

# Modelling of Innovation Statistics.

**Erik Andersson**

- ◆ Introduction
- ◆ Theory + computational methods
- ◆ Temporal aspects
- ◆ Validation of background error modelling
- ◆ Observation error correlation
- ◆ Summary and conclusions

## **Acknowledgement:**

**F. Bouttier, C. Cardinali, L. Isaksen, S. Saarinen**

# The new information

The innovations provide the new information to the assimilation.

The innovations ( $\mathbf{d}$ ) = The observed departures from the background

$$\mathbf{d} = \mathbf{y} - H\mathbf{x}_b$$

If the distribution of the data in time is accounted for, then

$$\mathbf{d} = \mathbf{y} - HM\mathbf{x}_b$$

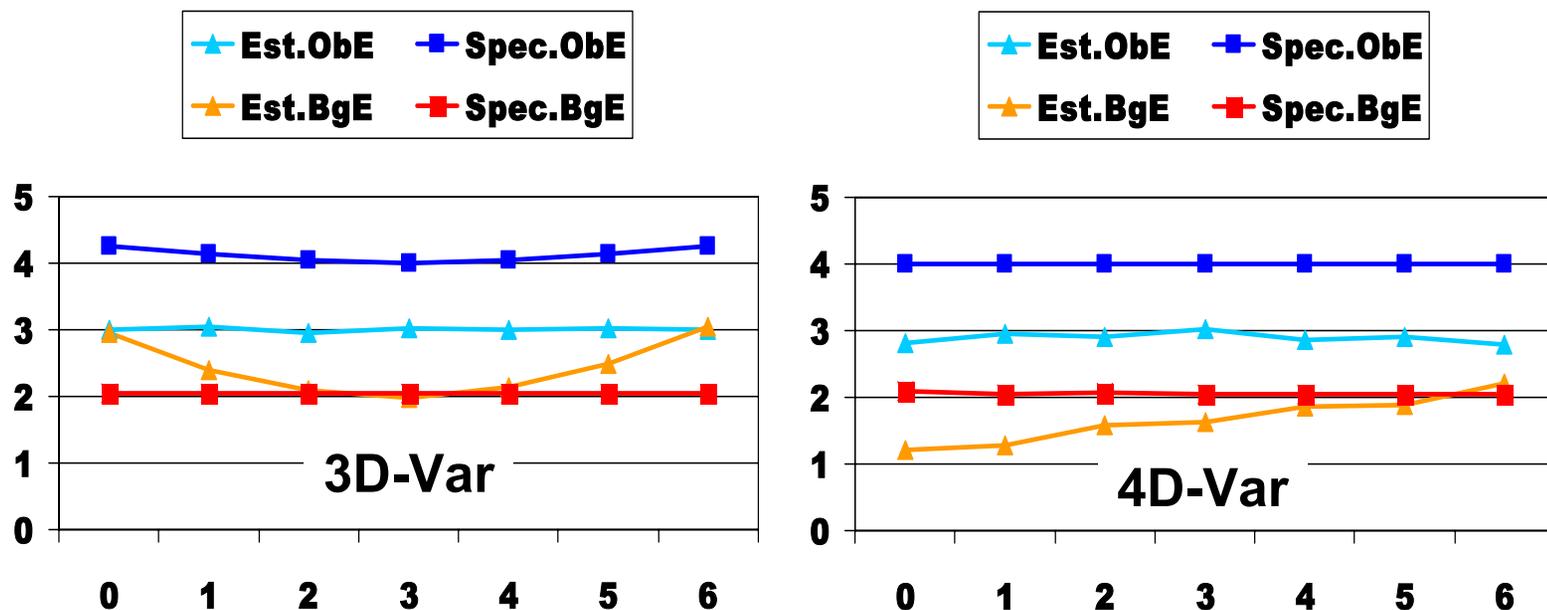
**The calculations of the innovations are carried out as accurately as practically possible:**

- ◆ We use the full non-linear forecast model  $M$ , at highest affordable resolution (T511)
- ◆ A large effort has been put on developing  $H$  to closely mimic the real observation (e.g. RTTOV)

# Temporal evolution of innovations

Heikki Järvinen (*Tellus* 2001) studied the innovations within 3D-Var and 4D-Var with a 6-hour assimilation window. Used Hollingsworth-Lönnberg (*Tellus* 1986) de-correlation method to isolate Obs and Bg errors.

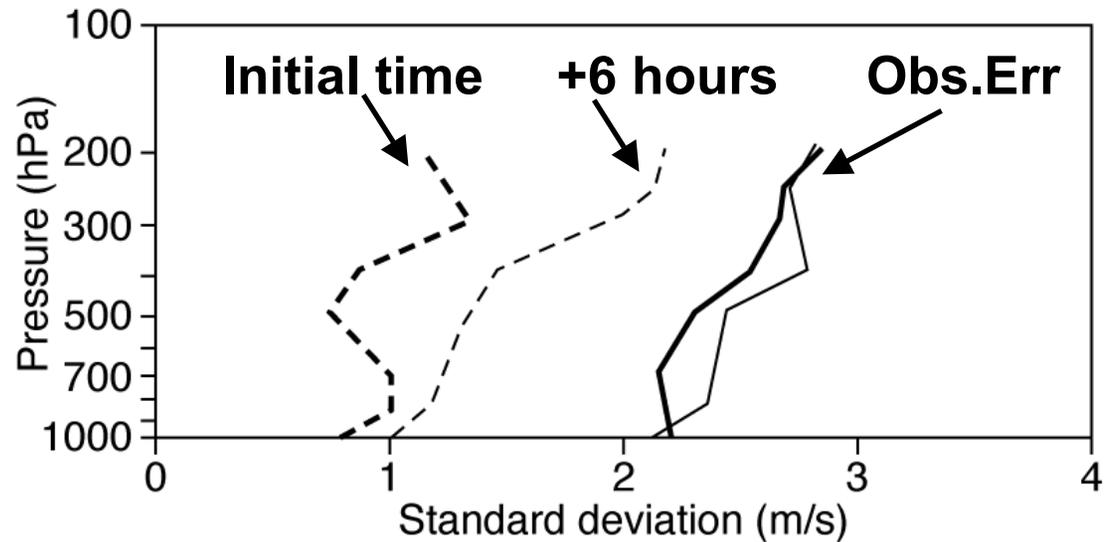
## Aircraft data, N.Amer, 200 hPa



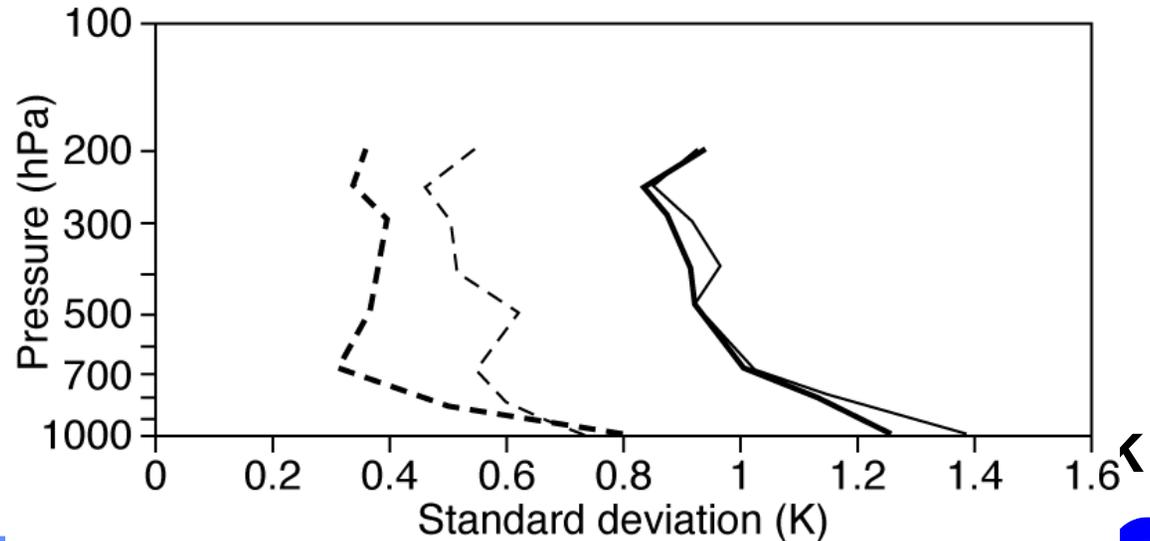
# Temporal evolution of innovations

## Over 6 hours for Aircraft data, North America

Wind component



Temperature

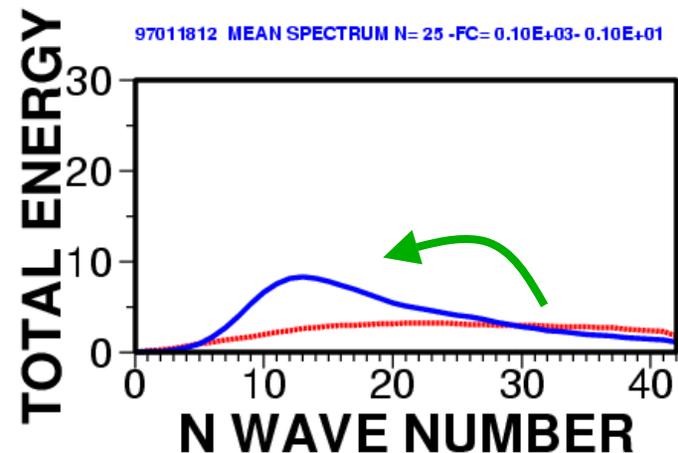
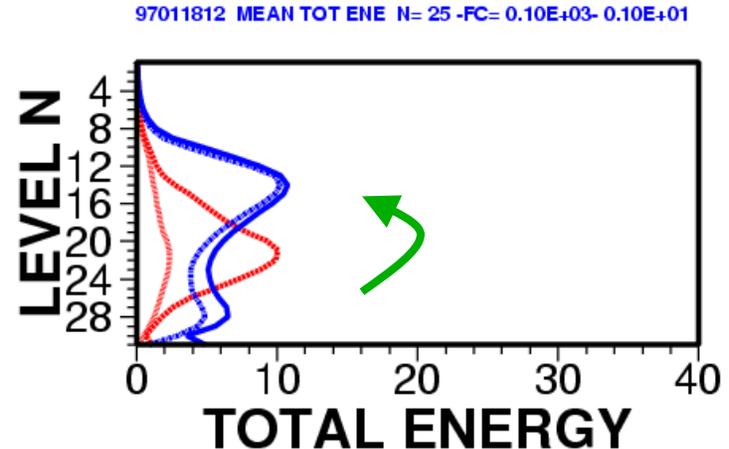


# Average SV energy distribution for 18-20 Jan 1997

SV1:25 average vertical distribution at **initial time** of the kinetic (**dotted**, x100) and total (**solid**, x100) energy, and the corresponding **final time** distributions. The bottom figure shows the SV1:25 average total energy spectrum at **initial** (x100) and at **final time**.

Note the SV typical **upward and upscale energy transfer/growth**, and the transformation **from initial potential to mainly final kinetic energy**.

**(R. Buizza)**



# The errors

Unfortunately, the observations  $\mathbf{y}$ ,  
the background  $\mathbf{x}_b$ ,  
the model  $M$  and  
the observation operators  $H$   
are all affected by errors.

