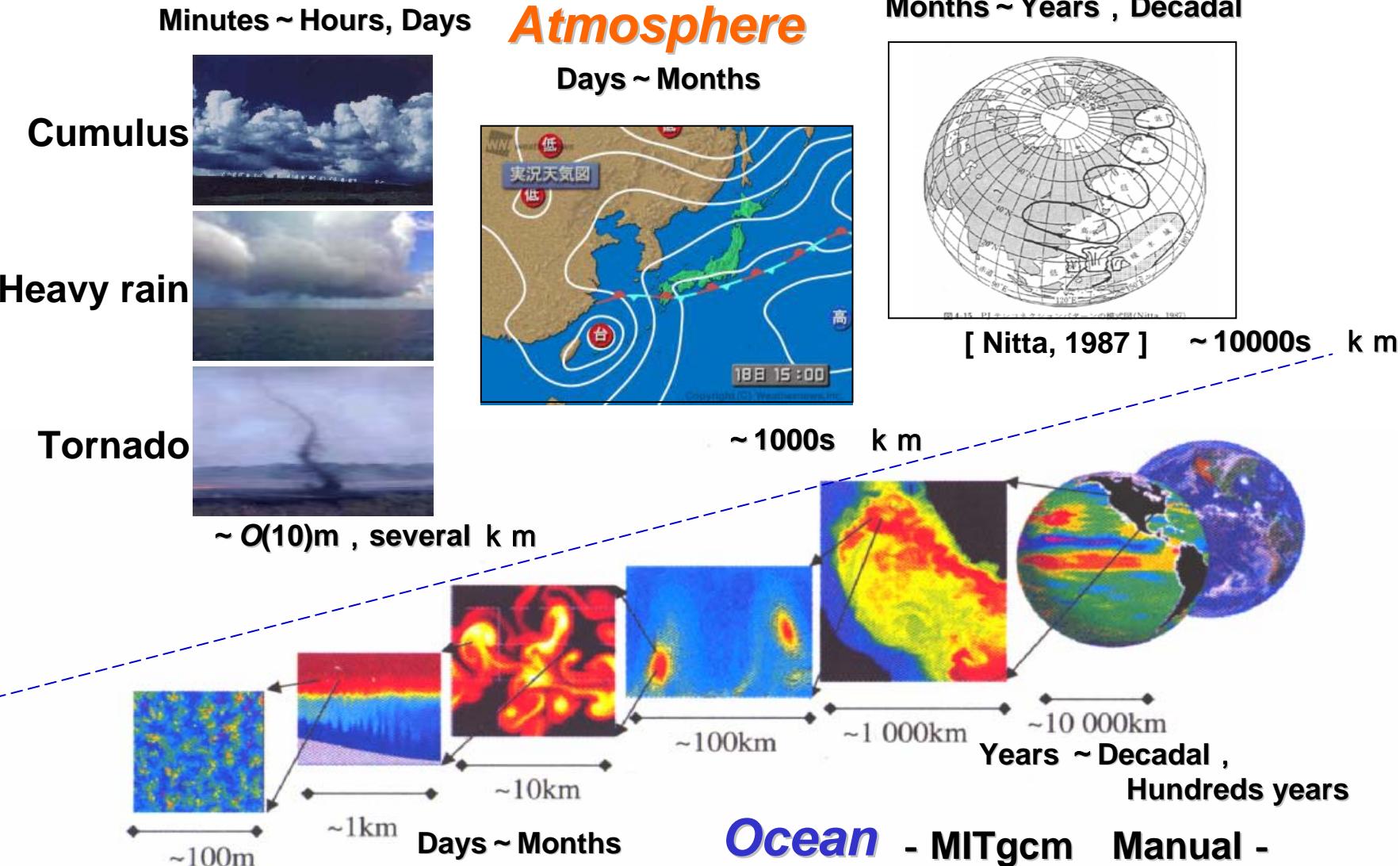


# **Multi-scale coupled atmosphere-ocean GCM and simulations**

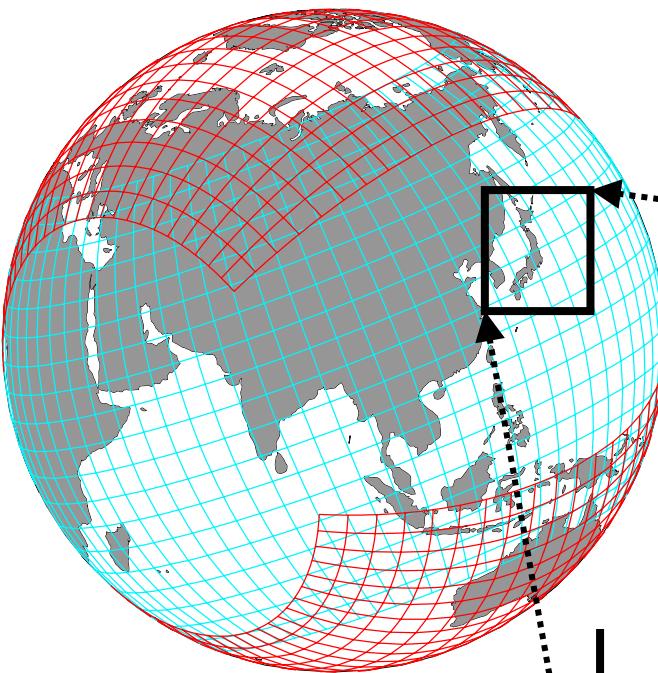
**Keiko Takahashi, Xindong Peng, Ryo Onishi , Mitsuru Ohdaira,  
Koji Goto, Hiromitsu Fuchigami, Takeshi Sugimura**

**Earth Simulator Center  
JAMSTEC**

# Time/Space scale

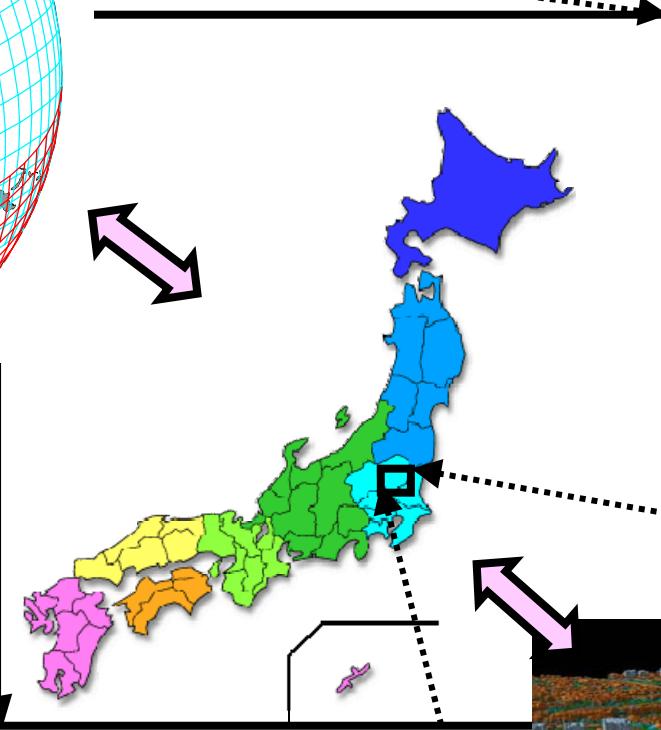


# Scalability



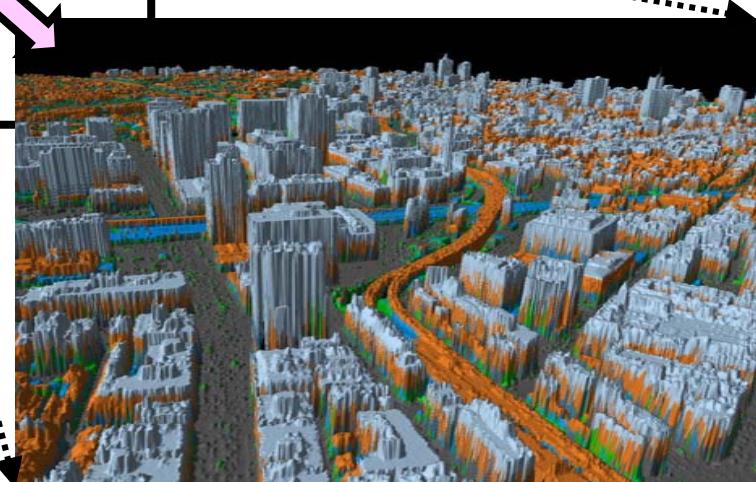
## Seasonal ~ Annual Prediction

2 ~ 10 km for horizontal,  
100 vertical layers



## Urban Weather /Climate Prediction

1m ~ 100m for horizontal,  
200 vertical layers



( Data: Geographical Survey Institute)

Days ~ Weeks Prediction  
Local heavy Rain Prediction,  
etc.

100m ~ 2km for horizontal  
100 vertical layers

# Dynamical Core Framework

## - Atmosphere -

- **Atmosphere:** Fully compressible, non-hydrostatic equation set

**Continuity equation**  $\frac{\partial \rho'}{\partial t} + \frac{1}{G^{\frac{1}{2}} a \cos \varphi} \frac{\partial (G^{\frac{1}{2}} \rho u)}{\partial \lambda} + \frac{1}{G^{\frac{1}{2}} a} \frac{\partial (G^{\frac{1}{2}} \cos \varphi \rho v)}{\partial \varpi} - \frac{1}{G^{\frac{1}{2}} a} \frac{\partial (\rho w^*)}{\partial \tau^*} = 0$

**Momentum equation**  $\frac{\partial \rho u}{\partial t} + \frac{1}{G^{\frac{1}{2}} a \cos \varphi} \frac{\partial (G^{\frac{1}{2}} p')}{\partial \lambda} = -\nabla \bullet (\rho u \vec{v}) + 2f_r \rho v - 2f_\varphi \rho w + \frac{\rho v u \tan \varphi}{a} - \frac{\rho w u}{a} + F_\lambda$

$\frac{\partial \rho v}{\partial t} + \frac{1}{G^{\frac{1}{2}} a} \frac{\partial (G^{\frac{1}{2}} p')}{\partial \varphi} = -\nabla \bullet (\rho v \vec{v}) + 2f_\lambda \rho w - 2f_r \rho u - \frac{\rho u u \tan \varphi}{a} - \frac{\rho w v}{a} + F_\varphi$

$\frac{\partial \rho w}{\partial t} + \frac{1}{G^{\frac{1}{2}}} \frac{\partial p'}{\partial z^*} + \rho' \mathbf{g} = -\nabla \bullet (\rho w \vec{v}) + 2f_\varphi \rho u - 2f_\lambda \rho v + \frac{\rho u u}{a} + \frac{\rho v v}{a} + F_r$

**Pressure equation**  $\frac{\partial p'}{\partial t} + \nabla \bullet (p \vec{v}) + (\gamma - 1) p \nabla \bullet \vec{v} = (\gamma - 1) \kappa \nabla^2 T + (\gamma - 1) \Phi$

**State equation**  $p = \rho R T$

$$G^{\frac{1}{2}} = \frac{\partial z}{\partial z^*} = 1 - \frac{z^*}{H} \quad \text{is a metric term.}$$

# Framework (2)

- Ocean: in-compressive and hydrostatic equations with the Boussinesq approximation

$$\frac{\partial c}{\partial t} = -\mathbf{v} \cdot \nabla c + F_c$$

$$\frac{\partial T}{\partial t} = -\mathbf{v} \cdot \nabla T + F_T$$

$$0 = \nabla \cdot \mathbf{v} = \left( \frac{1}{r \cos \varphi} \frac{\partial u}{\partial \lambda} + \frac{1}{r \cos \varphi} \frac{\partial (\cos \varphi v)}{\partial \varphi} + \frac{1}{r^2} \frac{\partial (r^2 w)}{\partial r} \right)$$

$$\frac{\partial u}{\partial t} = -\mathbf{v} \cdot \nabla u + 2f_r v - 2f_\varphi w + \frac{vu \tan \varphi}{r} - \frac{wu}{r} - \frac{1}{\rho_0 r \cos \varphi} \frac{\partial P'}{\partial \lambda} + F_\lambda$$

$$\frac{\partial v}{\partial t} = -\mathbf{v} \cdot \nabla v + 2f_\lambda w - 2f_r u - \frac{uu \tan \varphi}{r} - \frac{wv}{r} - \frac{1}{\rho_0 r} \frac{\partial P'}{\partial \varphi} = +F_\varphi$$

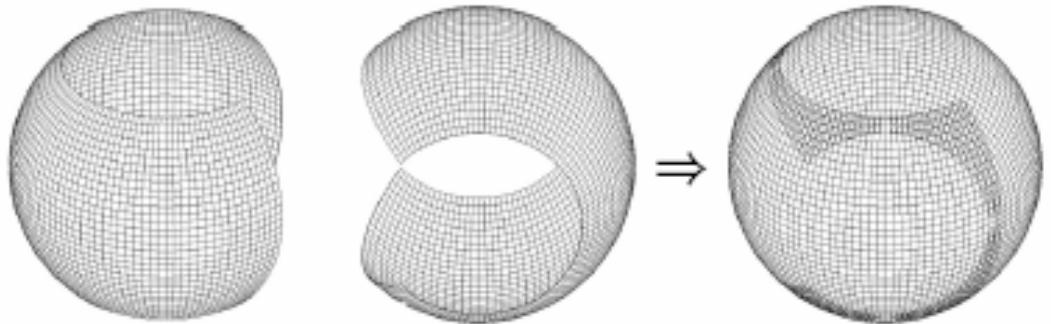
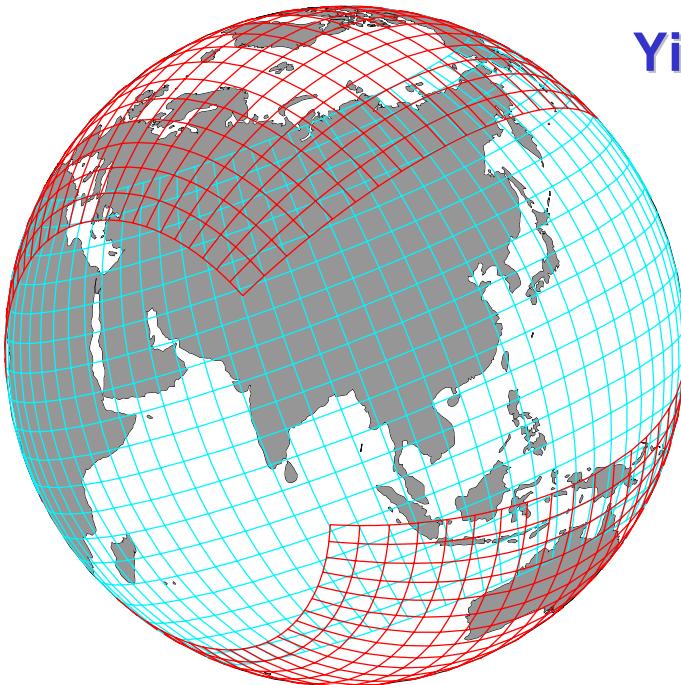
$$\frac{\partial w}{\partial t} = -\mathbf{v} \cdot \nabla w + 2f_\varphi u - 2f_\lambda v + \frac{uu}{r} + \frac{vv}{r} - \frac{1}{\rho_0} \frac{\partial P'}{\partial r} - \frac{\rho'}{\rho_0} \mathbf{g} + F_r$$

$$\frac{d}{dr} P_0 = -\rho_0 g(r)$$

$$\rho = \rho(T, c, P_0) \quad ( \text{: UNESCO scheme})$$

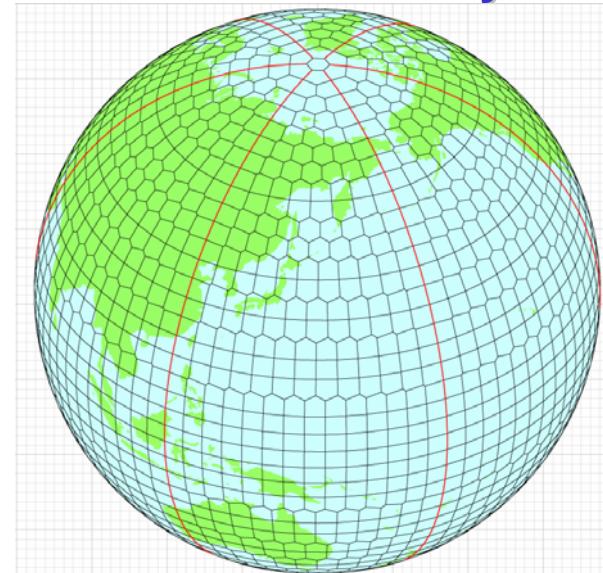
# Grid System

## Yin-Yang Grid System



## New Reduced Grid System

- Orthogonal coordinates.  
( same as the lat-lon geometry )
- No polar singularity.
- Relax of CFL condition.
- The same grid structure of N and E component.
- Easy to nest.
- High parallelization.
- But need to take care of conservation law.

