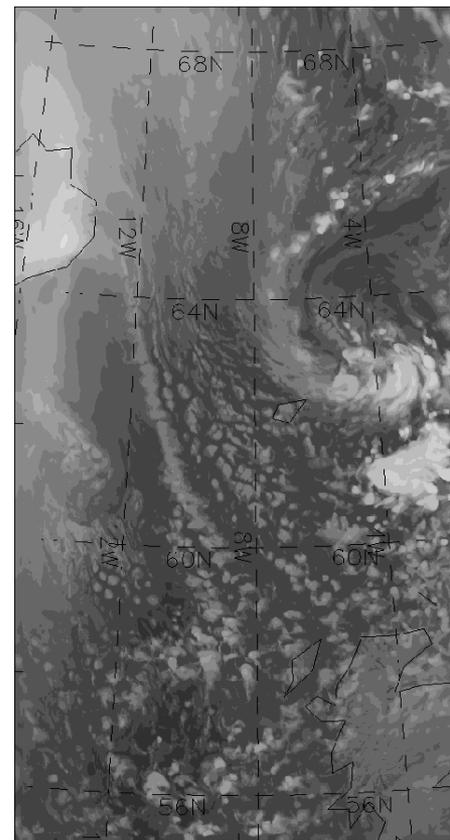
MODIS ch31



2 4 6 8 10

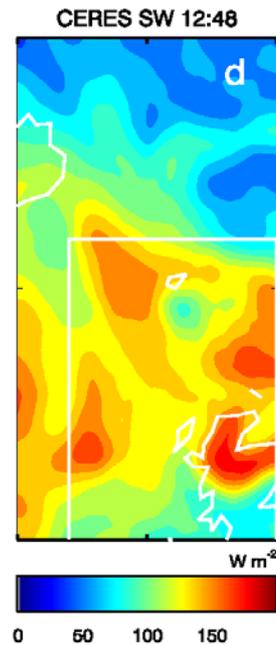
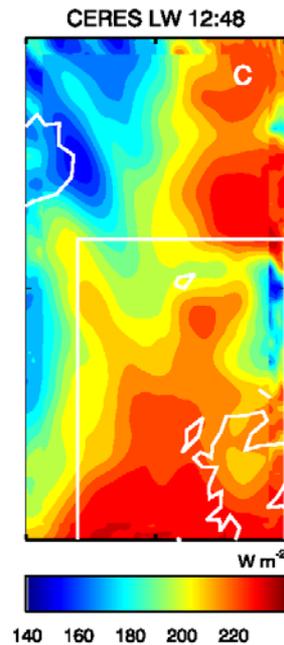
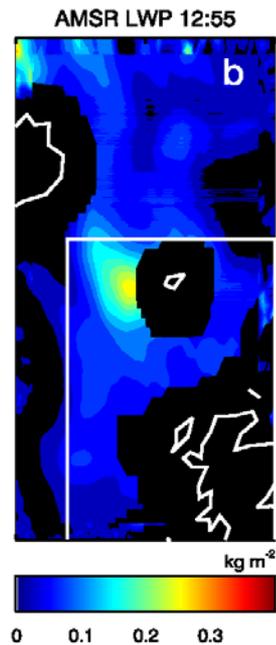
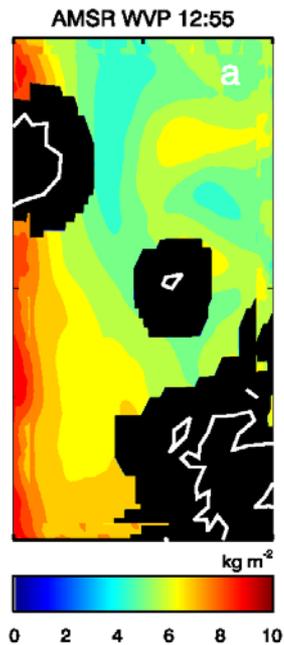
$W m^{-2} sr^{-1}$

LW flux TOA Control (h)

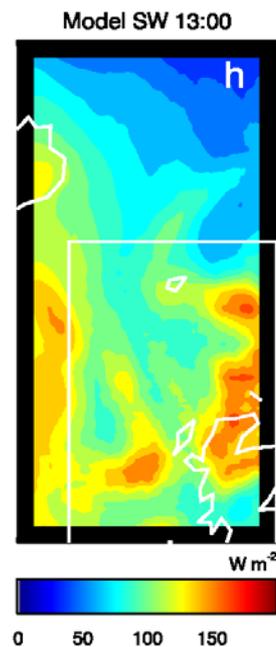
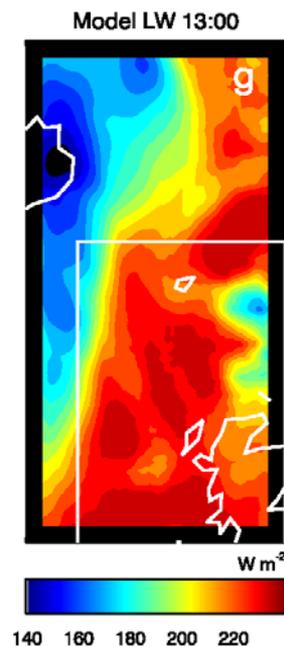
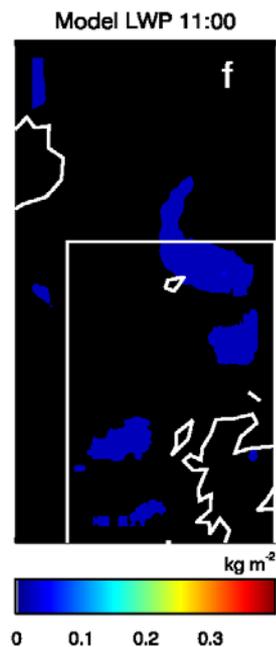
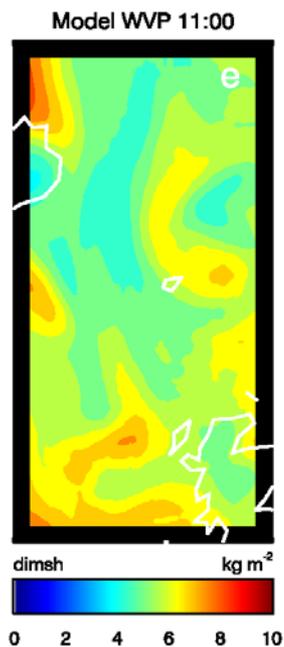