Let  $\hat{\mathbf{x}}$  denote 'the truth', and  $\boldsymbol{\varepsilon}$  the error, then

$$\mathbf{y} = \hat{\mathbf{y}} + \boldsymbol{\varepsilon}_o \quad \langle \boldsymbol{\varepsilon}_o, \boldsymbol{\varepsilon}_o^T \rangle = \hat{\mathbf{O}} \quad \text{Observation error covariance}$$

$$\mathbf{x}_b = \hat{\mathbf{x}}_b + \boldsymbol{\varepsilon}_b \quad \langle \boldsymbol{\varepsilon}_b, \boldsymbol{\varepsilon}_b^T \rangle = \hat{\mathbf{B}} \quad \text{Background error}$$

$$M\hat{\mathbf{x}}_{(t=0)} = \hat{\mathbf{x}}_{(t=T)} + \boldsymbol{\varepsilon}_q \quad \langle \boldsymbol{\varepsilon}_q, \boldsymbol{\varepsilon}_q^T \rangle = \hat{\mathbf{Q}} \quad \text{Model error}$$

$$H\hat{\mathbf{x}}_{(t)} = \hat{H}\hat{\mathbf{x}}_{(t)} + \boldsymbol{\varepsilon}_f \quad \langle \boldsymbol{\varepsilon}_f, \boldsymbol{\varepsilon}_f^T \rangle = \hat{\mathbf{F}} \quad \text{Representativity error}$$

# The innovation covariance

The innovation covariance can be written

$$\langle \mathbf{d}, \mathbf{d}^T \rangle = \hat{\mathbf{H}} \hat{\mathbf{P}}^f \hat{\mathbf{H}}^T + \hat{\mathbf{O}} + \hat{\mathbf{F}} - (\hat{\mathbf{H}} \hat{\mathbf{X}}^T + \hat{\mathbf{X}} \hat{\mathbf{H}}^T)$$

with

**(Joiner and Dee, QJ 2000)**

$$\hat{\mathbf{P}}^f = \hat{\mathbf{M}} \hat{\mathbf{B}} \hat{\mathbf{M}}^T + \hat{\mathbf{Q}}$$

**Confusion surrounding ‘*Model error*’:**

- ◆ **Q = Model error, due to imperfections in  $M$**
- ◆ **MBM<sup>T</sup> = Predictability error, due to evolution of errors in the initial conditions**
- ◆ **P<sup>f</sup> = MBM<sup>T</sup> + Q = Forecast error**
- ◆ **B = Bg-error = Initial condition error**

# 4D-Var approximations

In our 4D-Var the 'true' co-variances are approximated:

$$\hat{\mathbf{O}} + \hat{\mathbf{F}} \approx \mathbf{R}$$

R diagonal

$$\hat{\mathbf{X}} \approx 0$$

No cross co-variances

$$\hat{\mathbf{Q}} \approx 0$$

Perfect model assumption

$$\hat{\mathbf{H}}\hat{\mathbf{P}}^f\hat{\mathbf{H}}^T \approx \mathbf{H}\mathbf{M}\mathbf{B}\mathbf{M}^T\mathbf{H}^T$$

Tangent linear obs. operators

Tangent linear forecast model

$$\text{4D-Var: } \mathbf{H}\mathbf{M}\mathbf{B}\mathbf{M}^T\mathbf{H}^T + \mathbf{R}$$

TL dynamics?

TL physics?

$$\text{3D-Fgat: } \mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R}$$

Truncation?

$$\text{OI: } \mathbf{B}_o + \mathbf{R}$$

B given through  $\mathbf{J}_b$ -modelling

# Which data are useful?

At this point we can conclude that:

◆ An analysis scheme which models innovation covariances well is a good analysis scheme. We put our effort on

- ◆ Characterizing background error – but still simplified
- ◆ Using the forecast model to evolve errors – but  $Q=0$
- ◆ Developing accurate observation operators – but  $R=O+F\%$

Conversely:

◆ Observations whose innovations are easily modelled, are useful observations

- ◆ Un-correlated with the background
- ◆ Un-correlated with other observations
- ◆ Accurately characterized by  $H$

# Validation

From samples of innovations we can compute

$$\langle \mathbf{d}, \mathbf{d}^T \rangle$$

If we knew how to diagnose

$$\mathbf{HMBM}^T \mathbf{H}^T + \mathbf{R}$$

in the full 4D-Var system,

then the two could be compared, and some shortcomings due to the modelling assumptions might become apparent.

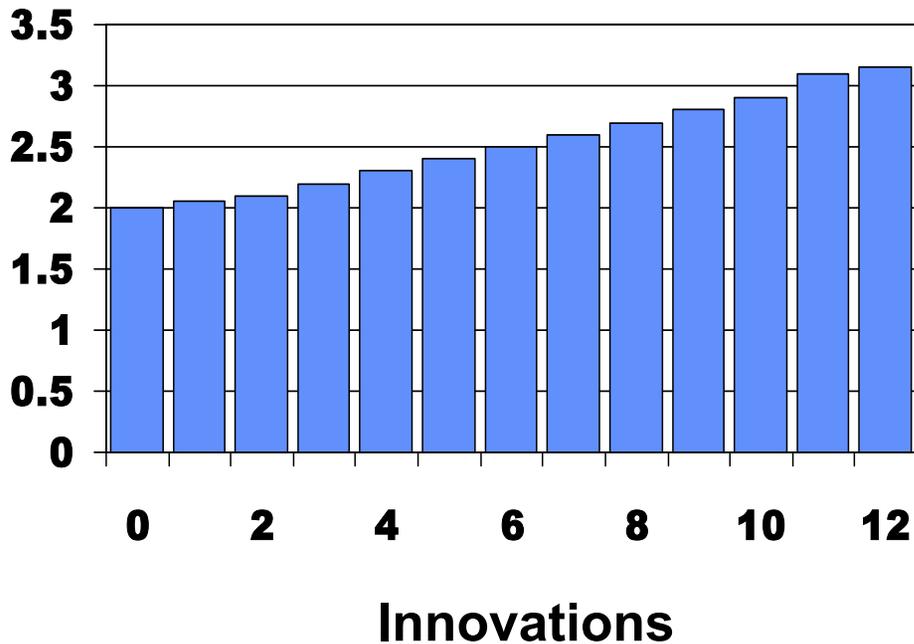
**Discrepancies could be due to H, M, B, R or Q !!!**

In a 'well-tuned' system:  $\langle \mathbf{d}, \mathbf{d}^T \rangle \cdot \mathbf{HMBM}^T \mathbf{H}^T + \mathbf{R}(+\mathbf{Q})$

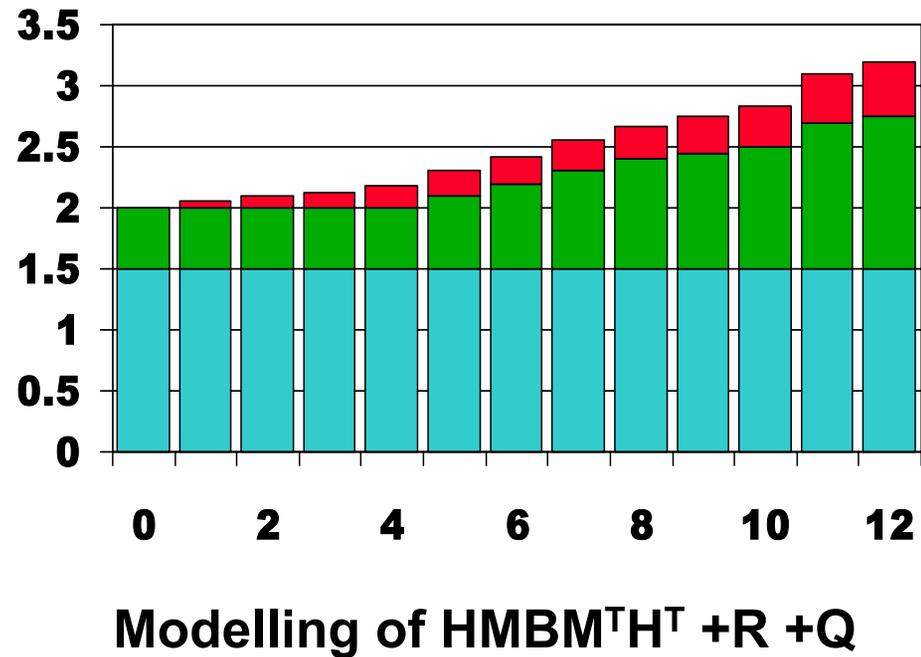
# What we expect...

(Fictitious example, for illustration only)

■ Obs-Bg



■ R ■ MBM ■ Q



# Diagnosing $\mathbf{HMBM}^T \mathbf{H}^T$ in 4D-Var

Due to its definition, the 4D-Var control-variable  $\mathbf{P}$  is a standard multivariately normal quantity, i.e.:

$$\boldsymbol{\chi} \sim \mathcal{N}(0, \mathbf{I})$$

and:  $\mathbf{L}\boldsymbol{\chi} = \boldsymbol{\delta\mathbf{x}} \sim \mathcal{N}(0, \mathbf{B})$

$$\mathbf{H}\mathbf{M}\boldsymbol{\delta\mathbf{x}} \sim \mathcal{N}(0, \mathbf{HMBM}^T \mathbf{H}^T)$$

**Randomization:** Generate a random sample of  $N$  vectors,  $\mathbf{P}^{(N)}$ , with zero mean and unit variance, then

$$\boldsymbol{\chi}^{(N)} (\boldsymbol{\chi}^{(N)})^T \equiv \mathbf{I}^{(N)}$$

$$\mathbf{L}\boldsymbol{\chi}^{(N)} (\mathbf{L}\boldsymbol{\chi}^{(N)})^T \equiv \mathbf{B}^{(N)}$$

Similarly, an estimate of  $\mathbf{HMBM}^T \mathbf{H}^T$ , can be obtained from

$$\mathbf{H}\mathbf{M}\mathbf{L}\boldsymbol{\chi}^{(N)} (\mathbf{H}\mathbf{M}\mathbf{L}\boldsymbol{\chi}^{(N)})^T \equiv \mathbf{HMBM}^T \mathbf{H}^T \quad (\text{M. Ehrendorfer})$$

# Diagnosing $\mathbf{HMBM}^T \mathbf{H}^T$ in 4D-Var

We compute

$$\mathbf{HMBM}^T \mathbf{H}^T \approx \frac{1}{N} \sum_{i=1}^N \mathbf{HML} \chi_i (\mathbf{HML} \chi_i)^T$$

for a sample of  $N=100$  vectors, accumulating variance-contributions for the diagonal elements only.

The number of estimated diagonal elements = the number of used observations (. 3,500,000).

The uncertainty in the resulting randomization estimate is about 3% .

In the following, we will see results of such calculations with  $N=100$ , for:

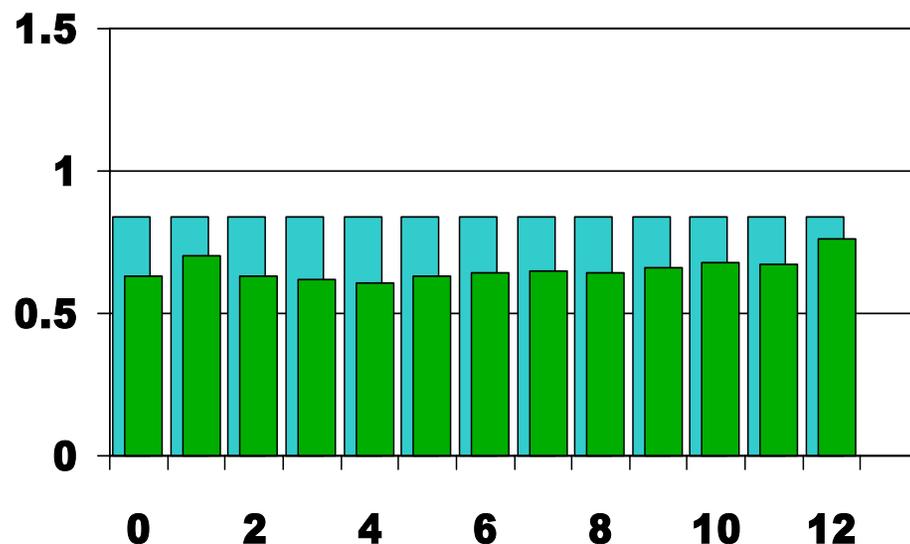
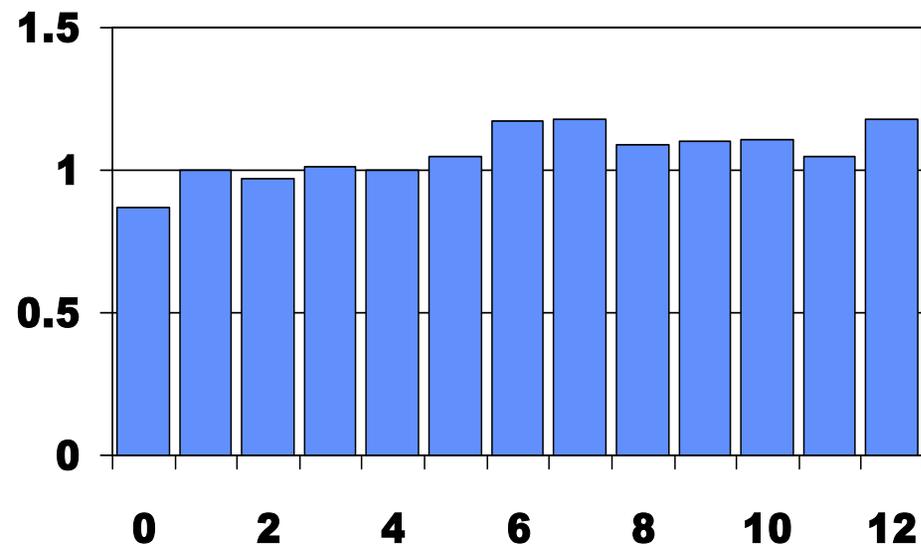
- ◆ A few H of used data in 4D-Var
- ◆ Current L (that is,  $\mathbf{J}_b$ ), and some earlier versions

