

# The Global Nonhydrostatic Atmospheric Model MPAS: Preliminary results from uniform and variable-resolution mesh tests

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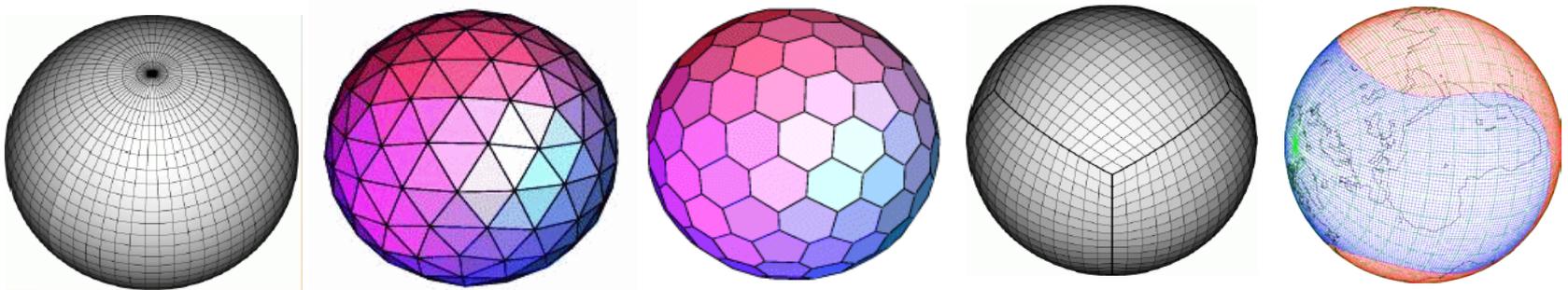
*Based on unstructured centroidal  
Voronoi (hexagonal) meshes using C-  
grid staggering and selective grid  
refinement.*

Jointly developed, primarily by NCAR  
and LANL/DOE, for weather,  
regional climate, and climate applications

MPAS infrastructure - NCAR, LANL, others.  
MPAS - Atmosphere (NCAR)  
MPAS - Ocean (LANL)  
MPAS - Ice, etc. (LANL and others)

# Primary drivers for global dynamical core development

- (1) Scalable solvers needed for the new computer architectures.
- (2) Nonhydrostatic global atmospheric models needed for cloud-permitting simulations ( $\Delta x \sim \text{kms}$ ).

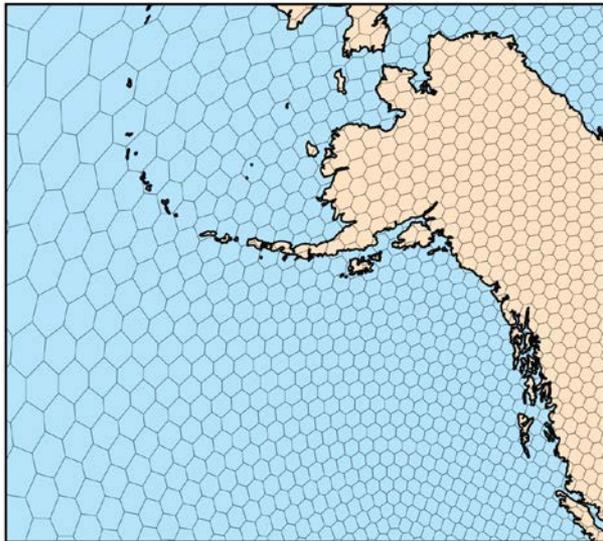
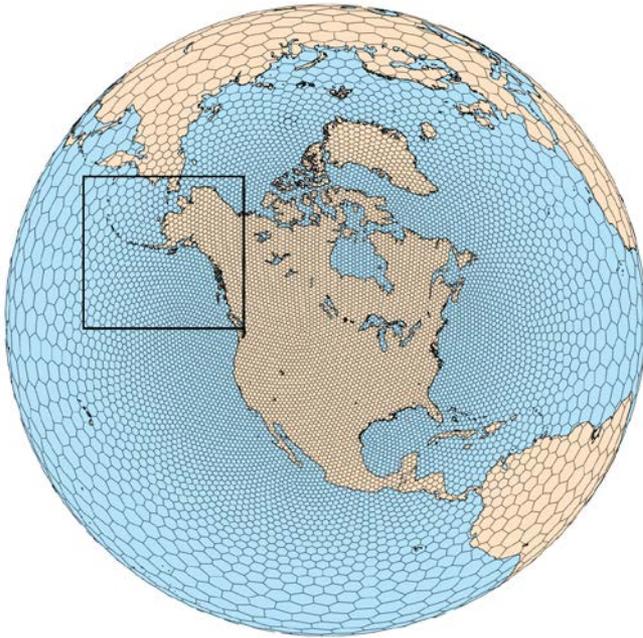


Newton Institute, PDEs on the sphere, major core development efforts:

- (1) FV methods on icosahedral (hexagonal) meshes.
- (2) FV, SE, and DG methods on the cubed sphere.

Most of these models are using horizontally-explicit integration techniques to facilitate scaling to  $10^5$ - $10^6$  processors or application to accelerators.

# MPAS: C-Grid Spherical Centroidal Voronoi Meshes



## Unstructured mesh

Mesh generation uses a density function.

Uniform resolution – traditional icosahedral mesh.

## Centroidal Voronoi

Mostly *hexagons*, some pentagons and 7-sided cells.

Cell centers are at cell center-of-mass.

Lines connecting cell centers intersect cell edges at right angles.

Lines connecting cell centers are bisected by cell edge.

## C-grid

Solve for normal velocities on cell edges.

## Equations

Fully compressible

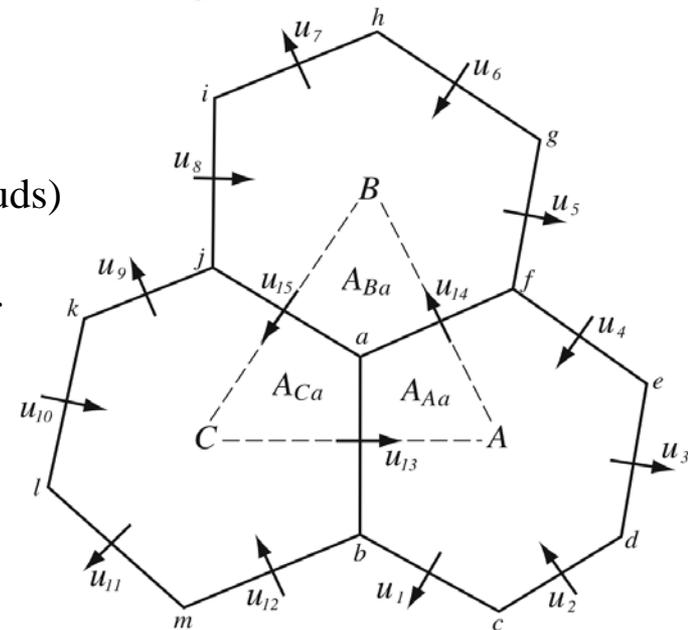
nonhydrostatic equations

(*explicit* simulation of clouds)

## Solver Technology

Integration scheme similar to WRF.

WRF-NRCM physics



# MPAS Nonhydrostatic Atmospheric Solver

## Nonhydrostatic formulation

### Equations

- Prognostic equations for coupled variables.
- Generalized height coordinate.
- Horizontally vector invariant eqn set.
- Continuity equation for dry air mass.
- Thermodynamic equation for coupled potential temperature.

### Time integration scheme

As in Advanced Research WRF -  
Split horizontally-explicit  
vertically-implicit Runge-Kutta (3rd order)

Variables:

$$(U, V, \Omega, \Theta, Q_j) = \tilde{\rho}_d \cdot (u, v, \dot{\eta}, \theta, q_j)$$

Vertical coordinate:

$$z = \zeta + A(\zeta) h_s(x, y, \zeta)$$

Prognostic equations:

$$\begin{aligned} \frac{\partial \mathbf{V}_H}{\partial t} = & -\frac{\rho_d}{\rho_m} \left[ \nabla_\zeta \left( \frac{p}{\zeta_z} \right) - \frac{\partial \mathbf{z}_H p}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H \\ & - \mathbf{v}_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \mathbf{v}_H}{\partial \zeta} - \rho_d \nabla_\zeta K - eW \cos \alpha_r - \frac{uW}{r_e} + \mathbf{F}_{V_H}, \end{aligned}$$

$$\begin{aligned} \frac{\partial W}{\partial t} = & -\frac{\rho_d}{\rho_m} \left[ \frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v} W)_\zeta \\ & + \frac{uU + vV}{r_e} + e(U \cos \alpha_r - V \sin \alpha_r) + F_W, \end{aligned}$$

$$\frac{\partial \Theta_m}{\partial t} = -(\nabla \cdot \mathbf{V} \theta_m)_\zeta + F_{\Theta_m},$$

$$\frac{\partial \tilde{\rho}_d}{\partial t} = -(\nabla \cdot \mathbf{V})_\zeta,$$