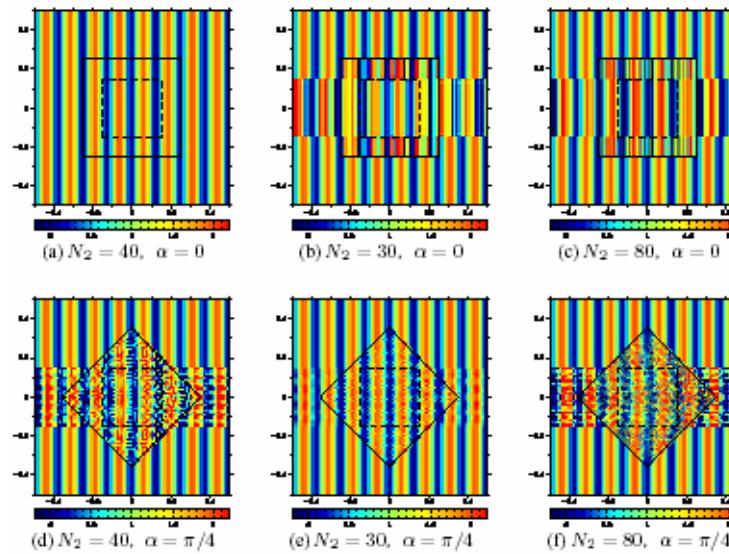
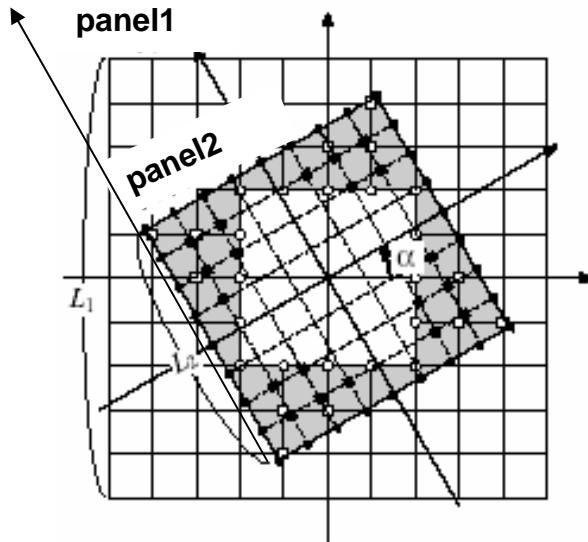


# Issues on overset grid systems

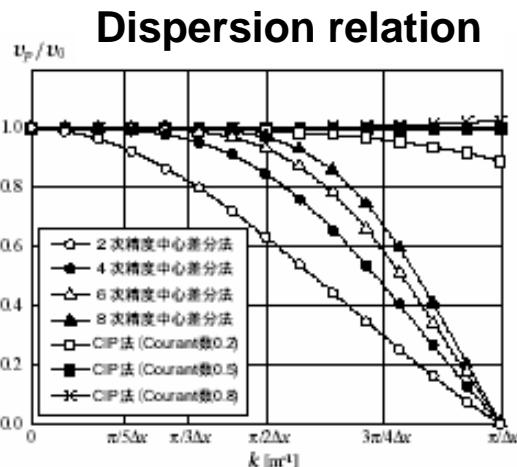
1. How are wave propagation characteristics  
on Yin-Yang grid system?
2. Is it necessary to high ordered computational schemes?
3. How long is integration possible on Yin-Yang grid system?
4. Which conservation scheme is suitable  
on Yin-Yang grid system?

# Wave propagation characteristics on overset grid system

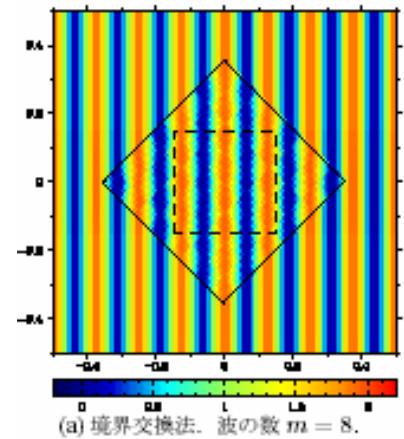
(Takeshi Sugimura)



High order computational schemes  
and interpolation are required.



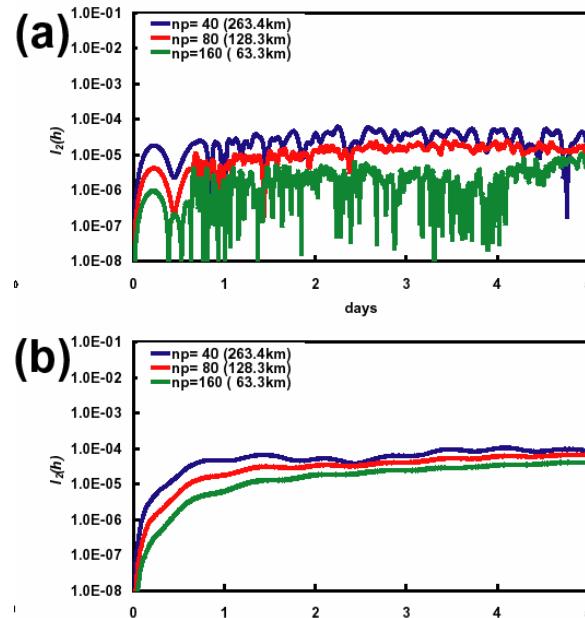
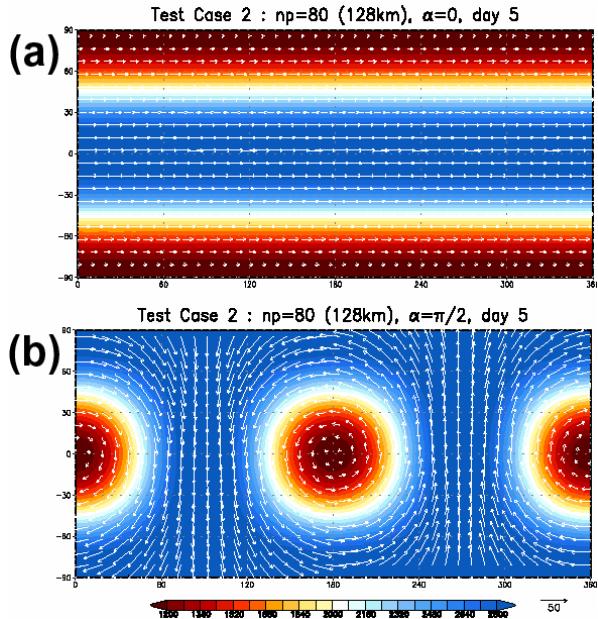
Dispersion relation is important  
to avoid errors on interface of  
overset grid system.



# How long is integration possible ?

(Mitsuru Ohdaira, ESC)  
(Kenji Komine, ESC)

## *Test Case 2 : Global Steady State Nonlinear Zonal Geostropic Flow*

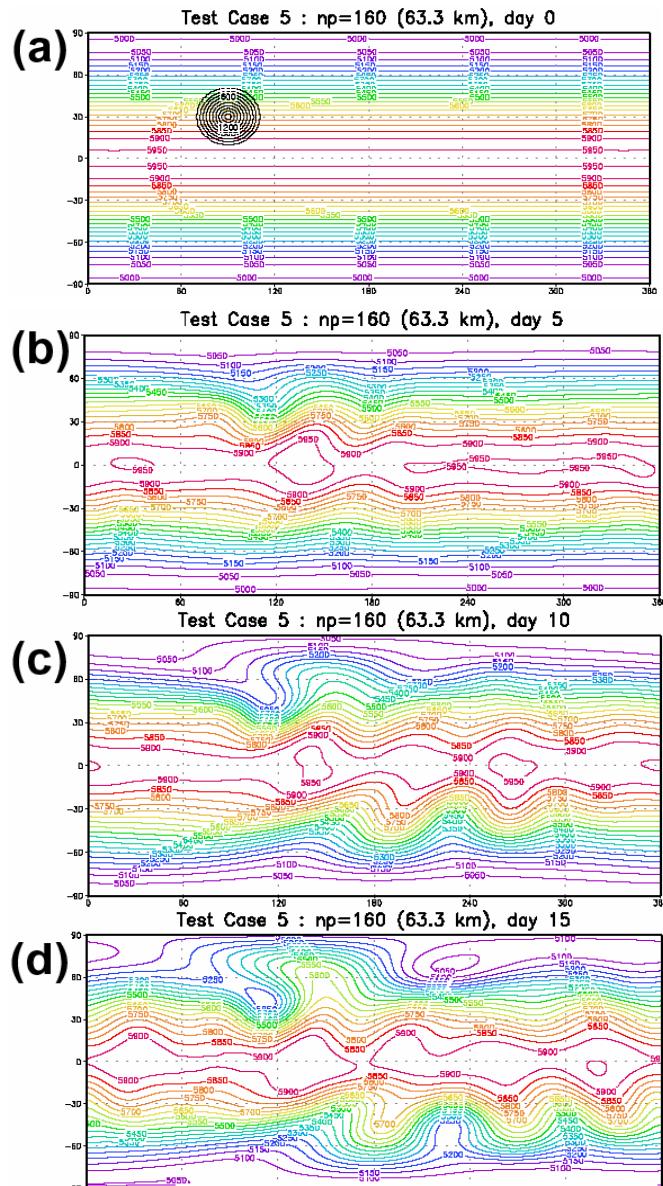


The solid body rotation field is maintained. The 2nd-order accuracy is maintained.

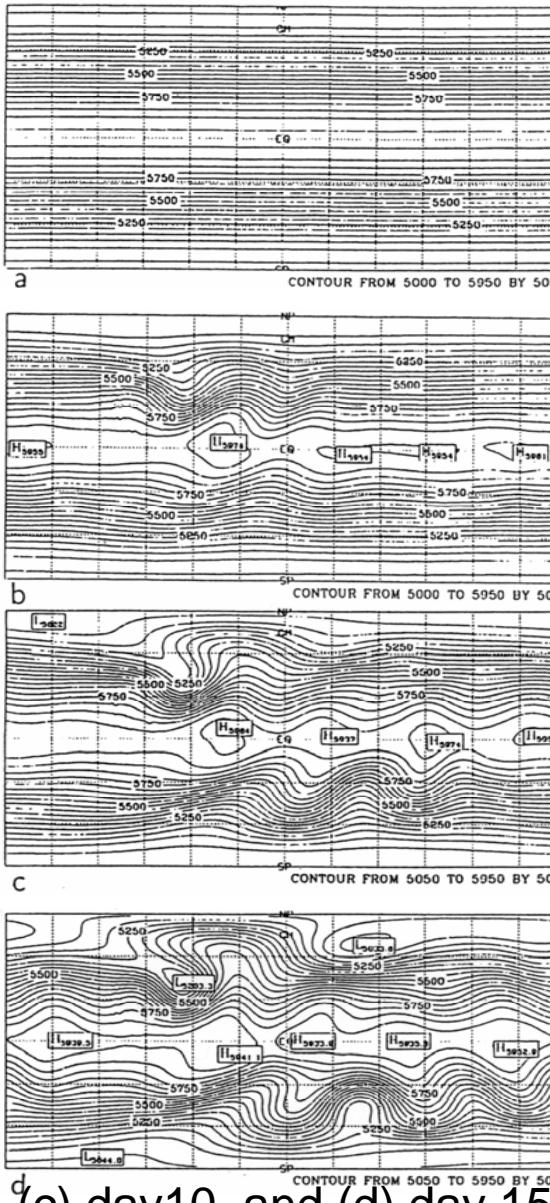
- At least 3 years integration with the 2nd-order accuracy is maintained.

# *Test Case 5 : Zonal Flow over an Isolated Mountain*

## Simulation Results



## Results with spectral scheme

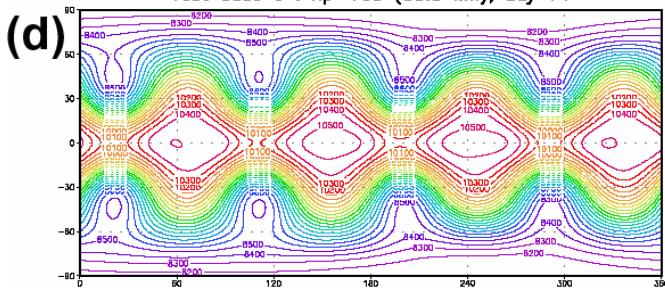
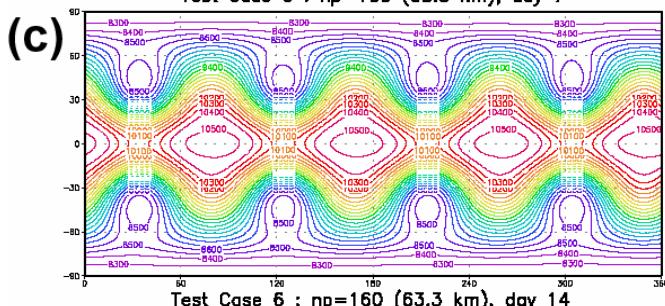
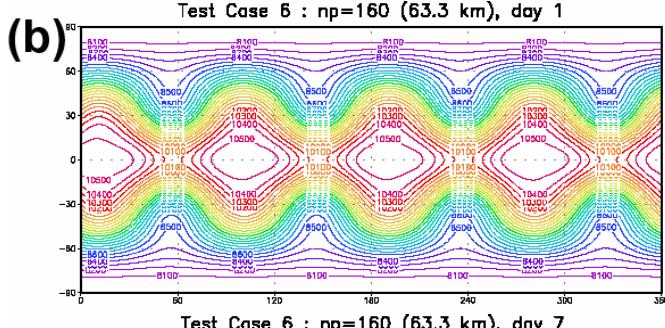
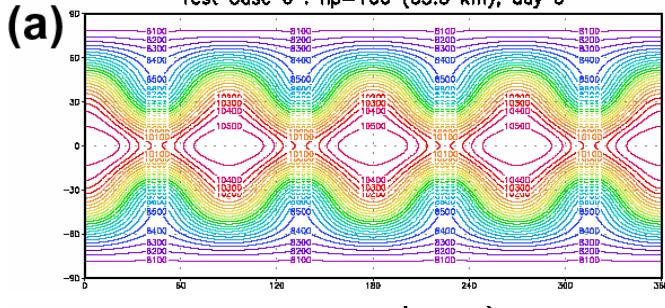


R. Jakob, J. J. Hack and  
D. L. Williamson,  
Solutions to the Shallow  
Water Test Set Using  
the Spectral Transform  
Method.,  
NCAR/TN-388+STR, 1993

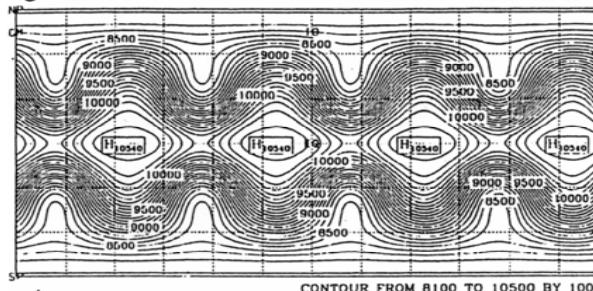
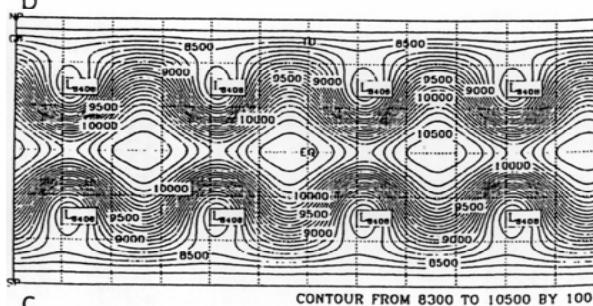
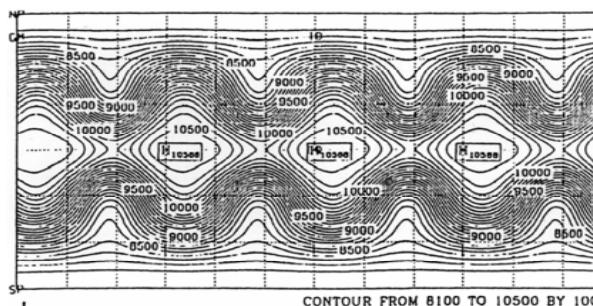
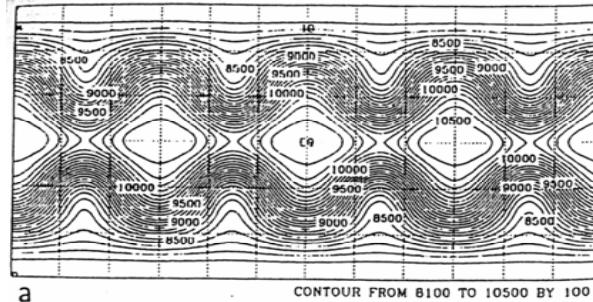
# Test Case 6 : Rossby-Haurwitz Wave