# DRIBU, North Atlantic

## Surface pressure data (hPa)

Obs-Bg

R HMBMH



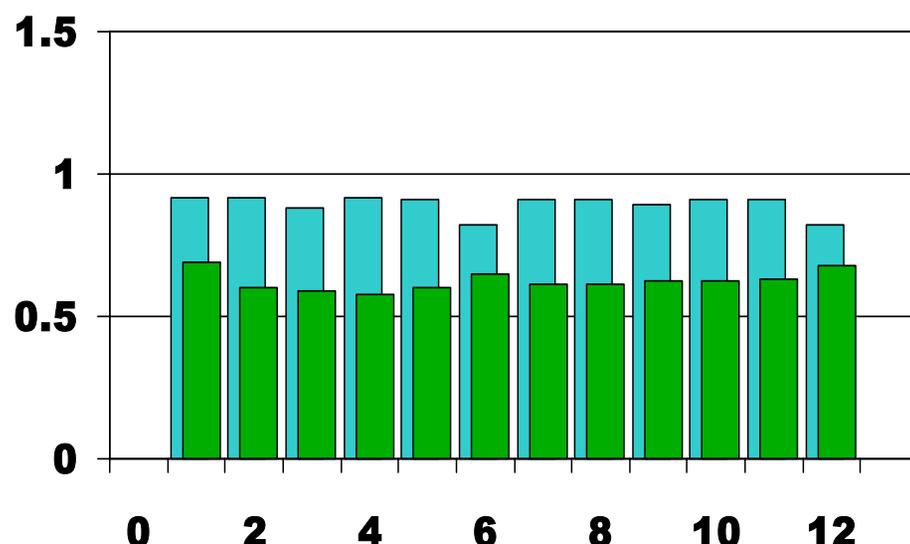
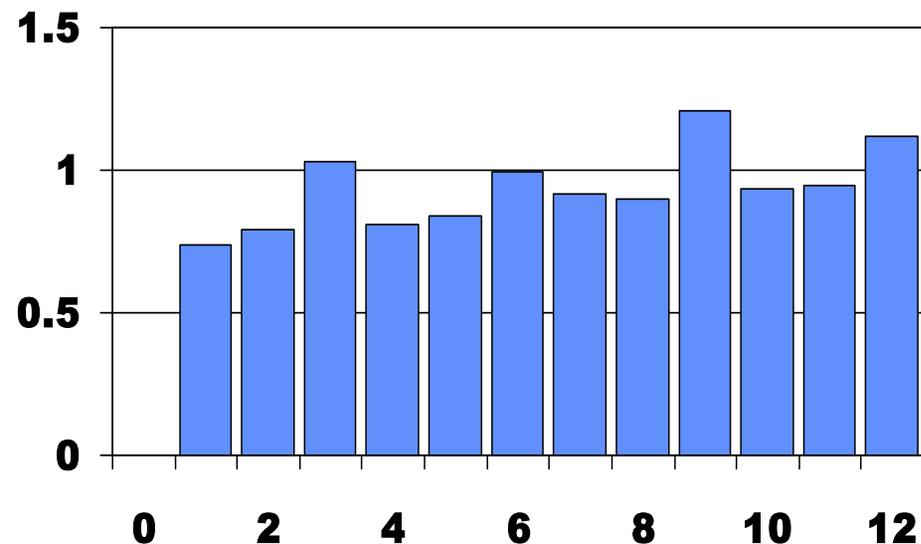
20030205-12 to 20030211-12,

About 600 data per bin

# SYNOP, North Atlantic Surface pressure data (hPa)

Obs-Bg

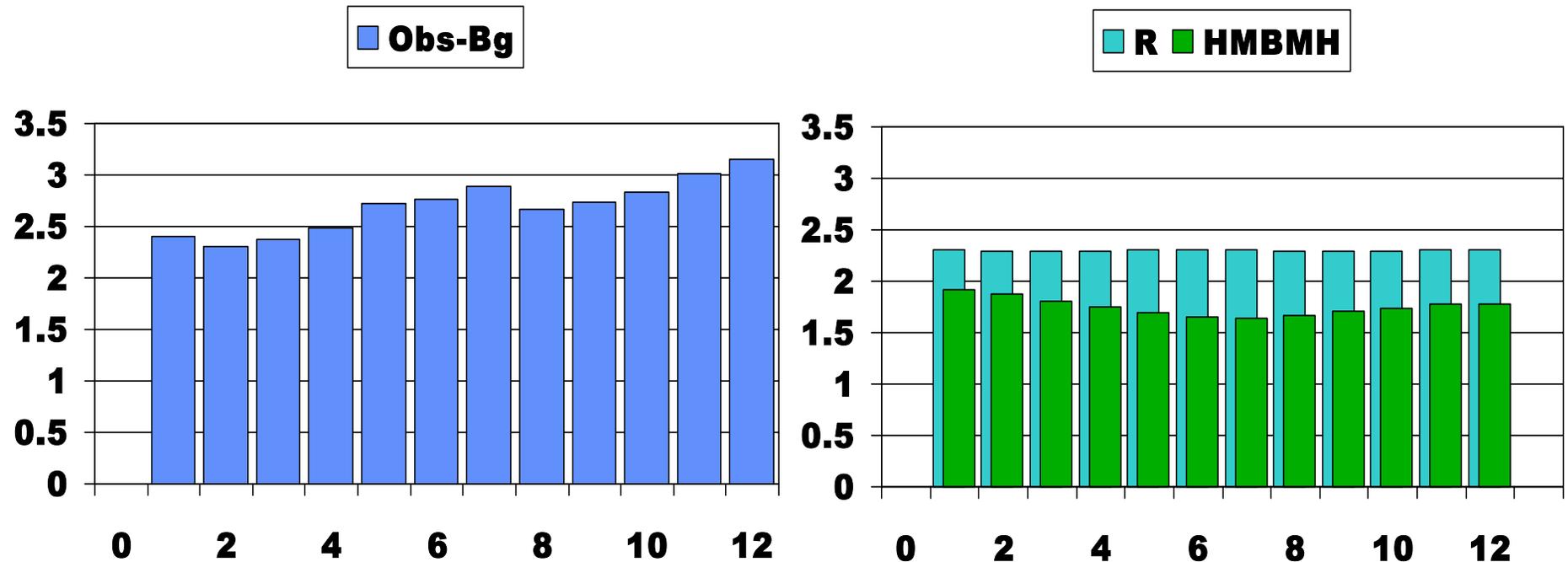
R HMBMH



20030205-12 to 20030211-12,

About 4,000 data per bin 6-hourly, 2,000 3-hourly, 1,000 1-hourly

# American wind profilers, U-component (m/s), 300-200 hPa

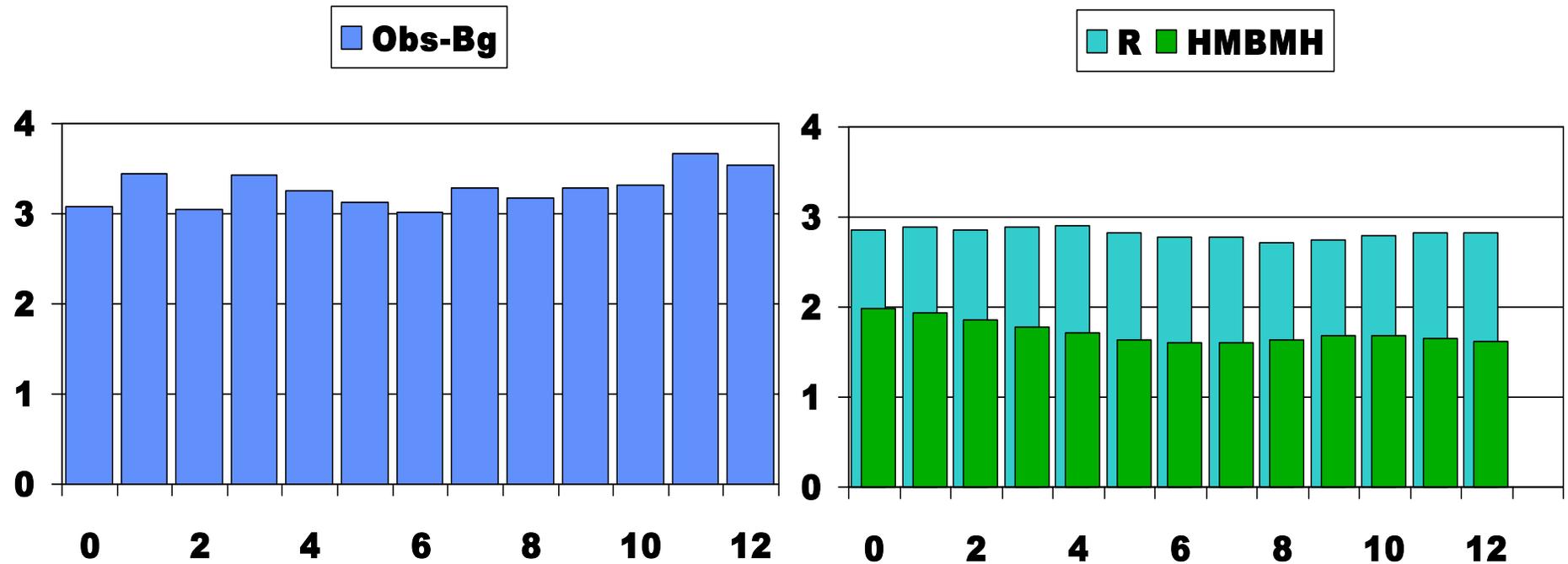


20030205-12 to 20030211-12,

About 12,000 data per bin

# Aircraft data, North Atlantic

## U-component (m/s), 300-200 hPa



20030205-12 to 20030211-12,

About 2,200 - 4,500 data per bin

# Used data (Sept 2003)

## Conventional

- ◆ **SYNOP**  
Surf.Press, Wind-10m, RH-2m
- ◆ **AIREP**  
Wind, Temperature
- ◆ **SATOB AMVs**  
Meteosat, GOES, MODIS
- ◆ **DRIBU**  
Surf.Press, Wind-10m
- ◆ **TEMP**  
Wind, Temp, Humidity profiles
- ◆ **DROPSONDE**  
Wind and Temp profiles
- ◆ **PILOT, Am+Eu+Jp Profilers**  
Wind profiles
- ◆ **PAOB**  
Surface pressure proxy