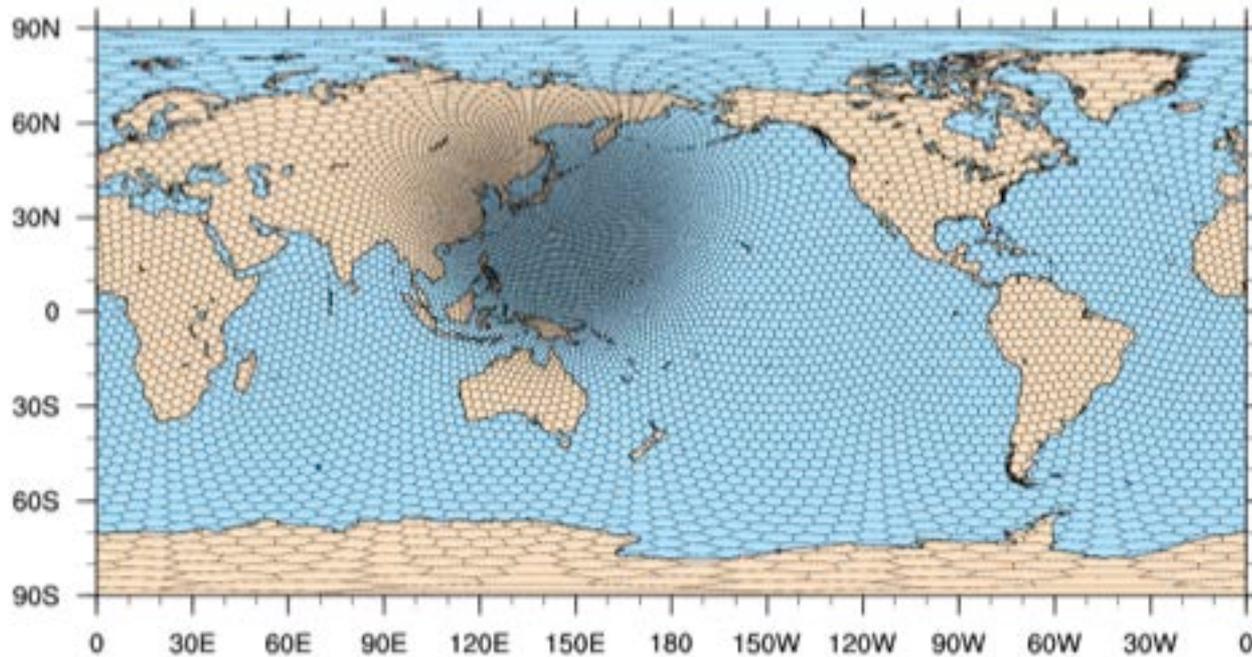
$$\frac{\partial Q_j}{\partial t} = -(\nabla \cdot \mathbf{V} q_j)_\zeta + \rho_d S_j + F_{Q_j},$$

Diagnostics and definitions:

$$\theta_m = \theta [1 + (R_v/R_d)q_v] \quad p = p_0 \left( \frac{R_d \zeta_z \Theta_m}{p_0} \right)^\gamma$$

$$\frac{\rho_m}{\rho_d} = 1 + q_v + q_c + q_r + \dots$$

# Variable Resolution Meshes



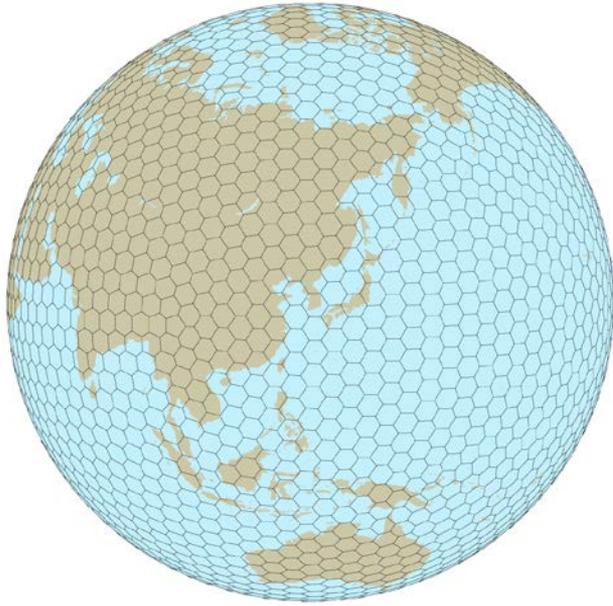
Applications: Regional climate, weather prediction  
(address problems with one-way nesting)

Static refinement: Obvious next step for applications

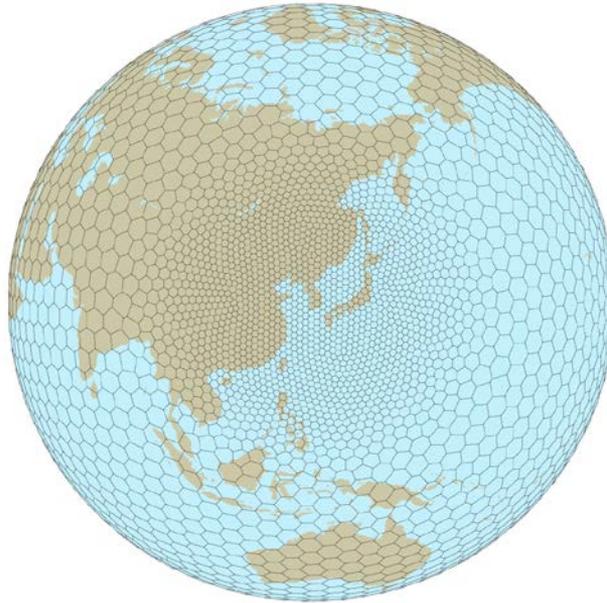
Smooth conforming meshes: Motivations

- (i) unstructured mesh looks the same everywhere
- (ii) preserve the accuracy of the numerics
- (iii) minimize wave-reflection problems  
at mesh density transitions

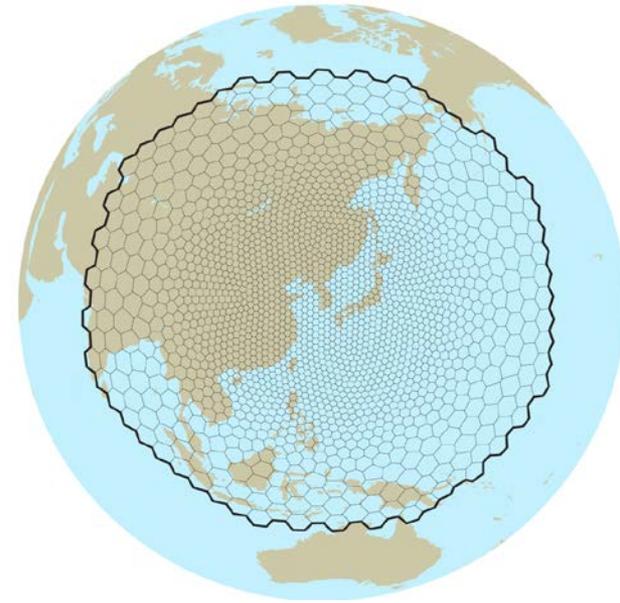
# MPAS Global Mesh and Integration Options



Global Uniform Mesh



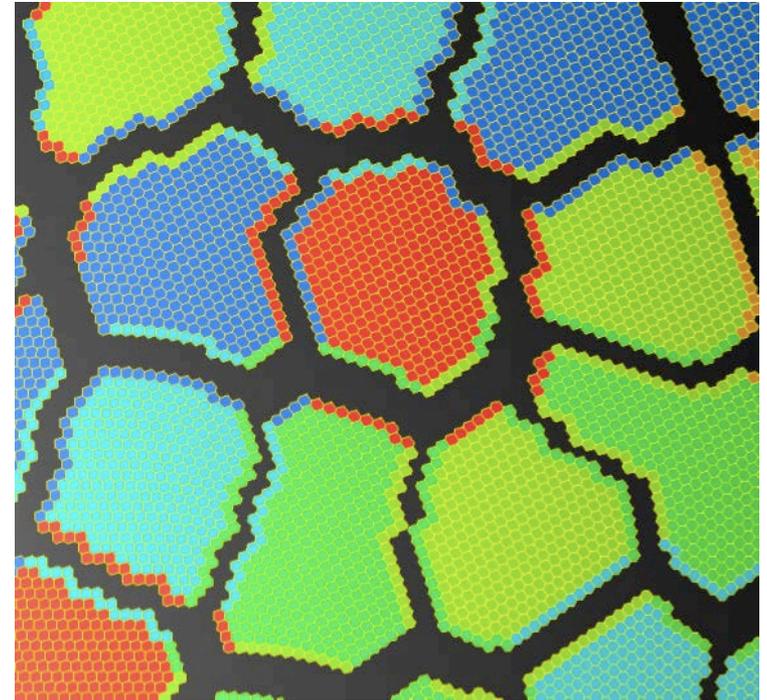
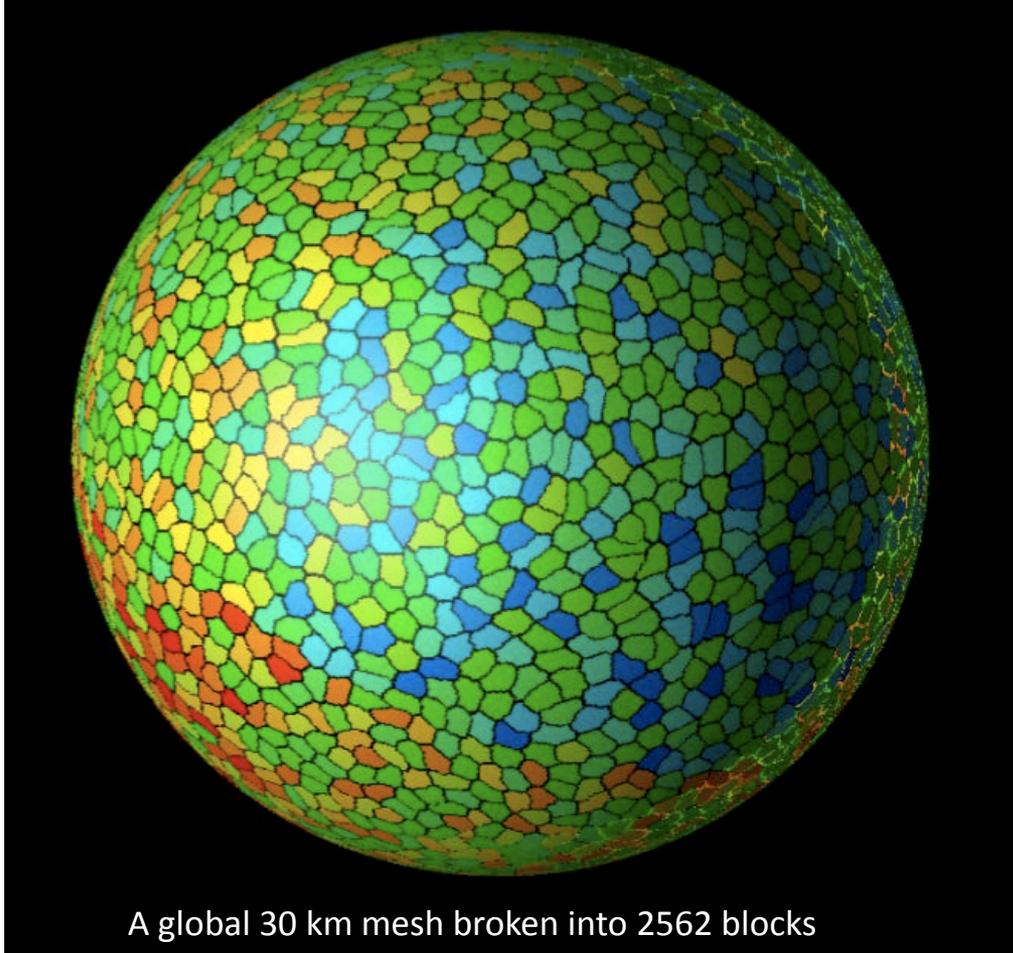
Global Variable Resolution Mesh