100 140 180 220 260

$W m^{-2}$



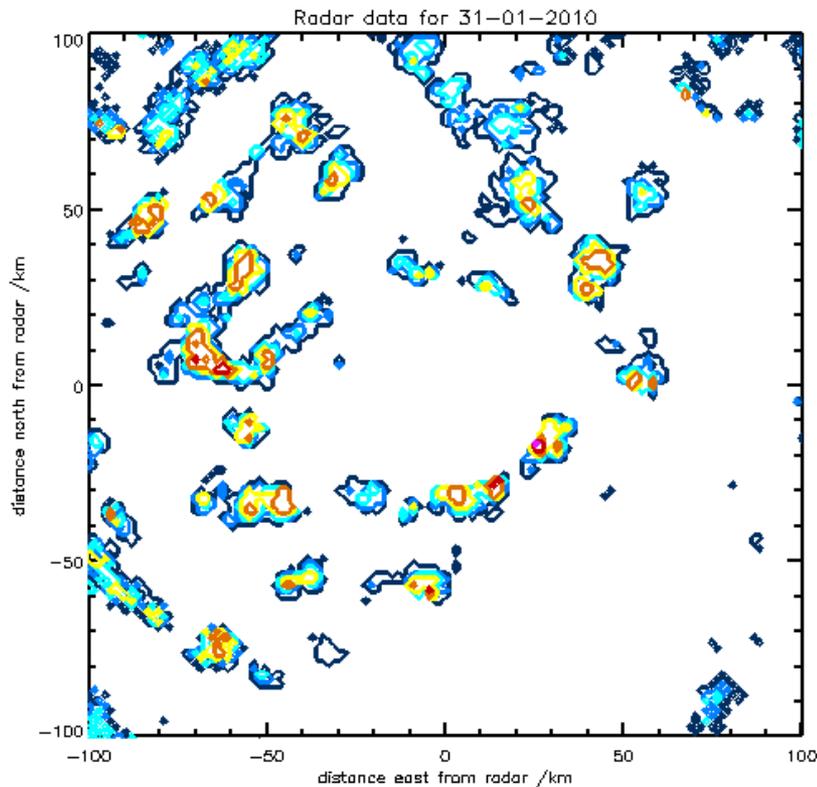
Microwave Obs



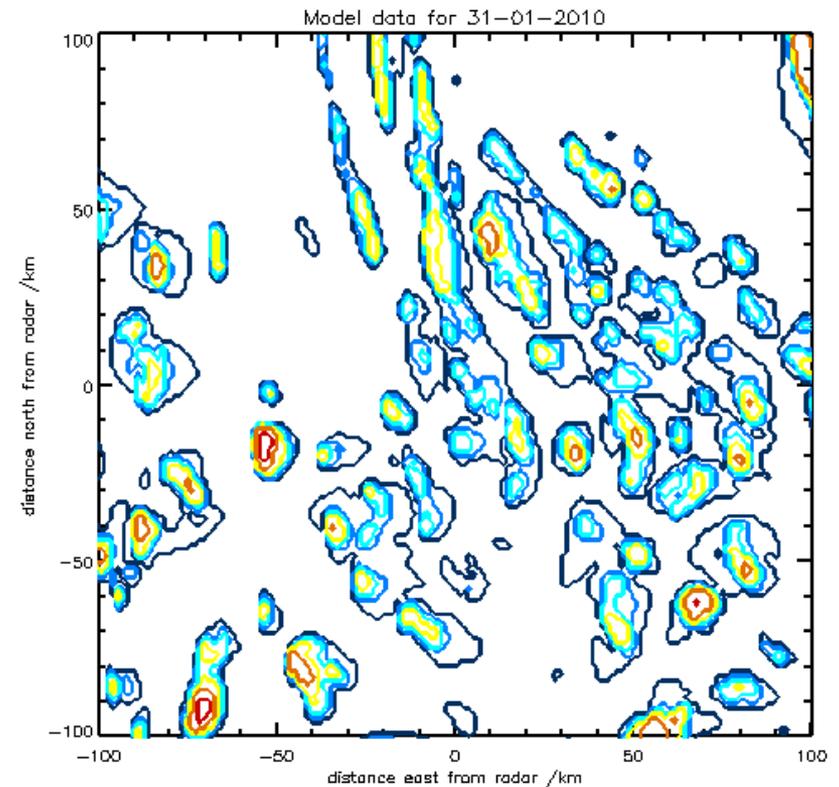
Control Model

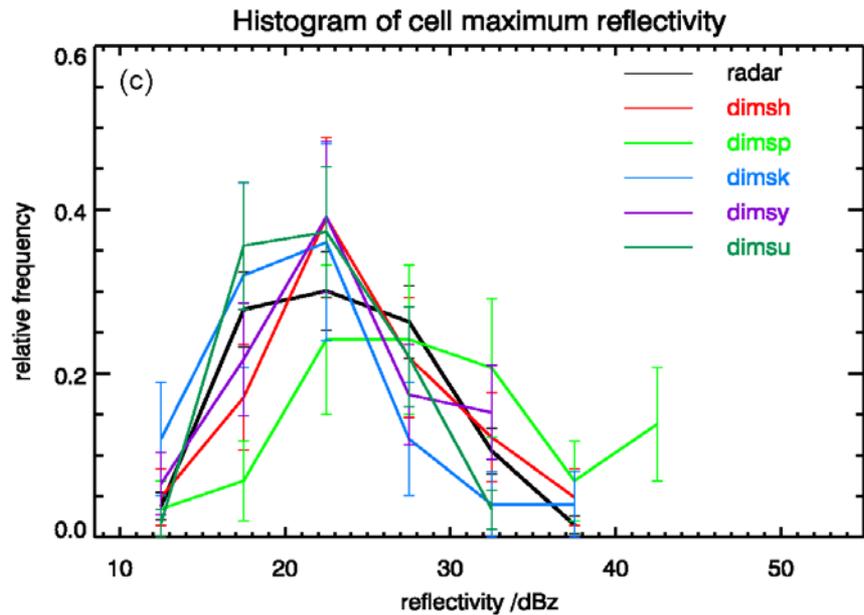
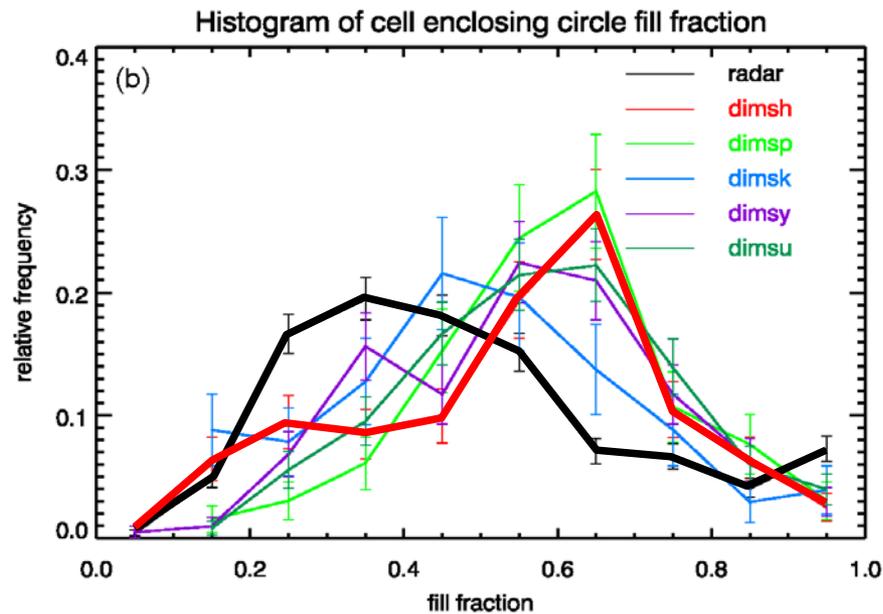
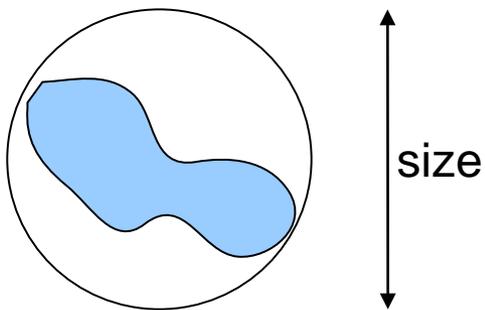
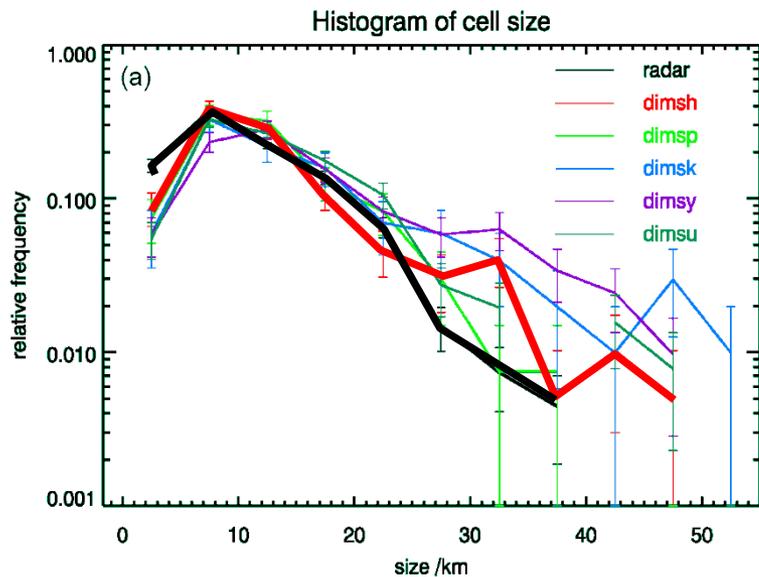
Reflectivity values for UM computed using model microphysics data, this reduces processing done on radar data and removes assumptions used when converting reflectivity to rain-rate

Radar



Model





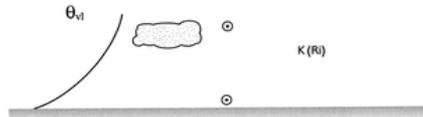
Main short comings of the control model:

Not enough cloud cover in the stratiform region

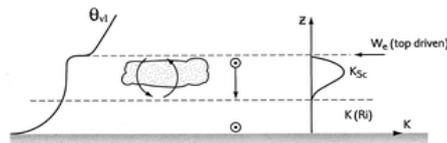
Not enough liquid water

Experiment	Sh. dom. BL.	Tnuc=-18C	AcE=0.1	No ice	PSD	3dSmag
dimsh						
dimsp				✓		
dimsq			✓			
dimsn						✓
dimsk	✓					
dimsi	✓		✓			
dimsz	✓	✓				
dimsy	✓	✓	✓			
dimsu	✓	✓	✓		✓	
dimsw		✓	✓			✓

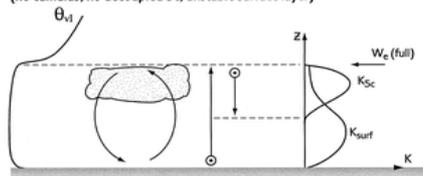
(a) I. Stable boundary layer, possibly with non-turbulent cloud (no cumulus, no decoupled Sc, stable surface layer)



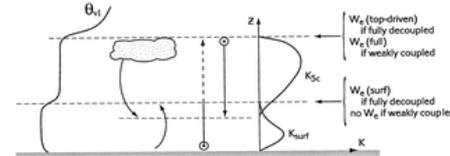
(b) II. Stratocumulus over a stable surface layer (no cumulus, decoupled Sc, stable surface layer)



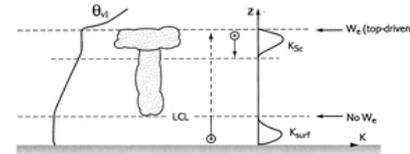
(c) III. Single mixed layer, possibly cloud-topped (no cumulus, no decoupled Sc, unstable surface layer)



(d) IV. Decoupled stratocumulus not over cumulus (no cumulus, decoupled Sc, unstable surface layer) Lock 2000



(e) V. Decoupled stratocumulus over cumulus (cumulus, decoupled Sc, unstable surface layer)



(f) VI. Cumulus-capped layer (cumulus, no decoupled Sc, unstable surface layer)

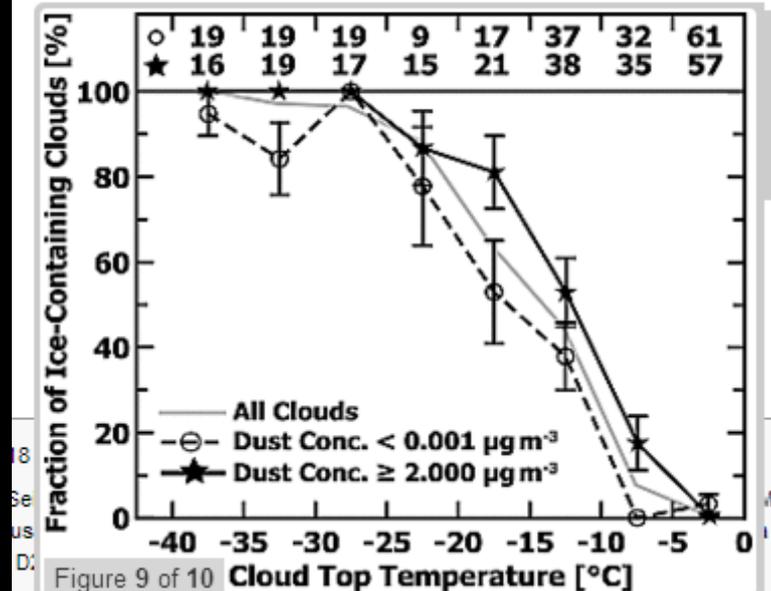
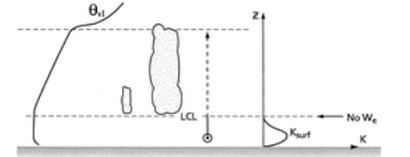
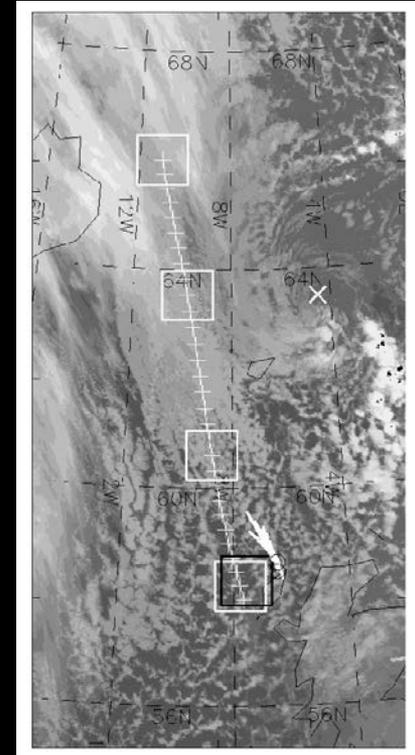
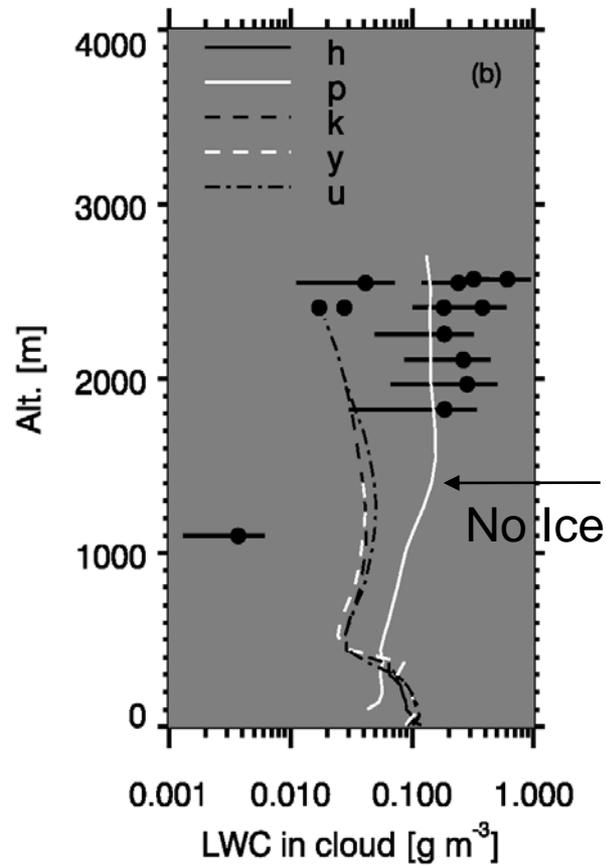
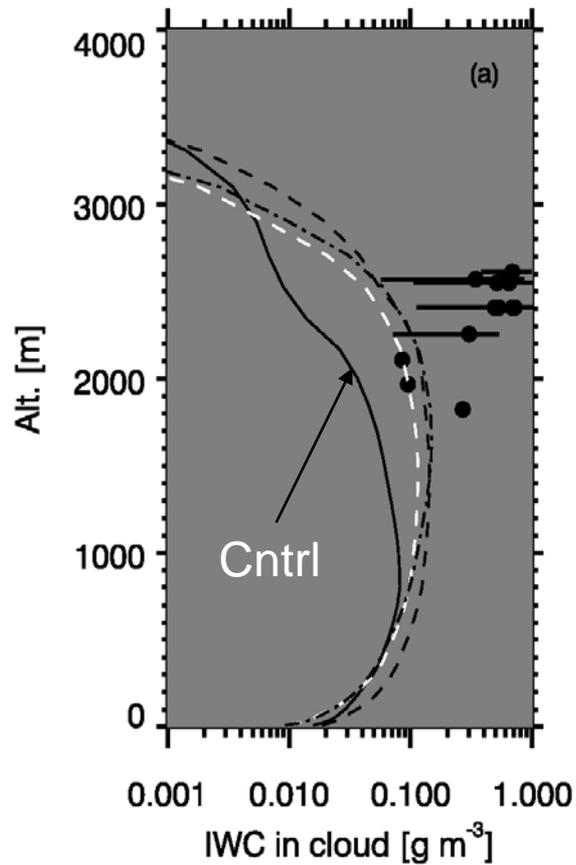
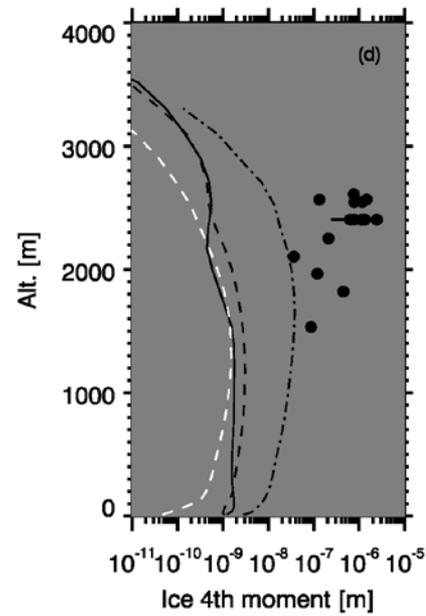
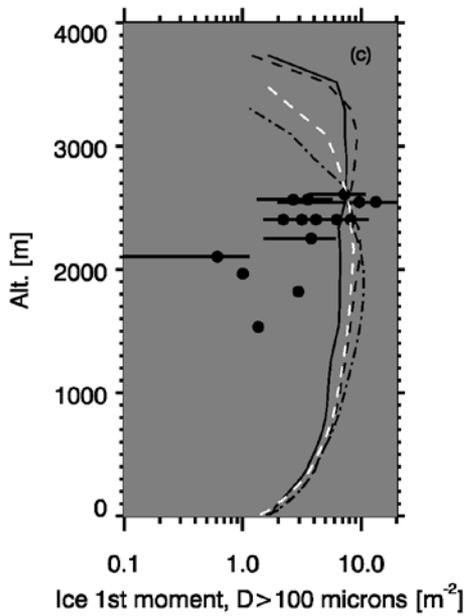
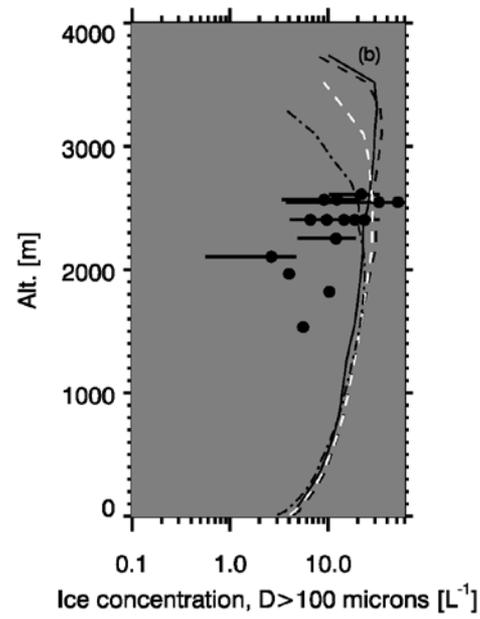
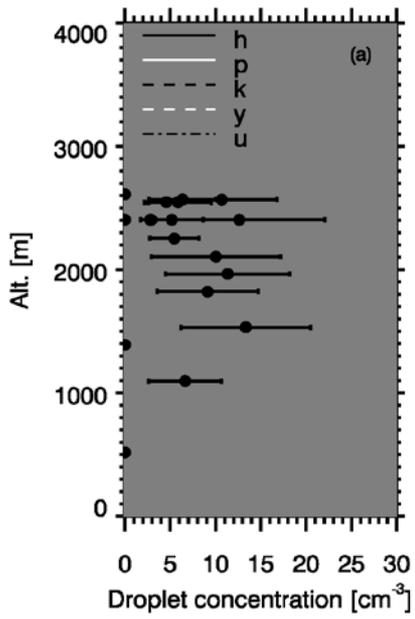