## Satellite

- ◆ **NOAA-15/16/17**  
HIRS, AMSU-A&B radiances
- ◆ **AQUA**  
AIRS and AMSU-A radiances
- ◆ **DMSP-13/14/15**  
SSMI radiances
- ◆ **Meteosat-5/7, GOES-9/10/12**  
Water Vapour radiances
- ◆ **QuikScat**  
Ambiguous winds
- ◆ **SBUV, (GOME), MIPAS**  
Layer ozone

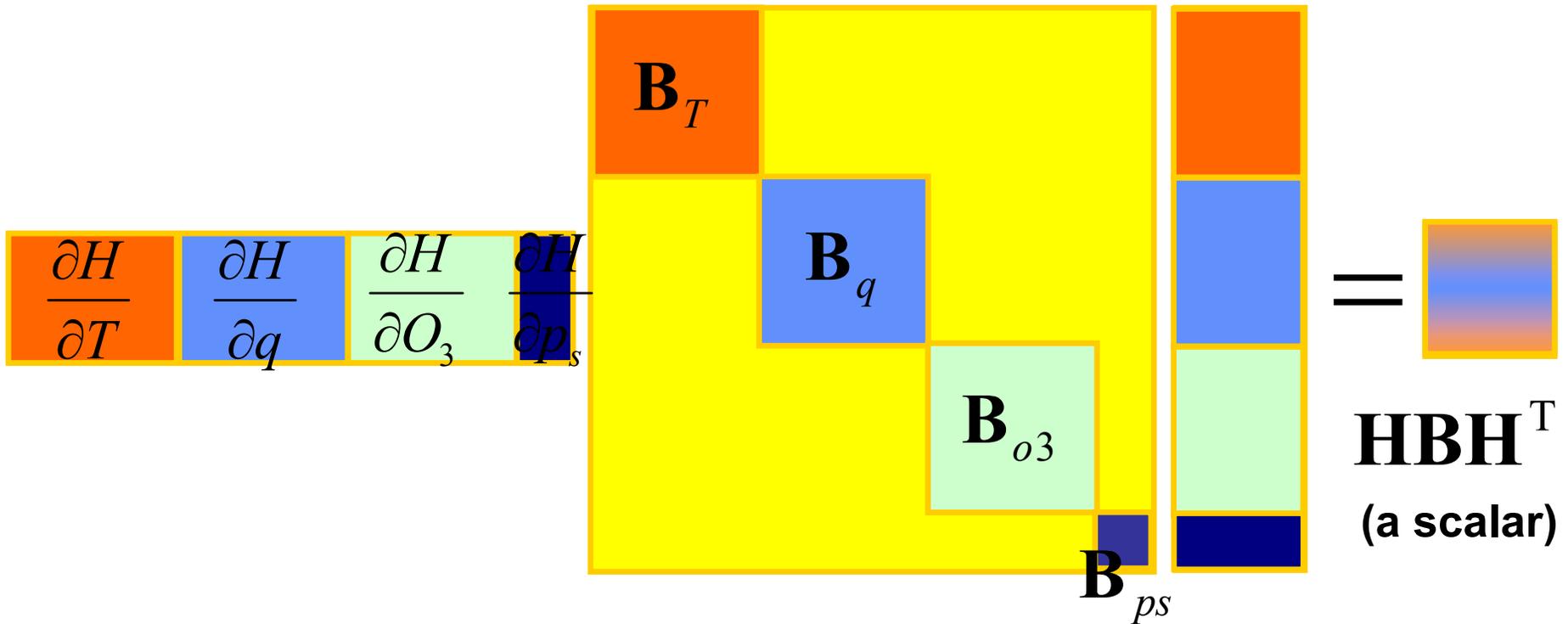
In preparation: MSG, SSMI/S, Cloud and precipitation data...

