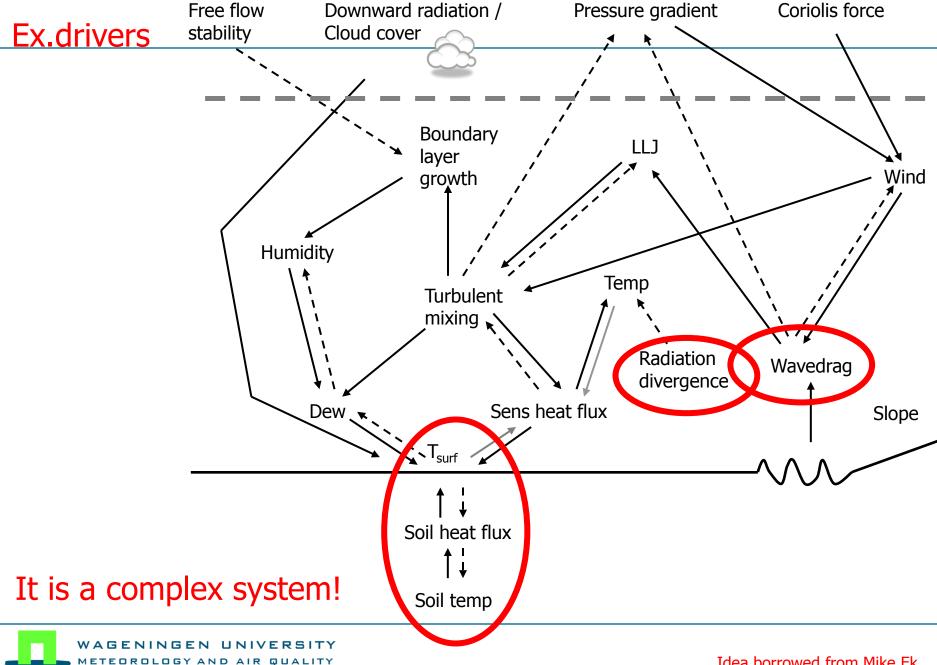
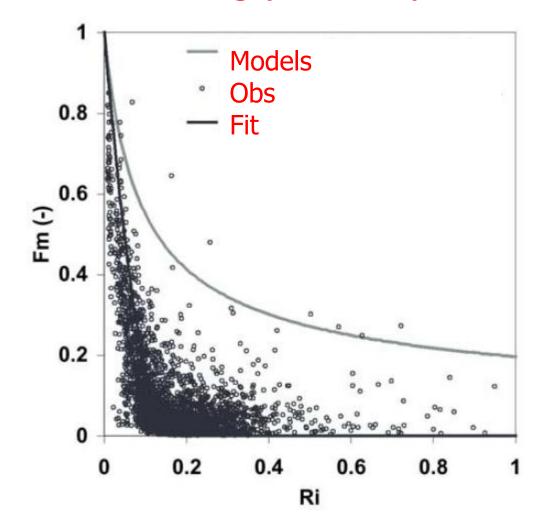
Stable boundary layer issues