Figure 9 of 10 Cloud Top Temperature [°C]

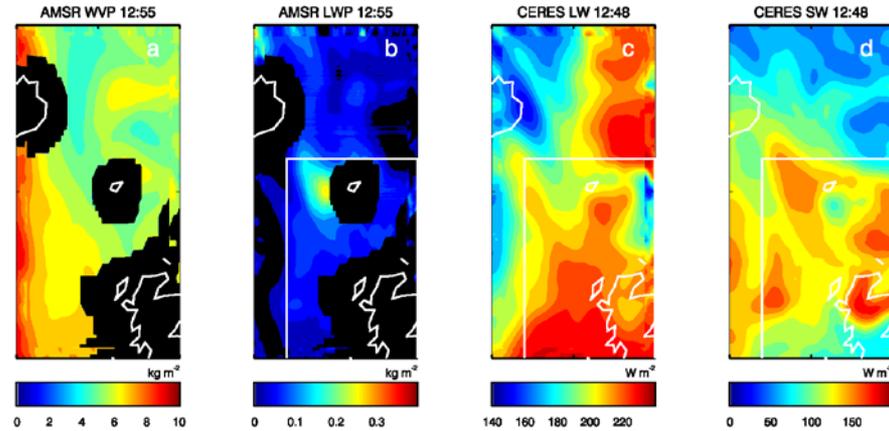


Experiment	Sh. dom. BL.	Tnuc=-18C	AcE=0.1	No ice	PSD	3dSmag
dimsh						
dimsp				✓		
dimsq			✓			
dimsn						✓
dimsk	✓					
dimsi	✓		✓			
dimsz	✓	✓				
dimsy	✓	✓	✓			
dimsu	✓		✓		✓	
dimsw		✓	✓			✓

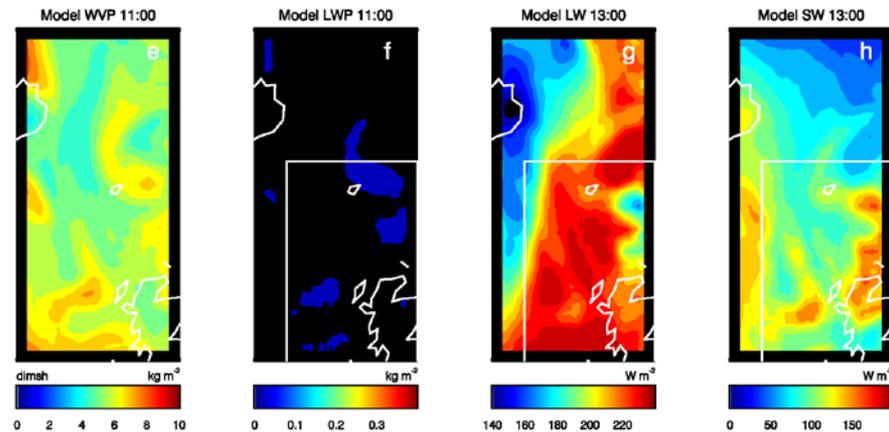


Experiment	Sh. dom. BL.	Tnuc=-18C	AcE=0.1	No ice	PSD	3dSmag
dimsh						
dimp				✓		
dimq			✓			
dimn						✓
dimk	✓					
dimi	✓		✓			
dimz	✓	✓				
dimy	✓		✓			
dimu	✓	✓	✓		✓	
dimw		✓	✓			✓

