The Dynamics of the Stratospheric Polar Vortex

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Outline

- Structure of the stratospheric polar vortex
- Phenomenology
 - vortex-vortex interactions during stratospheric sudden warmings
 - impacts (on trace gases & on troposphere)
- Some theoretical considerations

Zonal-mean Wind



Stratospheric Polar Vortices 50 mb (about 20 km) NH & SH





Cross-Section of Methane October







Consequence 1: Forecasting

 If the stratosphere has an impact on the tropospheric state 10-60 days in the future then there is potential to use this information for sub-seasonal forecasting



Baldwin et al. (2003)

Impact of Stratosphere 2



Charlton et al.

 Changing the stratospheric initial conditions results in a tropospheric impact 15-20 days into the run.

Effect of SH Ozone Depletion



 Recent trends in the Antarctic can be well simulated by forcing a model with ozone trends confined to the stratosphere

Gillett and Thompson (2003)

Two Dynamical Paradigms for Dynamical Variability of the Stratospheric Polar Vortex

- Wave, mean-flow interaction.
- Vortex-vortex interaction.

Some Textbook Quotes

- "Numerous observational studies confirm that enhanced propagation of planetary waves from the troposphere, primarily zonal wavenumber 1 and 2, is essential for the development of warmings."
- "Most of the dramatic mean-flow deceleration that occurs during a sudden warming is caused by amplification of quasi-stationary planetary waves in the troposphere followed by propagation into the stratosphere."
- "It is generally accepted that sudden warmings are an example of transient mean-flow forcing due to planetary wave driving."

An Introduction to Dynamic Meteorology, 4th edition (2004), James R Holton, Elsevier, p. 425.



The basic notion: the troposphere acts as a wave maker, and disturbances propagate quasi-linearly into the stratosphere where they "break".

EP Fluxes and Divergence



The Seductive Transformed Eulerian Mean Momentum Equation



Idealised 3D Vortex-Vortex Interactions in the Winter Stratosphere



FIG. 9. Evolution for $-\kappa_2/\kappa_1 = 0.8$, $z_2 - z_1 = 0$, at t = 0, 4, 8, 12, 16, and 20 days (from upper left to lower right; top view).



NH Dec/Jan 84/85: Geo Ht 10 hPa



NH Dec/Jan 84/85: PV 840K



Zonal-mean wind & polar cap temperature, 10 hPa, NH winter 1984/85



NH Dec/Jan 84/85: PV 450K



NH Jan 87: PV 840 K



NH Dec/Jan 84/85: PV isosurface

Lait's PV isosurface at 0000UT on 13-Dec-1984



The Polar Vortex: NH 2005/6

01 November 2005

Theta (K) Altitude (kn 47 km 2000 K 1800 K 45 km 1600 K 42 km 1400 K 40 km 1200 K 37 km 1000 K 33 km 900 K 31 km 800 K 29 km 700 K 600 K 26 km 23 km 550 K 22 km 500 K 20 km 450 K 18 km 400 K 15 km 380 K 14 km 360 K 340 K 13 km 10 km 320 K 300 K 7 km 3 km -180 180 190 200 210 220 230 240 250 260 270 Temperature

Courtesy of Lynn Harvey

SH Sep 2002: 850K PV



SH Oct 2002: 850K PV



SH 21 Sep 2002: PV 350K & p*





f mean sea level pressure for 00Z 21st September 20

Merger of anticyclones, SH, 10 & 13 Oct 1992, PV 1100K



Schematic of Top-Down Breakdown of SH Polar Vortex



Lahoz et al., QJRMS, 1996

Variability of the Polar Vortex

- Evolution of coherent vortical structures, involving strongly local, nonlinear dynamics (e.g. during vortex merger) and the interaction of anticyclones with the polar vortex.
- Deep, nonlinear evolution between axi-asysmmetric states in the upper troposphere and stratosphere.
- Possibility for instability of highly distorted polar vortex to finite-amplitude perturbations (e.g. cyclogenesis in the troposphere).
- Tropospheric wave maker & vertical propagation?
- Troposphere-stratosphere as a coupled system?

SPARC & IPY

- Characterise the structure and evolution of the (meteorological and chemical) of the stratospheric polar vortex (NH & SH).
- Archive of data or metadata at the SPARC Data Centre

End