Gert-Jan Steeneveld GABLS workshop Nov 2011, ECWMF

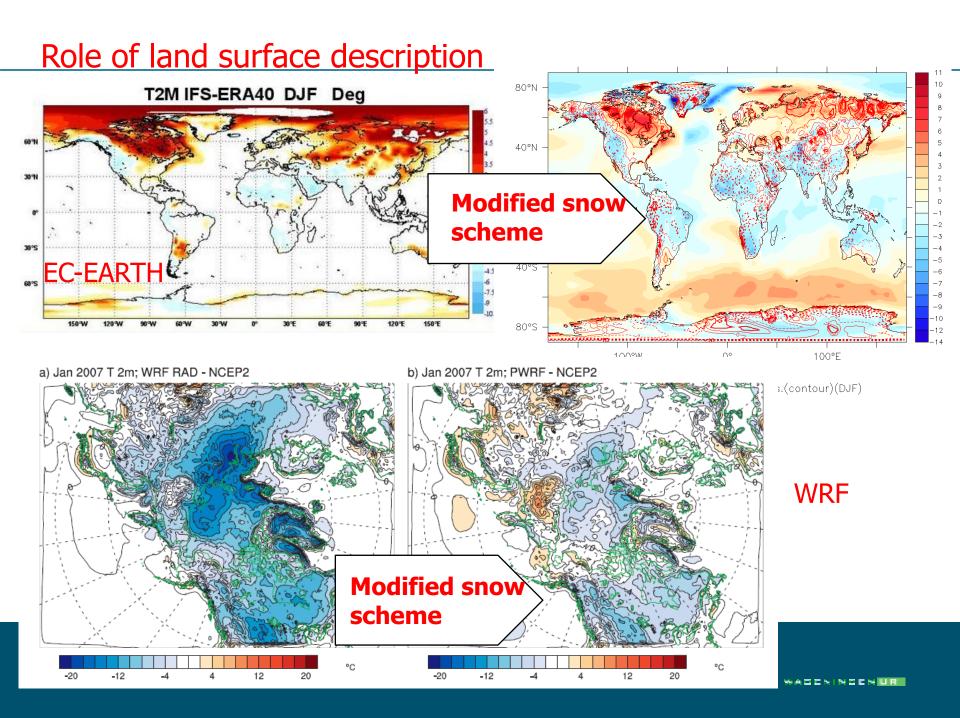


Turbulent mixing (CASES-99)

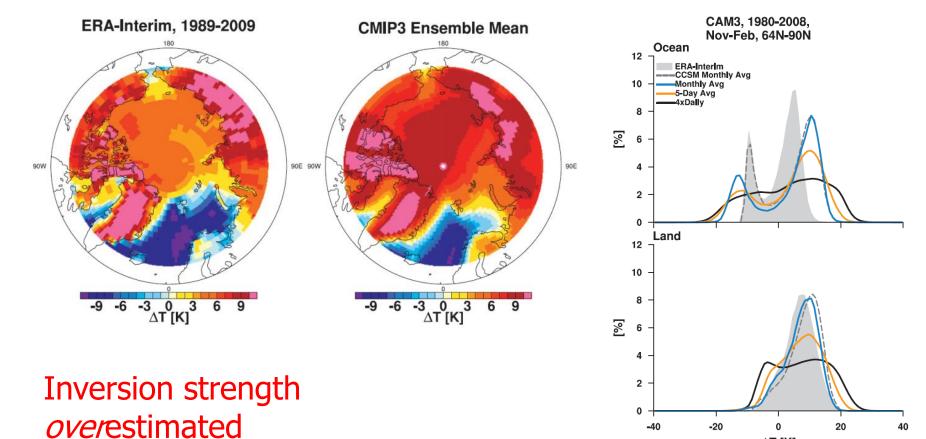


Do models still need long tail formulations?

Over land only?



Arctic climate



-40

-20

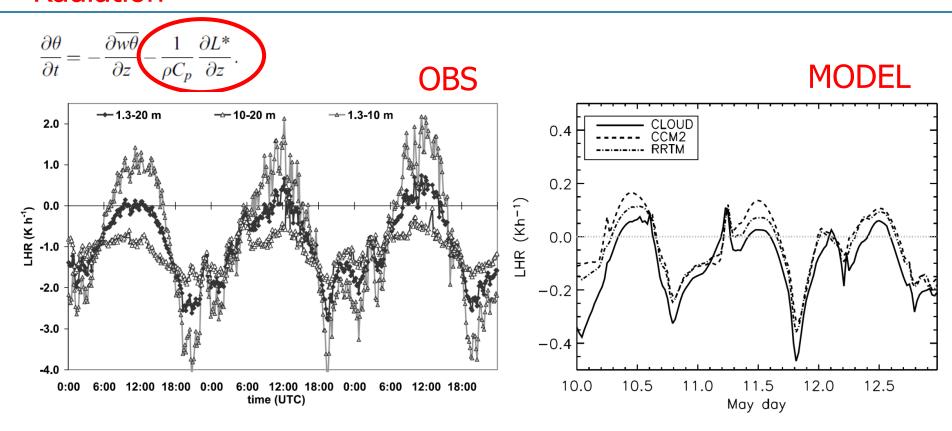
20

 $\Delta T [K]$

40

Medeiros, Brian, Clara Deser, Robert A. Tomas, Jennifer E. Kay, 2011: Arctic Inversion Strength in Climate Models. J. Climate, 24, 4733-4740. doi: 10.1175/2011JCLI3968.1