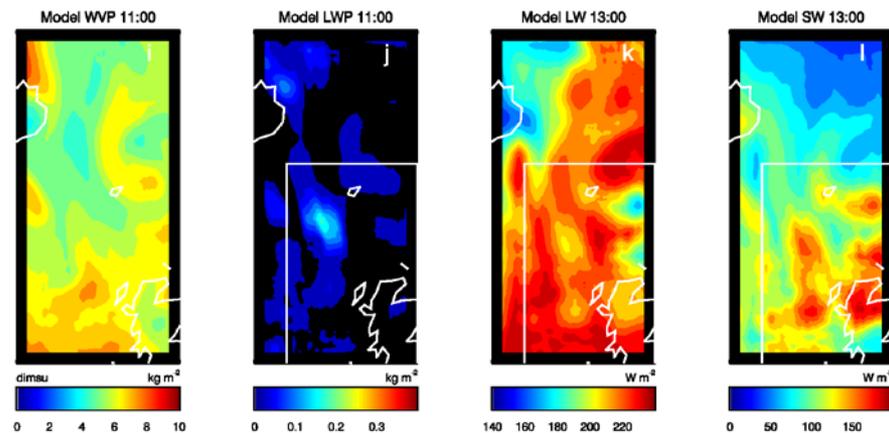
Microwave Obs

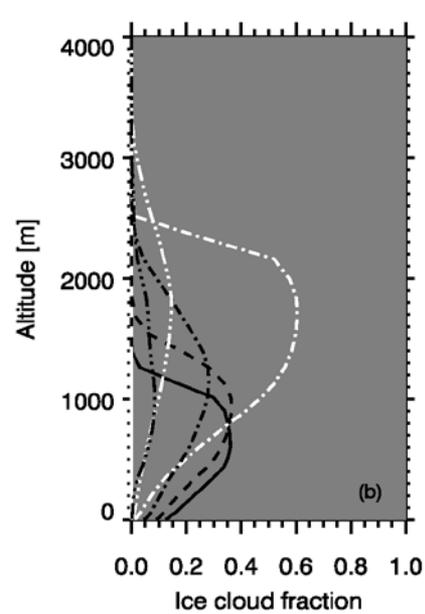
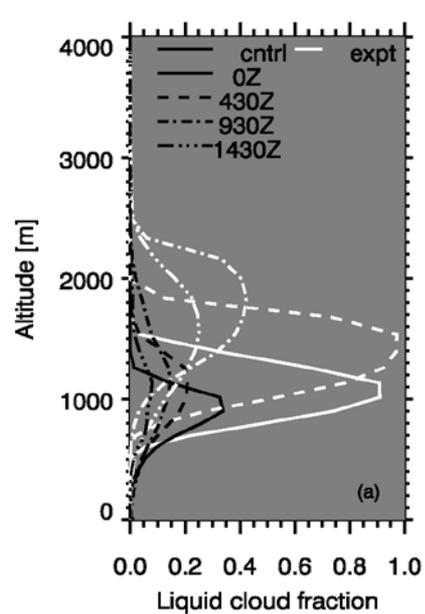
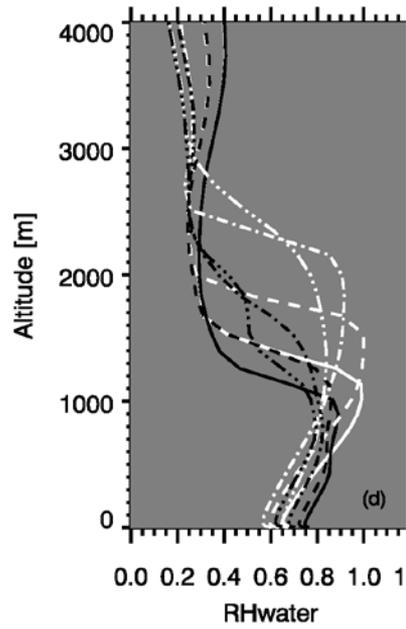
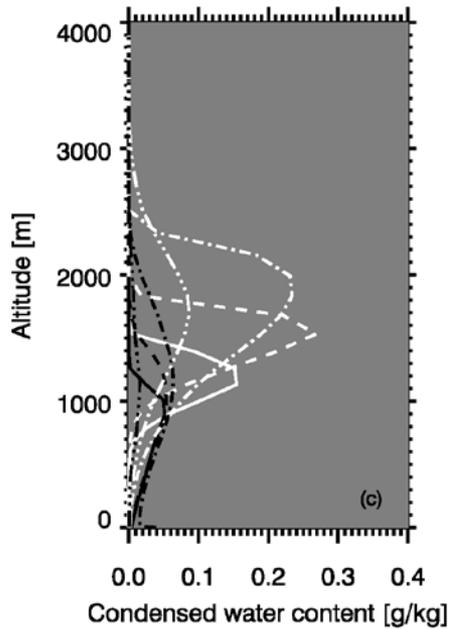
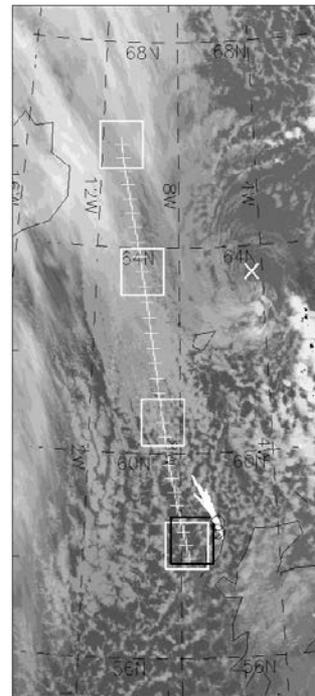
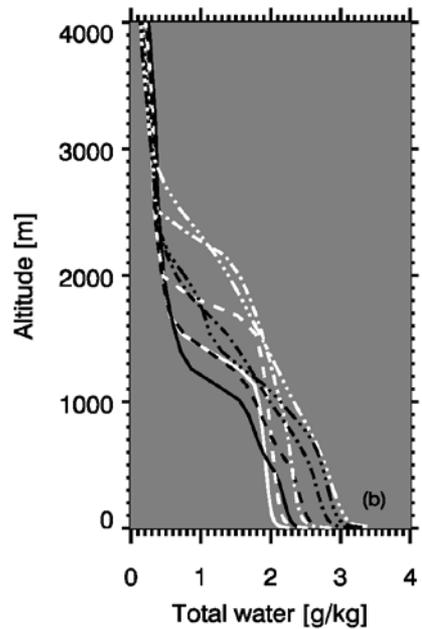
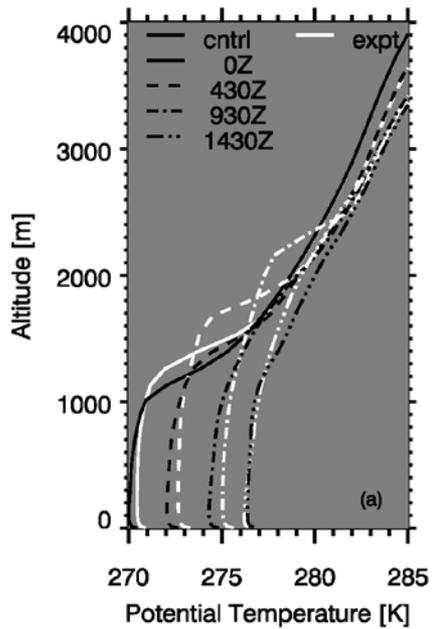


Control Model

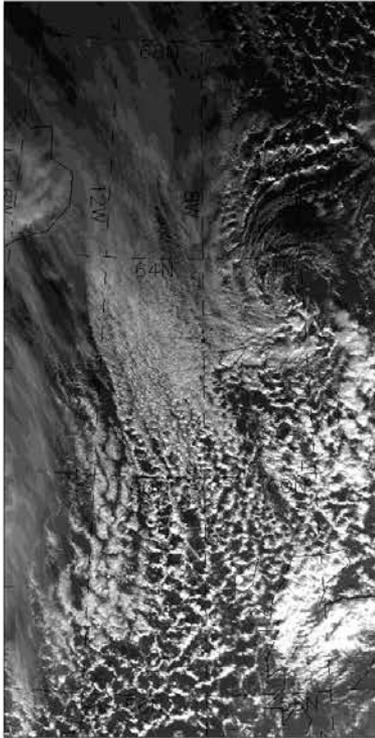


Modified Model
Sh dom BL
Thet=-18C
New PSD





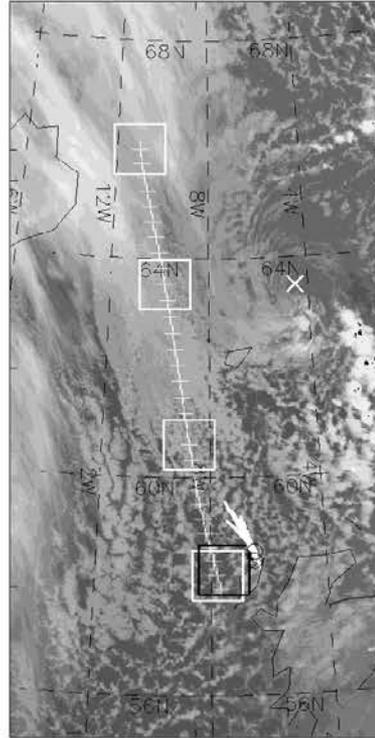
MODIS ch4



0 20 40 60 80

$W m^{-2} sr^{-1}$

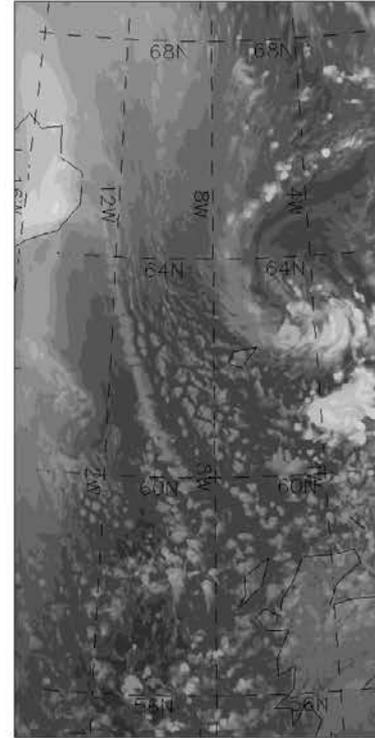
MODIS ch31



2 4 6 8 10

$W m^{-2} sr^{-1}$

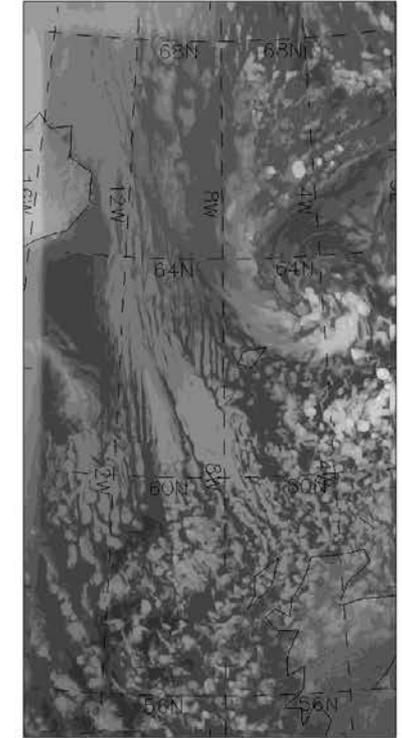
LW flux TOA Control (h)