Regional Mesh - driven by  
(1) previous global MPAS run  
(no spatial interpolation needed!)  
(2) other global model run  
(3) analyses

*Voronoi meshes will allow us to cleanly incorporate both downscaling and upscaling effects (avoiding the problems in traditional grid nesting) and to assess the accuracy of the traditional downscaling approaches used in regional climate and NWP applications.*

# Domain Decomposition for Parallel Processing



# Preliminary MPAS-ANH Scaling Results

## 120-km simulations:

- 40962 grid cells
- 93% efficiency on 240 cores (relative to 8 cores)
- 79% efficiency on 504 cores
- 63% efficiency on 912 cores

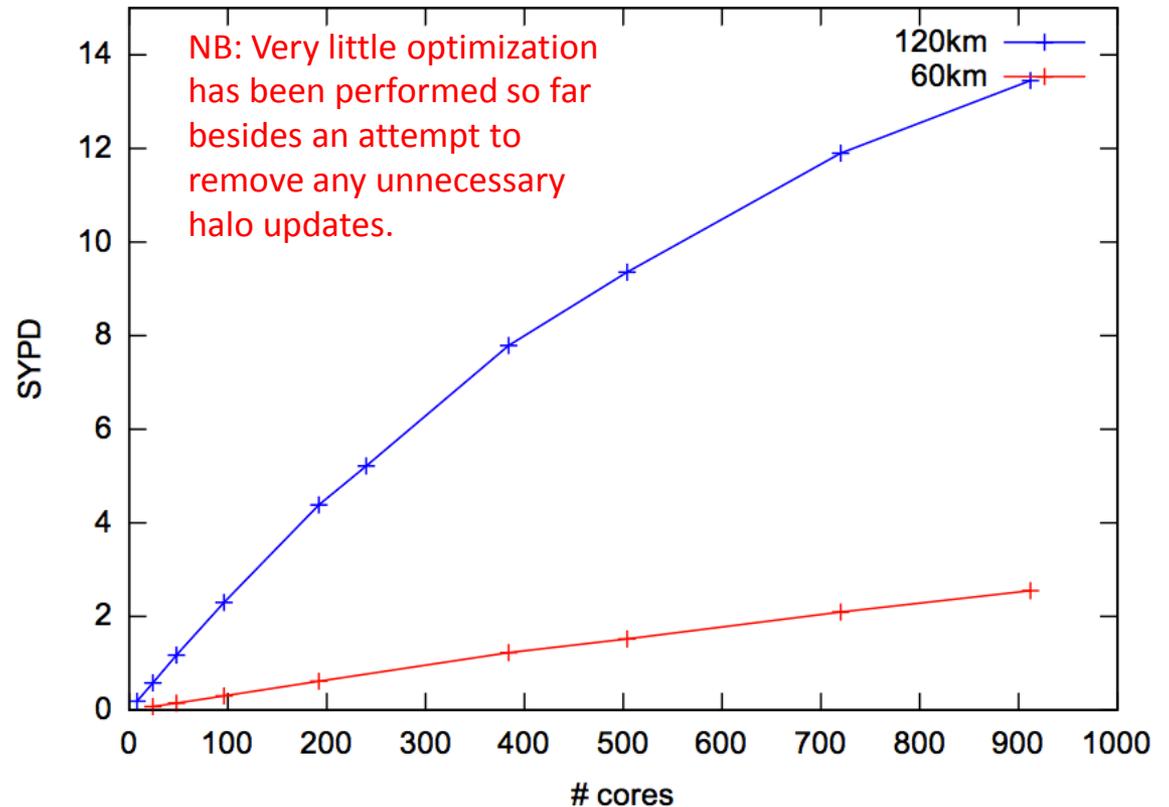
## 60-km simulations:

- 163842 grid cells
- 98% efficiency on 504 cores (relative to 24 cores)
- 91% efficiency on 912 cores

## Weak-scaling extrapolation:

- 60 km – 912 cores (91%)
- 30 km – 3,648 cores
- 15 km – 14,592 cores
- 7.7 km – 58,368 cores
- 3.8 km – 233,472 cores
- 1.9 km – 933,888 cores

MPAS-A non-hydrostatic core scalability



Simulation rate given for the dynamical core only

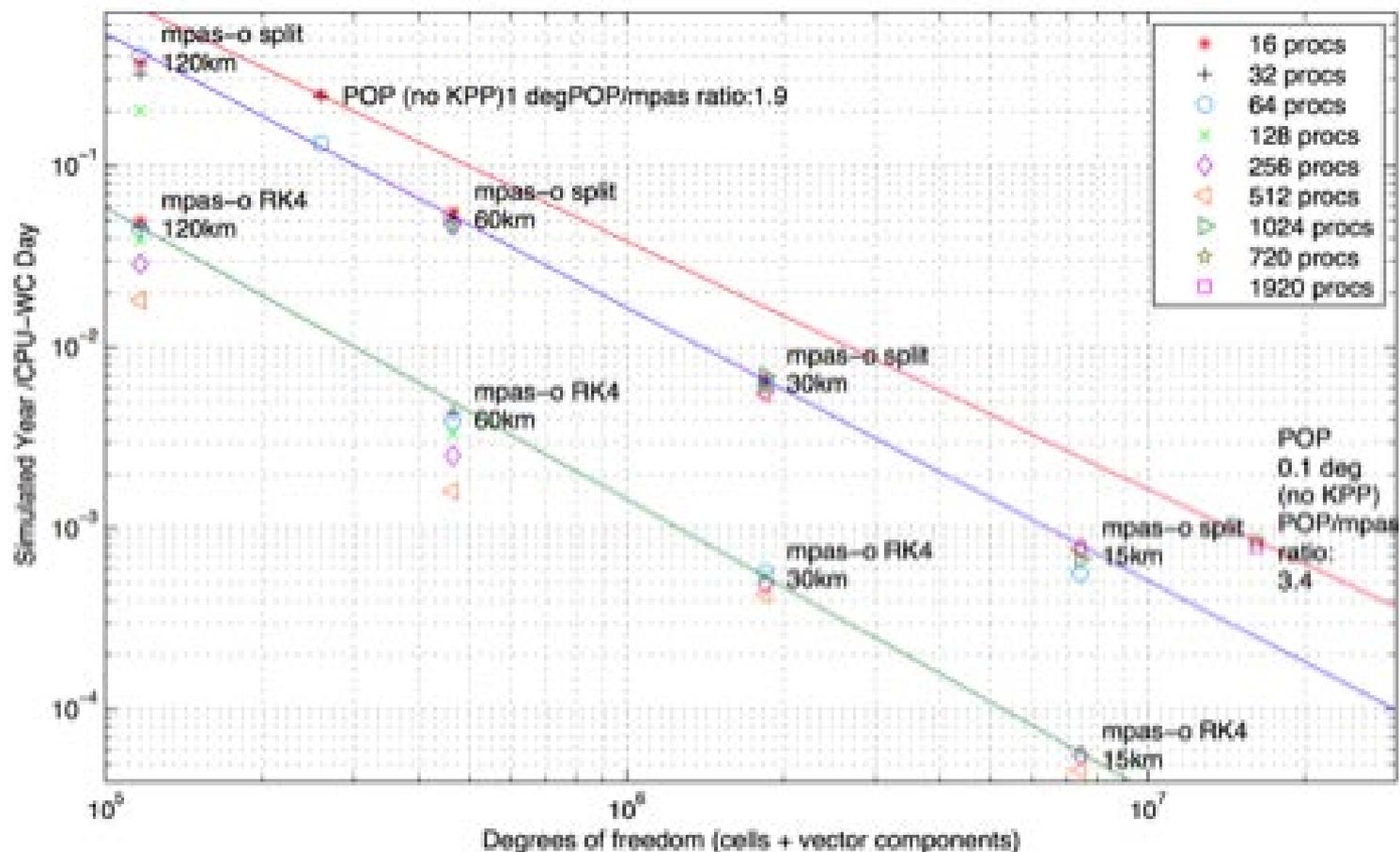
- 8 scalars w/positive-definite advection
- 41 vertical levels
- All runs on a Cray XT5m (lynx)
- MPI parallelism only; no OpenMP yet