## Simulation Results

Test Case 6 : np=160 (63.3 km), day 0

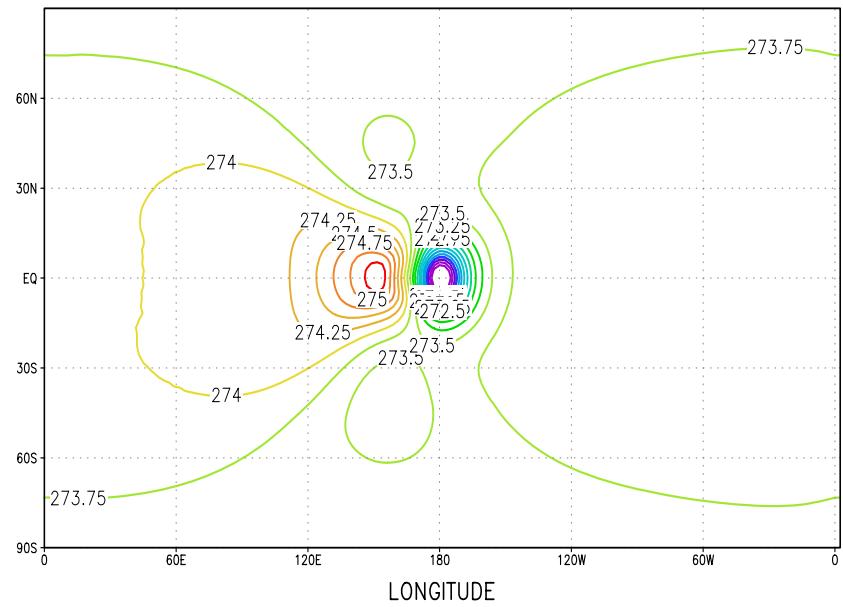
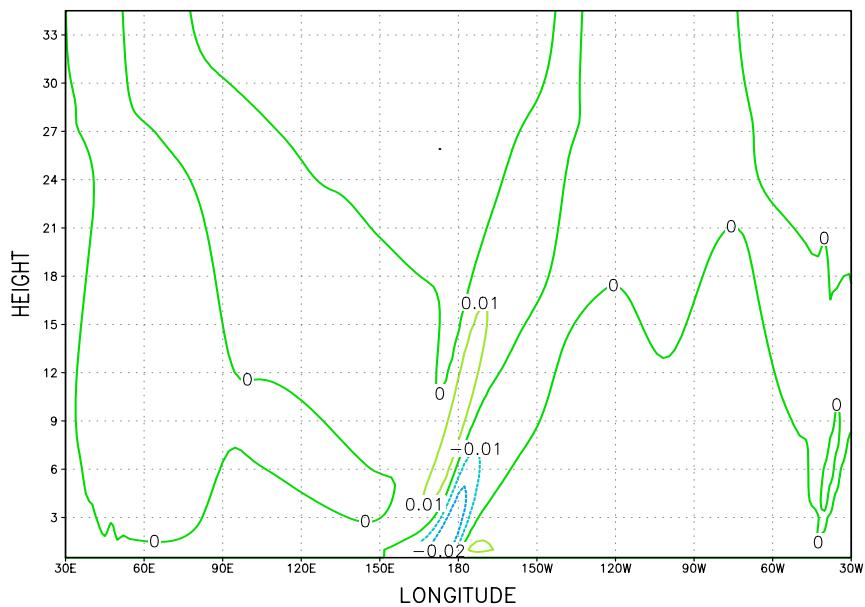
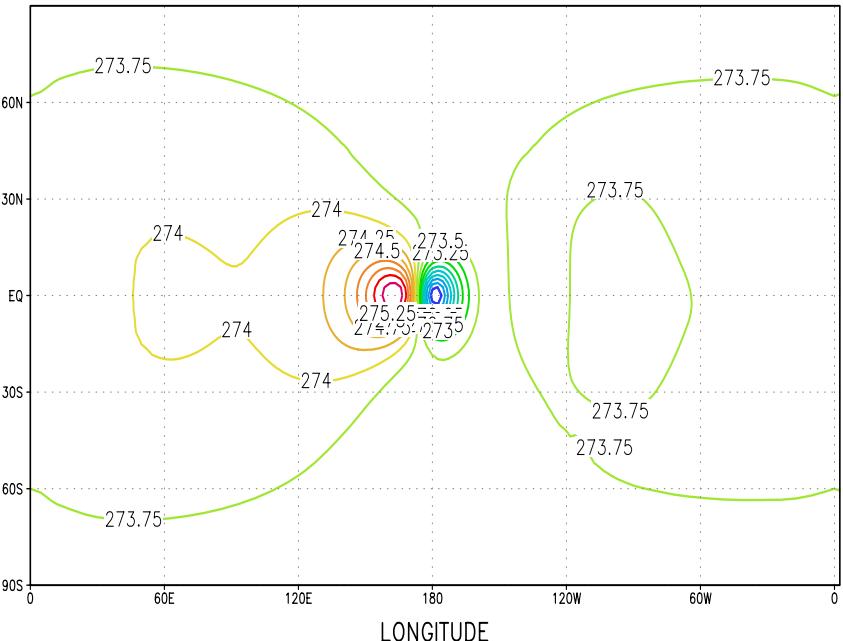
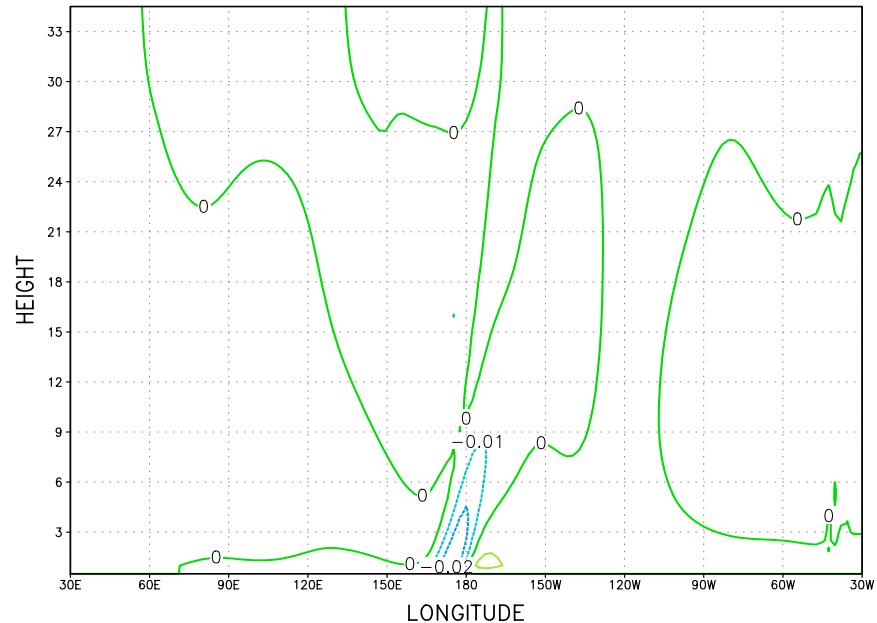


## Results with spectral scheme



Rossby-Haurwitz wave shape has been propagated from the west to the east without change from initial field after 14 days integration.

R. Jakob, J. J. Hack and  
D. L. Williamson,  
Solutions to the Shallow  
Water Test Set Using  
the Spectral Transform  
Method.,  
NCAR/TN-388+STR, 1993

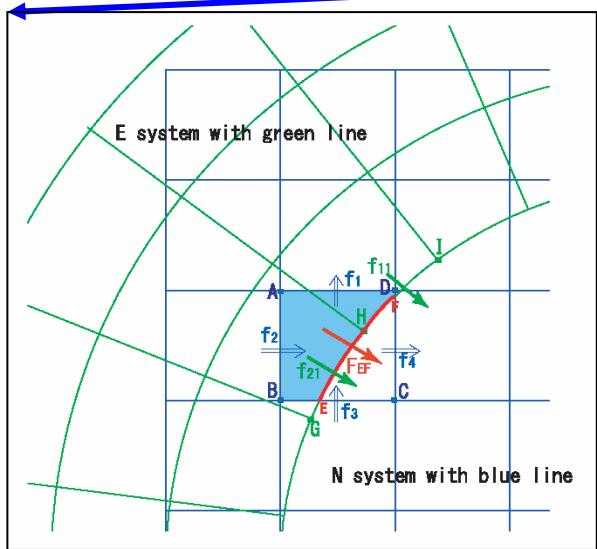


The  $\lambda$ -z cross section of vertical wind speed  $w$  (m s<sup>-1</sup>) along the equator after 12, 24hours.

Temperature (K) at 0km (bottom) levels  
after 12, 24hours.

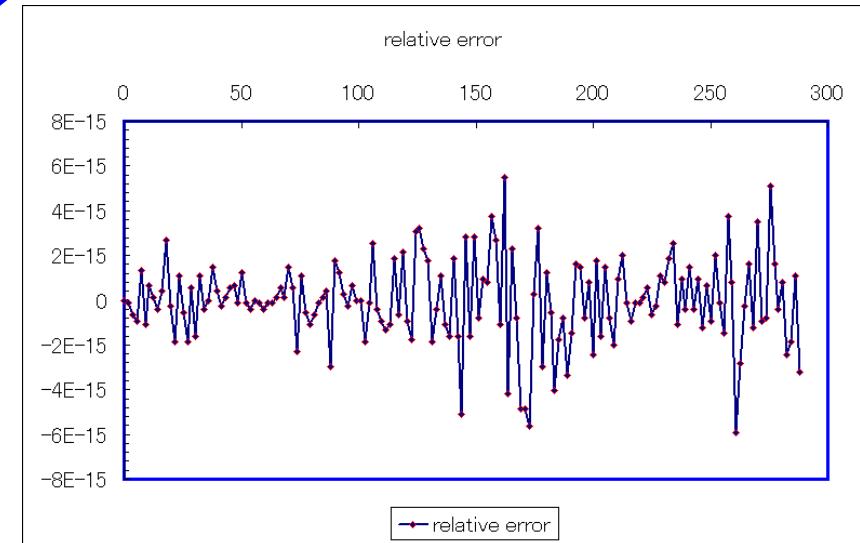
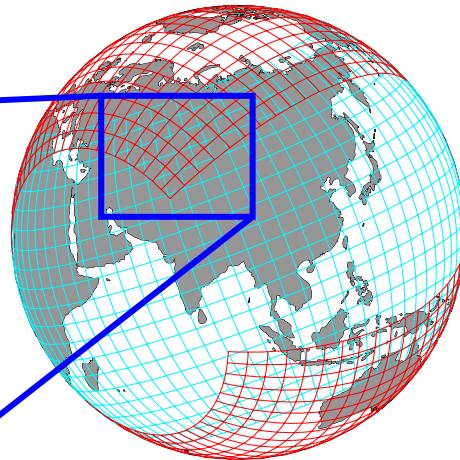
# Mass conserving numerical scheme

(Xindong Peng)



For flux  $F_{EF}$  on a circular arc EF shown as red circle is computed by the budget of fluxes  $f_N$  by on grid ABCD of N system and flux  $f_E$  estimated on a circular arc GHI of E system.

Computation all of fluxes  
on computational grids  
↓  
Correction for conserving



This conservative scheme, we have evaluated that time evolution of relative error of the mass has changed within the limit of rounding error.

# Implementation

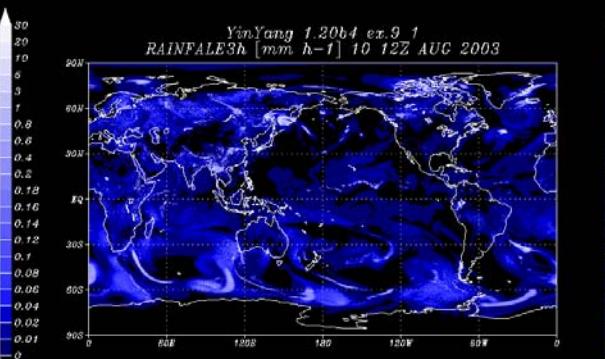
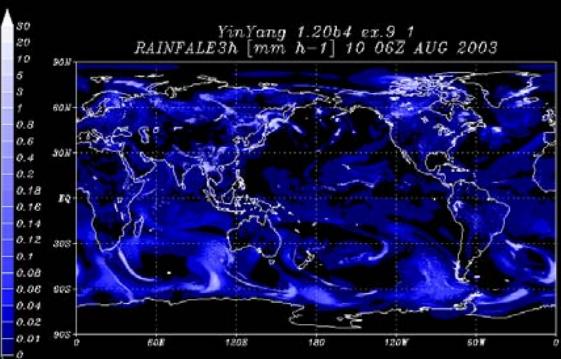
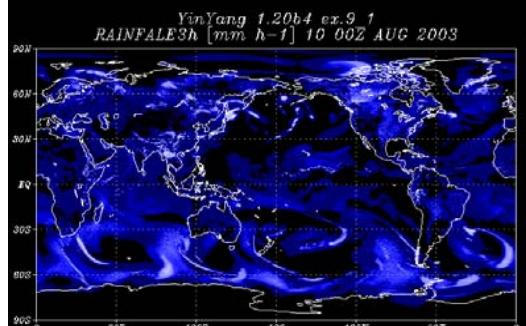
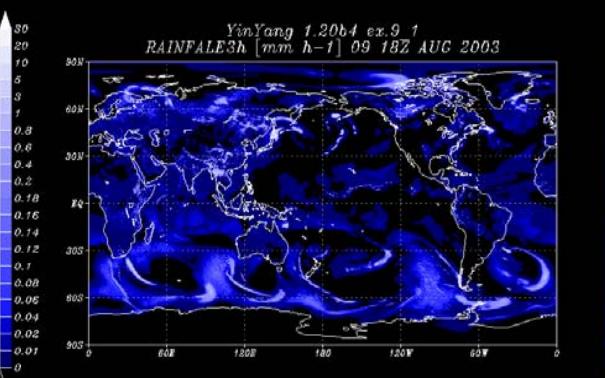
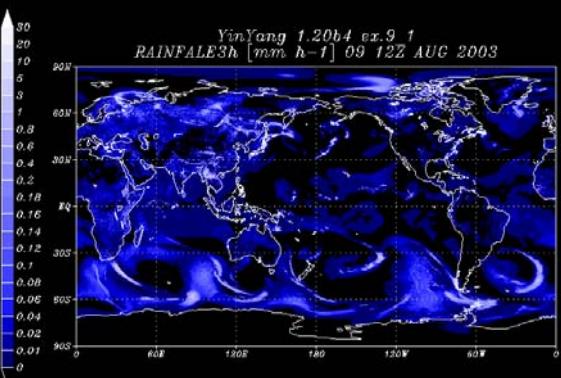
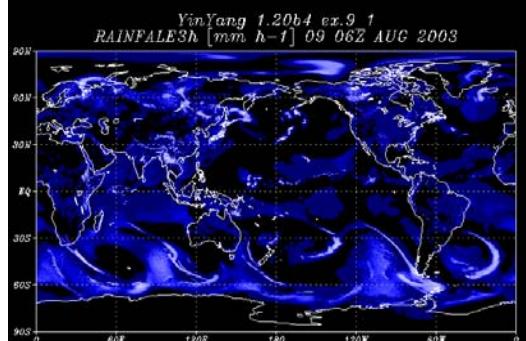
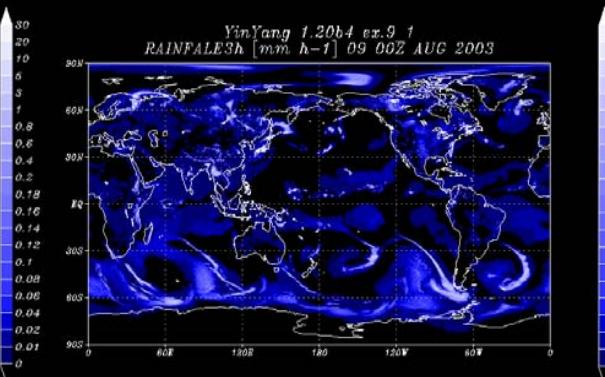
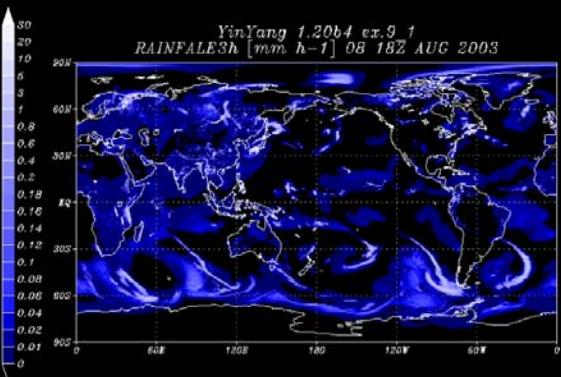
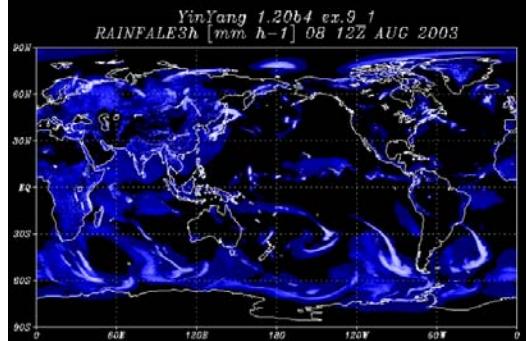
## Non-hydrostatic AGCM and OGCM