# One vertical column

**H**

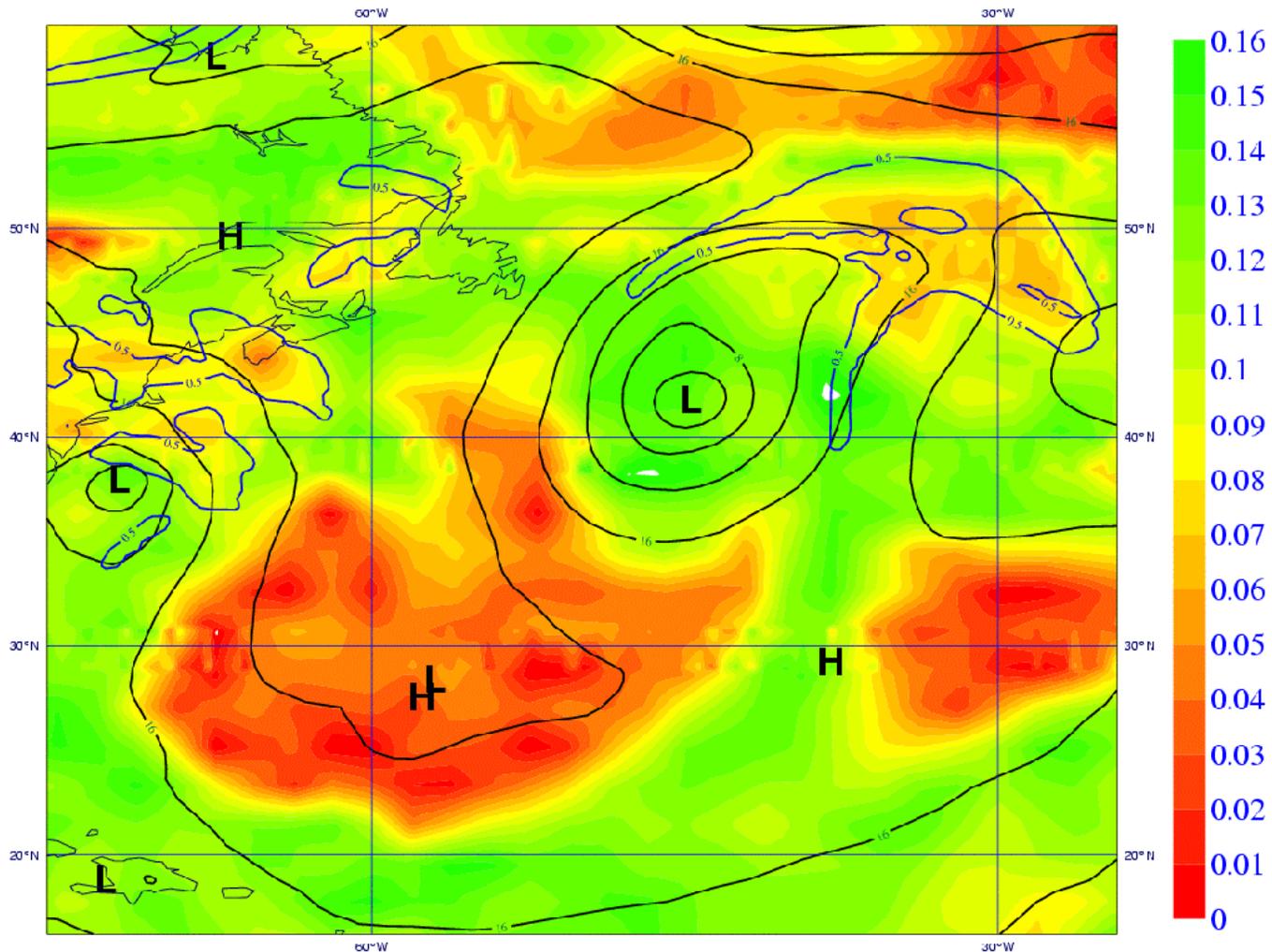
**B**

**H<sup>T</sup>**



# Relative Humidity $HBH^T$

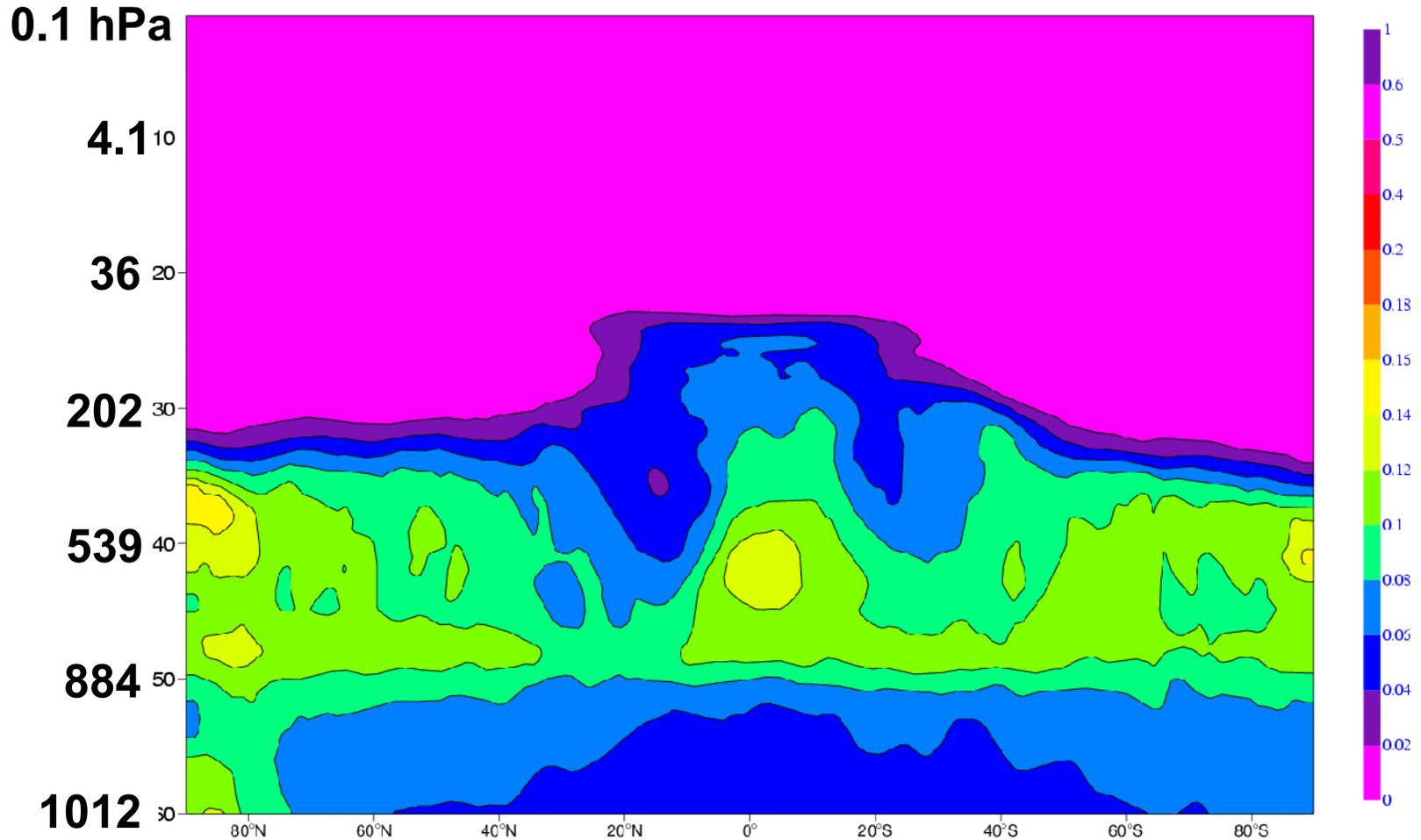
H is  $H(T,q,p)$



New humidity analysis formulation, E.Holm

# Zonal-mean Relative Humidity HBH<sup>T</sup>

H is  $H(T,q,ps)$



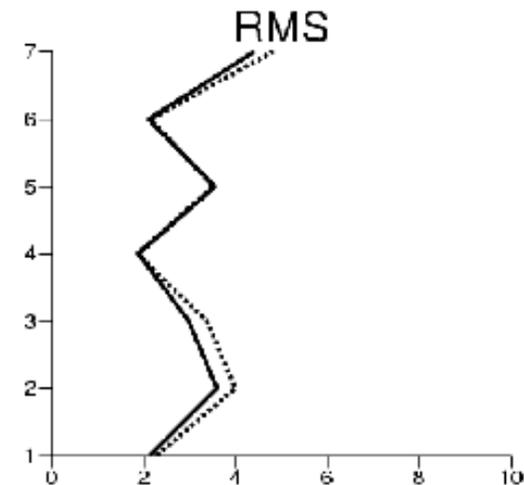
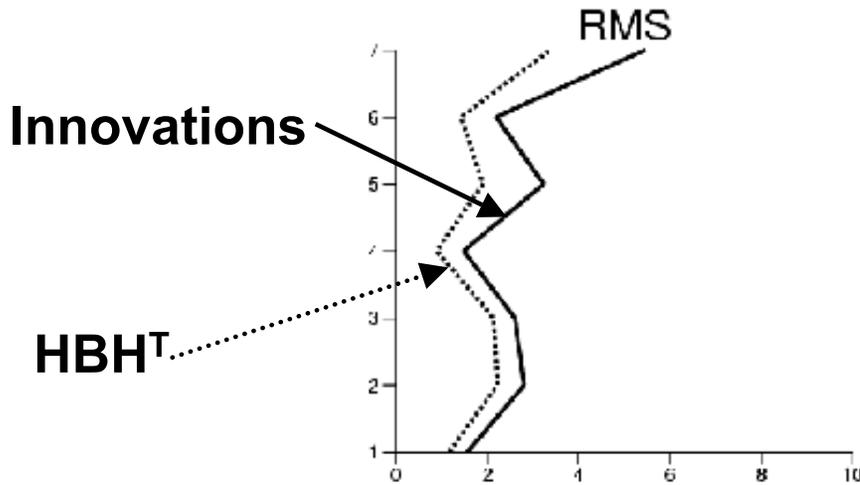
New humidity analysis formulation, E.Holm

# Comparison with innovation statistics

## Humidity-sensitive radiances

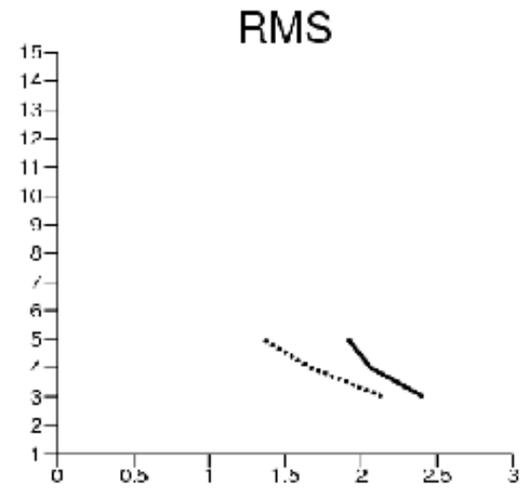
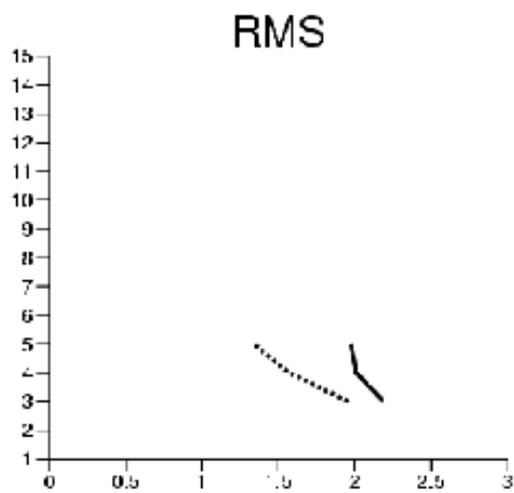
SSM/I-1C dmsp-13 SSMI Tb N.Hemis

SSM/I-1C dmsp-13 SSMI Tb Tropics



TOVS-1C noaa-17 AMSU-B Tb N.Hemis

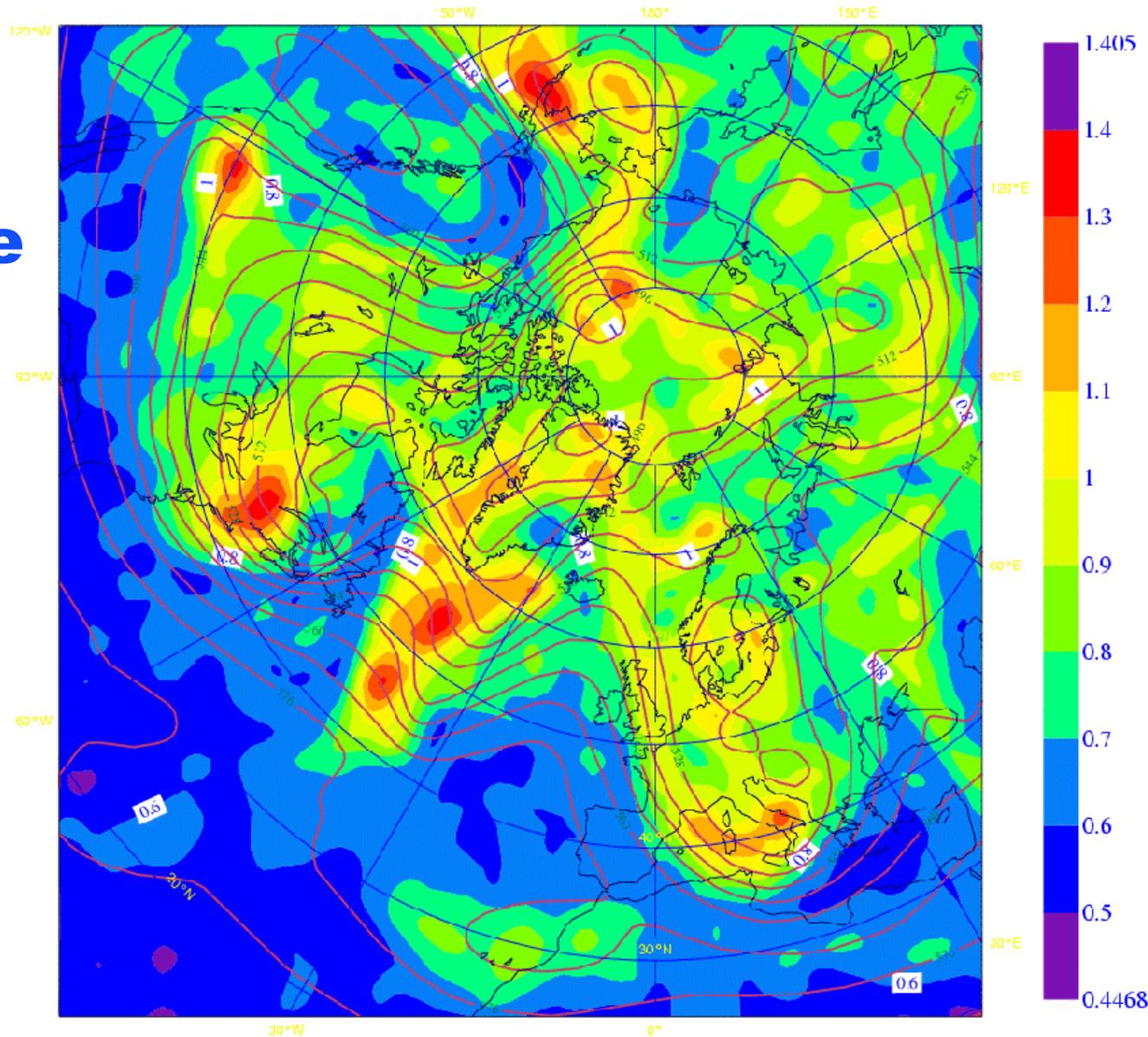
TOVS-1C noaa-17 AMSU-B Tb Tropics



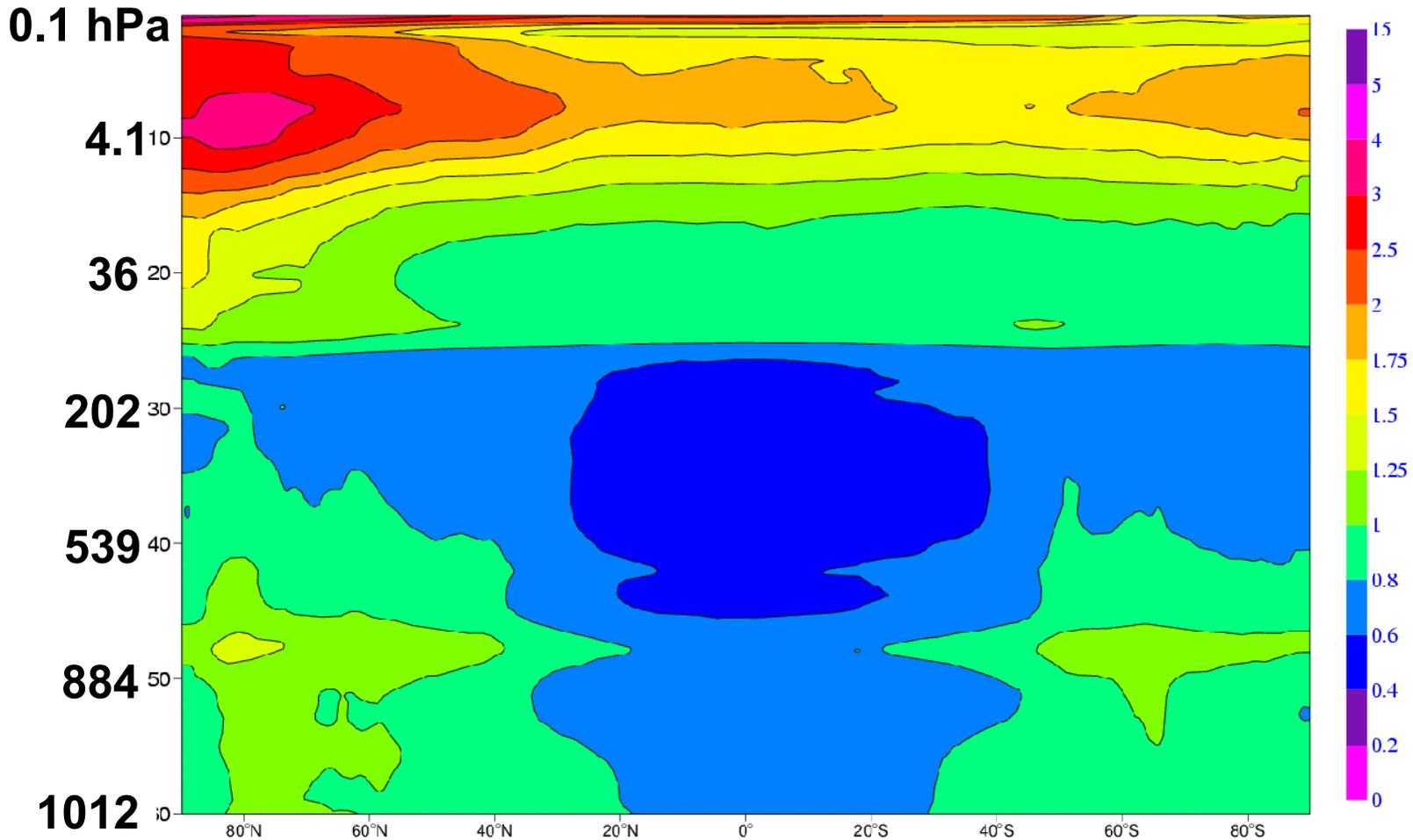
# Temperature HBH<sup>T</sup> (K)

T lev39 HBH<sup>T</sup>  
(shaded),

Z 500 hPa  
(contoured)