# Preliminary MPAS-Ocean Scaling Results

Hydrostatic ocean model

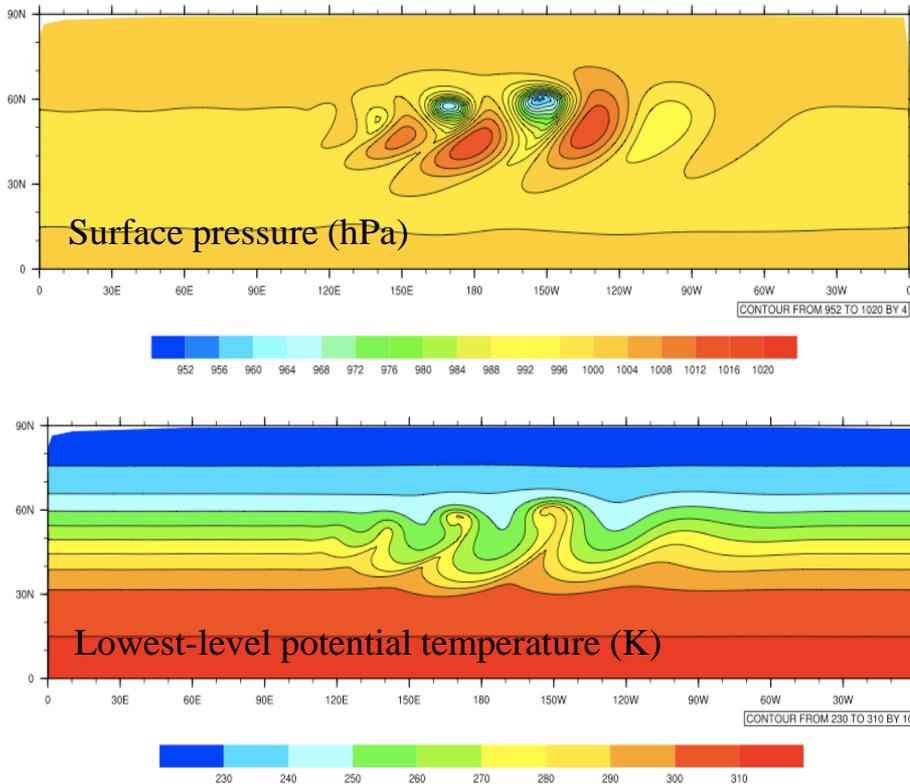
(1) Explicit RK4 integration

(2) Timesplit

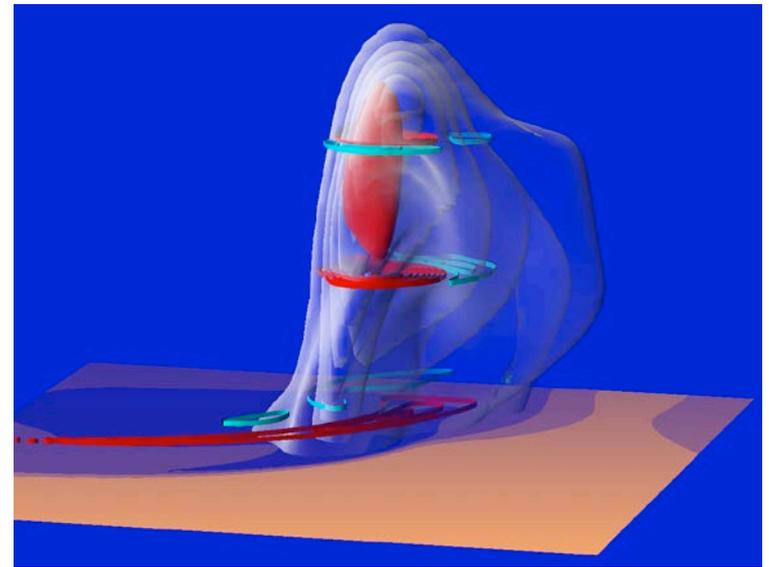


# Model for Prediction Across Scales: MPAS

*Global simulation tests*  
*Baroclinic waves*



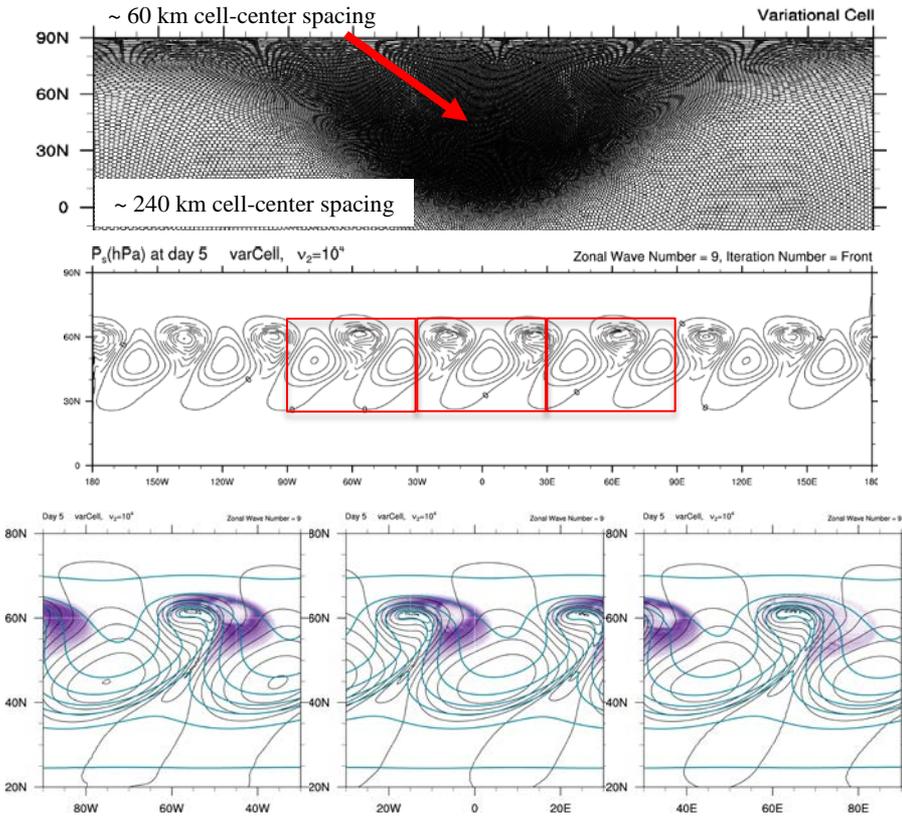
*Convection on a Cartesian plane*  
*Supercell thunderstorms*



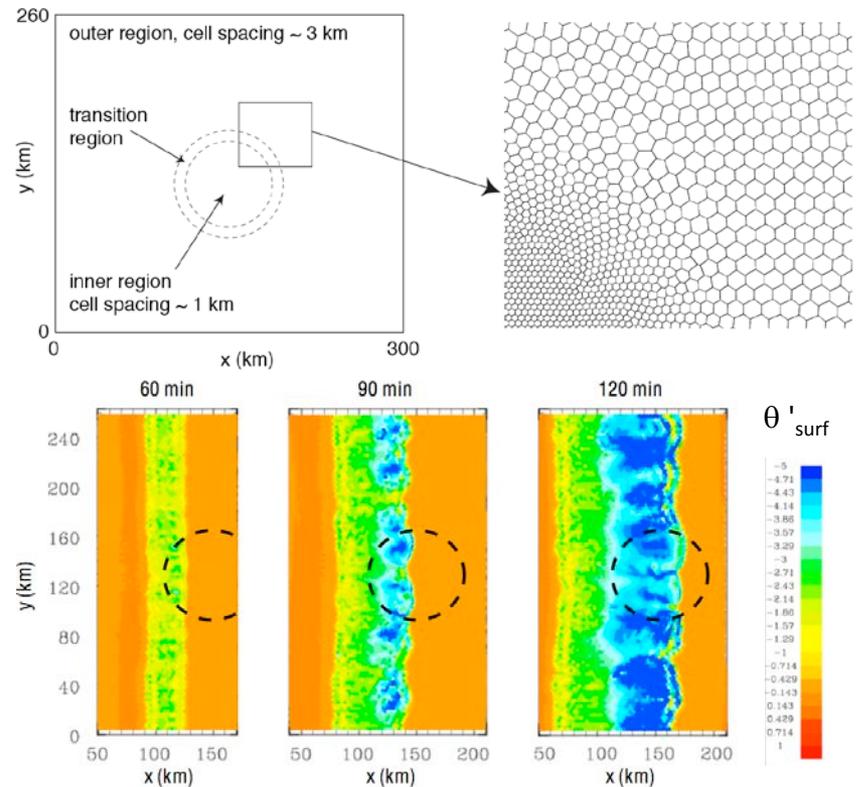
- W contours at 1, 5, and 10 km (c.i. = 3 m/s)
- 30 m/s W surface shaded in red
- Rainwater surfaces - transparent shells
- Surface temperature shaded on baseplane
- 500 meter mesh