		Non-hydrostatic AGCM	Non-hydrostatic OGCM
Equations System		Fully compressive N-S equations	non-hydrostatic incompressive N-S equations
Grid System		Yin-Yang grid system	Yin-Yang grid system
Discrimination	Space	Arakawa-C grid(horizontal), z*(vertical)	Arakawa-C grid(horizontal), z (vertical)
	Time	4th order Runge-Kutta	4th order Runge-Kutta
Advection terms		5th order flux form, CIP-CSLR	5th order flux form
not Advection terms		4th order flux form	4th order flux form
Sound wave		HEVI, HIVI	Implicit methods(2-dimensional, 3-dimensional )
Gravity wave		-	-
Microphysics		Qc, Qci, Qr, Qs, Qg	-
Cumulus Param.		Kain-Fritsch scheme	-
Turbulence		Smagorinsky scheme (static), dynamic Smagorinsky[LES]	Smagorinsky scheme (static), dynamic Smagorinsky[LES]
		Nesting systems(1 way,2 way)	Nnesting systems(1 way,2 way) Tide, Multi-grid Methods(Poison eq.)
Parallelization		2-dim. decompositon, inter nodes: MPI, intra nodes: micro-task	2-dim. decompositon, inter nodes: MPI, intra nodes: micro-task

- Kain-Fritsch cumulus parameterization ← more than 10km horizontal resolution
- Reisner et al. (1998) cloud microphysics ← less than 10km horizontal resolution
- Smagorinsky (1965) turbulence closure scheme ← for both atmosphere and Ocean components
- Monin-Obukhov similarity theory (Zhang and Anthes, 1982)
- Simple short and long-wave radiation scheme
- 0-D bucket model for land surface process
- Multi-grid methods for Poisson equation (Ocean)

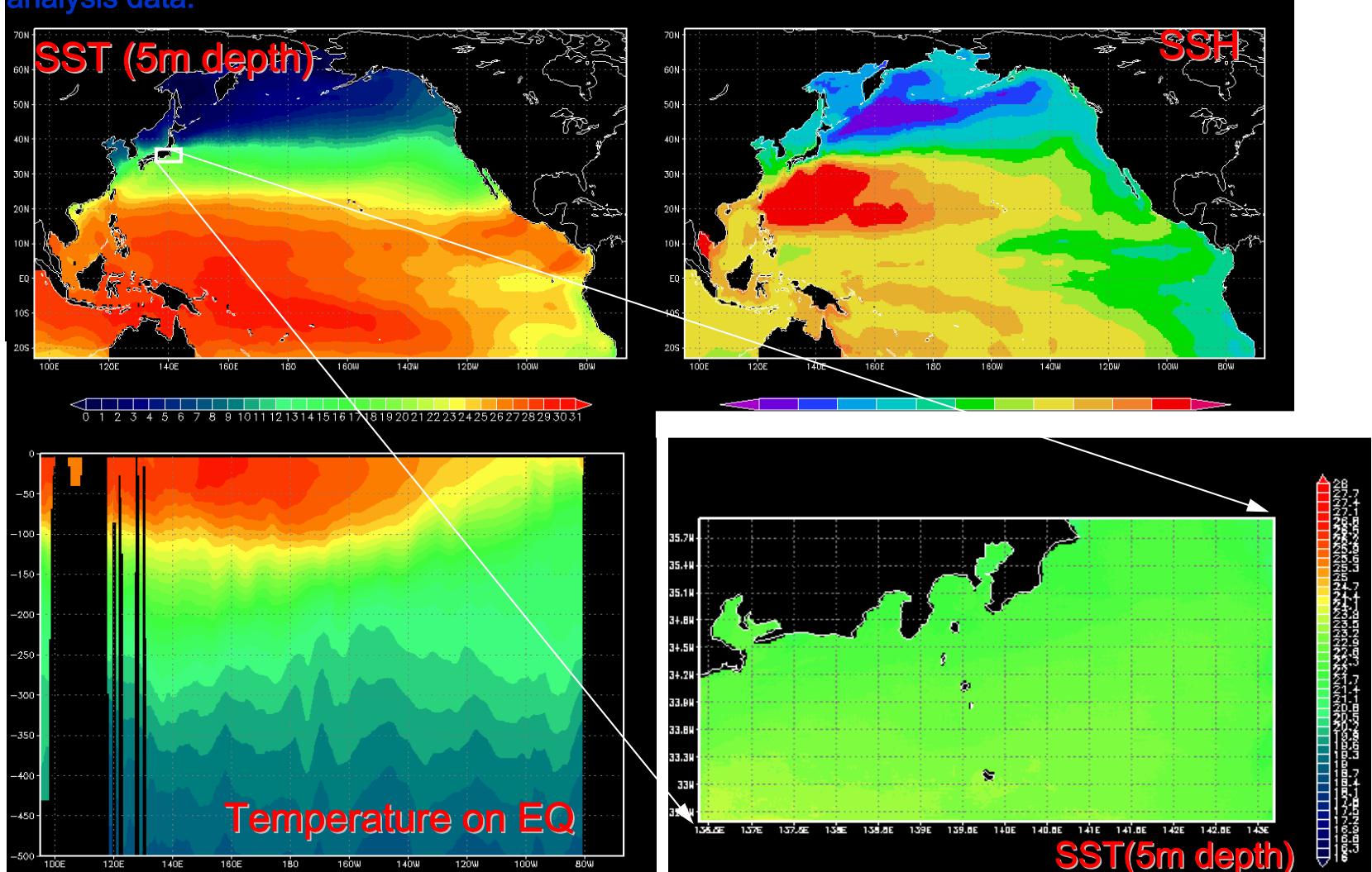
# Atmospheric Global Circulation (5.5km horizontal and 32 vertical)

## 48hours integration, micro cloud physics only



# North Pacific Ocean and Japan Region

- North Pacific Ocean : 50km(horizontal) , 40layers (vertical).  
January monthly data as initial data , surface boundary condition : NCEP re-analysis data,  
25 years integration without motion as initial condition.
- Japan Region : 2.5km (horizontal) , 40layers (vertical).  
Initial data: 1<sup>st</sup> Jan of the above results from North Pacific Ocean.  
boundary condition : above results from North Pacific Ocean、surface boundary : NCEP re-analysis data.

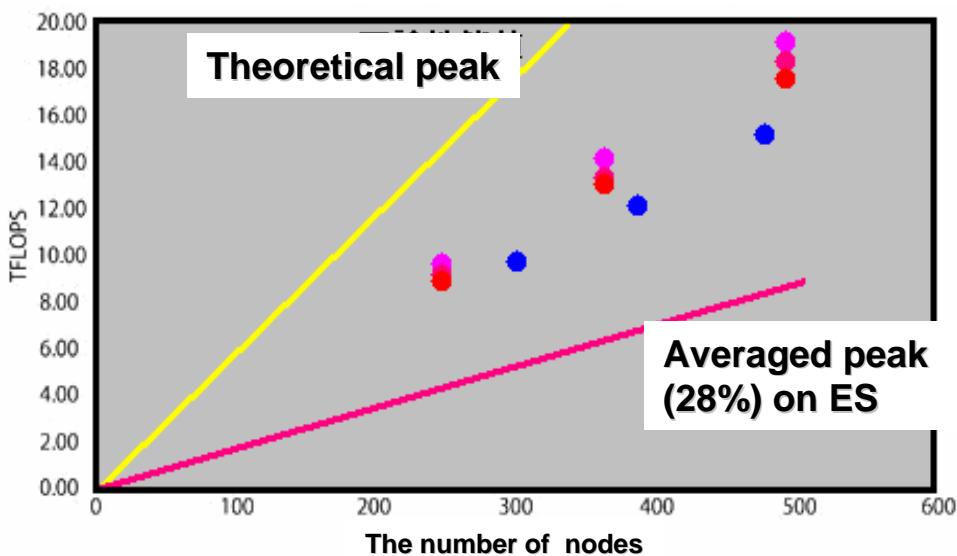


# Computational performance on the Earth Simulator

CASE	TPN	TAP	grid pts	Mflops/AP	Vector Length	V.OP ratio	Tflops	Peak ratio	Parallel efficiency	Speed up
C	512	4096	3,866,296,320	4166.7	229	99.3%	17.07	52.1%	90.0%	461.0
	384	3072		4273.8	229	99.3%	13.13	53.4%	92.3%	354.6
	256	2048		4401.9	229	99.3%	9.02	55.0%	94.8%	242.6
A	512	4096	2,882,764,800	4575.2	228	99.5%	18.74	57.2%	93.6%	479.1
	384	3072		4606.1	228	99.5%	14.15	57.6%	95.1%	365.2
	256	2048		4692.4	228	99.5%	9.61	58.7%	96.7%	247.5
RA	512	4096	2,882,764,800	4340.8	229	99.4%	17.78	54.3%	90.7%	464.4
	384	3072		4401.0	229	99.4%	13.52	55.0%	92.9%	356.6
	256	2048		4560.5	229	99.4%	9.34	57.0%	95.1%	243.5
O	498	3984	4,954,521,600	3629.3	240	99.3%	14.46	45.4%	80.6%	401.3
	398	3184		3568.5	240	99.3%	11.36	44.6%	83.8%	333.7
	303	2424		3986.8	240	99.3%	9.66	49.8%	87.2%	264.2
	207	1656		4234.3	240	99.3%	7.01	52.9%	90.9%	188.2

C: Coupled AOGCM, A: AGCM,

RA: regional Atmos. O: OGCM

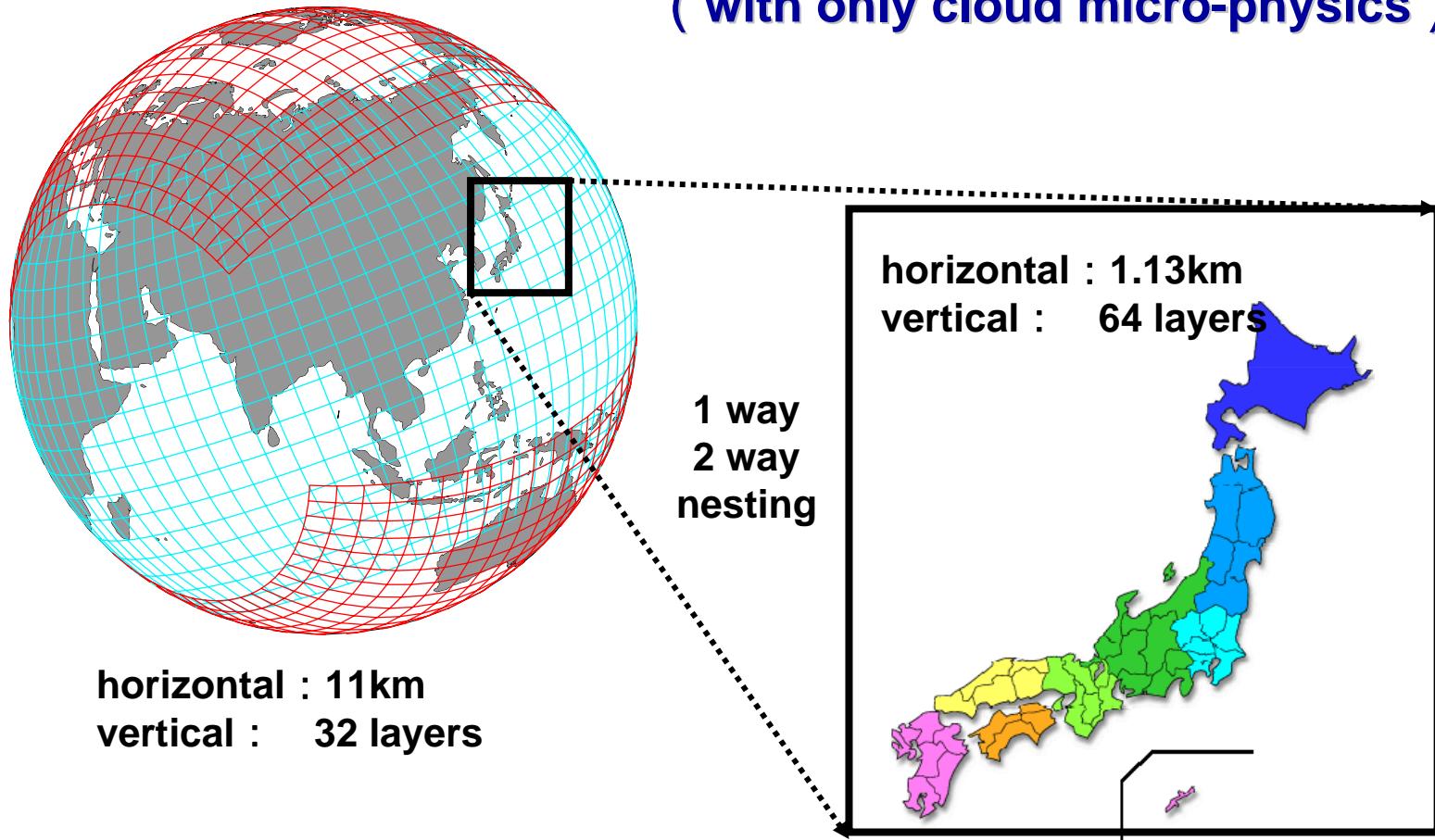


Global atmosphere: 5.5km for horizontal  
32 vertical layers  
⇒ 3.0 hrs on 512 nodes  
for 72 hrs integration

Regional Coupled A-O model:  
Global atmosphere: 11km for horizontal  
32 vertical layers  
+Coupled: 2.78 km Japanese region  
⇒ 1.5 hrs on 512 nodes  
for 120 hrs integration