100 140 180 220 260

$W m^{-2}$

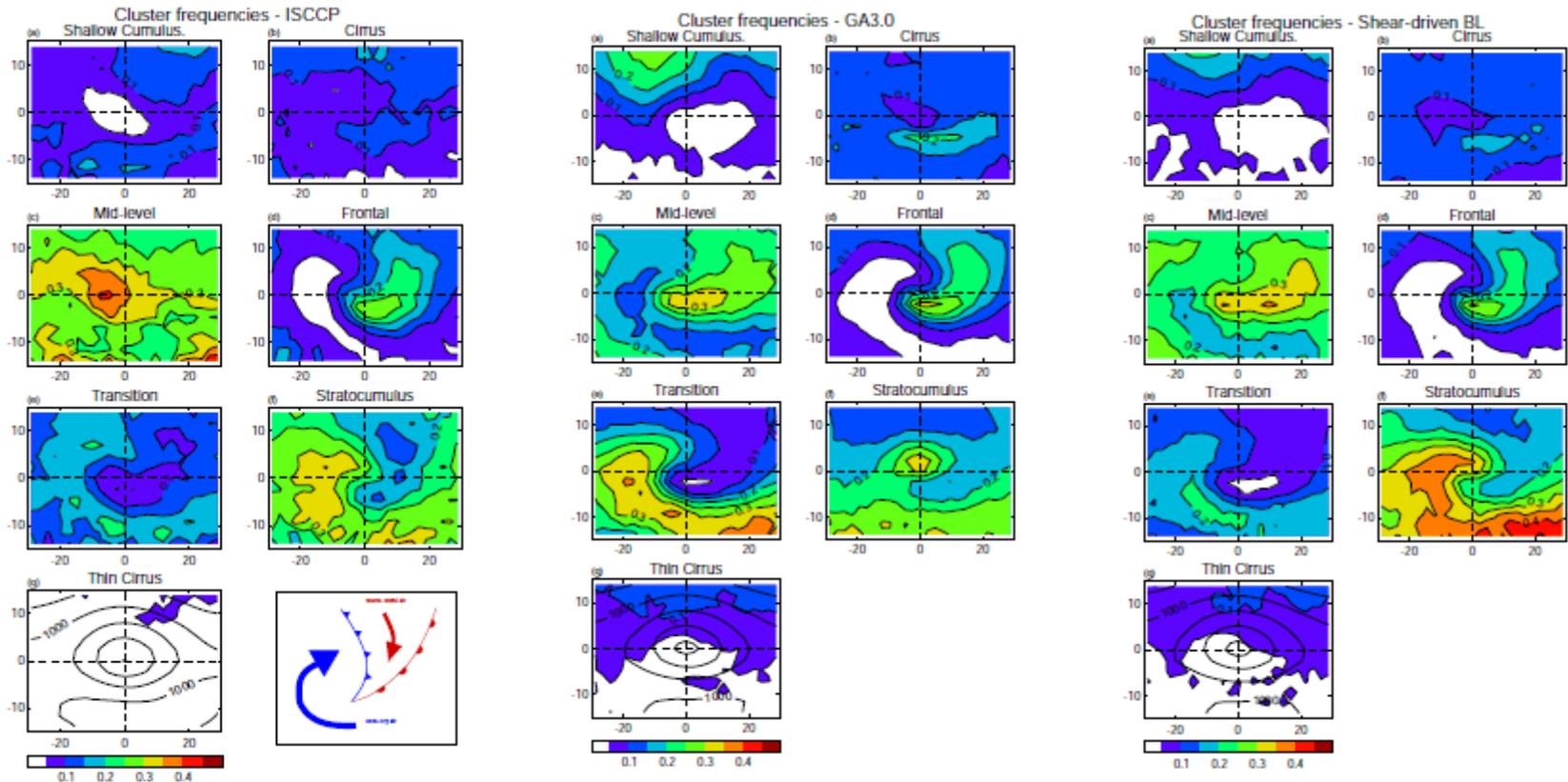
LW flux TOA Sensitivity u



100 140 180 220 260

$W m^{-2}$

Effect in a climate model

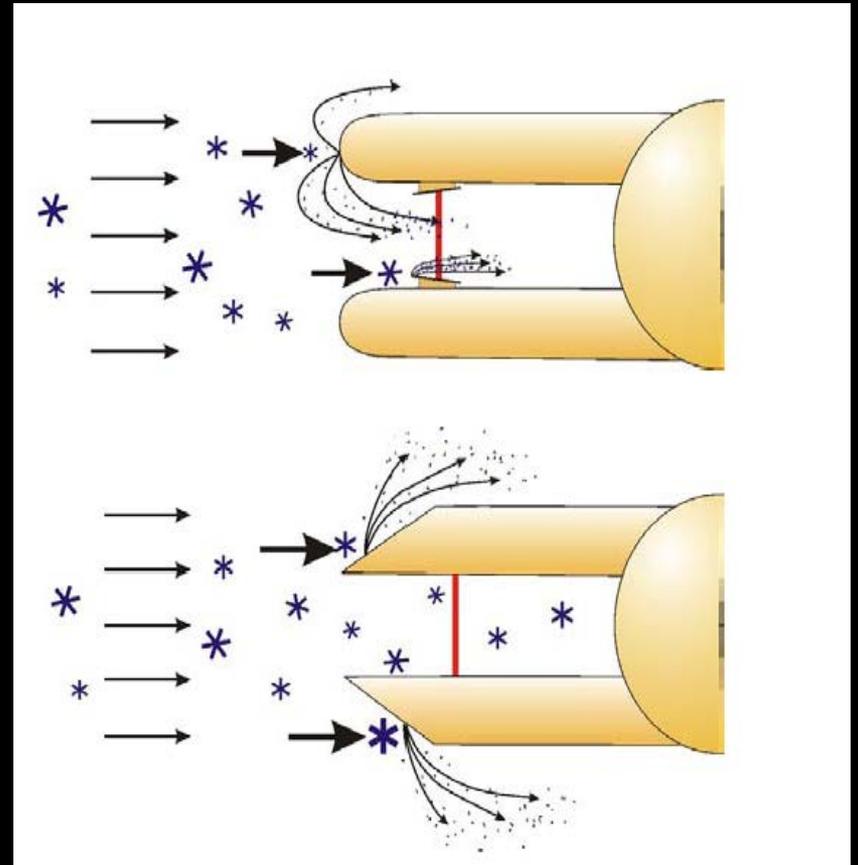


Satellite

Old

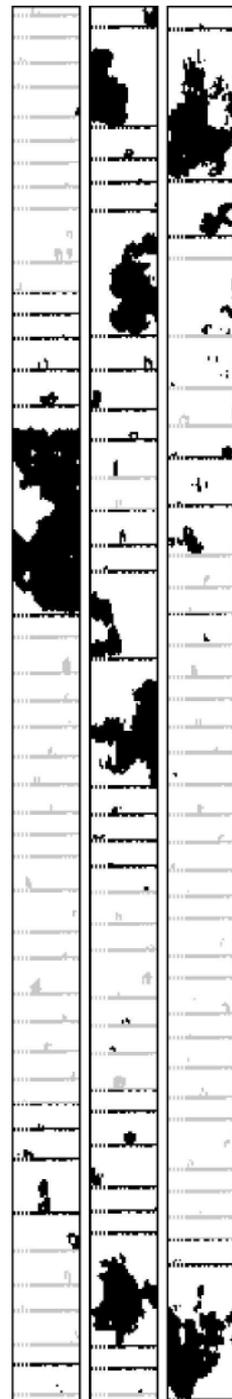
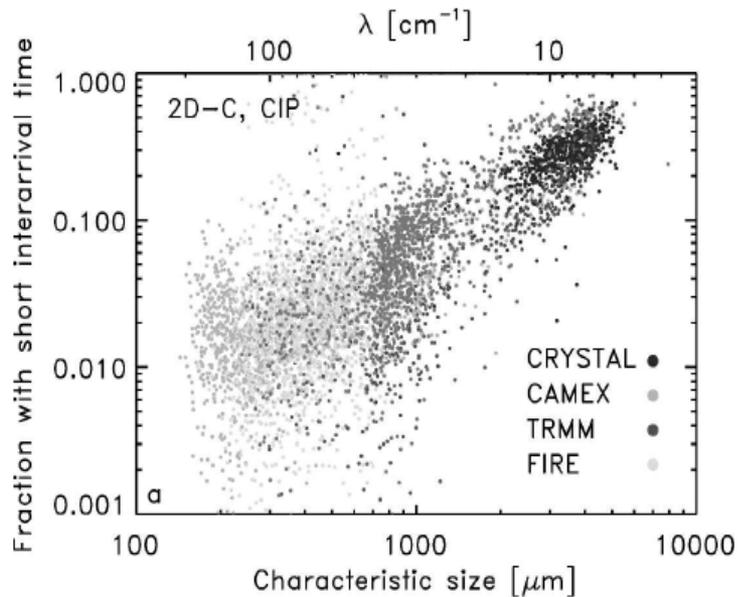
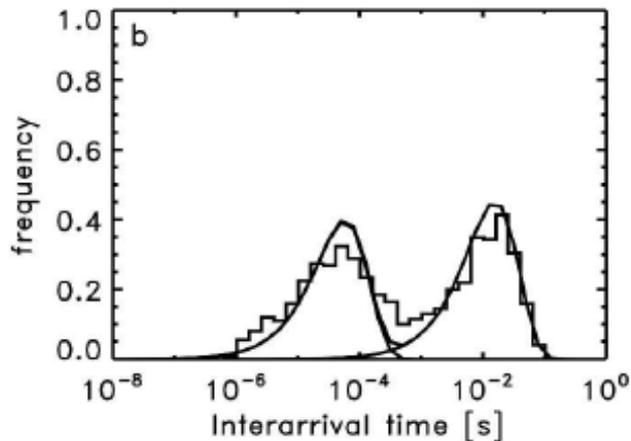
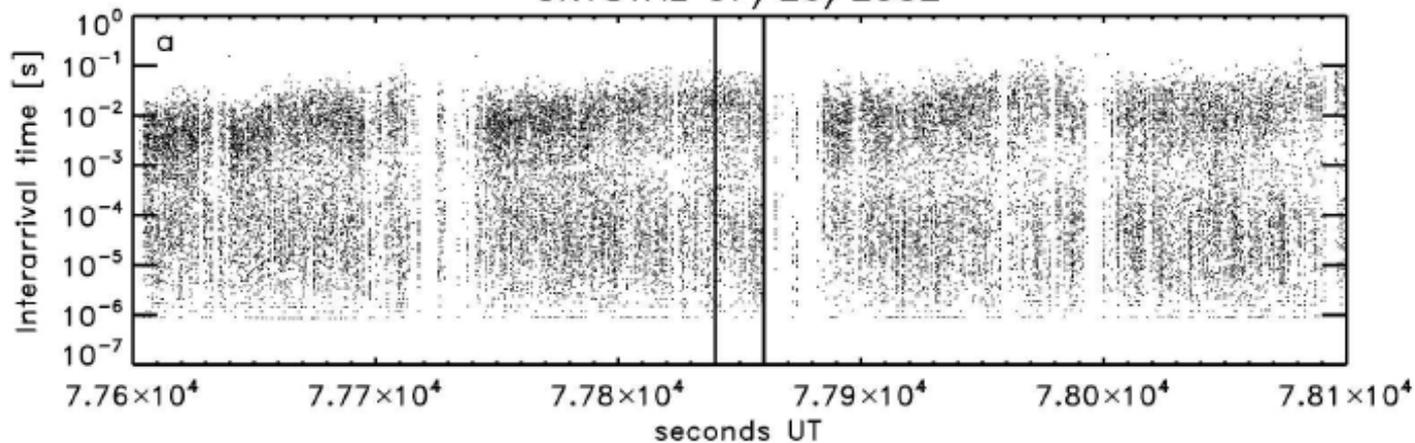
New