# Model for Prediction Across Scales: MPAS

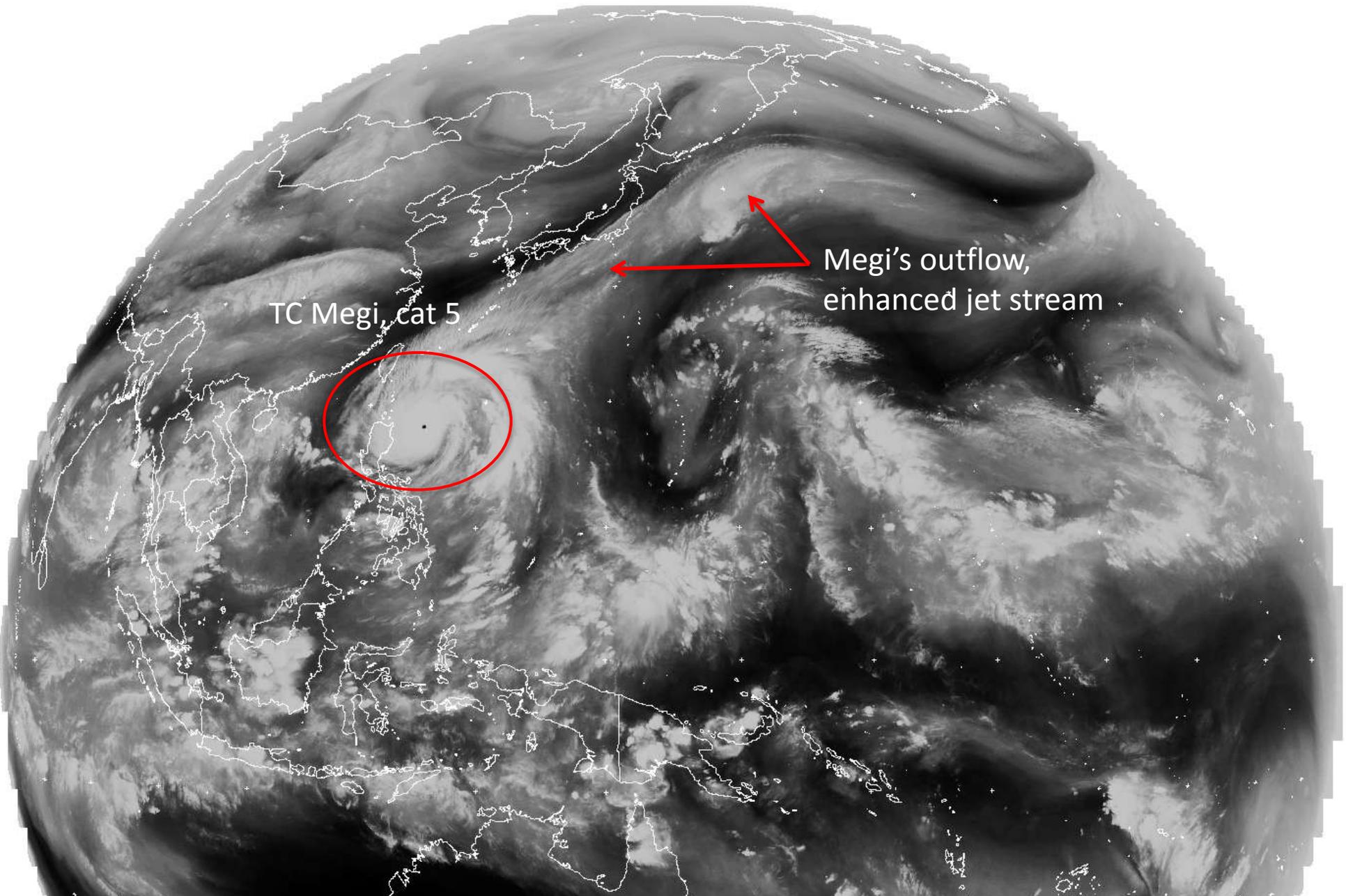
*Global variable-resolution  
moist baroclinic waves*



*Squall-lines on a Cartesian plane  
using a variable-resolution mesh*



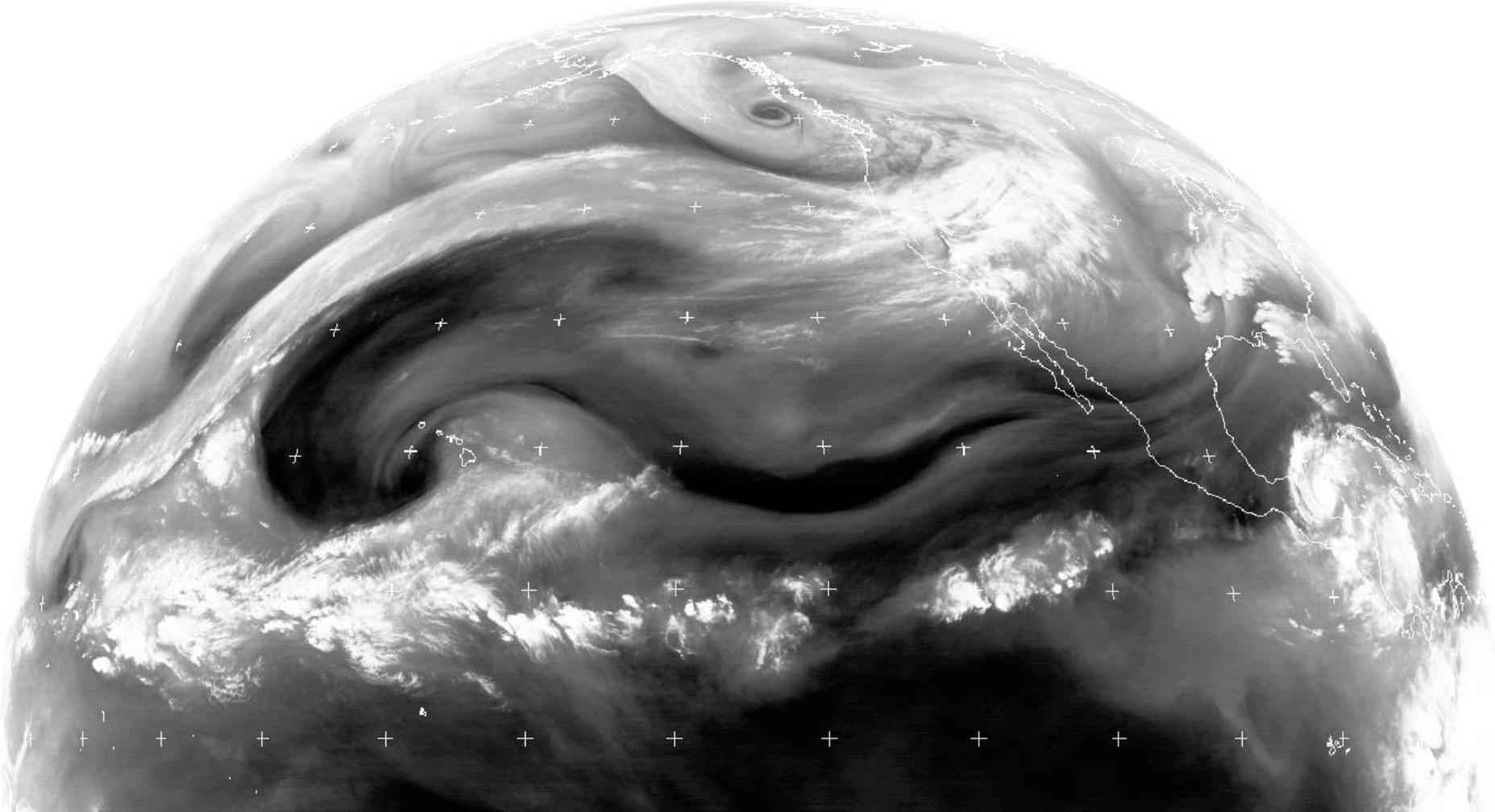
mtsatsat, 2010-10-17, 12 UTC, vapor channel