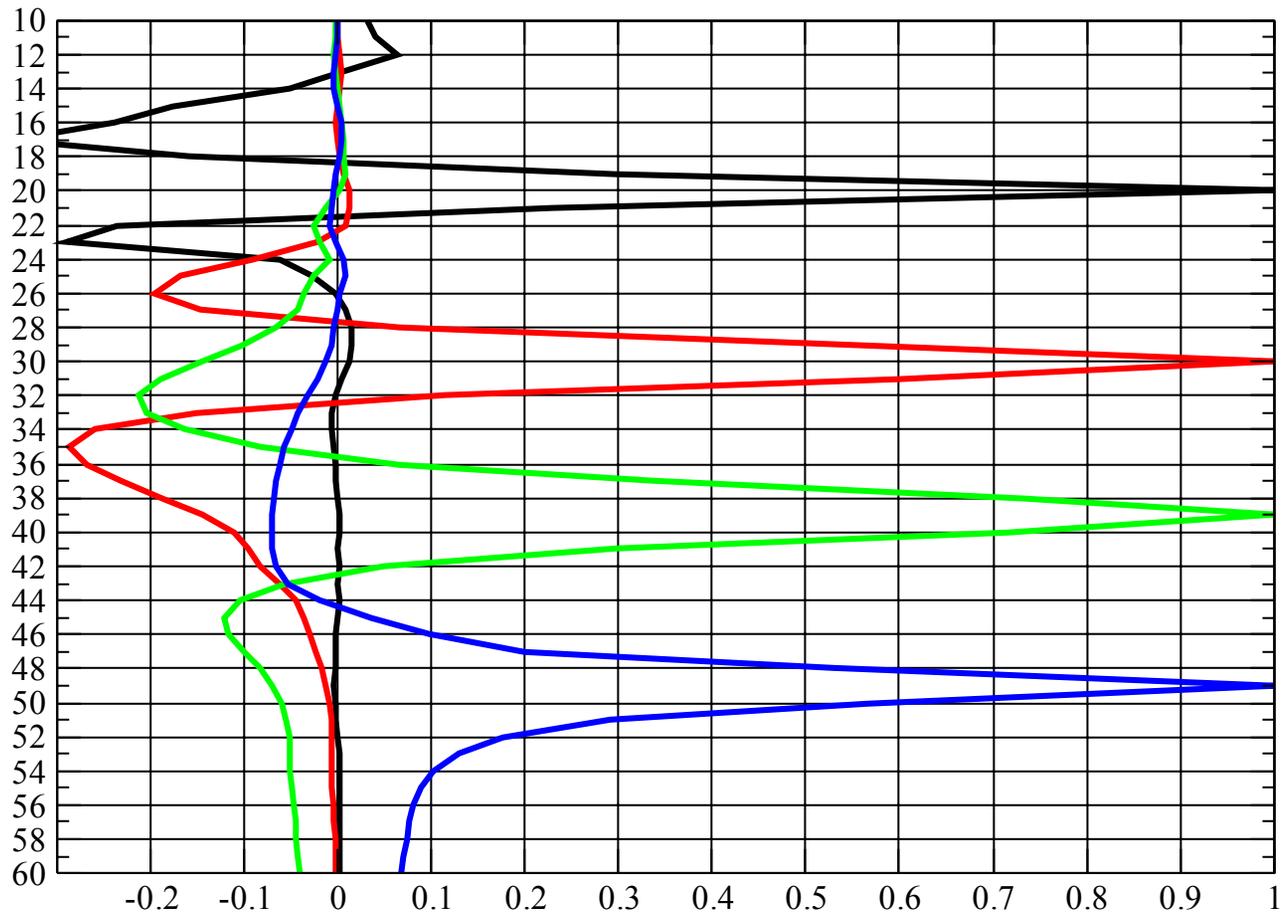
# Temperature HBH<sup>T</sup> (K) zonally averaged cross-section



# BG-error correlations (temperature) from an Ensemble of 4D-Var assimilations

## Vertical Structure Functions for (full) Temperature

4dVar ensemble statistics



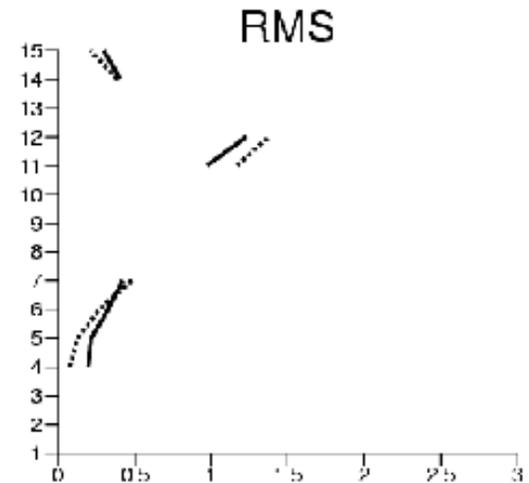
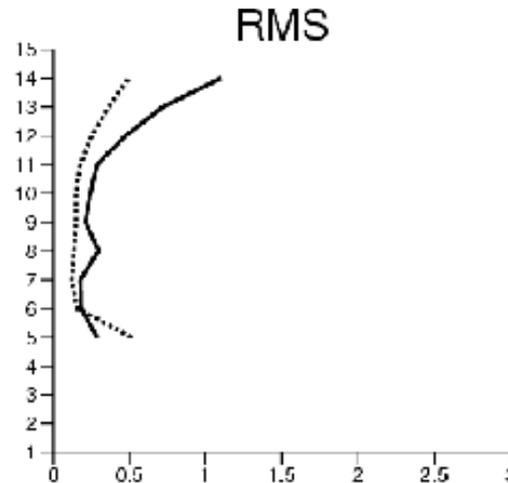
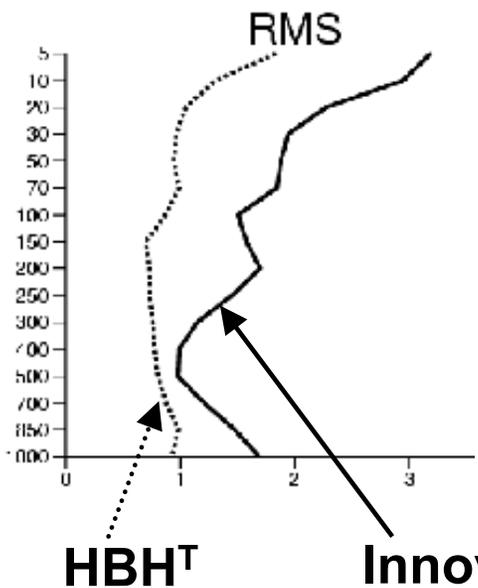
# Comparison with innovation statistics

## Temperature: radiosonde and radiances

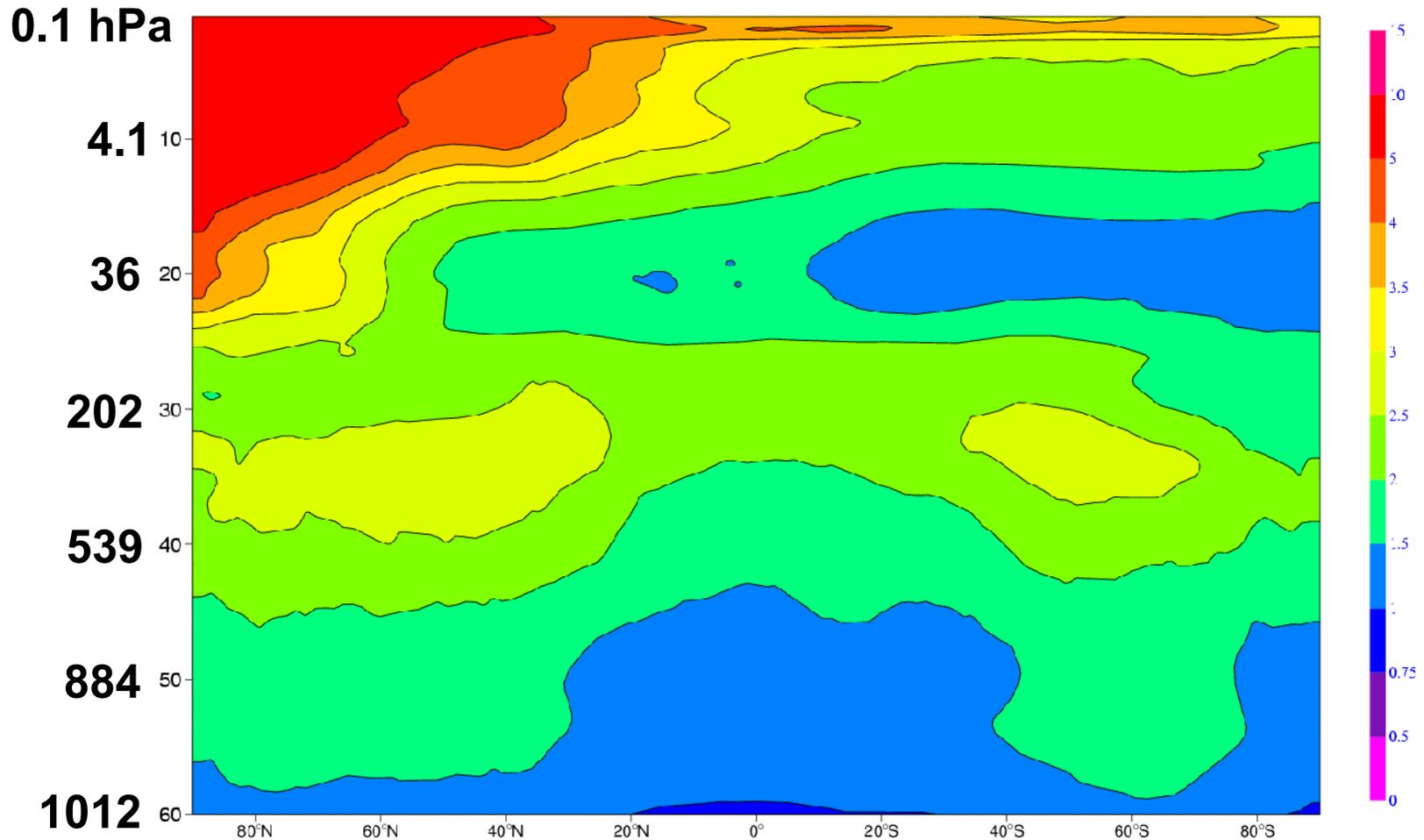
TEMP-T N.Hemis

TOVS-1C noaa-16 AMSU-A Tb N.Hemis

TOVS-1C noaa-16 HIRS Tb S.Hemis



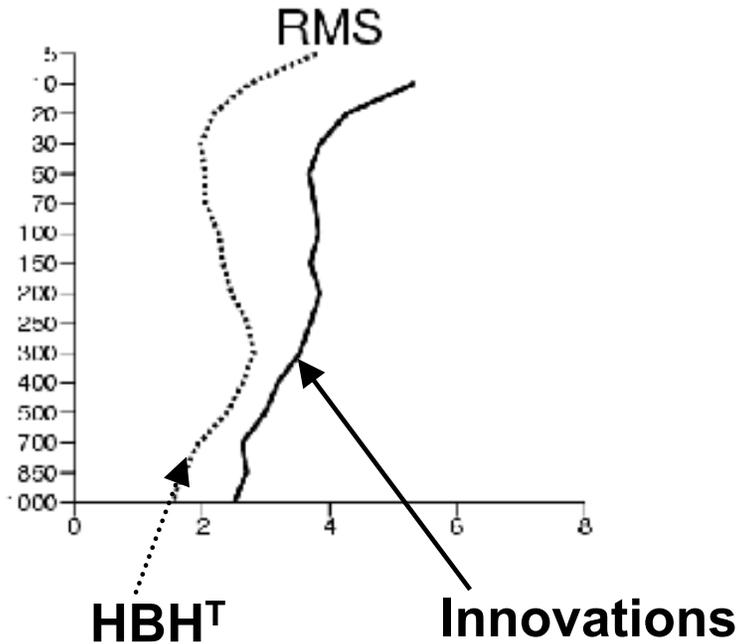
# Zonal-mean U-component HBH<sup>T</sup> (m/s)