# Experiments of Atmosphere Component

( with only cloud micro-physics )



horizontal : 11km  
vertical : 32 layers

1 way  
2 way  
nesting

horizontal : 1.13km  
vertical : 64 layers

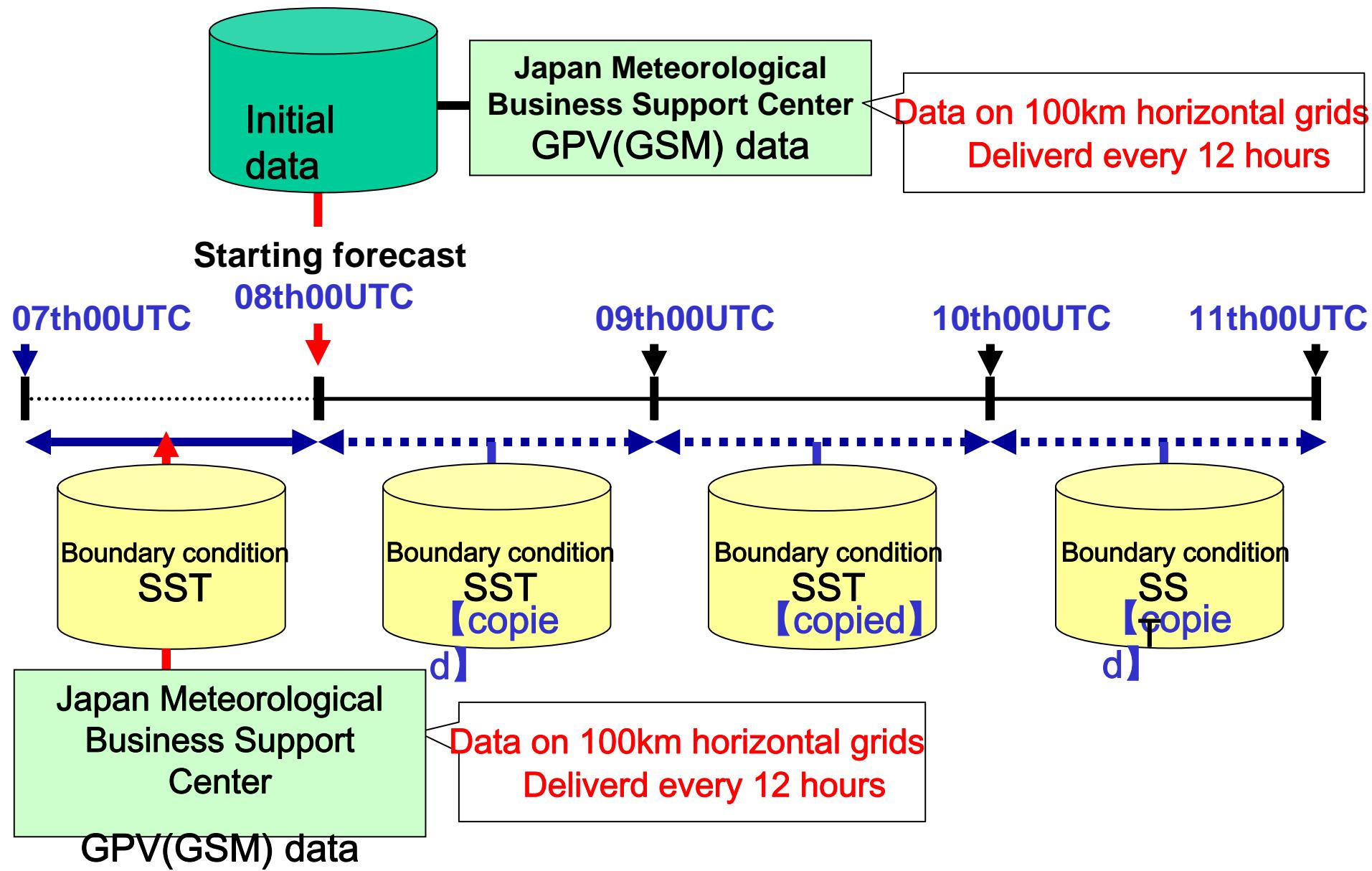
Global  
Simulations

Online  
1way, 2way  
nesting

Japanese region  
Simulations

# AGCM simulations : initial and boundary data

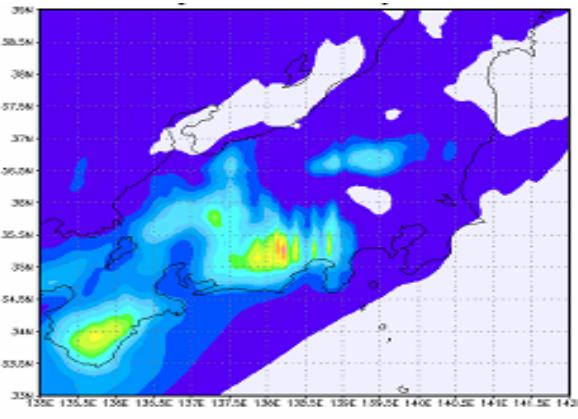
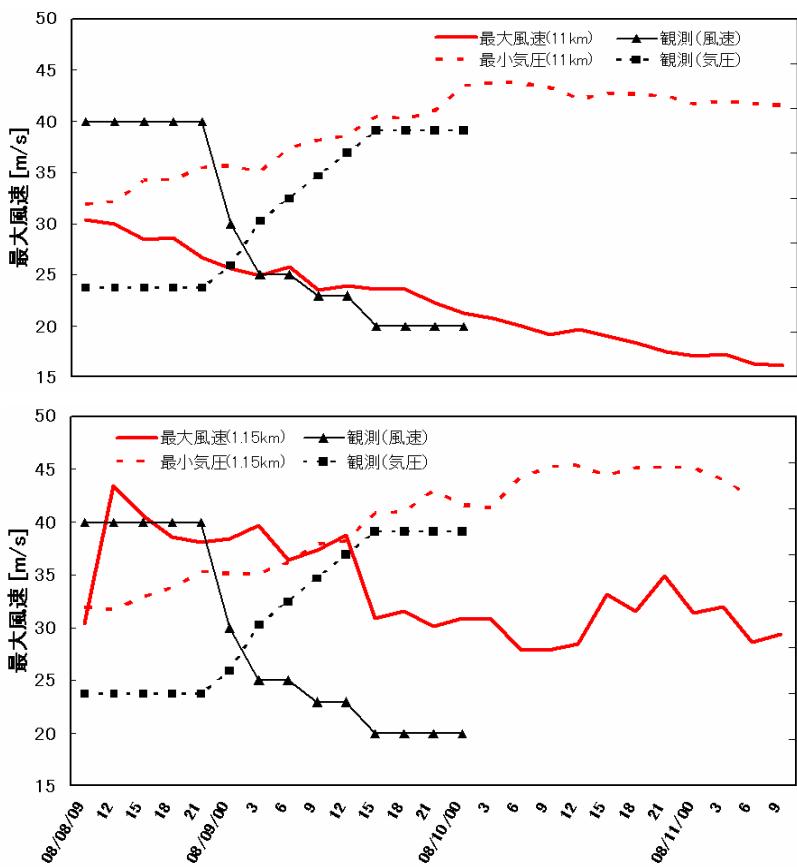
2003/8/8/00UTC ~ 11/8/00UTC: 72 hours forecasting



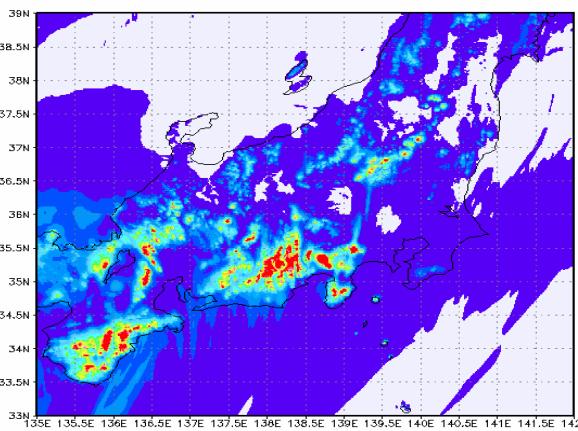
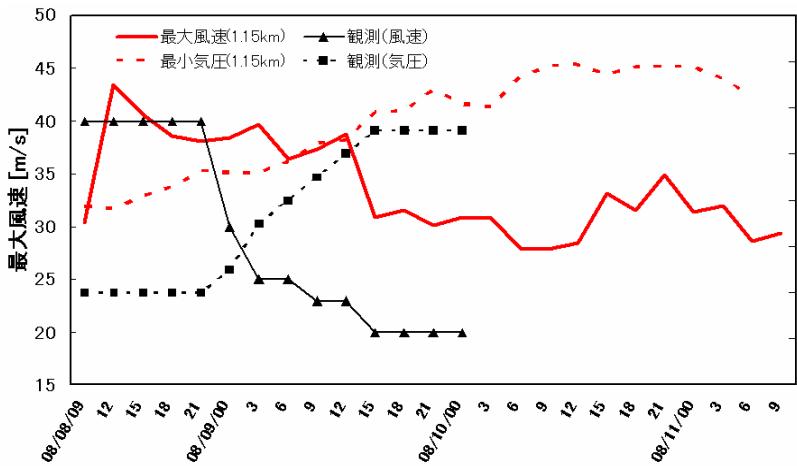
# Typhoon ETAU(2003) Intensity Atmosphere component, 1way nesting

00UTC 08 Aug – 00UTC 11 Aug, 2003

Global  
11km

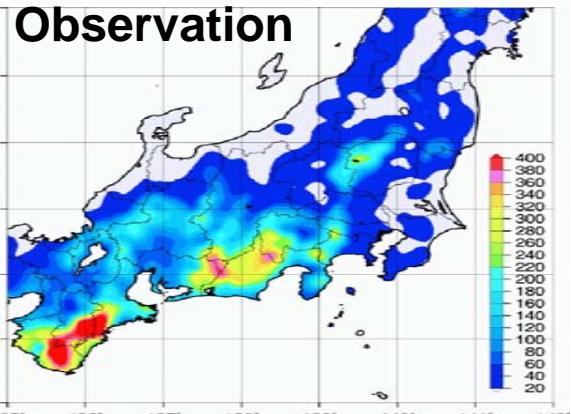


Nested  
Japanese  
Region  
1.13km



**Black: observation**  
dot line: min. surface pressure (hPa)  
solid line: max. wind speed (m/s)

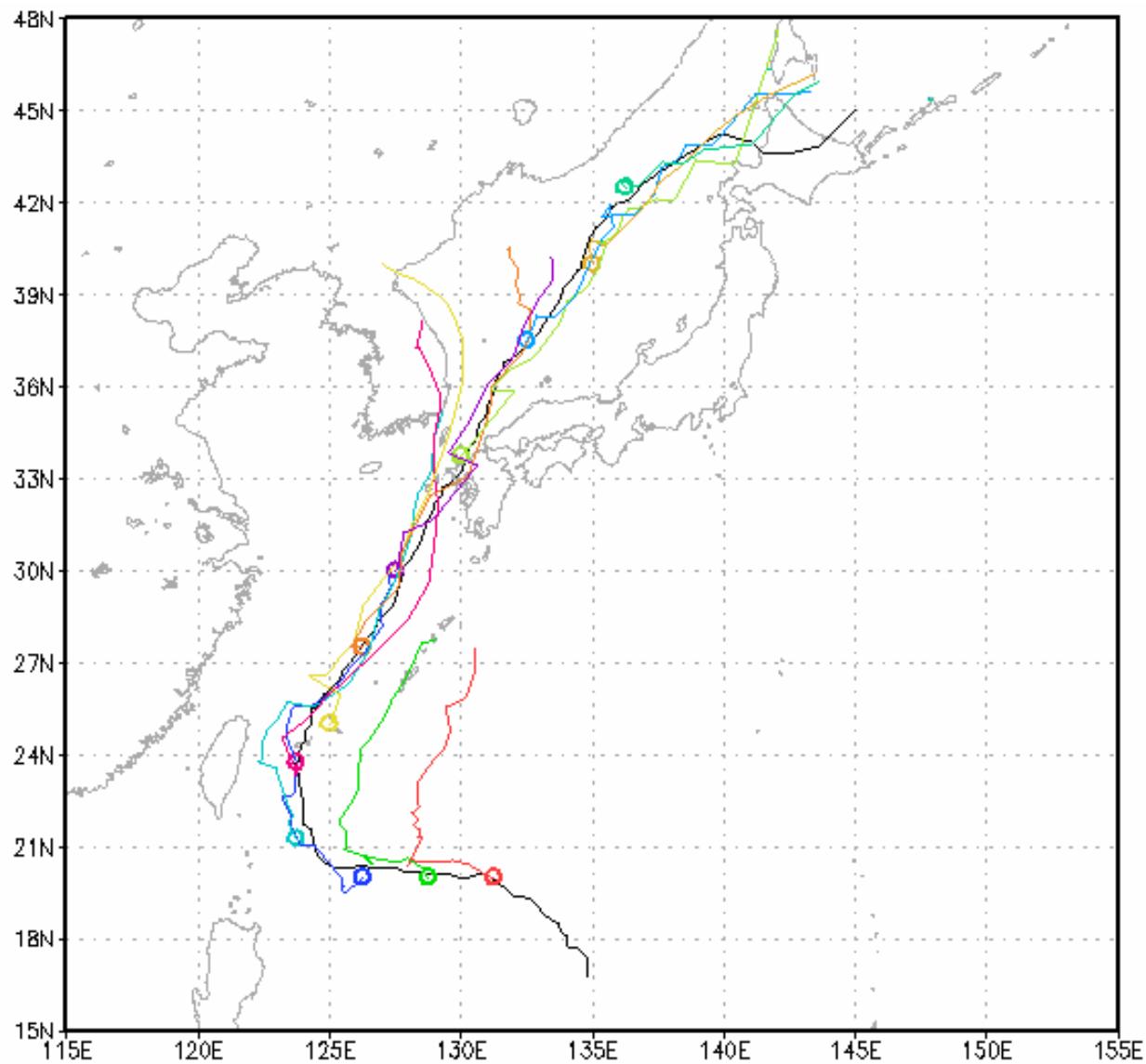
**Red: simulation results**  
dot line: min. SLP (hPa)  
solid line: max. wind speed (m/s)