Radiation



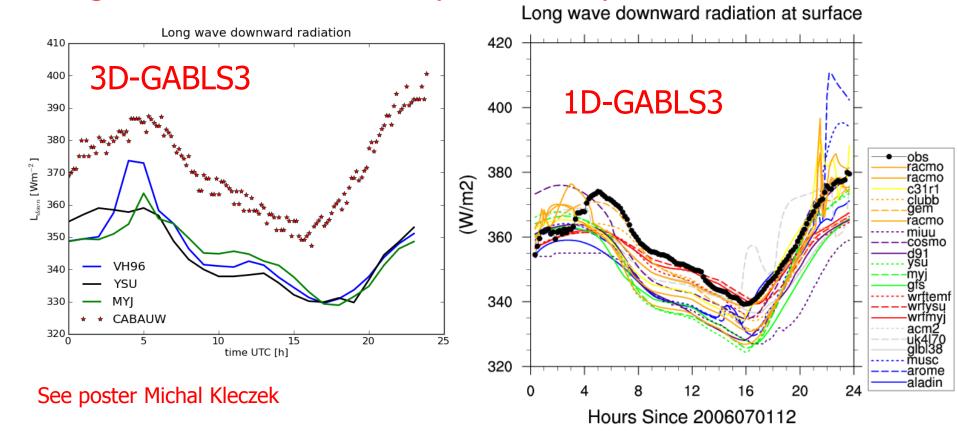
In weak wind conditions, radiation divergence is substantial contributor in nature, but not in atmospheric models. How to diagnose T2m?

Steeneveld, G.J., M.J.J. Wokke, C.D. Groot Zwaaftink, S. Pijlman, B.G. Heusinkveld, A.F.G. Jacobs, A.A.M. Holtslag, 2010: Observations of the radiation divergence in the surface layer and its implication for its parametrization in numerical weather prediction models. J. Geophys Res., 115, D06107, doi:10.1029/2009JD013074.

See also: Edwards et al 2009a,b



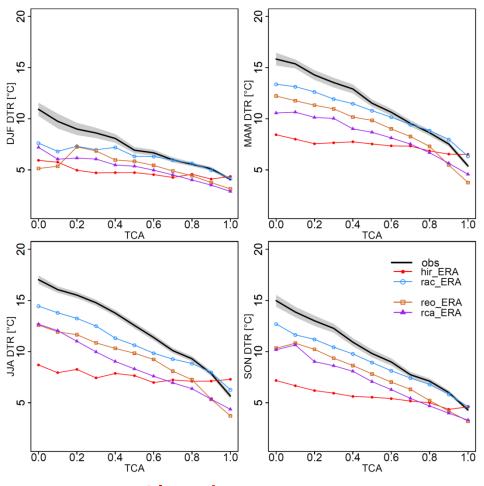
Long wave downward rad. systematically underestimated?



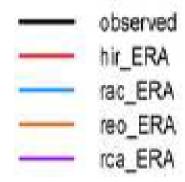
Also reported elsewere



Diurnal Temperature Range

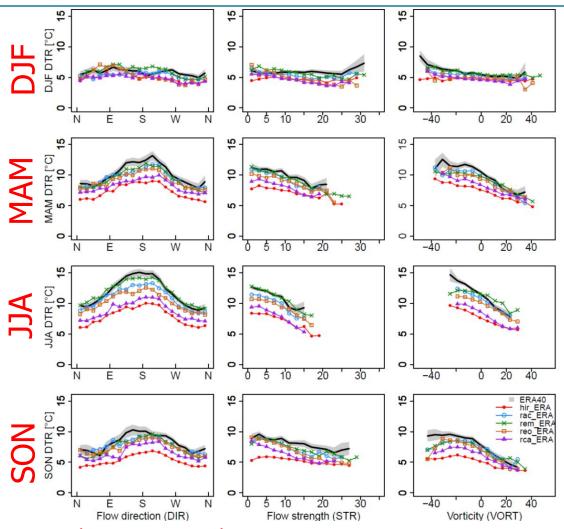


DTR usually, substantially underestimated by models (in this case for central Europe)