Aside part 1 - shattering



Korolev tips – from DMT site

CRYSTAL 07/26/2002



$$M_n = \sum D^n N(D)$$

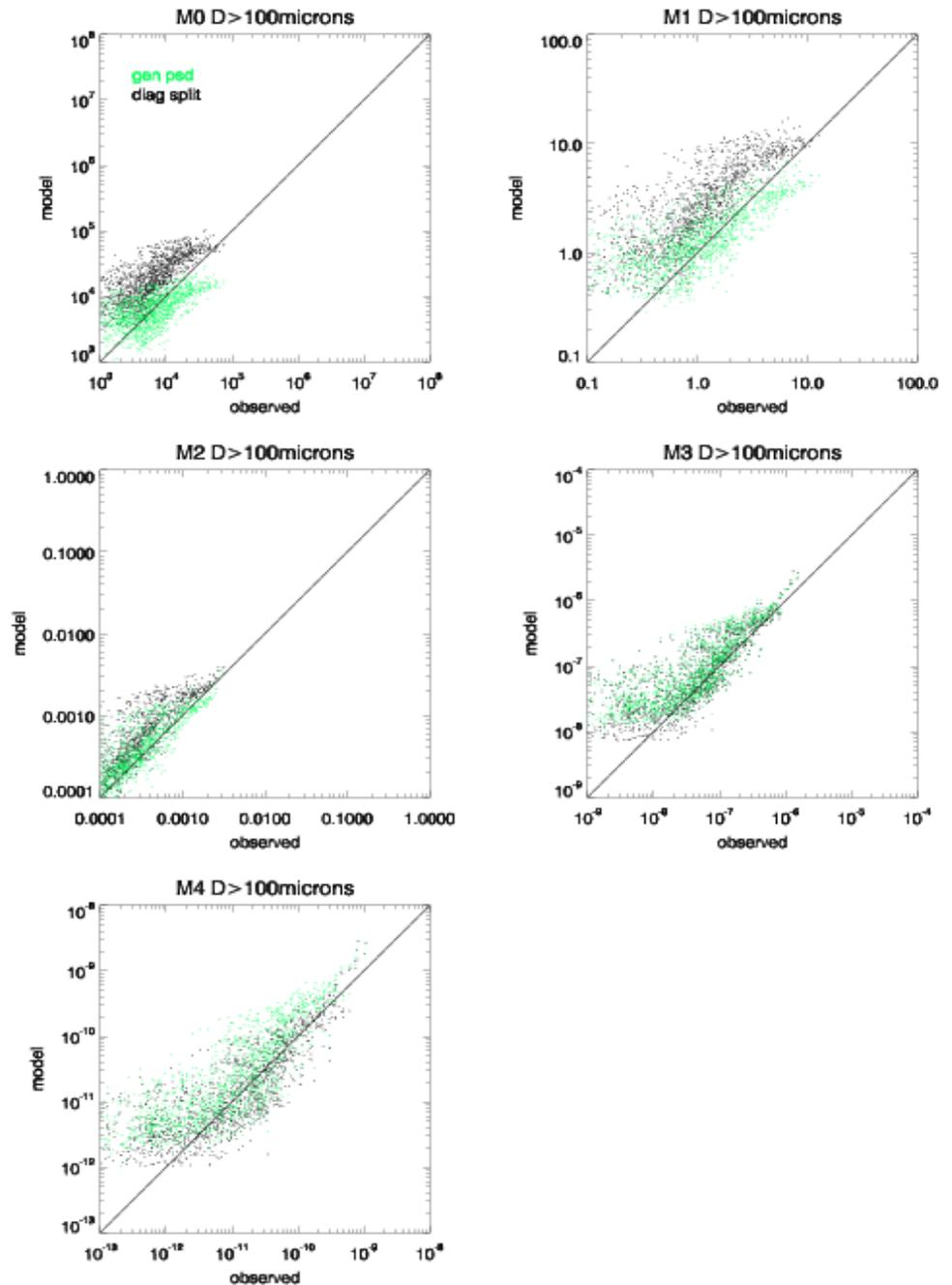
M0 = concentration

M1 proportional to
diffusional growth

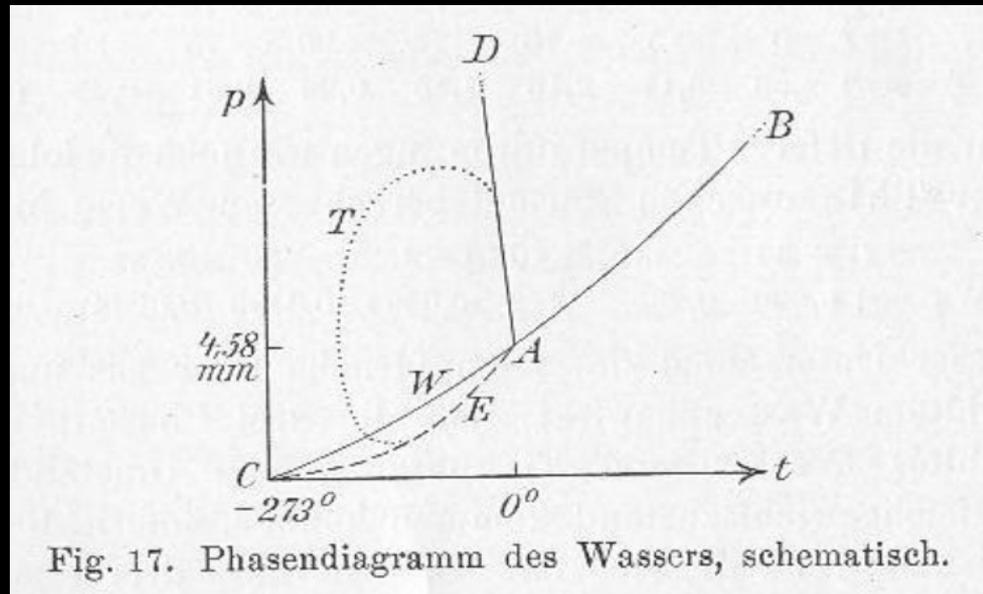
M2 proportional to
IWC

M2+a bit
proportional to snow
flux

M4 proportional to
radar reflectivity



Aside part 2 - mixed-phase



The difference in vapour pressure above ice and supercooled liquid means that under no circumstance will the 3 phases of water come into equilibrium at the same temperature.

If the 3 phases are in contact (through the vapour phase) then the vapour pressure will adjust to a value between the saturated vapour pressure over ice and liquid. The consequence must be that diffusional growth of ice will occur at the same time as liquid evaporates. This will continue until the liquid is used up.

Wegener 1911

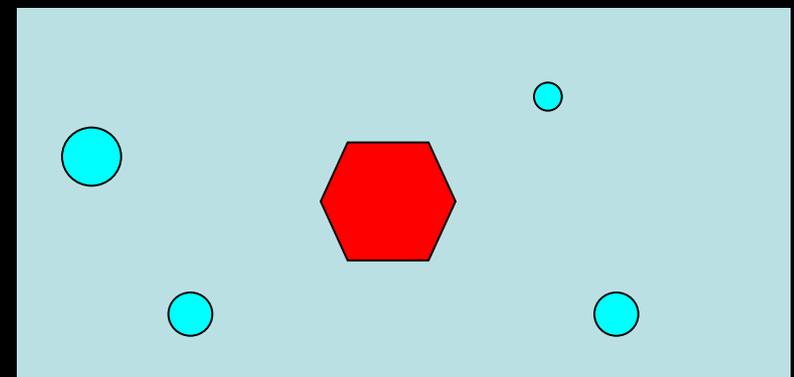
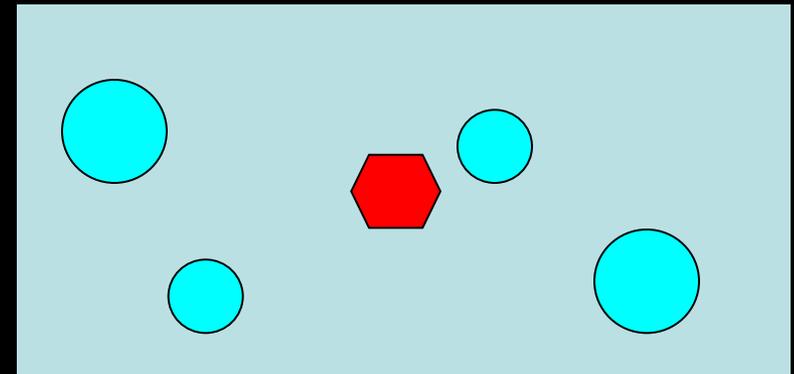
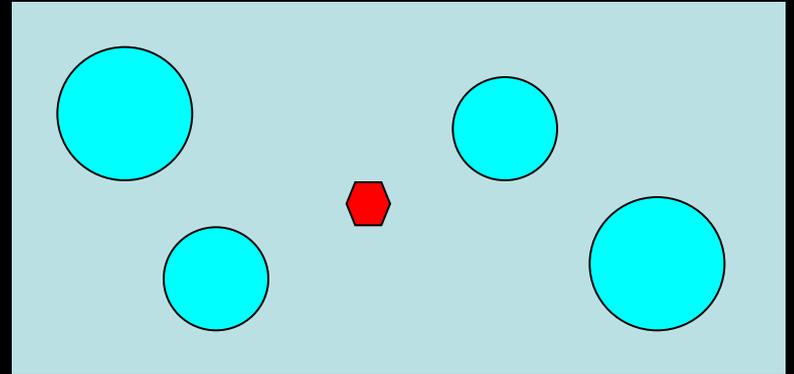
Bergeron 1935

On the physics of cloud and precipitation

- Review of precipitation formation ideas
 - Based on approach that clouds are colloids and precipitation is driven by colloidal instability.
- Recognizes that the result of 'neighbouring elements of different phase' has not been explored.
-

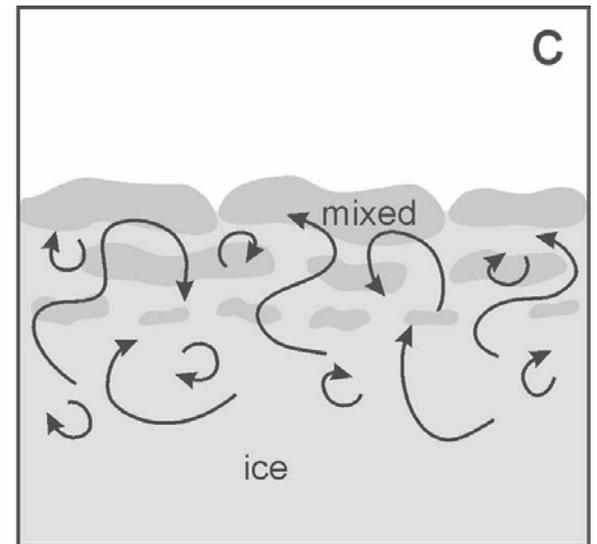
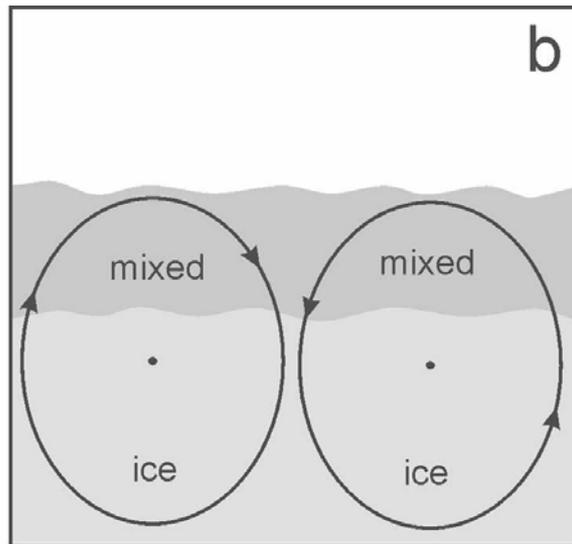
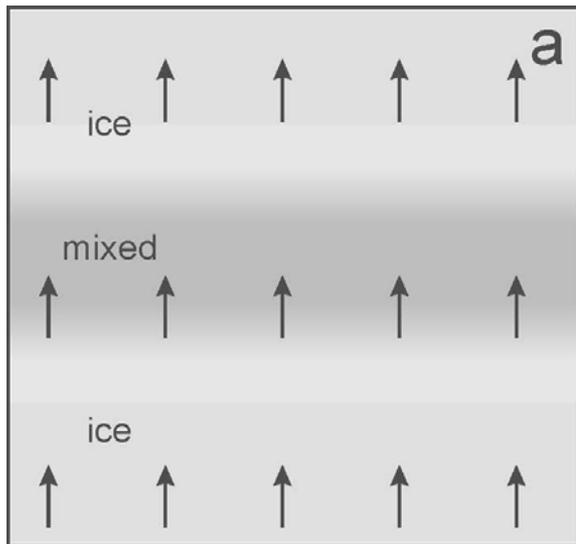
- Bergeron –

‘I thus presume that almost every real raindrop and all snowflakes originated around an ice crystal in the above mentioned way’

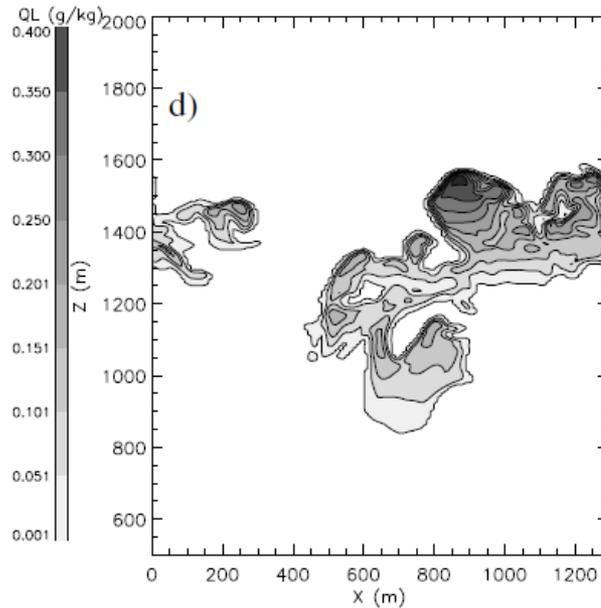
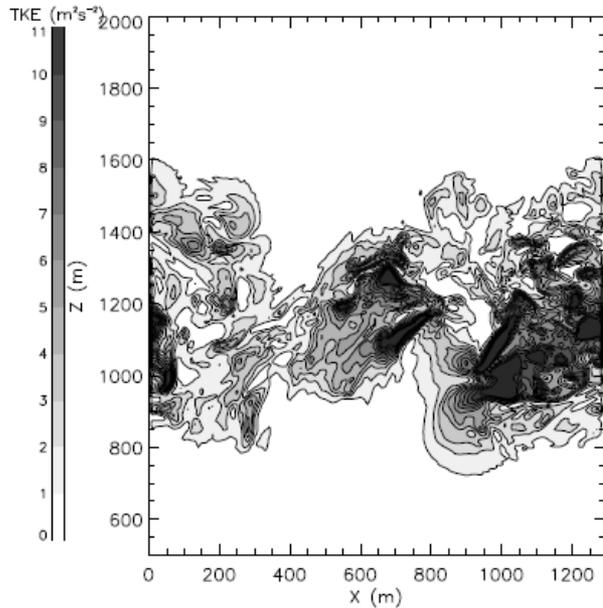
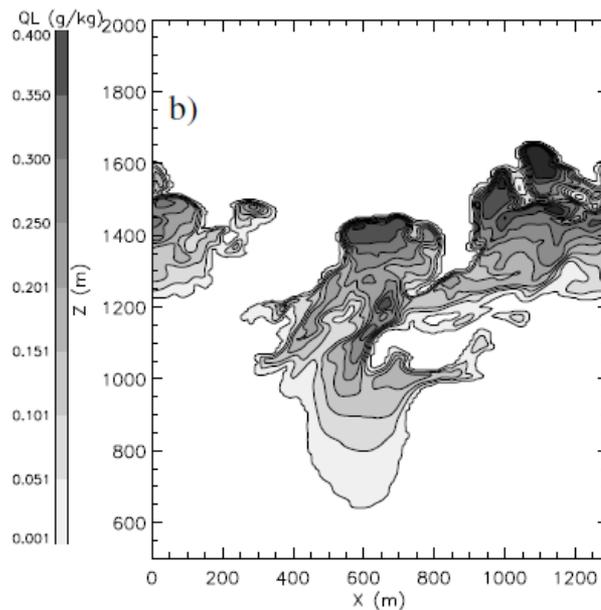
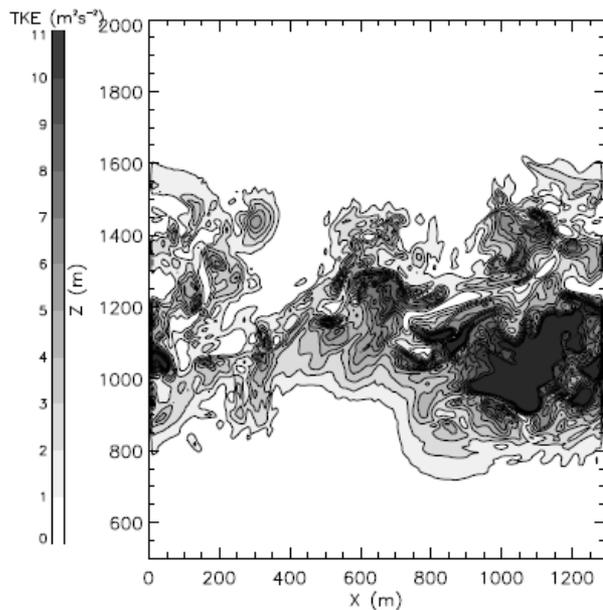