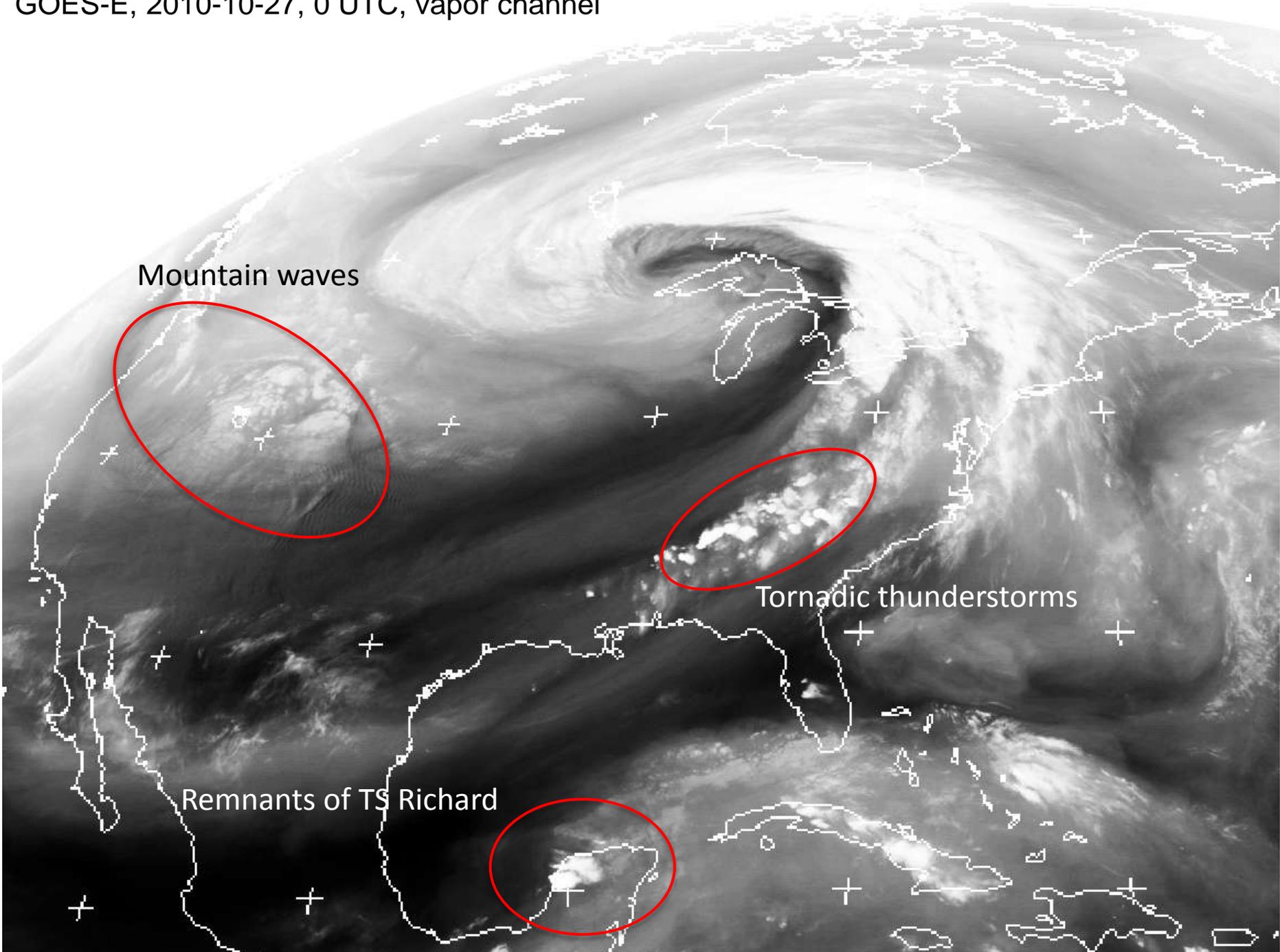
TC Megi, cat 5

Megi's outflow,  
enhanced jet stream

GOES-W, 2010-10-25, 0 UTC, vapor channel



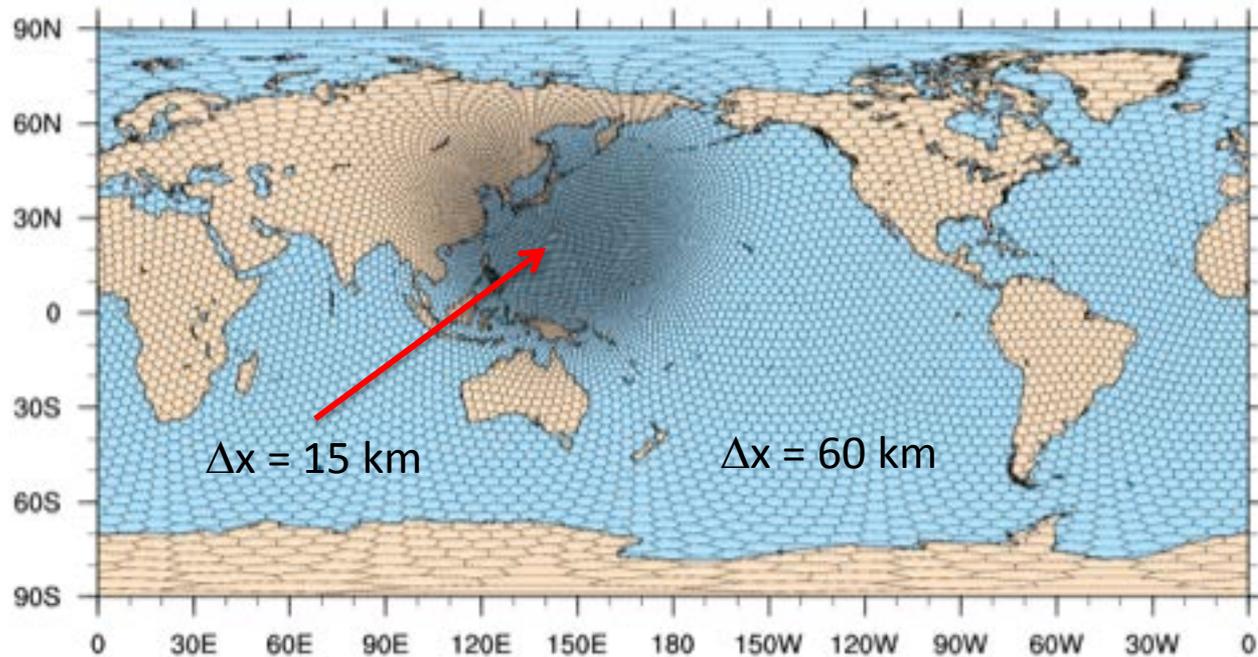
GOES-E, 2010-10-27, 0 UTC, vapor channel



# MPAS Forecast Tests

Current MPAS Physics:  
WSM6 cloud microphysics  
Kain\_Fritsch or Tiedtke convection  
Monin-Obukhov surface layer  
YSU pbl, Noah land-surface  
RRTMG lw and sw or CAM radiation.

MPAS mesh (4x finer than below), 41 levels

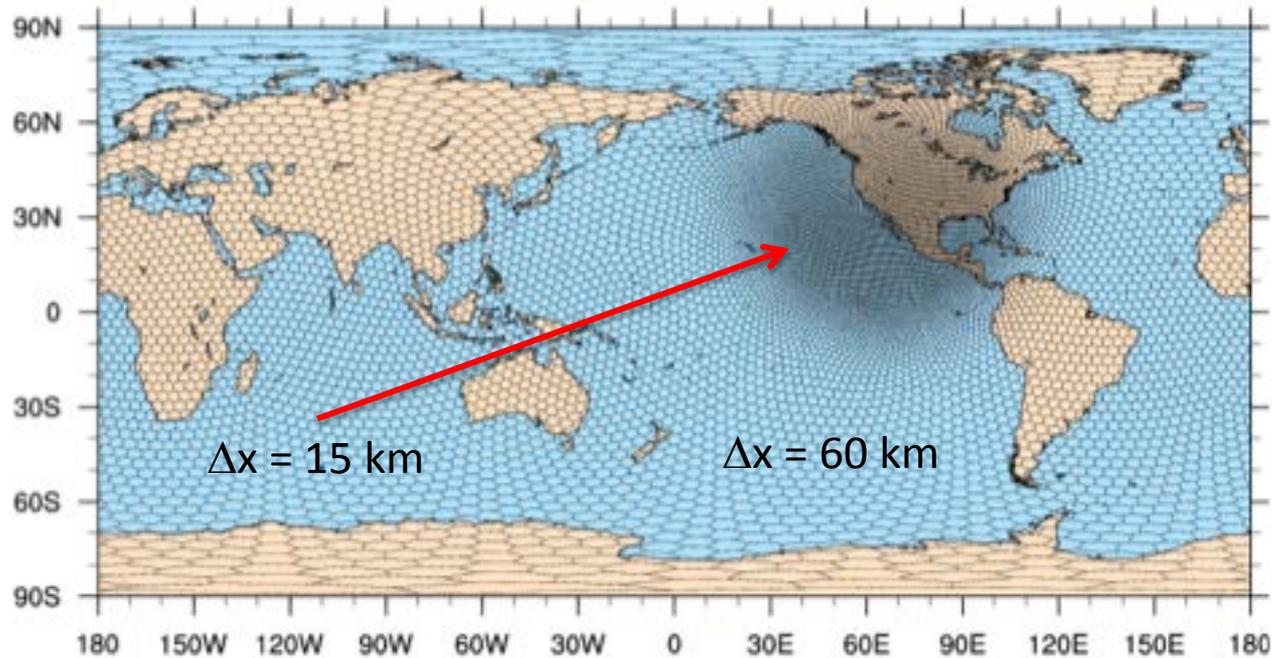


Western Pacific refinement

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RRTMG lw and sw or CAM radiation.

MPAS mesh (4x finer than below), 41 levels



Eastern Pacific refinement

# Variable Resolution Mesh Tests

$\Delta t$  is constant on the variable-resolution mesh.

Smagorinsky:  $K_h = c_s^2 l^2 |Def|$   
 $l^2$  scales with  $\Delta x^2$

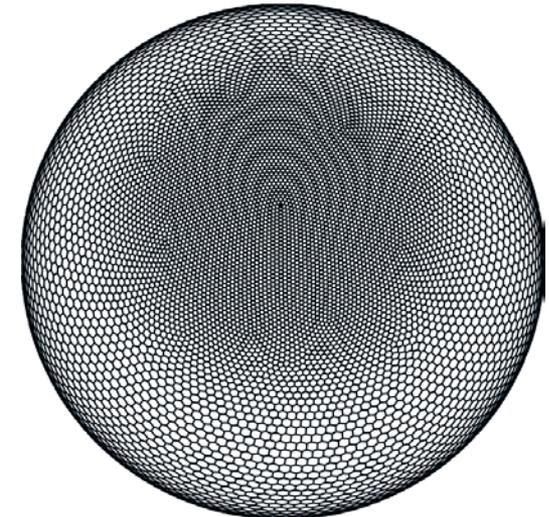
Viscosity and hyper-viscosity formulations:

$K_2 \nabla_{\zeta}^2 \phi$   $K_2$  scales with  $\Delta x^2$

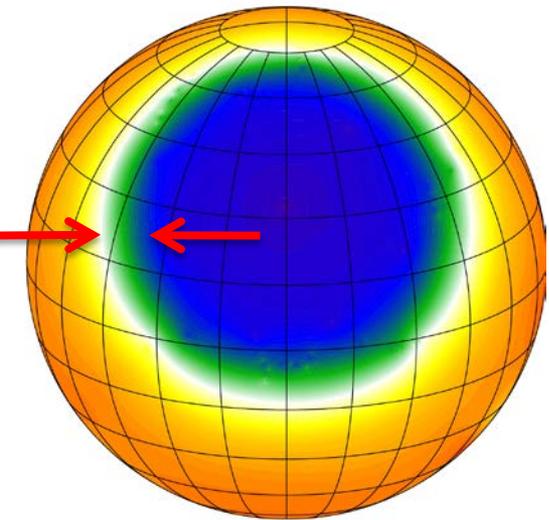
$K_4 \nabla_{\zeta}^2 (\nabla_{\zeta}^2 \phi)$   $K_4$  scales with  $\Delta x^4$

Locally  $2\Delta$  waves are damped at same rate.