Cloud cover

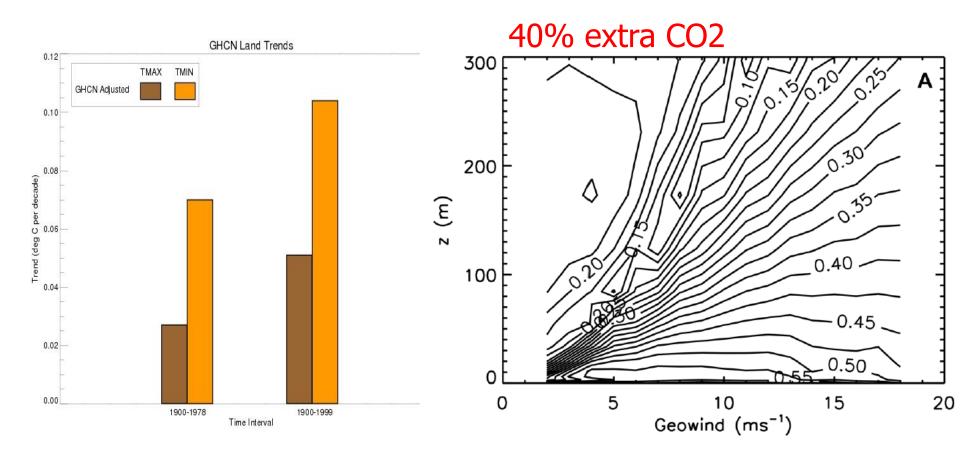
DIRECTION STRENGTH VORT



Underestimation DTR
Strongest for
-anticyclonic flow
-weak winds

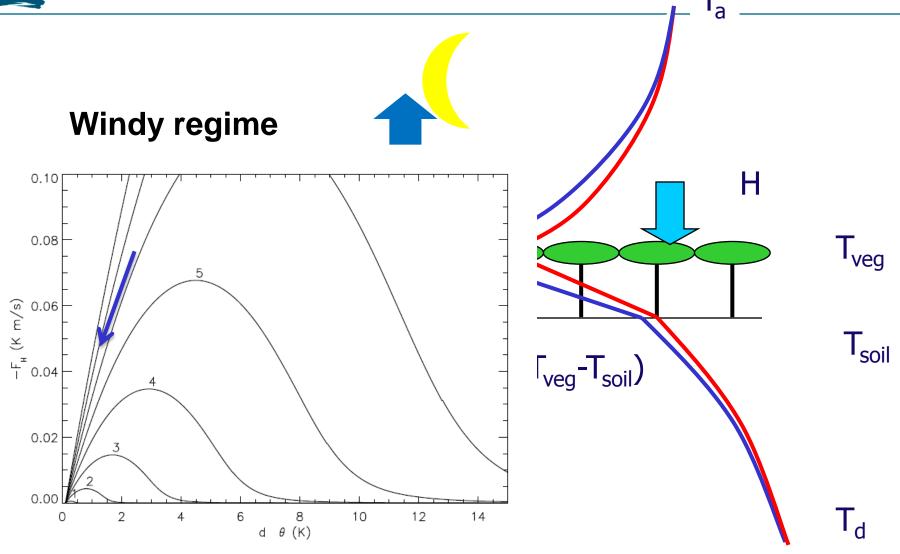
Kyselý, J., and E. Plavcová, 2011: Biases in the diurnal temperature range in Central Europe in an ensemble of regional climate models and their possible causes. *Clim. Dyn.*, in press.

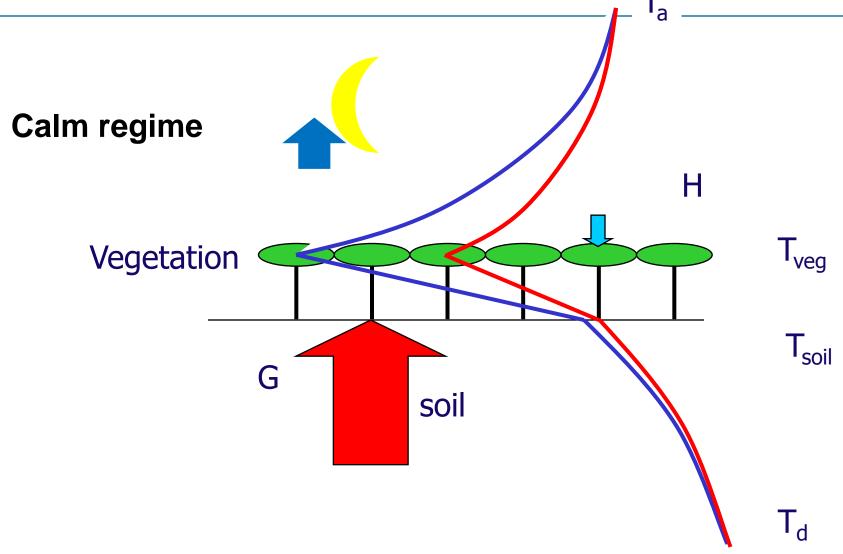
Role of Stable Boundary Layer in climate change