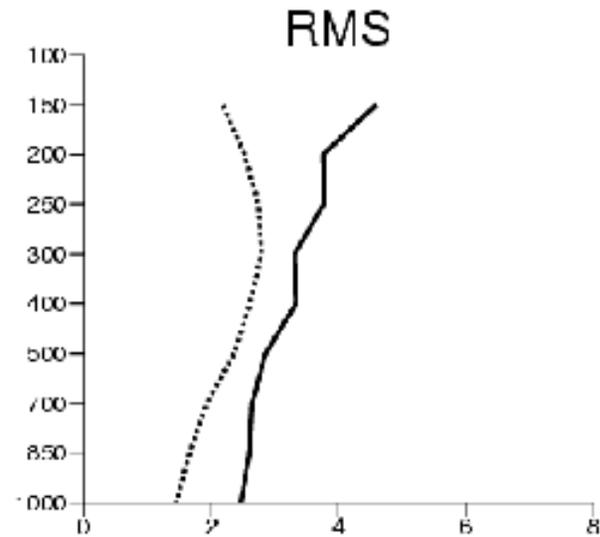
# Comparison with innovation statistics

## Wind: radiosonde and aircraft

TEMP-Uwind N.Hemis



AIREP-Uwind N.Hemis



# **Jb modelling developments**

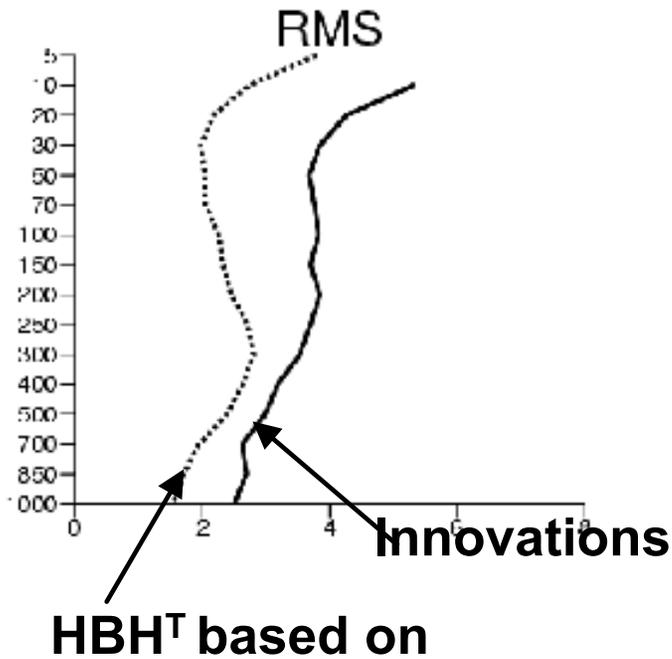
- ◆ **Base Jb statistics on an ensemble of 4D-Var assimilations, rather than the lagged forecast (NMC-) method which uses 48-24 hour forecast differences.**
  - ◆ Forecast errors become more large-scale (both vertical and horizontal) with time. Wind errors grow large particularly in upper troposphere.
  - ◆ Ensemble spread provides a more direct estimate of errors in short-range forecasts (M. Fisher).
  
- ◆ **Allow vertical correlations to vary with horizontal wave-number, and horizontal correlations to vary with vertical level (non-separability).**
  - ◆ This is a prevailing feature of short-range forecast error (Phillips 1986; Courtier et al., Andersson et al., Rabier et al. 1998)

# Jb modelling developments

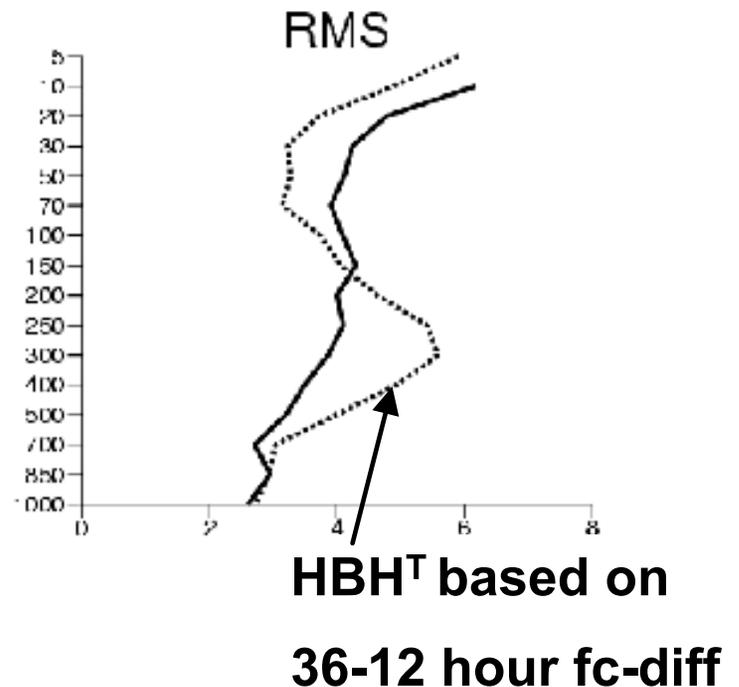
## DA-ensemble vs 'NMC-method'

	Ps (hPa) N.Hem
Innovations	0.92
HBH <sup>T</sup> DA-Ensemble	1.03
HBH <sup>T</sup> NMC-method	3.83

TEMP-Uwind N.Hemis  
used U



TEMP-Uwind N.Hemis  
used U

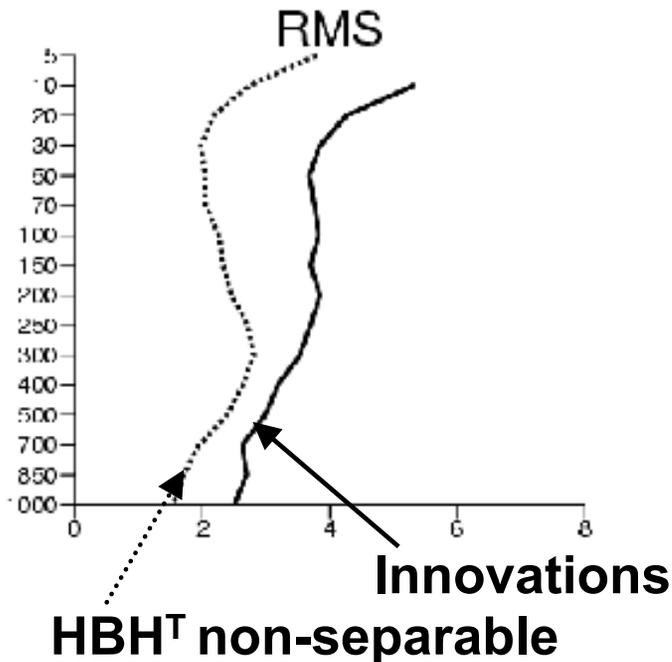


# Jb modelling developments

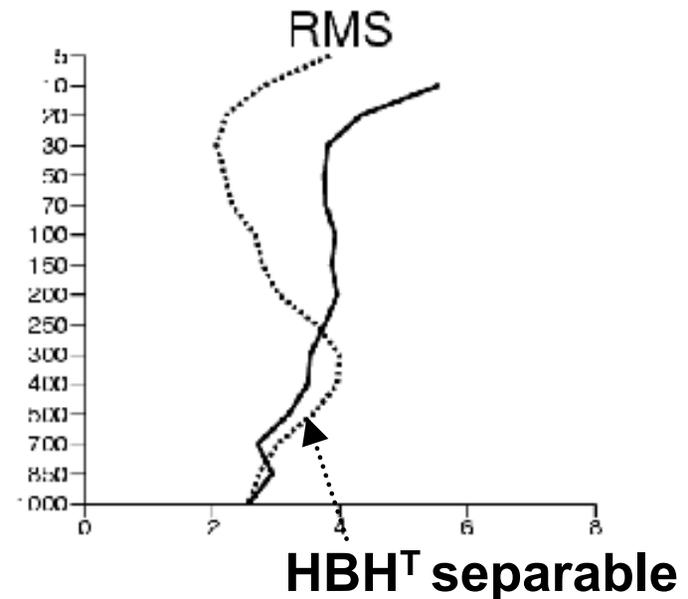
## Non-separable vs separable

	Ps (hPa) N.Hem
Innovations	0.92
HBH <sup>T</sup> Non-Sep	1.03
HBH <sup>T</sup> Sep	1.76