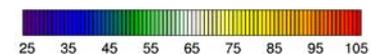
MPAS  
TR-10  
mesh



Transition-zone  
width



Cell to Cell Dist. (km)

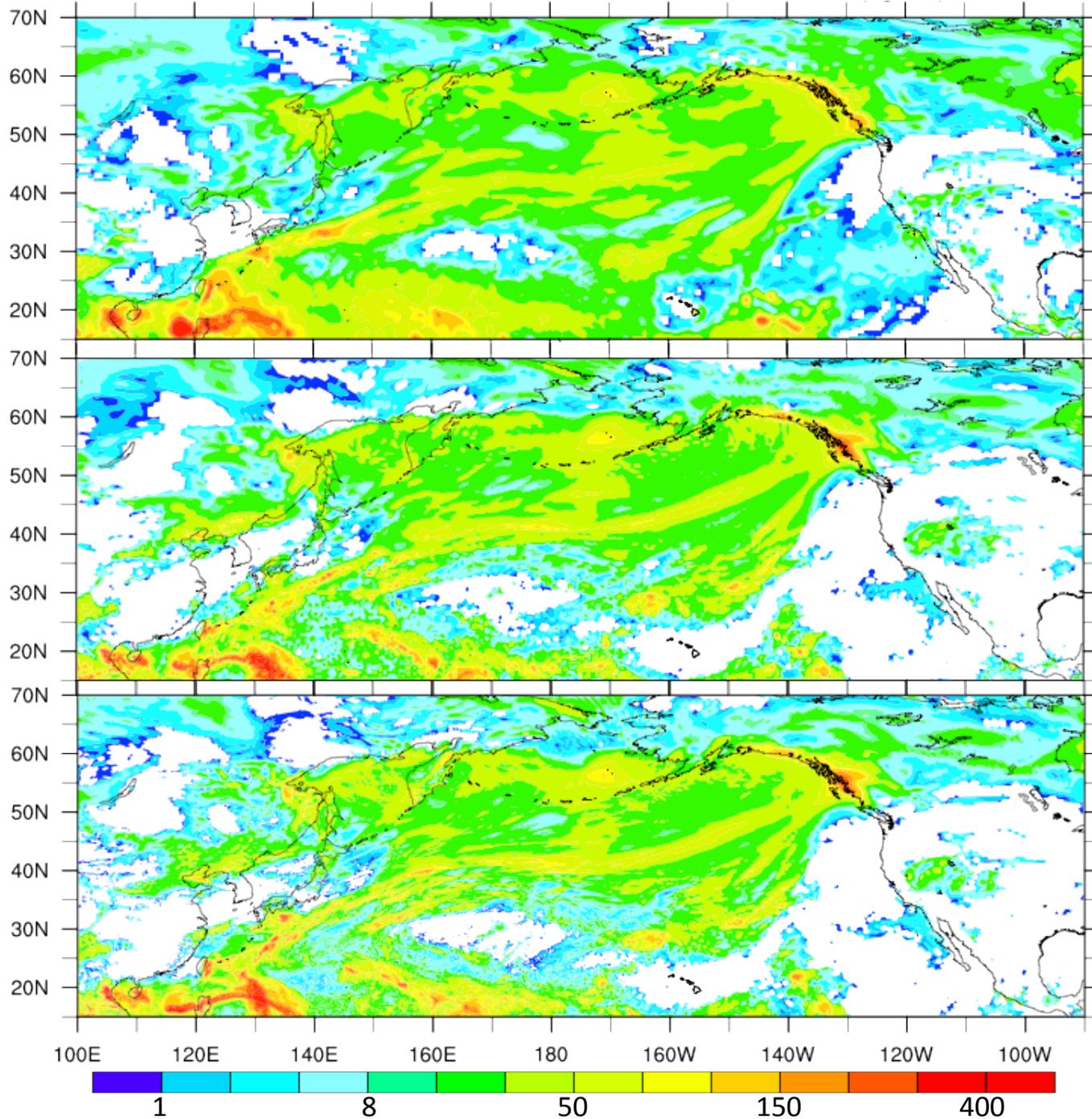


20 October 2010  
5 day accumulated  
precipitation (mm)

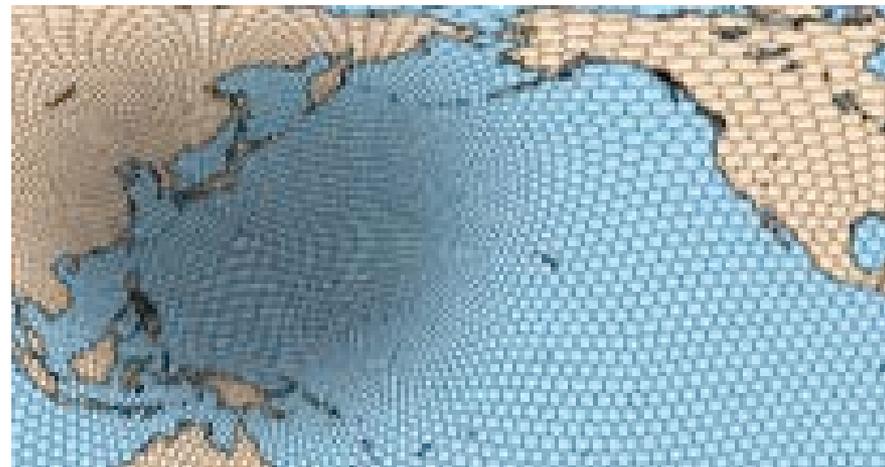
CFSR (~ 40 km)

MPAS (60 km)  
uniform resolution  
Smagorinsky

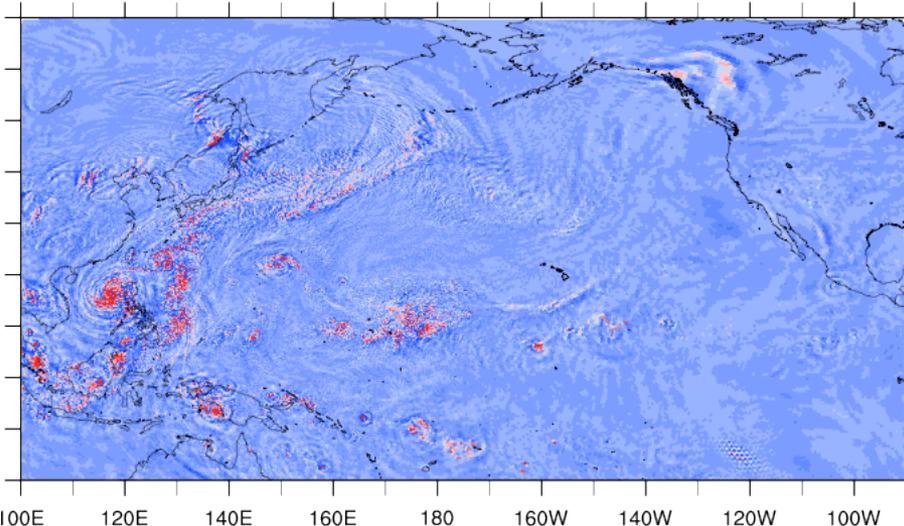
MPAS (60-15 km)  
variable resolution  
Western Pacific ref.  
Smagorinsky,  
( $\Delta x^2$  scaling)



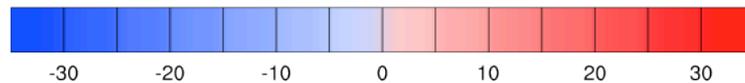
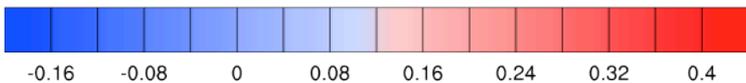
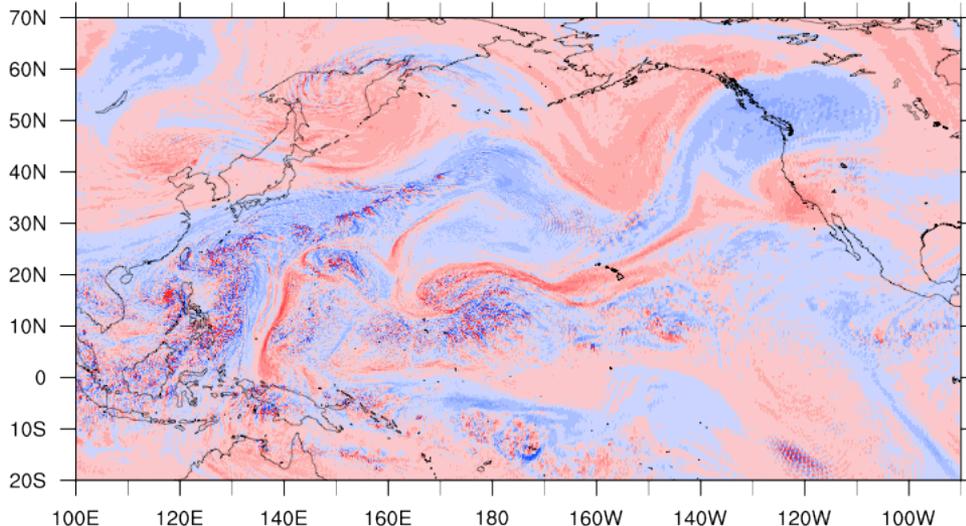
MPAS (60 – 15 km mesh)  
Western Pacific refinement  
15 October initialization  
Smagorinsky,  $\Delta x^2$  scaling



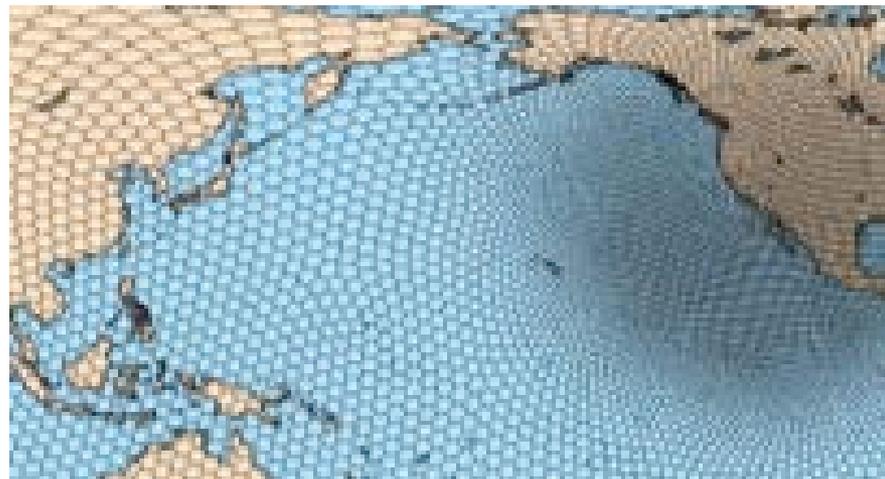
MPAS: 0000 UTC 19 OCT 2010 200 hPa VERTICAL VELOCITY ( $\text{m s}^{-1}$ )