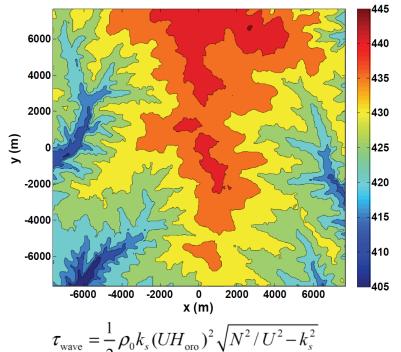
At what level do we find temperature change?

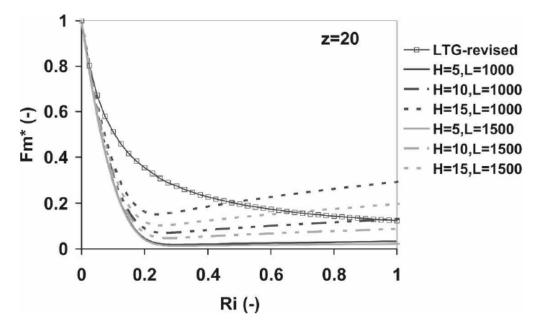




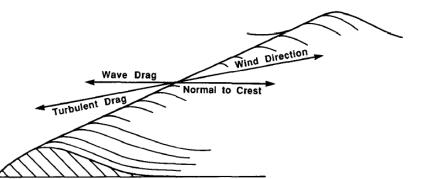


Wave drag



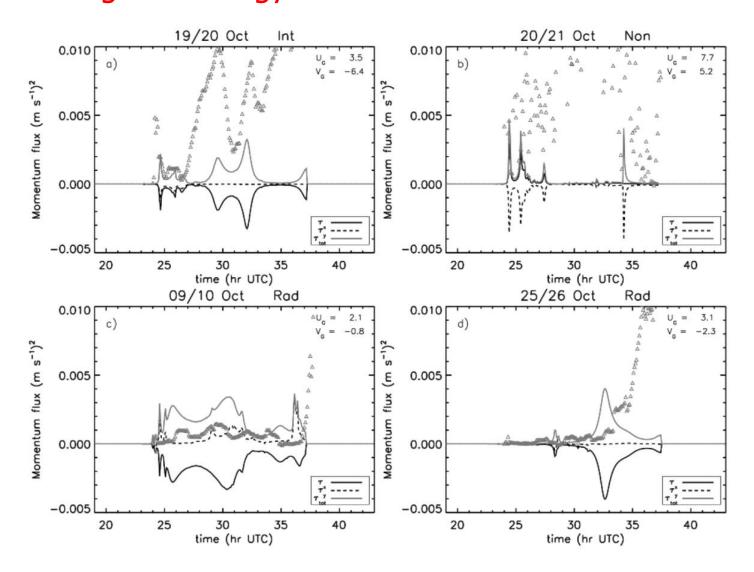


$$\tau_{\text{wave}} = \frac{1}{2} \rho_0 k_s (UH_{\text{oro}})^2 \sqrt{N^2 / U^2 - k_s^2}$$



$$F_m^* = \frac{(\tau_{\text{turb}} + \tau_{\text{wave}})/\rho}{(\kappa z)^2 \left(\frac{\partial U}{\partial z}\right)^2} = \frac{\tau_{\text{turb}}/\rho}{\left(\kappa z \frac{\partial U}{\partial z}\right)^2} + \frac{\tau_{\text{wave}}/\rho}{\left(\kappa z \frac{\partial U}{\partial z}\right)^2}$$

Wave drag climatology in CASES-99



Test role wave drag in 3D WRF

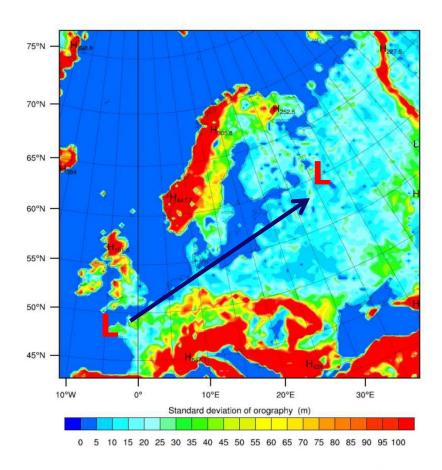


Figure 1: standard deviation of the orography [m] over Europe according to the USGS 2'x2' database.