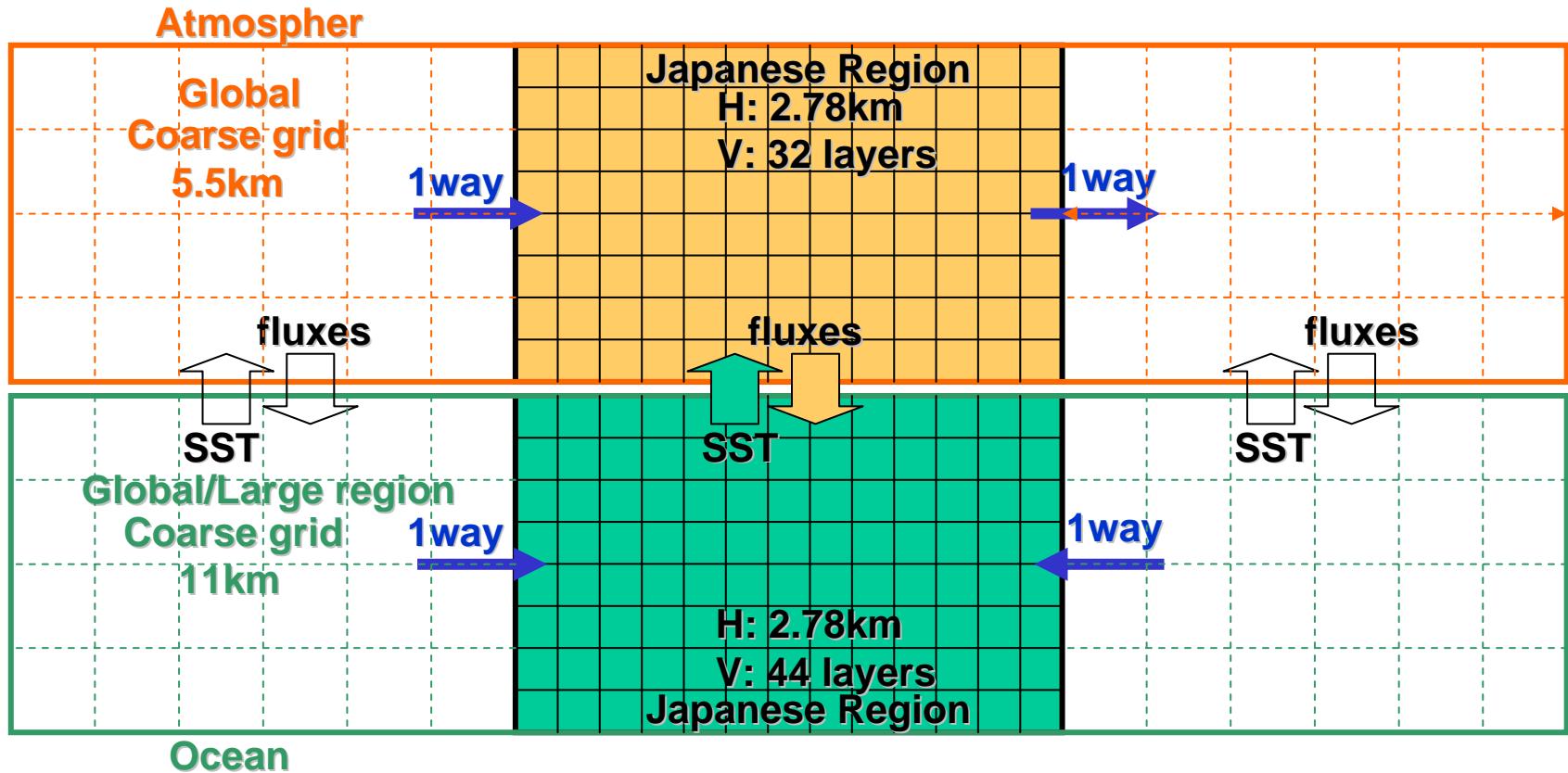
**Observation**

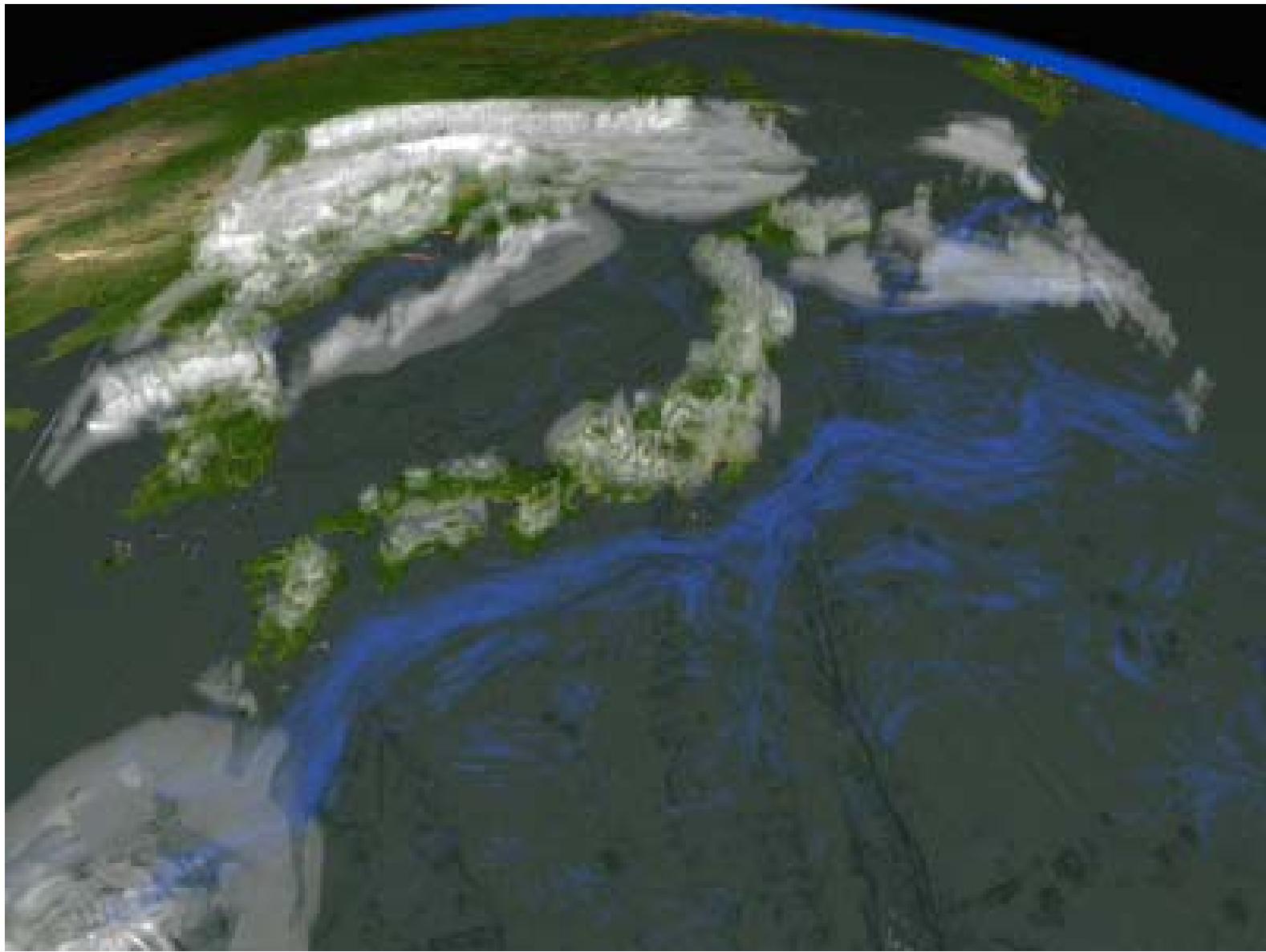


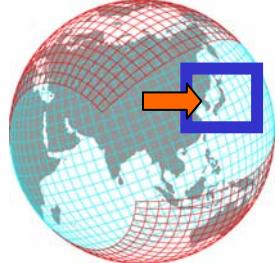
# Real time forecasting ShanShan, 12<sup>th</sup> Aug.-19<sup>th</sup> Aug. in 2006



# Regional Coupled Model

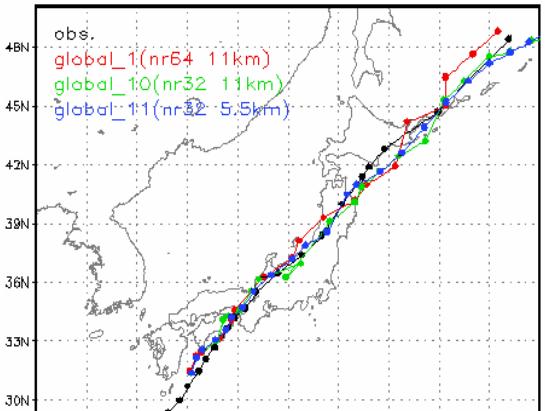






## Atmosphere Model

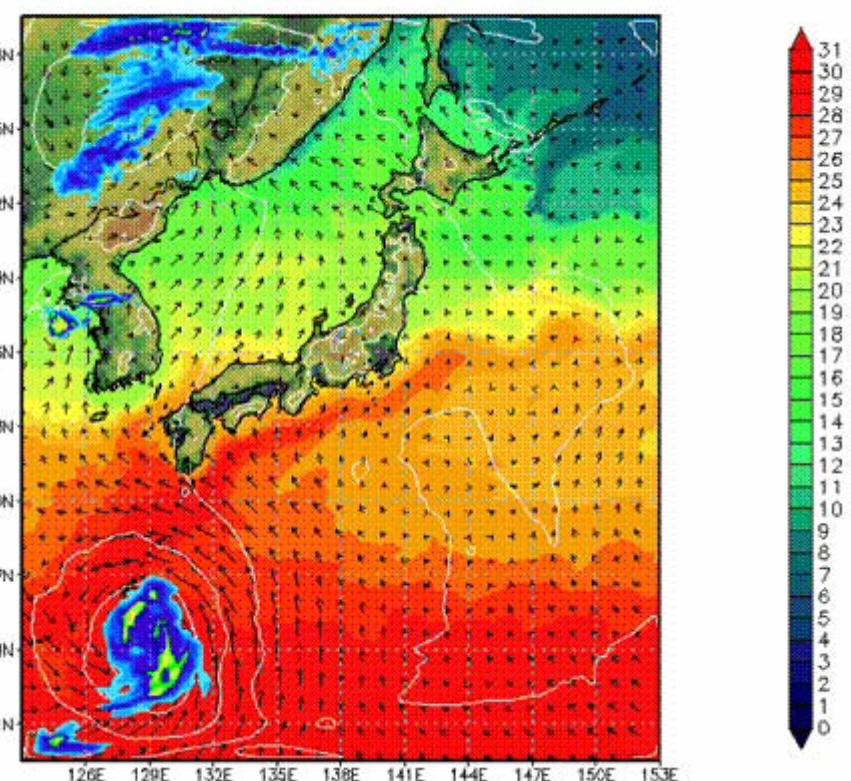
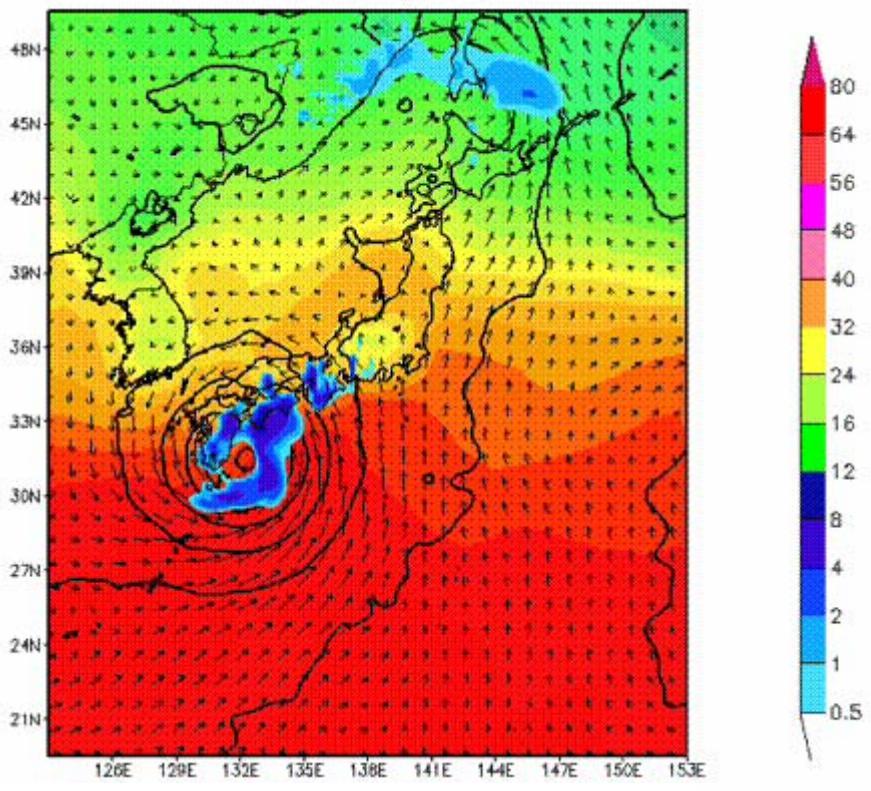
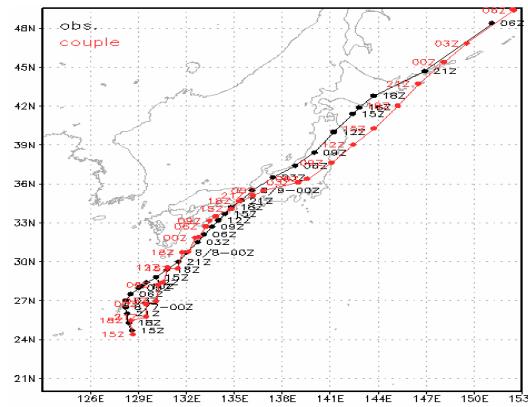
72hours forecast



Typhoon ETAU

## Coupled Atmosphere-Ocean Model

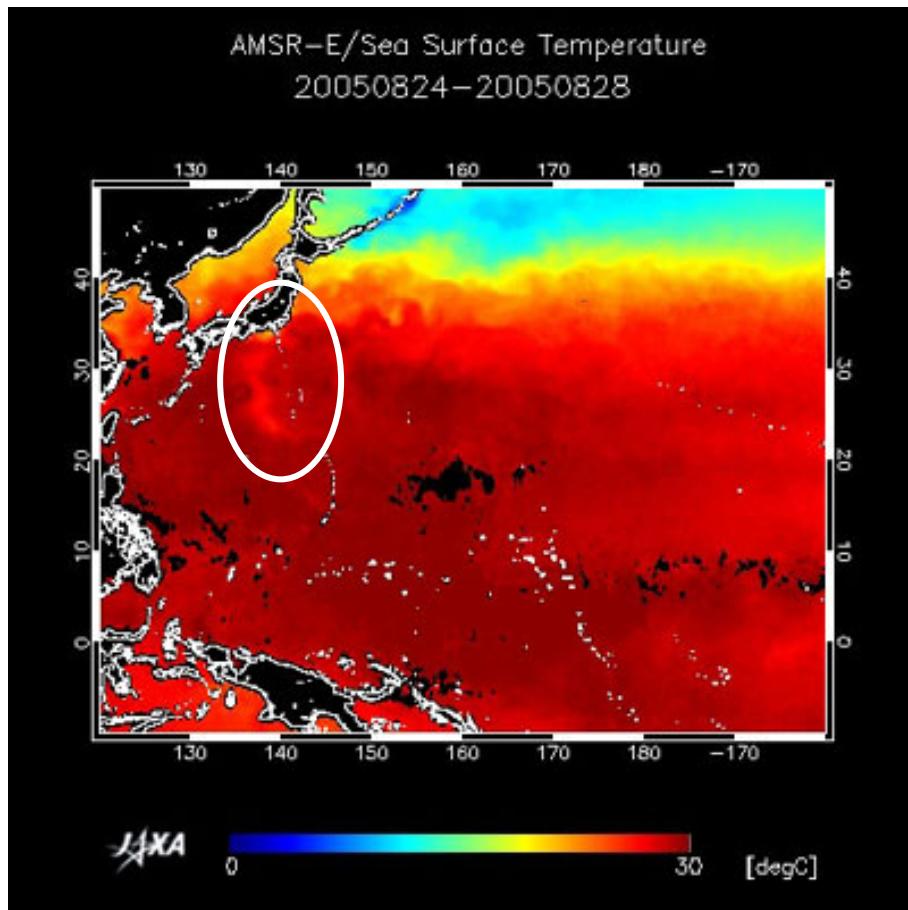
120 hours forecast



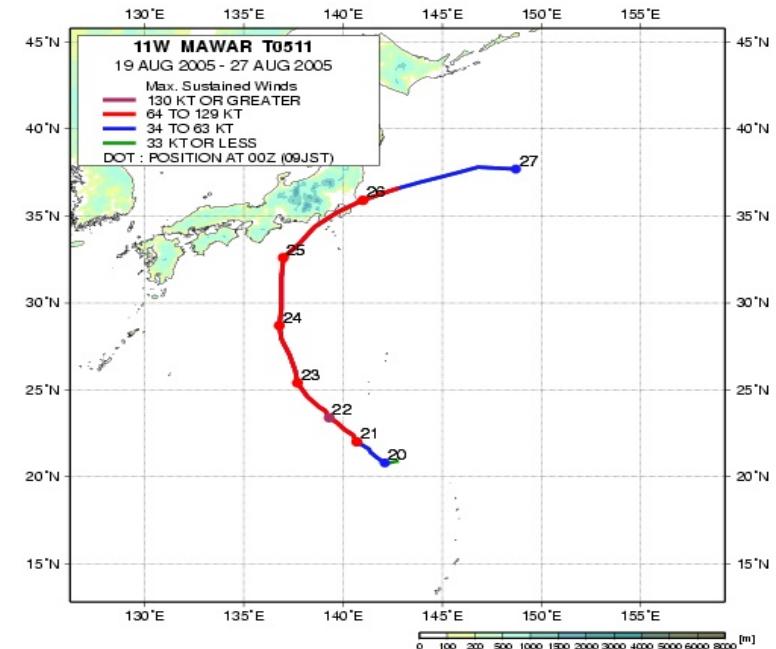
# Footmark of Typhoon 11

JAXA , <http://www.eorc.nasda.go.jp/imgdata/topics/2005/tp050922.html>

Observation data by AMSR-E in Aqua ( NASA)  
SST distribution averaged during 5 days  
from 24<sup>th</sup> August to 28<sup>th</sup> August in 2005