Dynamical effects

- The difference in vapour pressure above ice and supercooled liquid means that under no circumstance will the 3 phases of water come into equilibrium at the same temperature
- Heymsfield 1977: Activate liquid water
 $U_{\text{vert}} > \text{threshold}$
- Mazin 1986:
- Korolev et al.



$$\frac{1}{1 + S_i} \frac{dS_i}{dt} = a_i u - b_i B_0 M_1 S_i,$$



LEM
 $\Delta x = 10\text{m}$
 $\Delta z = 5\text{m}$
 $1.3 \times 1.3 \times 2.5\text{km}^3$

$T = -9, -6\text{C}$

$\text{IWC} = 0.1\text{ g/kg}$

Nice = 1, 10, 100
 L^{-1}

Shear:
 10 m/s over 50m
 15 m/s “ “

$$\frac{1}{1 + S_i} \frac{dS_i}{dt} = a_i u - b_i B_0 M_1 S_i,$$

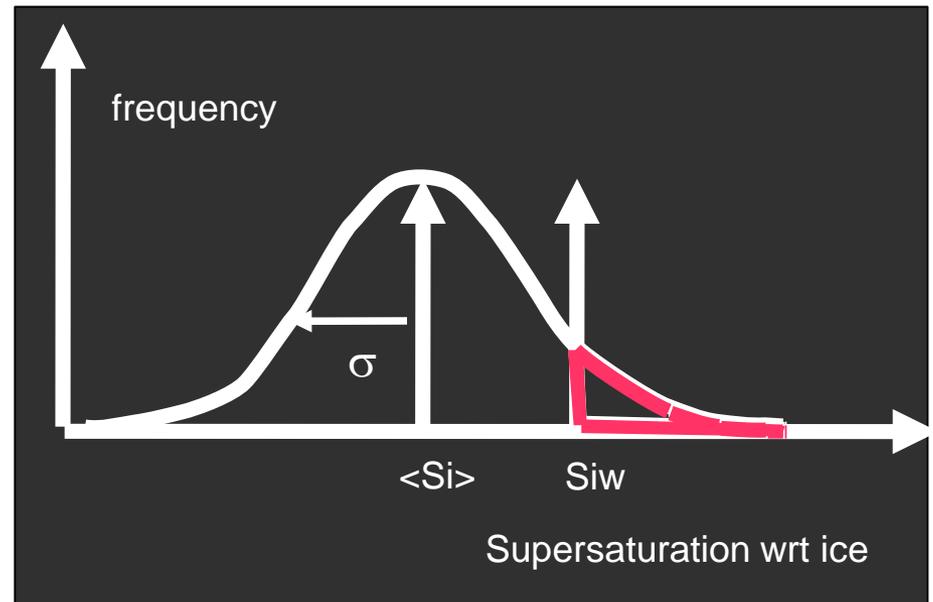
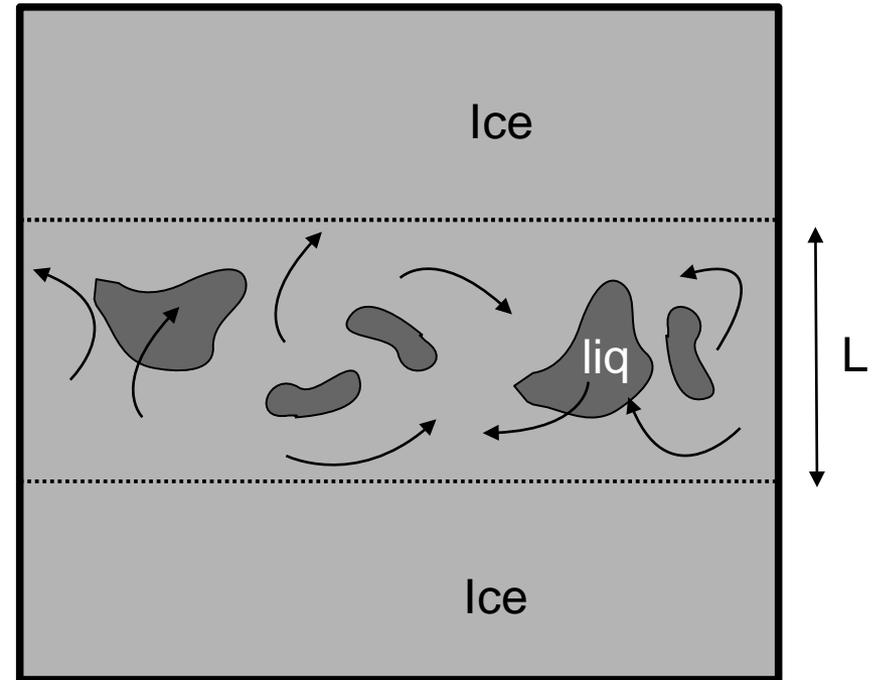
$$dS_i = (-(B + C)S_i + CS_E)dt + A\xi dt$$

$$A = a_i \sigma_w \tau_d^{\frac{1}{2}}$$

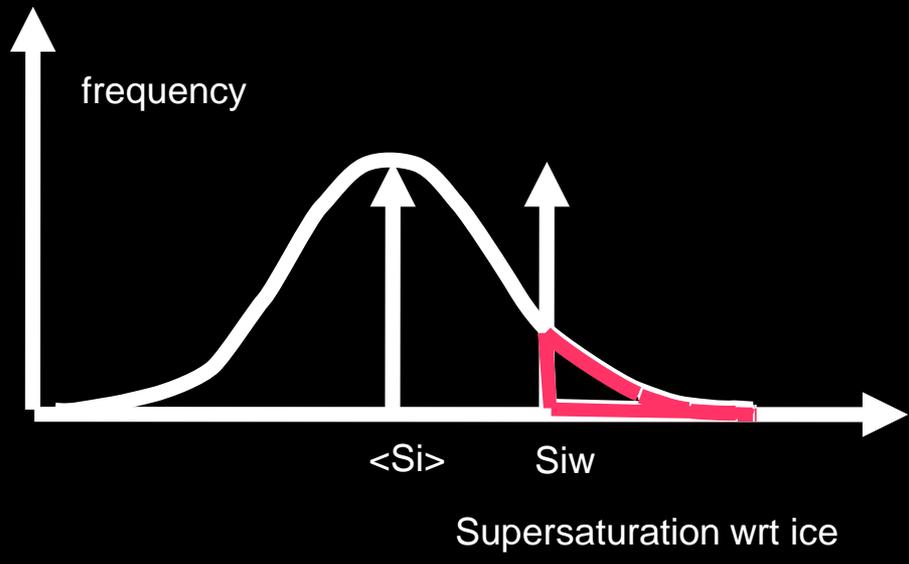
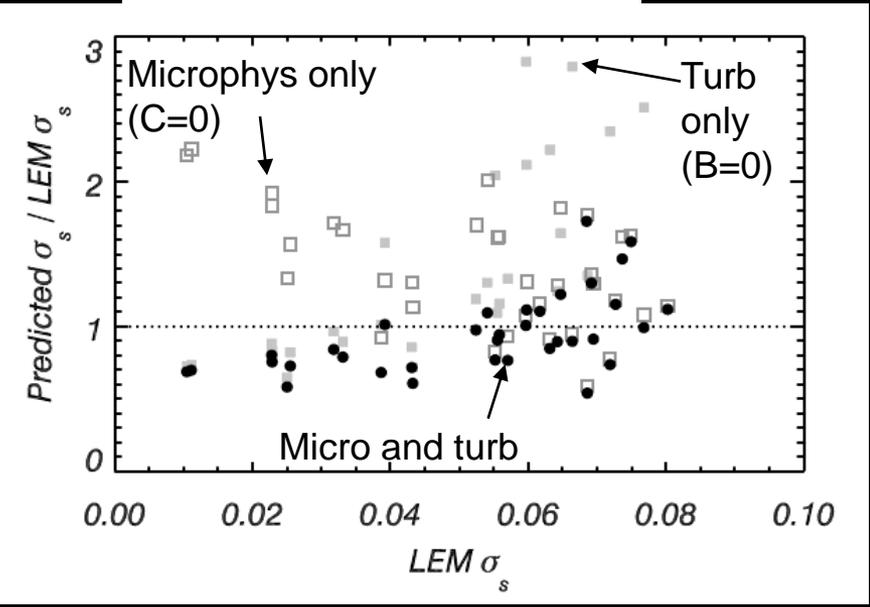
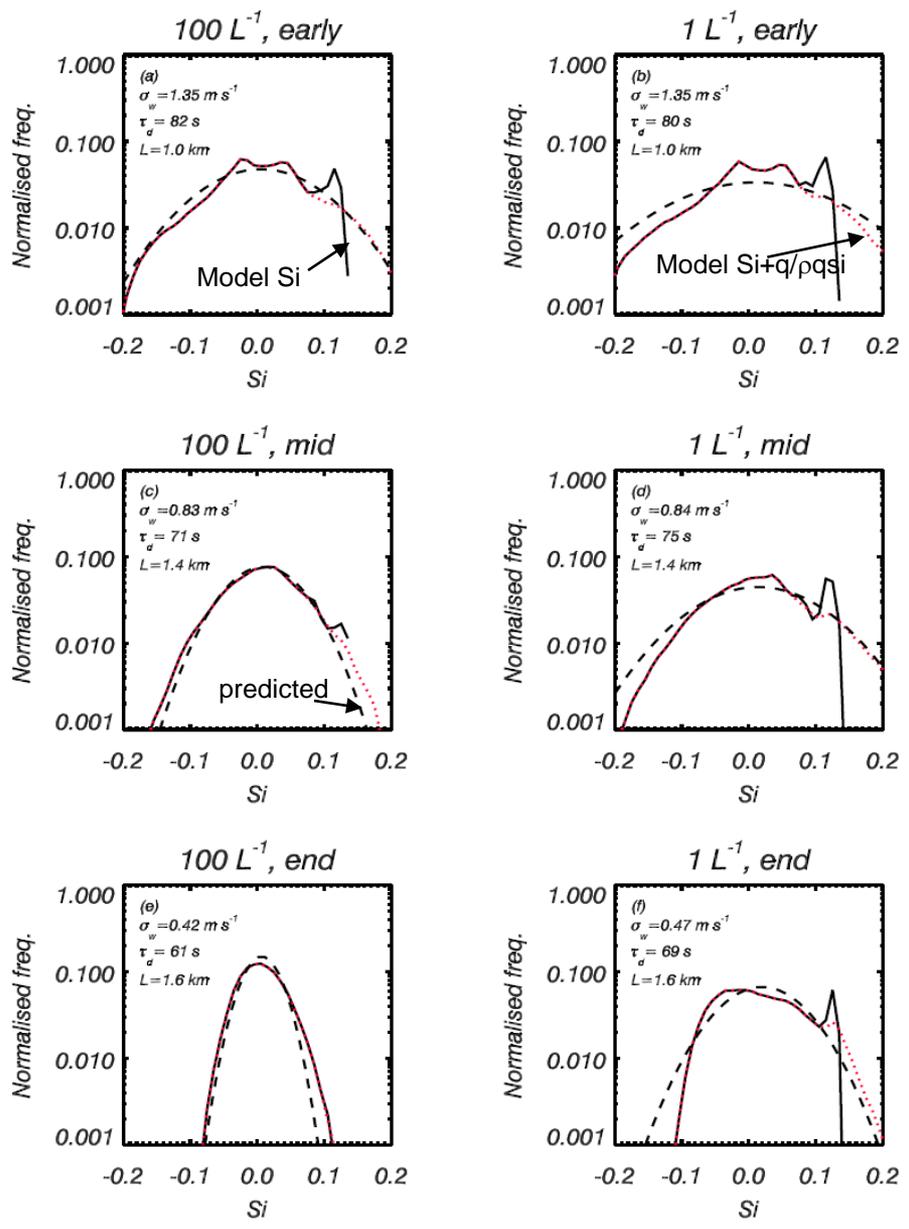
$B = b_i B_0 M_1$ microphys	and	$C = \left(\left(\frac{\epsilon}{L^2}\right)^{1/3}\right)$ homogenisation
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$$\langle S_i \rangle = S_E \frac{C}{B + C}$$