TEMP-Uwind N.Hemis  
used U



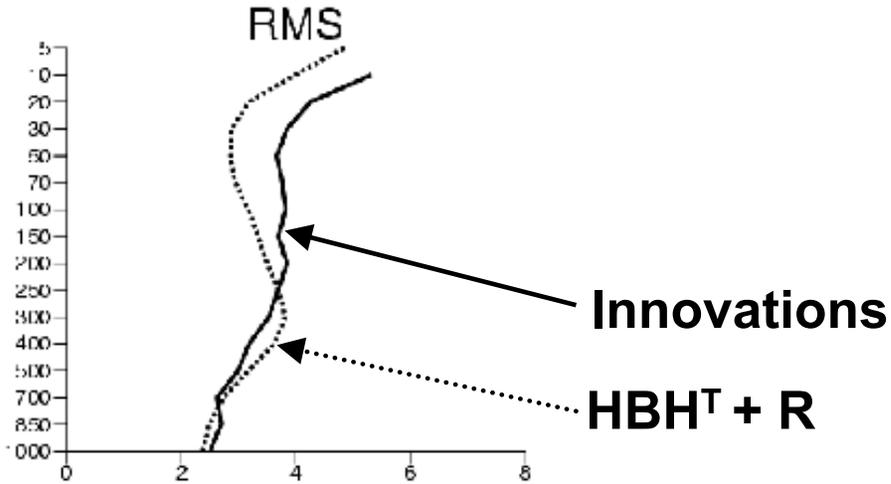
TEMP-Uwind N.Hemis  
used U



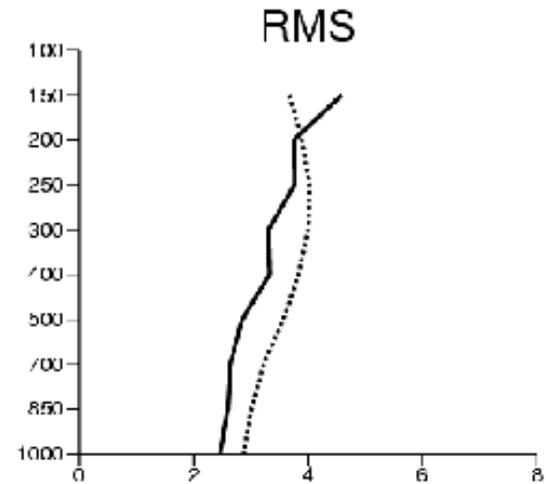
# HBH+R

## for a range of different data type

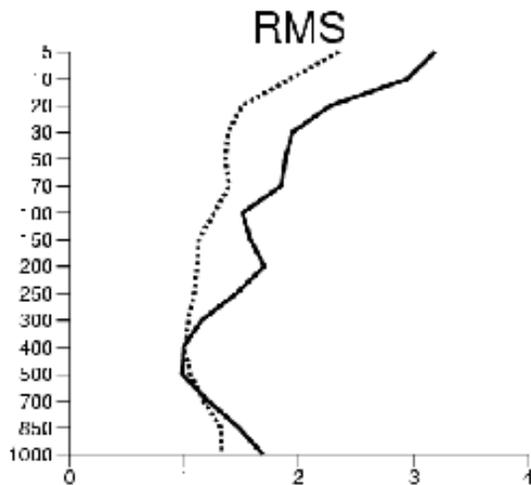
TEMP-Uwind N.Hemis  
used U



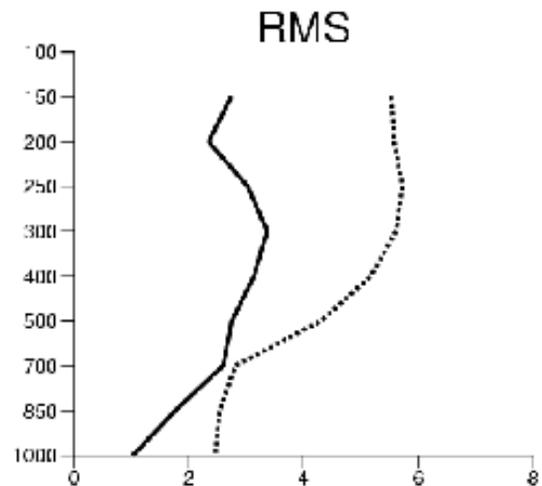
AIREP-Uwind N.Hemis  
used U



TEMP-T N.Hemis  
used T



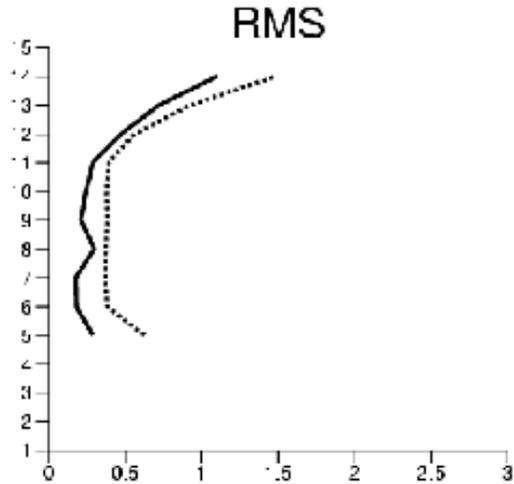
SATOB-Uwind N.Midlat  
used U



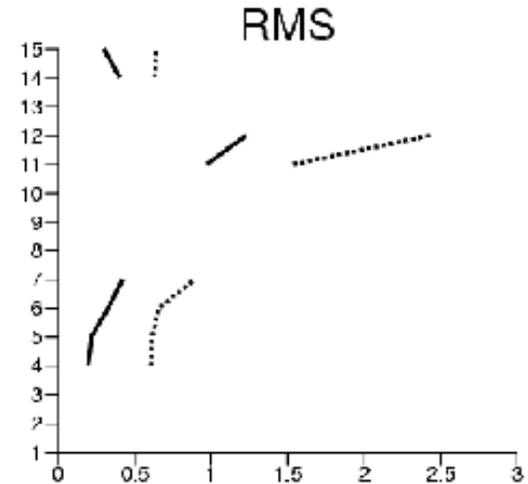
# HBH+R

## for a range of different data type

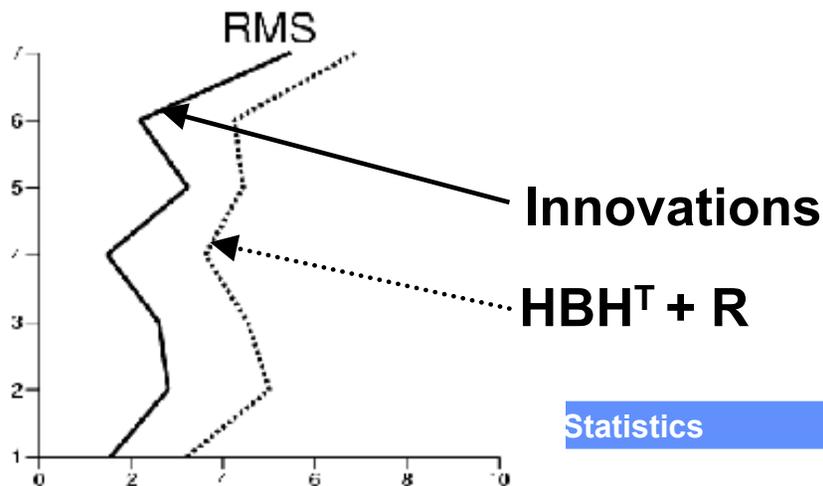
TOVS-1C noaa-16 AMSU-A Tb N.Hemis  
used Tb noaa-16 amsu-a



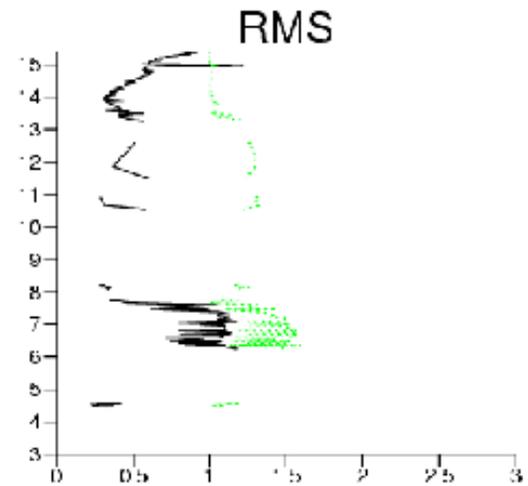
TOVS-1C noaa-16 HIRS Tb S.Hemis  
used Tb noaa-16 hirs



SSM/I-1C dmsp-13 SSM/I Tb N.Hemis  
used Tb dmsp-13 rad



AIRS NASA-1 Tb S.Hemis  
used Tb nasa-1 AIRS



# Observation error correlation

Has most recently been studied by Bormann et al. (*MWR* 2003) and Liu and Rabier (*QJ* 2002).

SATOB (or AMV=Atm. Motion Vectors) and radiosonde co-locations were studied, and error correlation functions were fitted:

$$R(r) = R_0 \left( 1 + \frac{r}{L} \right) e^{-r/L}$$

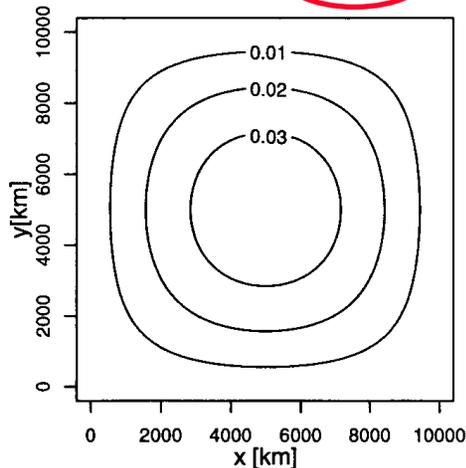
With  $R_0$  the intercept and  $L$  the length-scale.

Found  $L=190$  km.

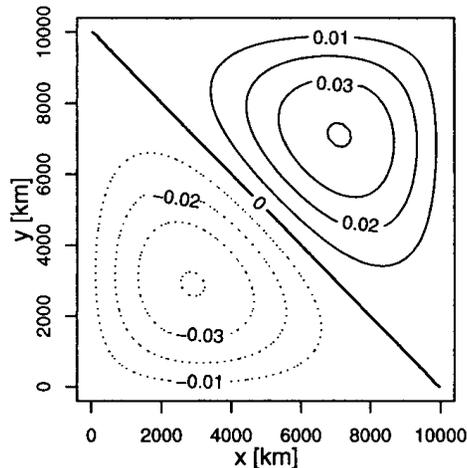
Eigenvectors of  $R$  were computed for an idealized observational data set with regular 200 km spacing.

# Eigenvectors of R for idealized observational dataset

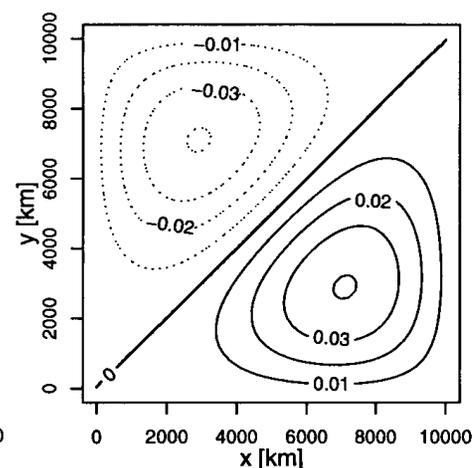
Eigenvector 1,  $\sqrt{e\lambda}=4.094$



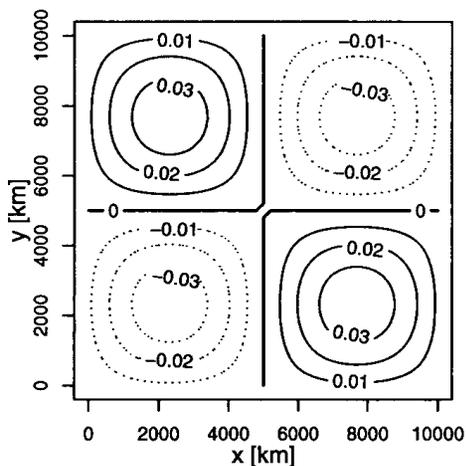
Eigenvector 2,  $\sqrt{e\lambda}=4.048$



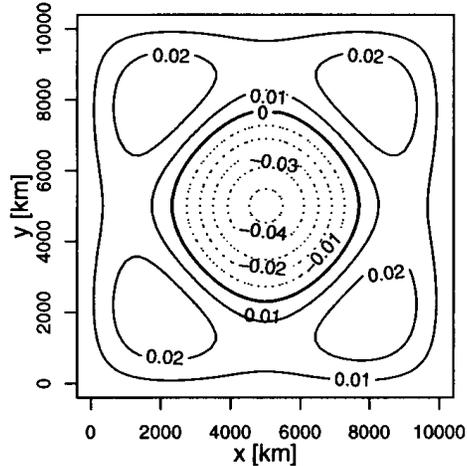
Eigenvector 3,  $\sqrt{e\lambda}=4.048$



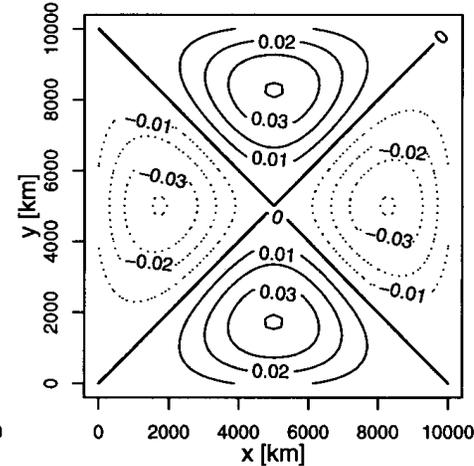
Eigenvector 4,  $\sqrt{e\lambda}=4.002$



Eigenvector 5,  $\sqrt{e\lambda}=3.972$



Eigenvector 6,  $\sqrt{e\lambda}=3.972$



**$L=190$  km.  
200 km  
spacing.**

# Summary and Conclusions (1)

- ◆ A method to diagnose the modelled evolution of background error within 4D-Var has been developed.
  - ◆ The modelling of innovations within 4D-Var has been studied, and compared to actual innovation statistics.
  - ◆ Discrepancies can be due to deficiencies in the specification of **B, R, H, M or Q.**
- 
- The evolution of  $MBM^T$  is not as expected.
  - There seems to be insufficient projection of **B** onto growing modes – I.e. there is insufficient flow-dependence in **B.**
  - Comparison with EnKF could be performed by replacing the vectors  $\delta\mathbf{x} = \mathbf{L}\chi$  with vectors obtained from ensemble differences. Also for DA-ensemble.

## Summary and Conclusions (2)

- It has not been possible to identify the model-error  $Q$  contribution to the innovations, at this stage.
- Observation errors are specified far too large for many satellite data types.
- Taking account of observation error correlations within 4D-Var, would now seem important.
- DA-ensemble has provided noticeable improvement over the lagged-forecast (NMC) method.
- The separability assumption (in B-modelling) is not appropriate for joint analysis of stratosphere + troposphere.

## Summary and Conclusions (3)

- There is insufficient regional variation in B. Tropopause height and Boundary Layer height variations are poorly represented. **Wavelet-Jb.**
- Current T, RH (and Z) BgErrors show marked flow dependence. This should be validated against innovations.
- **The rapid baroclinic error growth within the 12-hour assimilation window is currently underrepresented, probably due to the (relatively) static nature of B.**