Tracking of typhoon 11  
From database in JAXA/EORC

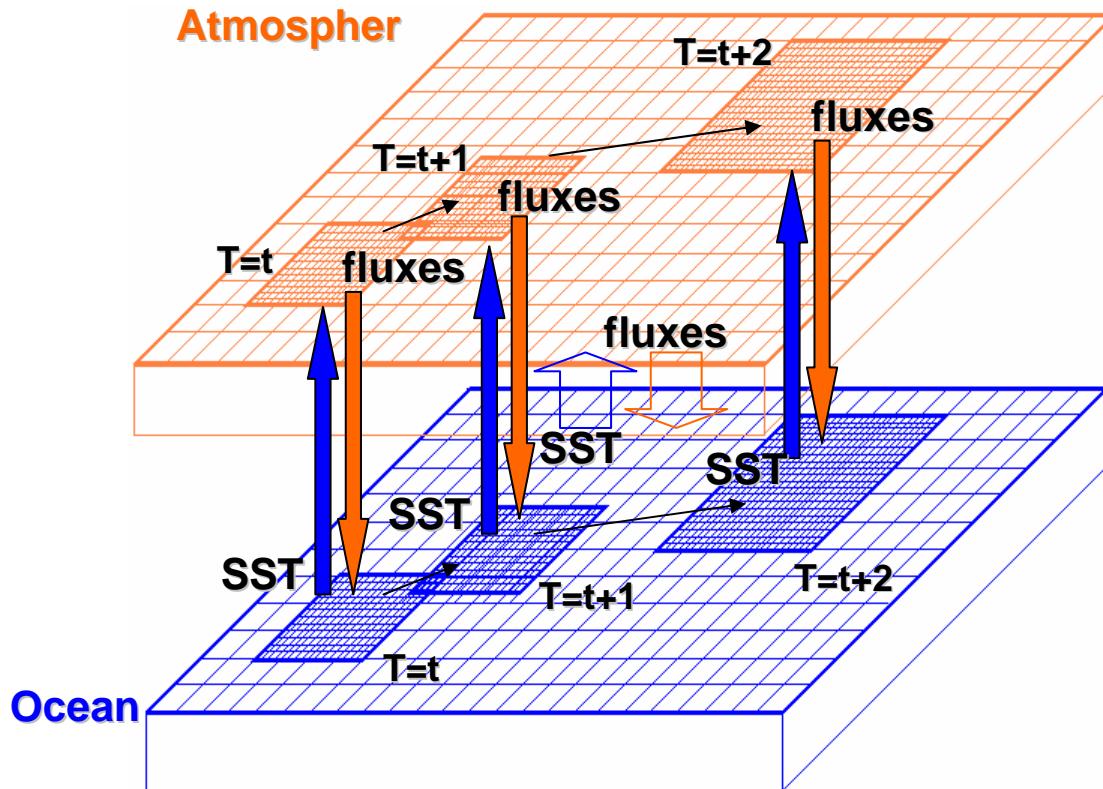




**Photo by Prof. Hasegawa in 1955**

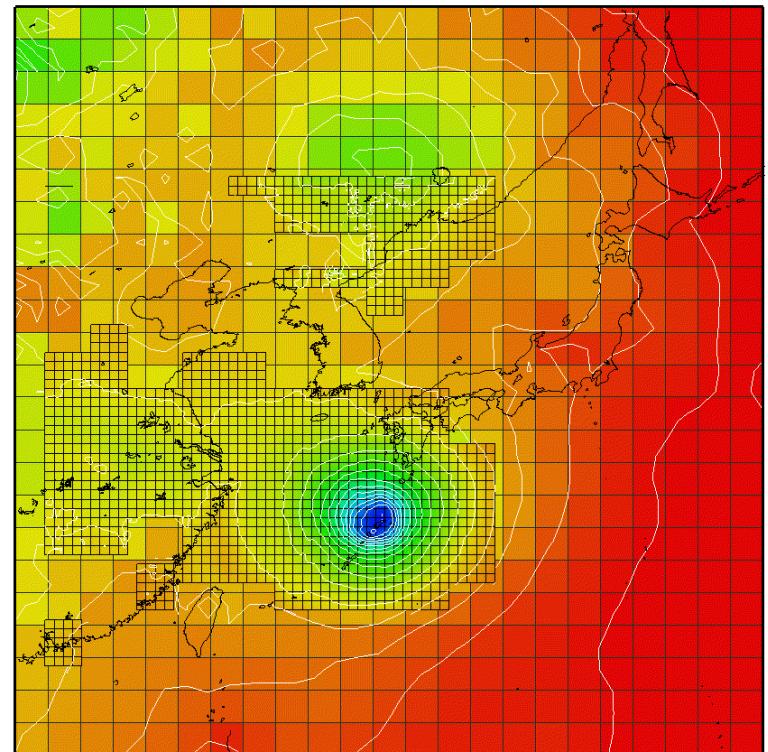
# Improvements in Model Development

Adaptive Mesh Refinement (AMR)  
for both Atmosphere and Ocean components



# Characteristics:

- **High Performance Computing**
  - No Overhead Computation for Moving Grid
  - Ultra High Parallelization
- **Multiple fine meshed regions are available**
- **Only Horizontal Refinement**
- **Refinement Criterion:**
  - low surface layer pressure
  - vorticity
  - gradient of physical parameters

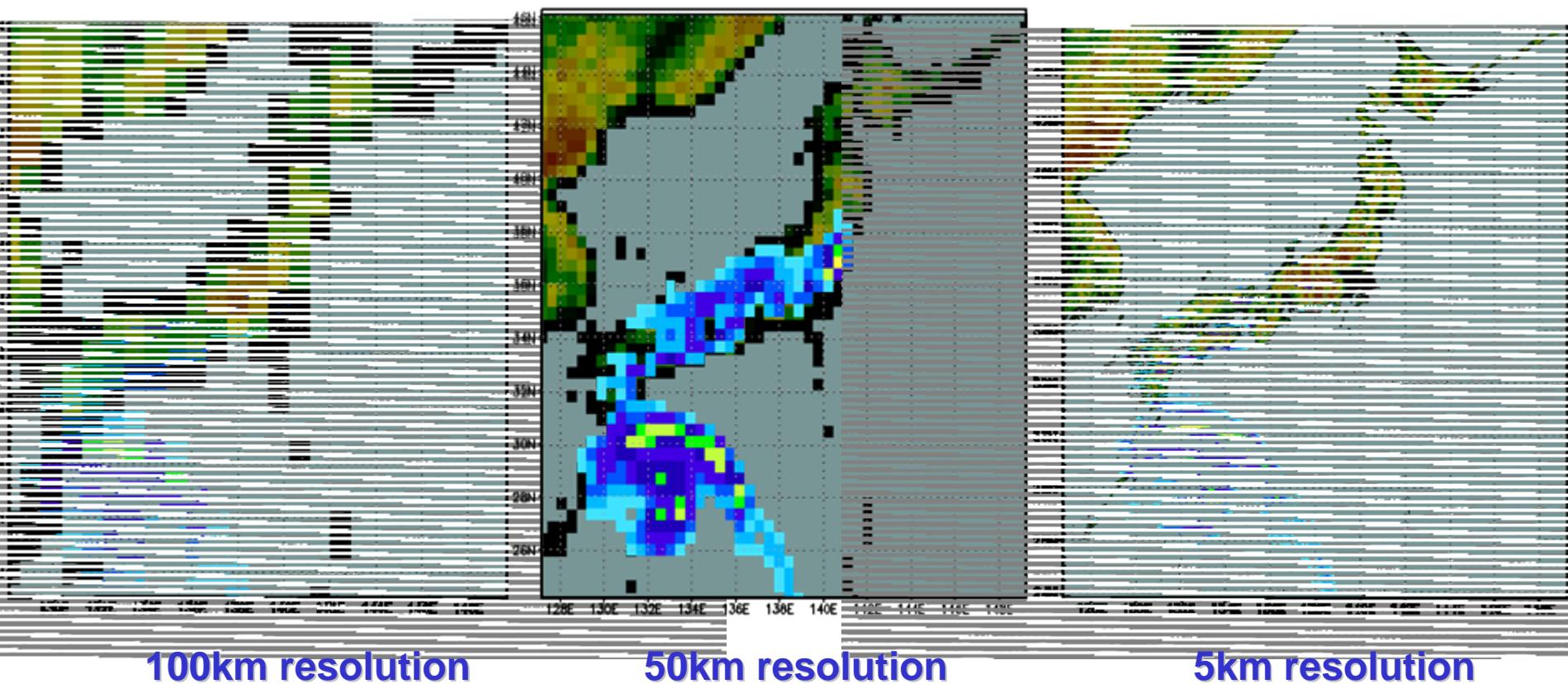


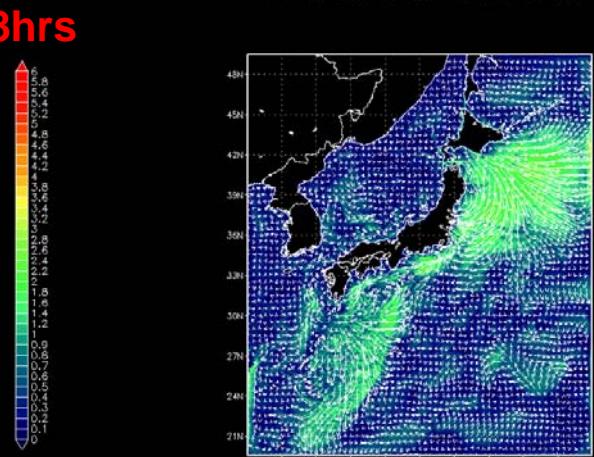
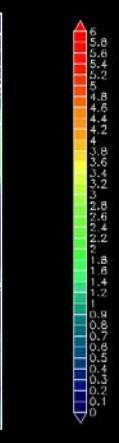
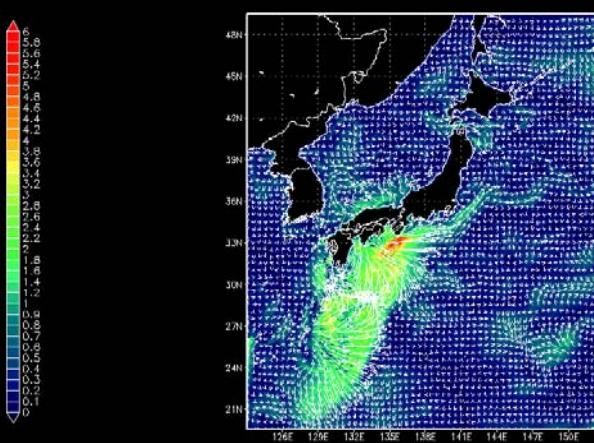
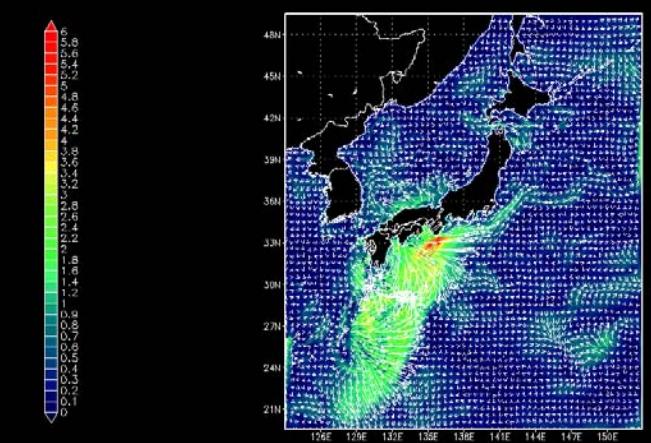
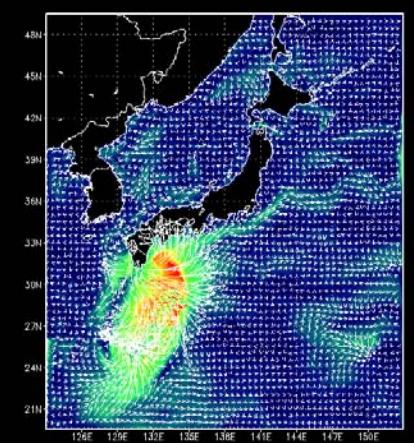
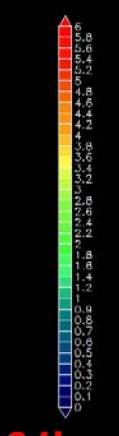
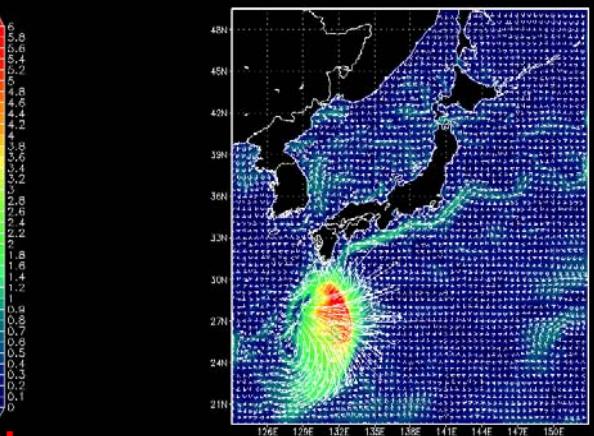
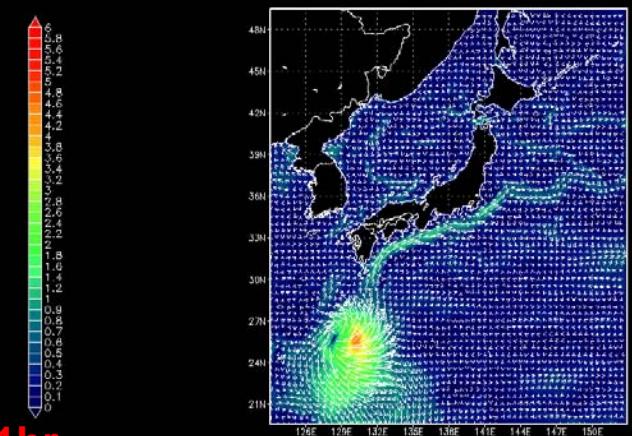
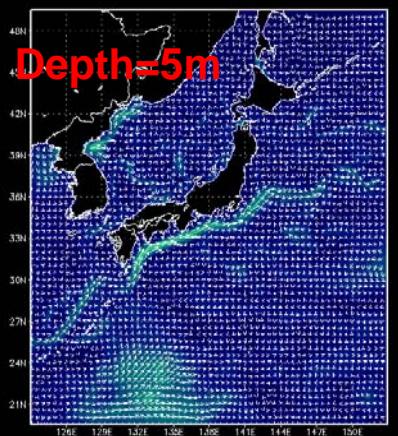
## **Near Future Work**

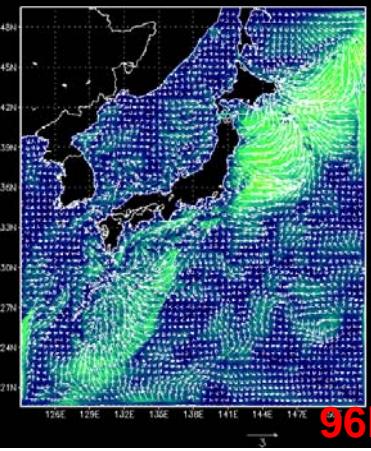
- Much more validation experiments  
for short term severe events.**
- Further accurate descritization schemes  
-Cost tuning with CIP-CSLR.**
- Progress of cloud micro-physics, radiation, interfaces  
-Effect of turbulence**
- Further longer integration validation experiments  
-Seasonal forecasting with ensemble methods**
- How to exchange information among  
different scale/physics phenomena ?**

**Thank you.**

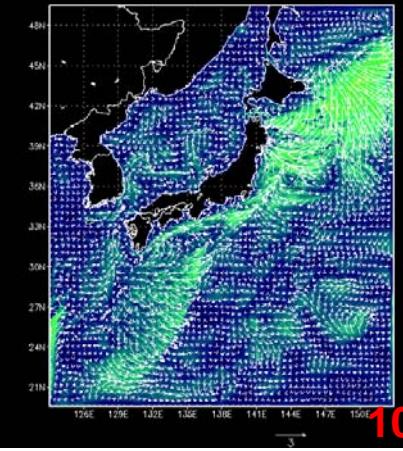
# High resolution and Precipitation distribution