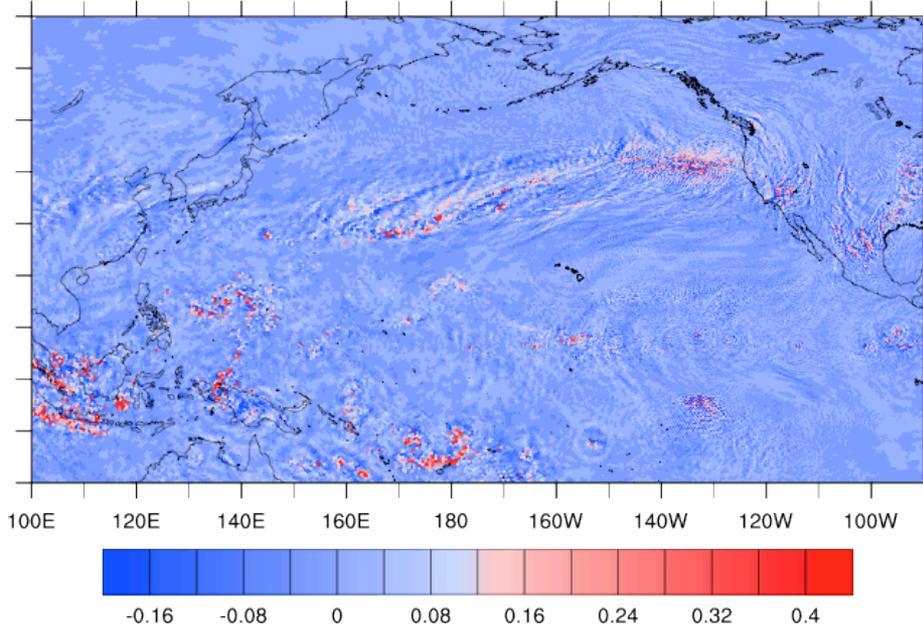
MPAS: 0000 UTC 19 OCT 2010 200 hPa VORTICITY ( $\times 1.e-05 \text{ s}^{-1}$ )



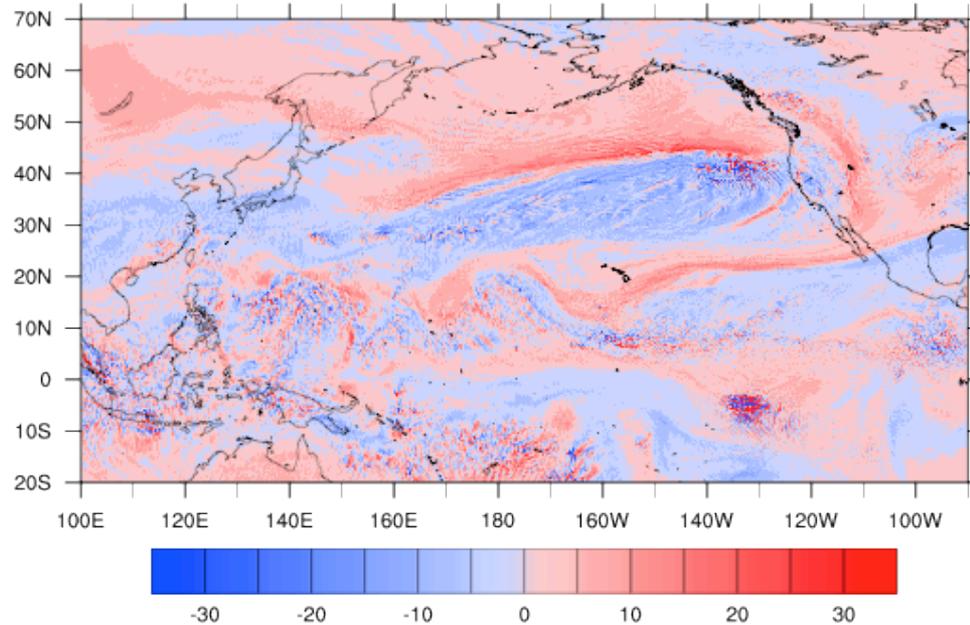
MPAS (60 – 15 km mesh)  
Eastern Pacific refinement  
21 October initialization  
Smagorinsky,  $\Delta x^2$  scaling



MPAS: 0000 UTC 24 OCT 2010 200 hPa VERTICAL VELOCITY ( $\text{m s}^{-1}$ )



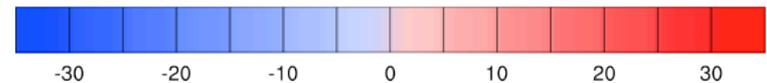
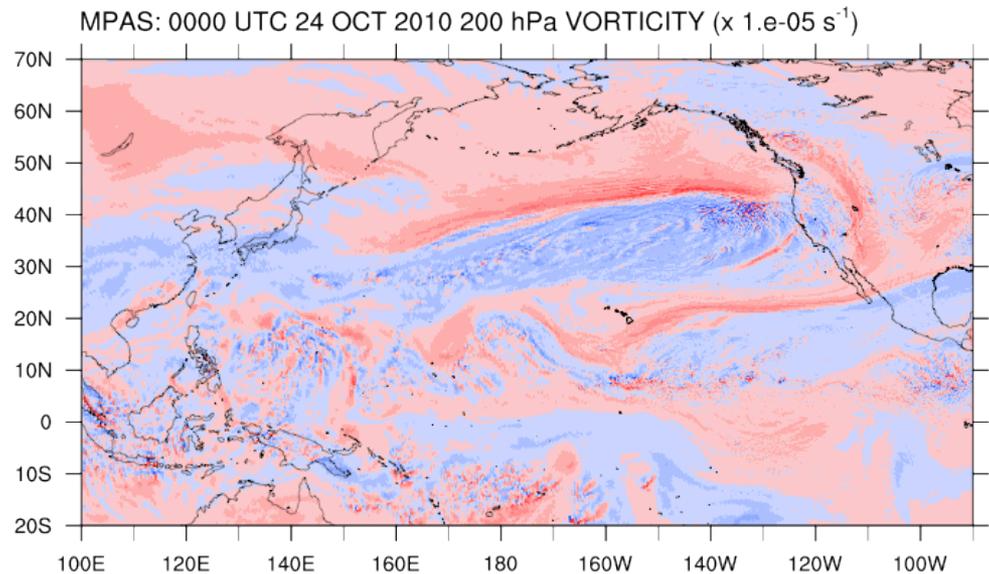
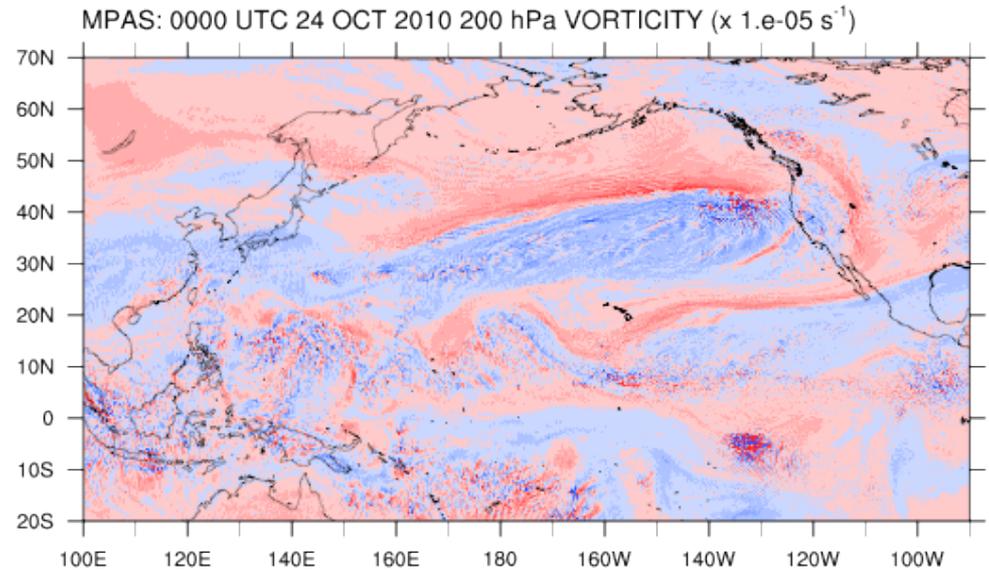
MPAS: 0000 UTC 24 OCT 2010 200 hPa VORTICITY ( $\times 1. \text{e-}05 \text{ s}^{-1}$ )



# MPAS (60 – 15 km mesh) Eastern Pacific refinement 21 October initialization

East-Pac mesh ( $\Delta x = 60\text{-}15\text{ km}$ )  
Smagorinsky,  $\Delta x^2$  scaling

East-Pac mesh ( $\Delta x = 60\text{-}15\text{ km}$ )  
Smagorinsky,  $\Delta x^2$  scaling;  
background  $K_4 = 2 \times 10^{10} \text{ m}^4 \text{ s}^{-1}$   
(15 km mesh value,  $\Delta x^4$  scaling)



# Summary

Preliminary results show adequate scaling for both ocean and atmospheric cores.  
(MPI-only configuration)

Variable-resolution mesh simulations in both idealized and full-physics NWP configurations suggest that the smooth mesh transitions are viable for applications.

The model filters, appropriately scaled by the mesh density, appear to be behaving properly.

*Current and future work:*

- Further optimization (MPI, OpenMP implementation).
- I/O (PIO) testing and development.
- Parameterizations (physics) need attention (scale-aware physics).
- Continued testing, hydrostatic – nonhydrostatic regime mesh transition.
- Applications.