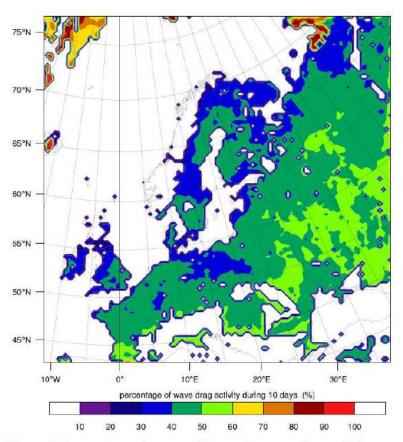


Figure 4: Percentage of orgoraphic gravity wave drag activity per location, during ten days. Areas with a standard deviation larger than 100 [m] are excluded.

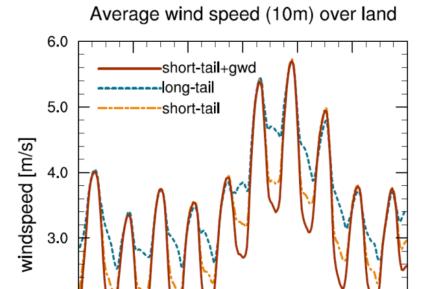


Figure 7: Average horizontal wind speed [ms⁻¹]at 10 m. over land during the different runs.

120

runtime [h]

160

200

240

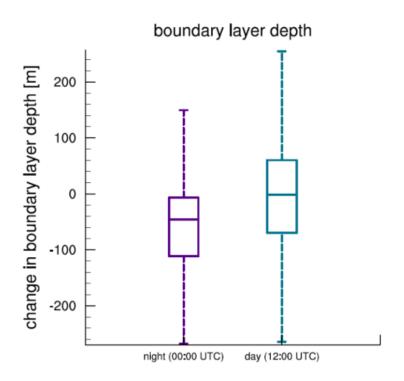


Figure 9: Change in boundary-layer depth [m] between LT run and ST+GWD run, during the night (left) and during the day (right).

Need further confirmation!

80

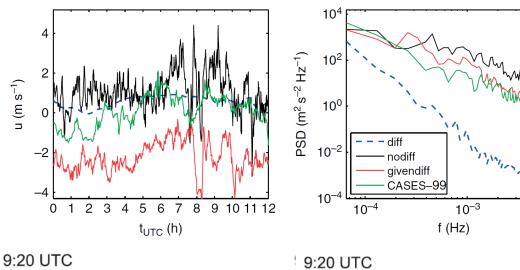


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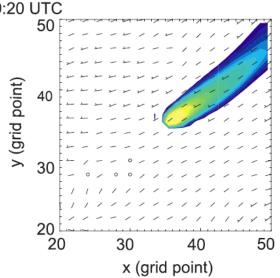
2.0

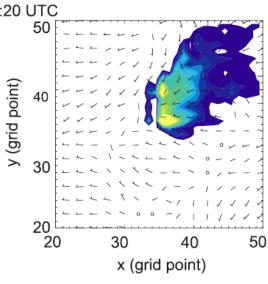
1.0

Horizontal diffusion



Parametrization of horizontal diffusion very important to allow for small scale motions and meandering





Modelled CO concentration





Not covered, though relevant

Observational issues

Drainage flows

Submeso motions

Intermittent turbulence

Land surface heterogeneity



Long tails vs short tails

Land surface coupling

Arctic climate

Mainly for low winds speeds:

Diurnal cycle of temperature Radiation divergence Small scale wave drag Horizontal diffusion

Further reading

- Steeneveld, G.J., A.A.M. Holtslag, C.J. Nappo and L. Mahrt, 2008: Exploring the role of small-scale gravity wave drag on stable boundary layers over land, *J. Appl. Meteor. Clim.*, 47, 2518–2530.
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- Prabha, T.V., Hoogenboom, G., Smirnova, T.G., 2011: Role of land surface parameterizations on modeling cold-pooling events and low-level jets, Atmos. Res., 99, 147-161.

What are the issues

Turbulence
Radiation
Land-surface interaction
Waves
Complex terrain
DTR
Meandering motions

Meandering motions

Horizontal diffusion (Belusic & Guttler paper)
Intermittency?
Predictability issues?