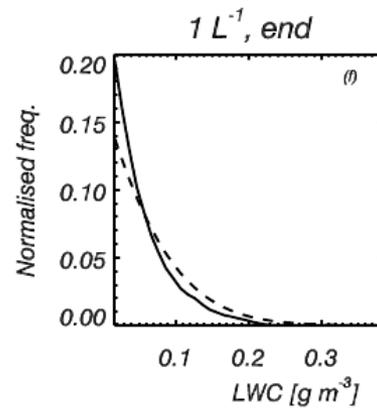
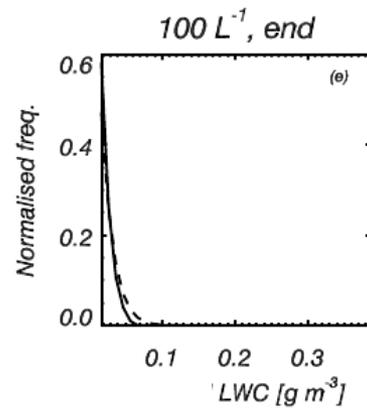
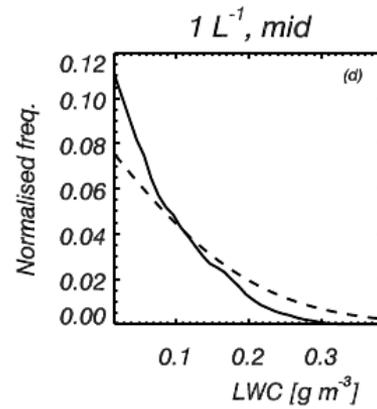
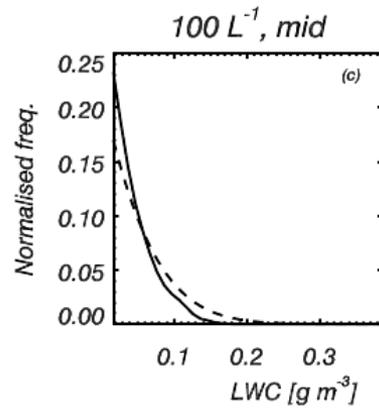
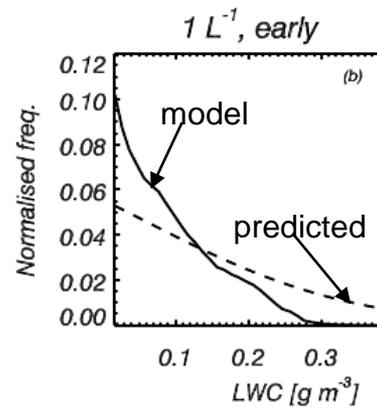
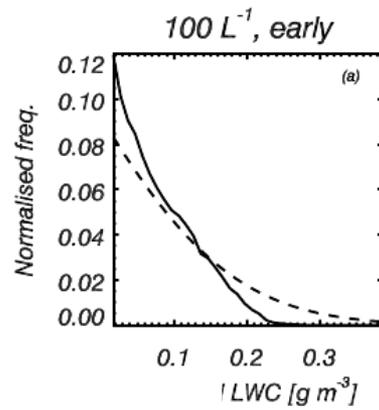
$$\sigma_S^2 = \frac{a_i^2 \sigma_w^2 \tau_d}{2(b_i B_0 M_1 + \left(\frac{\epsilon}{L^2}\right)^{1/3})}$$

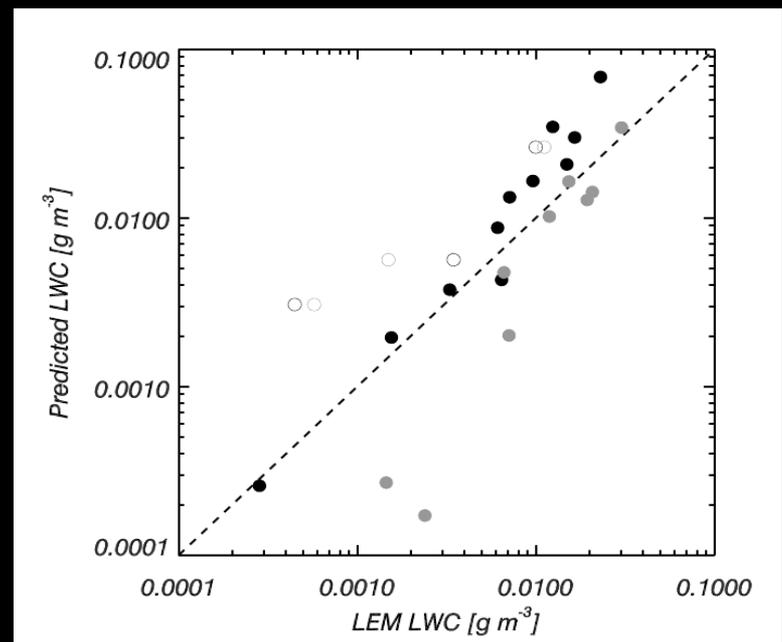
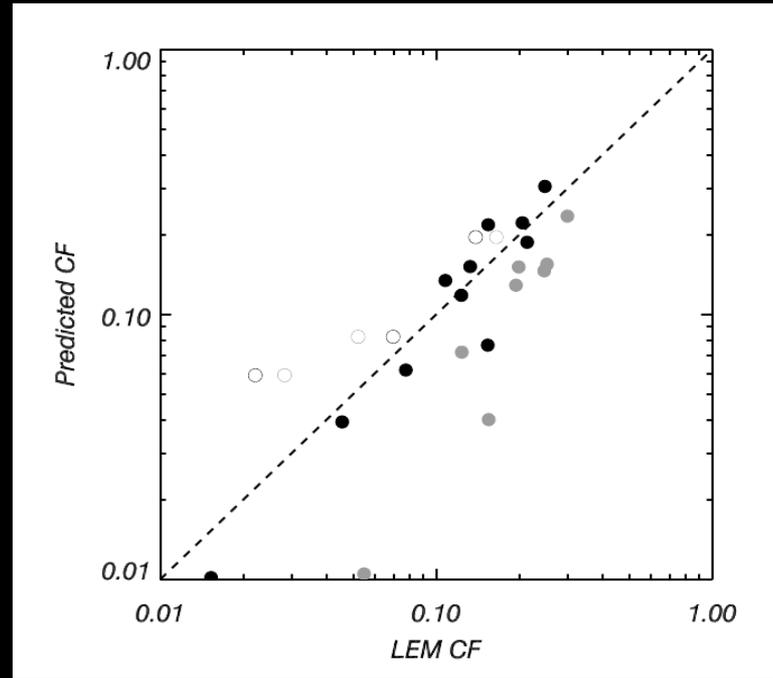
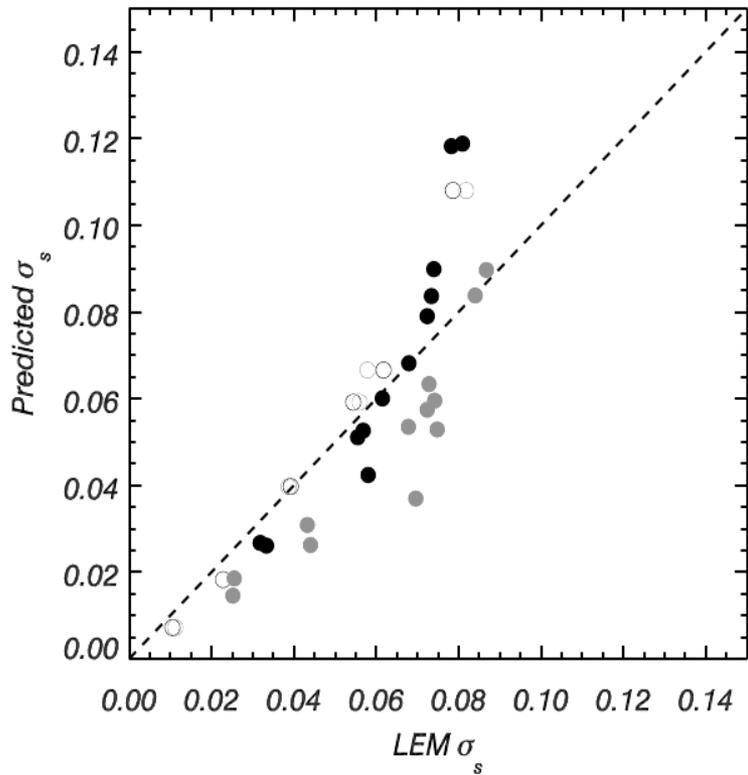


$$B = b_i B_0 M_1 \text{ and } C = \left(\frac{\epsilon}{L^2}\right)^{1/3}.$$



Supercooled liquid water distribution





Given turbulence characteristics (TKE, L) and ice cloud properties, we can estimate supercooled liquid cloud fraction (mixed-phase fraction) and the mean liquid water content of the turbulent region



General issues

Need to reassess treatment of mixed-phase –
physically based subgrid treatment

Need to understand if existing PSDs
representations are still adequate in the light of
shattering

Need to re-examine het. ice nucleation
representation – do we need to represent IN
prognostically

...