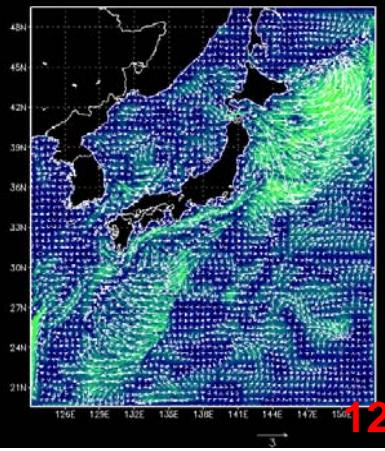




96hrs



108hrs

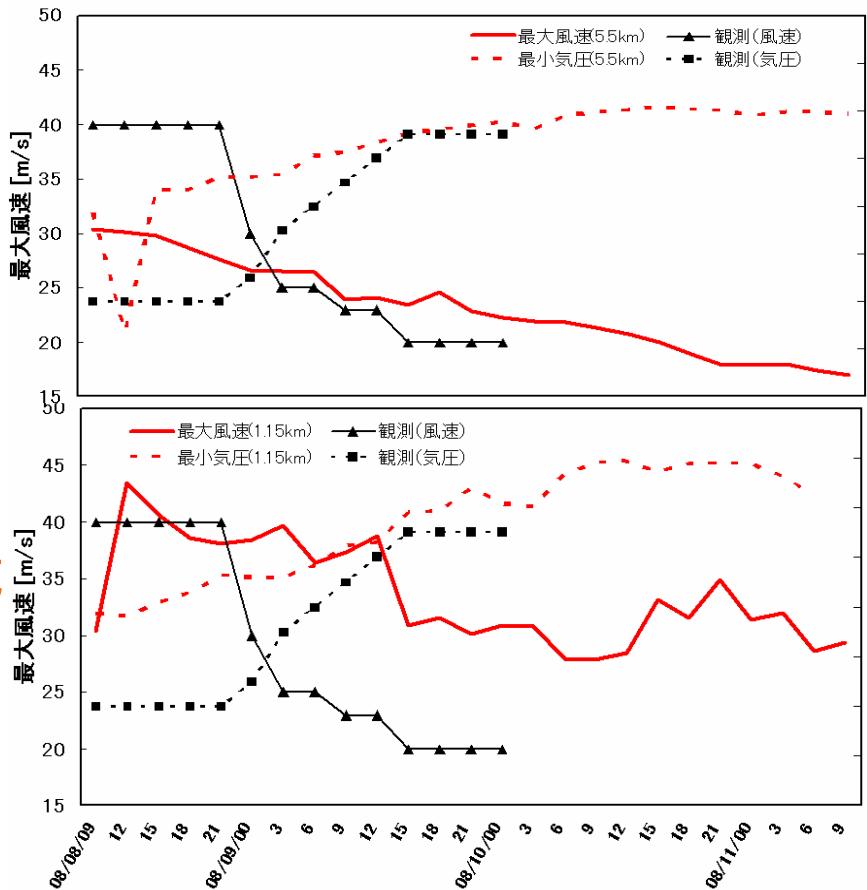


120hrs

# 台風強度予測シミュレーション

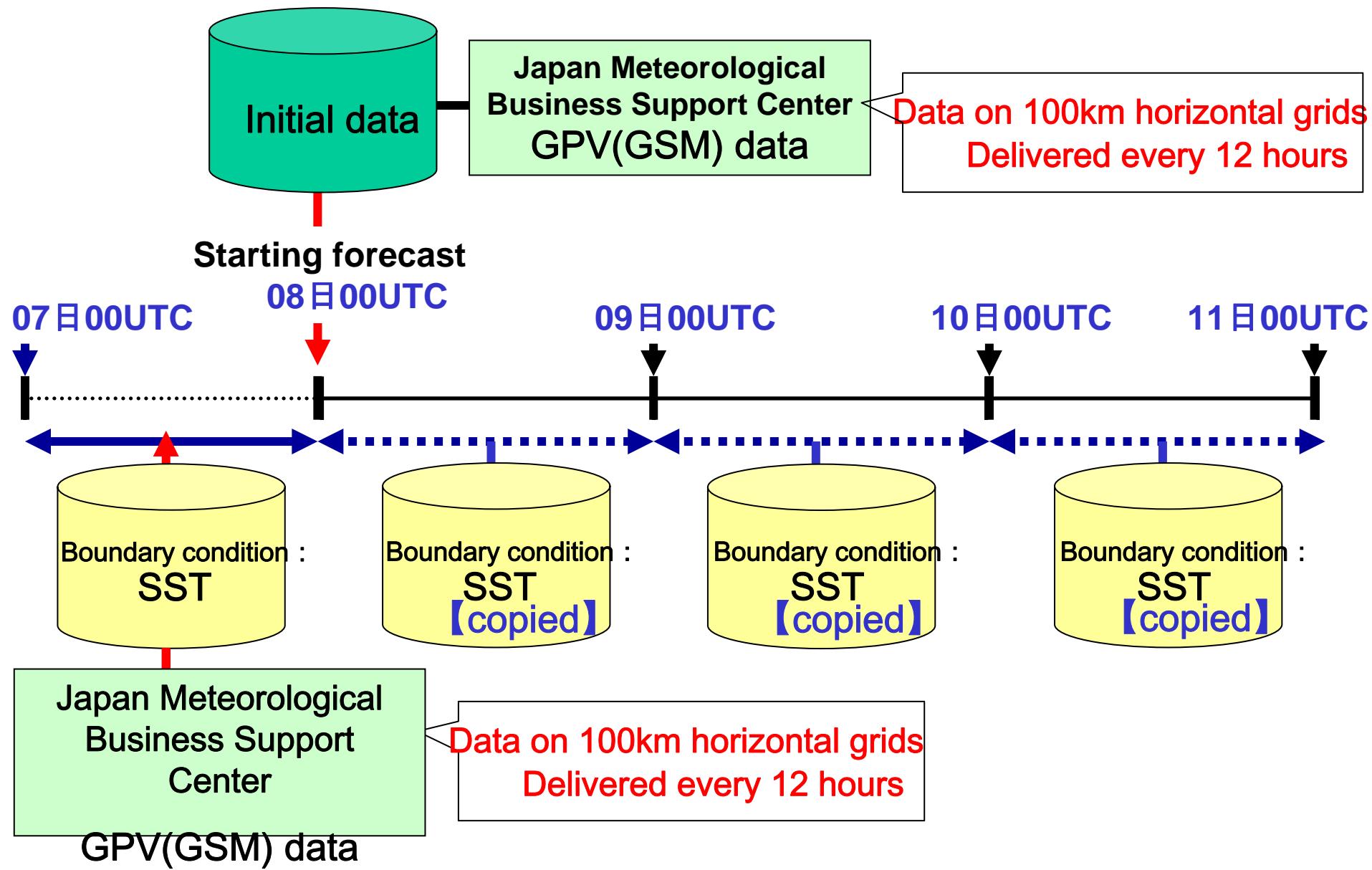
## - 解像度による違い -

全球大気  
5.5km

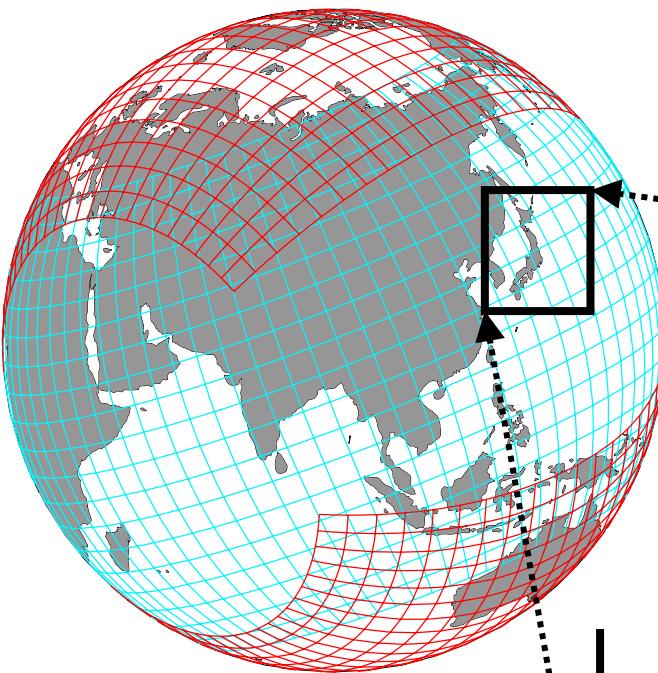


# AGCM simulations : initial and boundary data

2003/8/8/00UTC ~ 11/8/00UTC: 7 2 hours forecasting

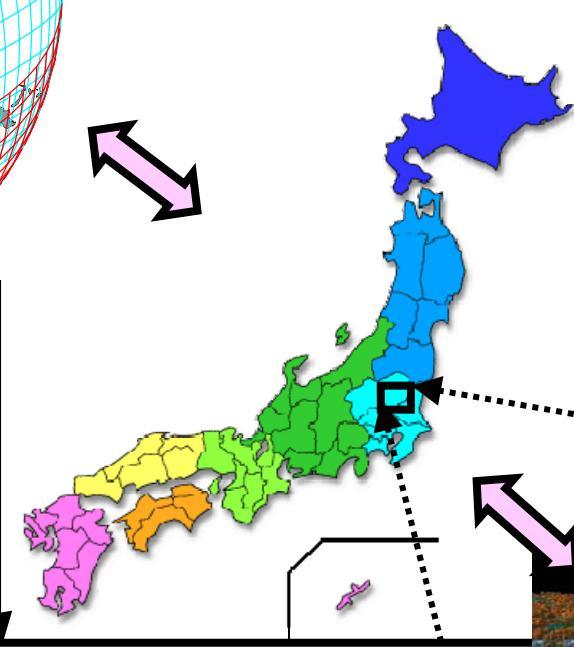


# Scalability



## Seasonal ~ Annual Prediction

5-40 km for horizontal,  
100 vertical layers



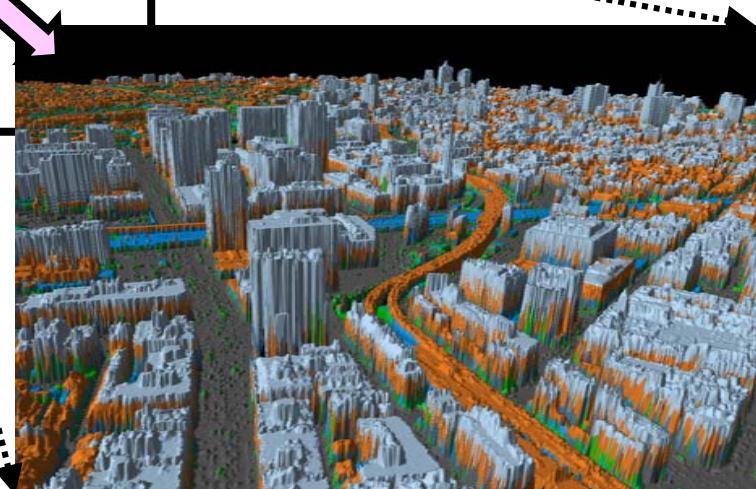
## Urban Weather /Climate Prediction

10m ~ 2km for horizontal,  
200 vertical layers

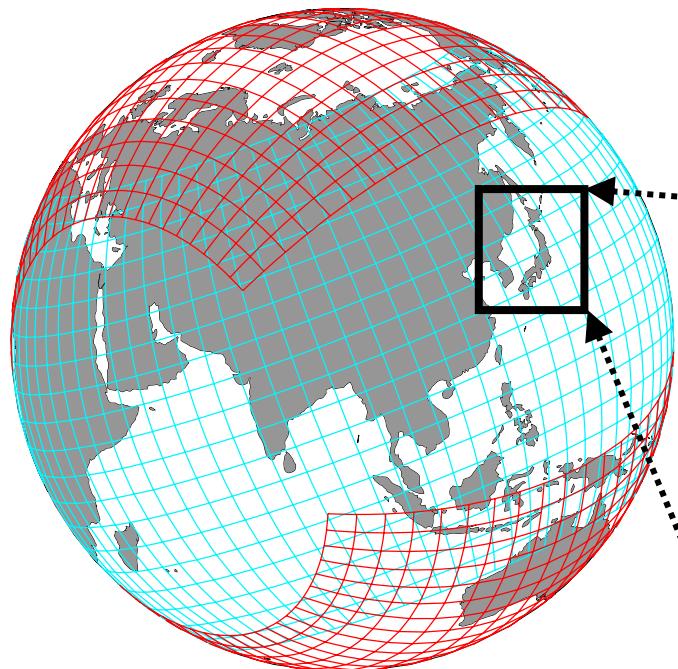
( Data: Geographical Survey Institute)

Days ~ Weeks Prediction  
Local heavy Rain Prediction,  
etc.

1 ~ 5km for horizontal  
100 vertical layers



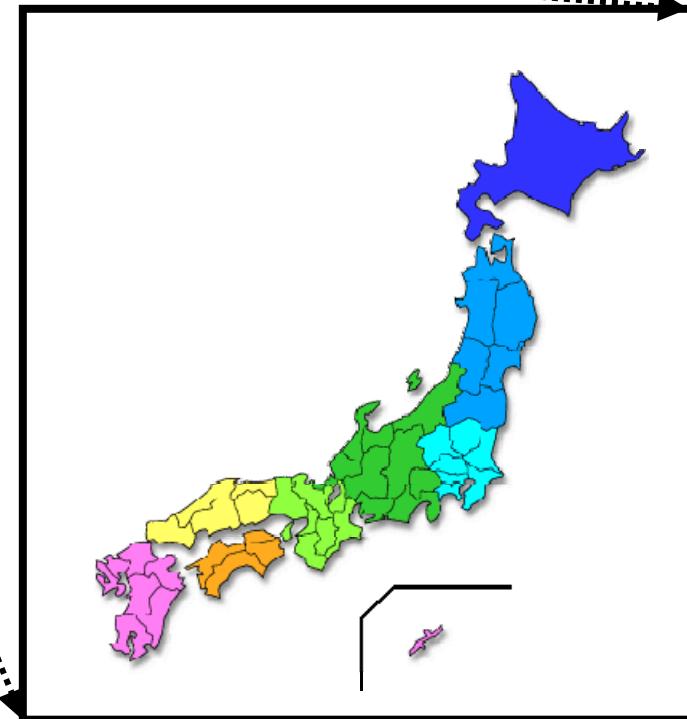
# Scalability



## Global

**Months ~ Years Prediction**

5-20 km for horizontal,  
100 vertical layers



## Reagional on the sphere

**Days ~ Weeks Prediction**

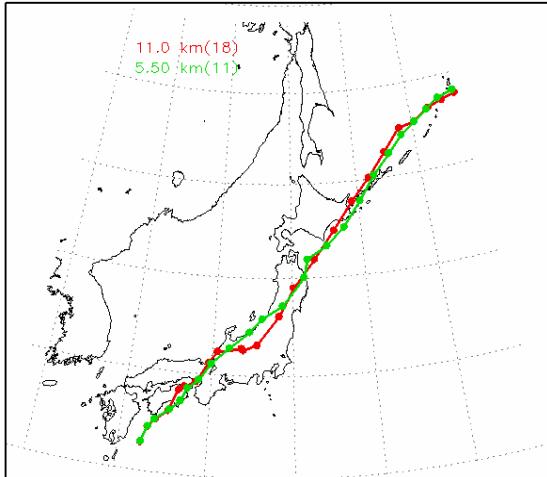
**Local heavy Rain Prediction, etc.**

1 ~ 5km for horizontal  
100 vertical layers

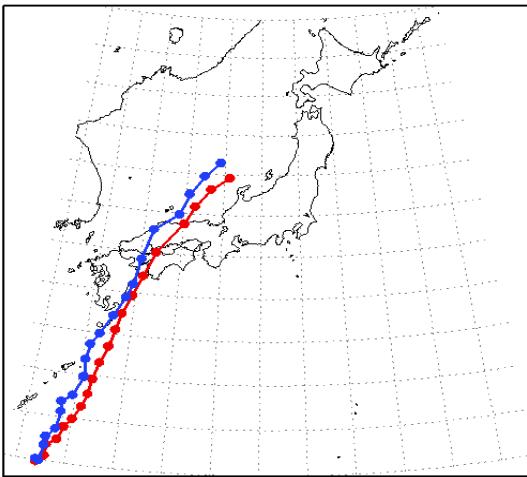
# Prediction of Typhoon tracking

## - simulation results -

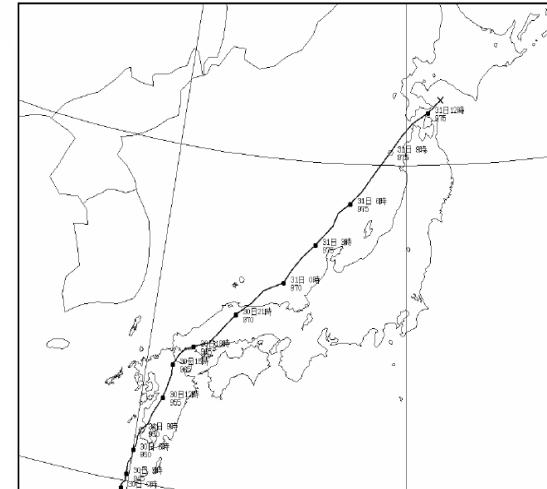
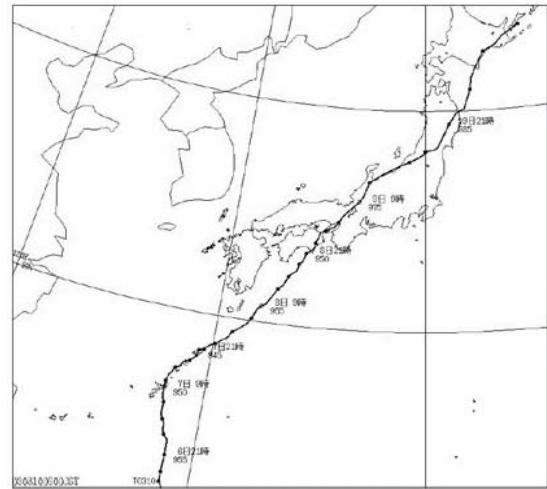
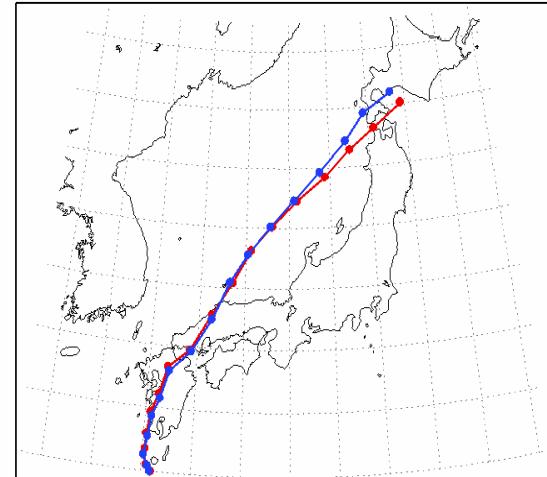
00UTC 08 Aug –  
00UTC 11 Aug, 2003



00UTC 19 Jun –  
00UTC 22 Jun, 2004



00UTC 29 Aug –  
00UTC 01 Sep, 2004



— Global 11km for horizontal, — Global 5.5km for horizontal

# Model Evaluation by Case Study

## Summer Case

- 48hr Global 60-km 32-lvls simulation
- Start at 00UTC Aug.8, 2003
- KF2 + large-scale condensation

06UTC Aug.8 2003

